

# ***A History of Coupled Climate Modeling***

**Peter R. Gent**

**Senior Scientist**

**National Center for Atmospheric Research**

## Climate Calculations with a Combined Ocean-Atmosphere Model

SYUKURO MANABE AND KIRK BRYAN

*Geophysical Fluid Dynamics Laboratory, ESSA, Princeton University, Princeton, N. J.*

13 March 1969 and 6 May 1969

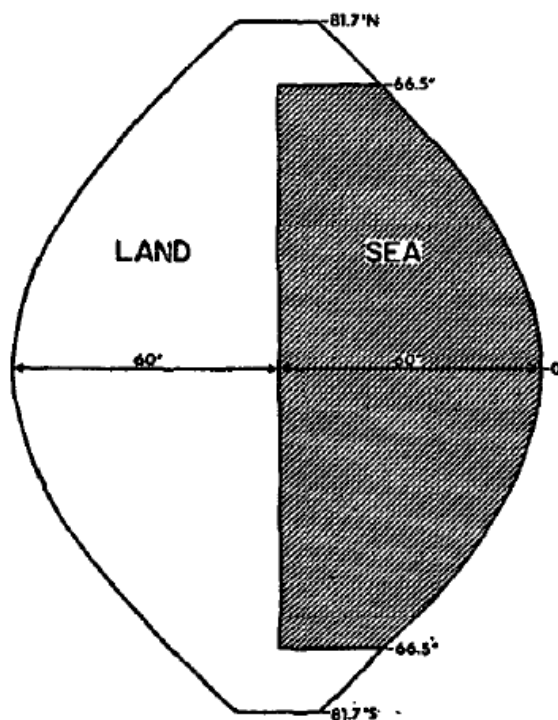


FIG. 1. Ocean-continent configuration of the model.

The first coupled ocean/atmosphere model with idealized continents was run at the GDFL: results published in JAS in 1969.

5° x 5° Horizontal grid:  
9 levels in atm, 5 in ocean.

# A Global Ocean-Atmosphere Climate Model. Part I. The Atmospheric Circulation

SYUKURO MANABE, KIRK BRYAN AND MICHAEL J. SPELMAN

*Geophysical Fluid Dynamics Laboratory/NOAA, Princeton University, Princeton, N. J. 08540*

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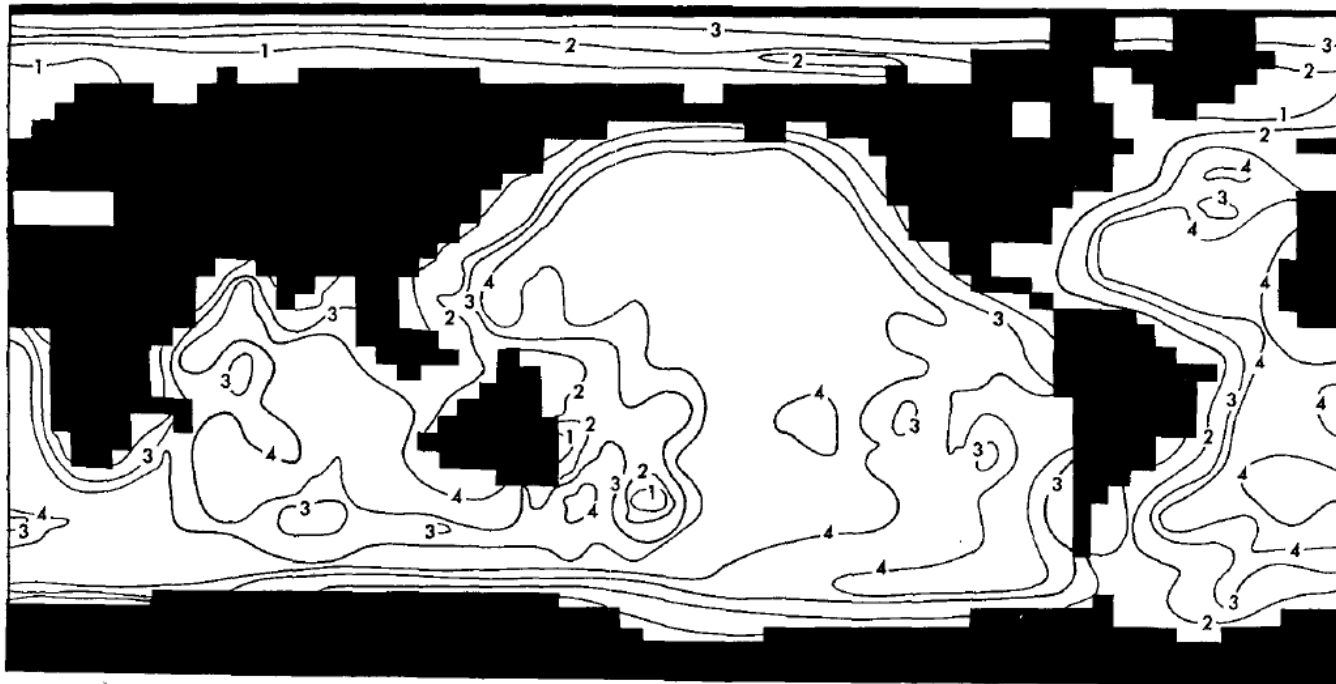


FIG. 2. The smoothed topography of the ocean model. Depth is given in kilometers.

5° x 5°  
Hor grid

9 levels  
in Atm, 5  
in Ocean

JPO  
1975

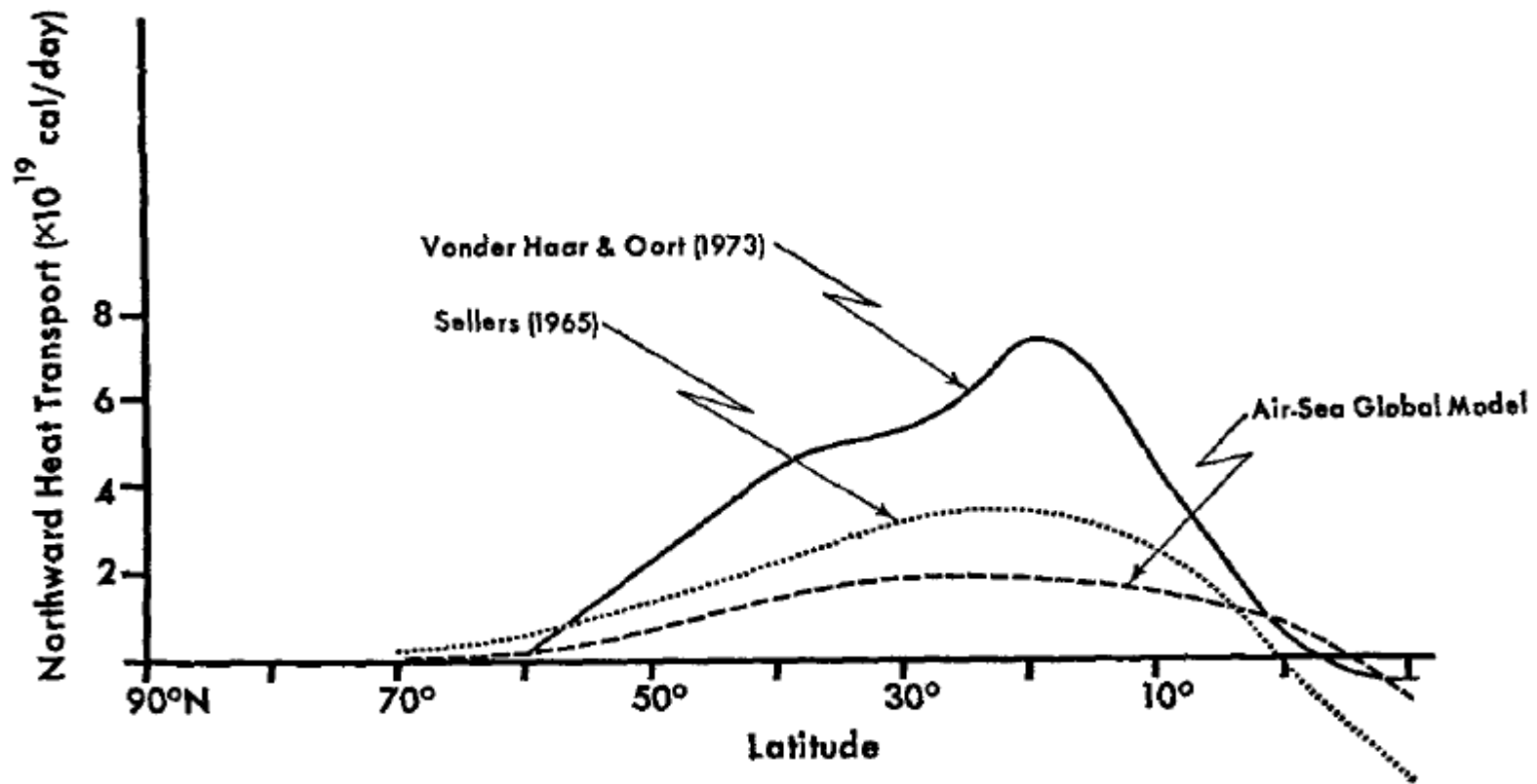


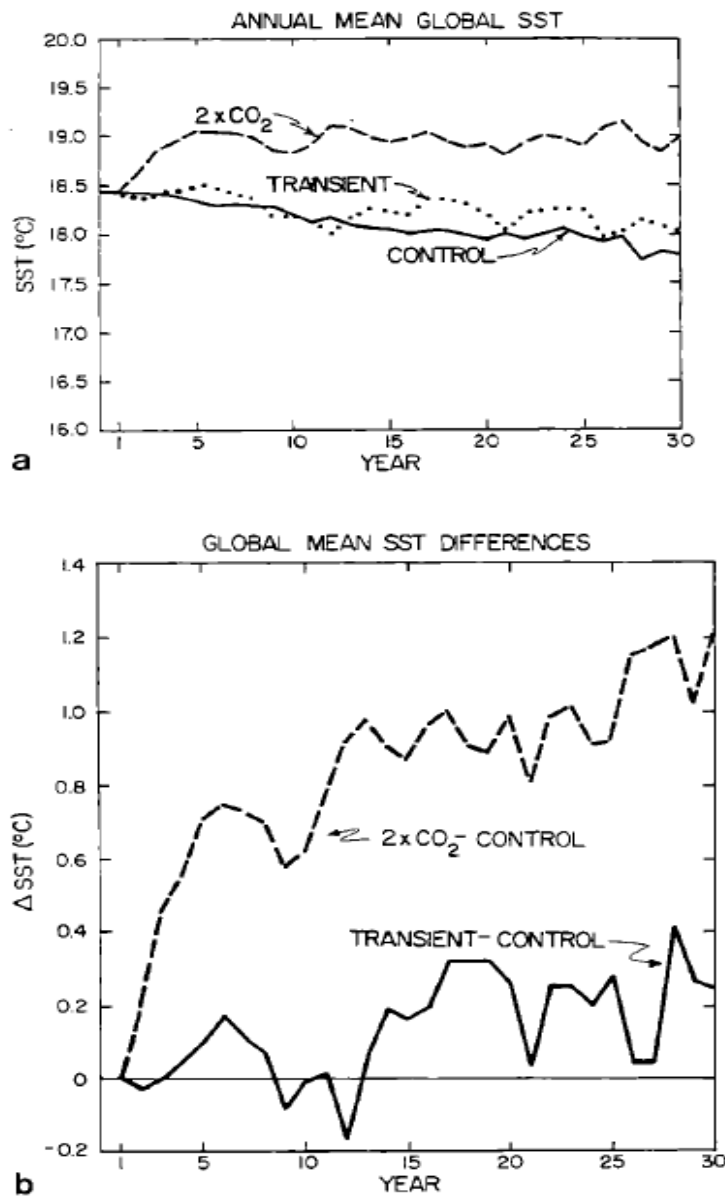
FIG. 12. Heat transport by the ocean model in a northward direction compared to estimates based on heat balance calculations.

The conservation equations for potential temperature ( $\theta$ ) and salinity ( $S$ ) under conditions of stable stratification are

$$(\theta, S)_t + \mathbf{v} \nabla (\theta, S) + w(\theta, S)_z = \kappa(\theta, S)_{zz} + A_H \nabla^2 (\theta, S). \quad (8)$$

JPO  
1975





**Fig. 4. a** Time evolution of globally averaged ocean surface temperature (°C) for 1×CO<sub>2</sub>, 2×CO<sub>2</sub>, and transient CO<sub>2</sub> experiments; **b** time evolution of ocean surface temperature differences from part (a) (°C)

Washington and Meehl (1989).

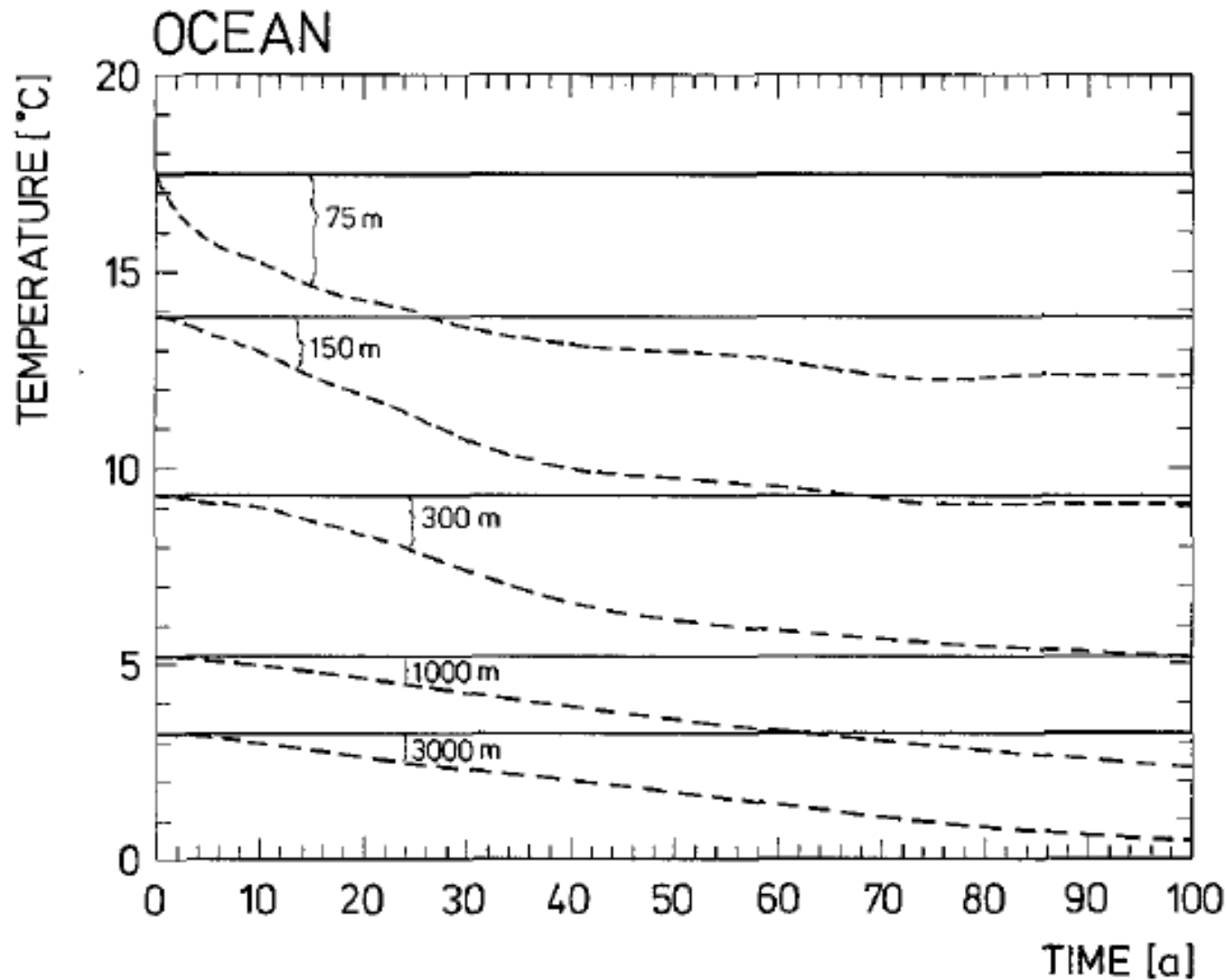
Atm 4.5°x7.5°x9  
Ocn 5°x5°x4 levels

In the control run, the global-average SST is reducing at a rate of 0.02°C/year.

If maintained, this means a 2°C SST fall in 100 year run.

# Flux Corrections or Adjustments

- Atmosphere component run with SST obs.
- Ocean component run with observed wind stress and atmosphere surface variables.
- Fluxes of heat and fresh water in these runs subtracted to form the flux corrections.
- Or coupled run with very strong relaxation back to observed SST and surface salinity.
- These corrections are totally unphysical.



**Fig. 14.** Temporal evolution of the global mean temperatures of the ocean in the coupled model, without flux correction (*dashed line*) and with flux correction (*solid line*)

Sausen et al. (1988).

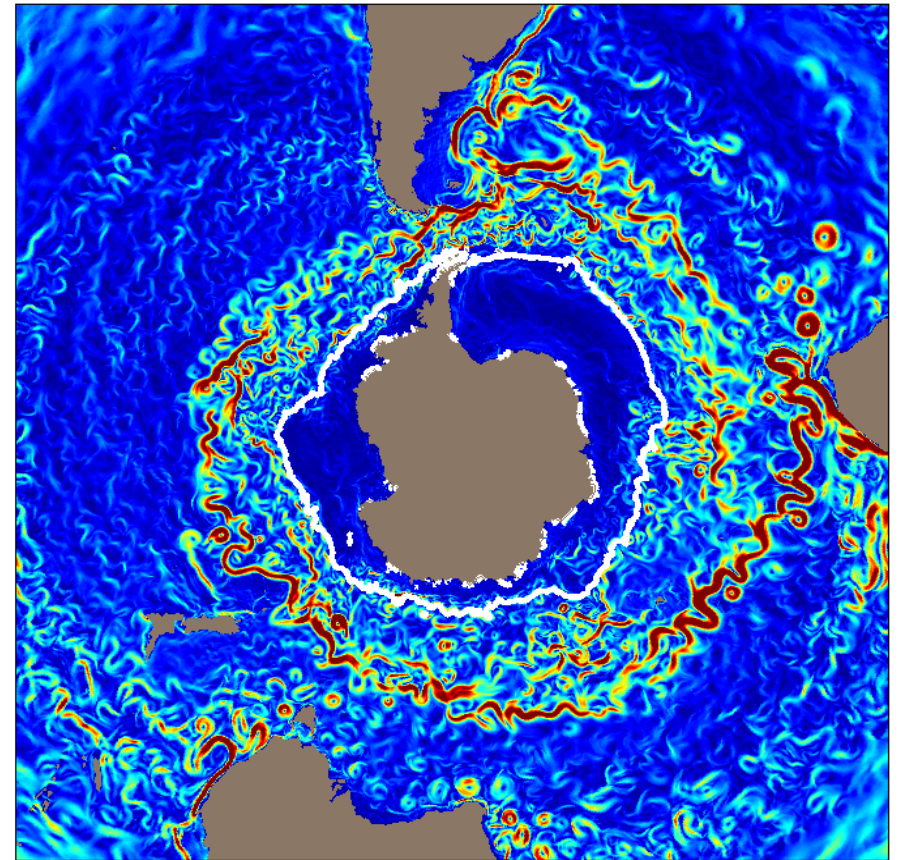
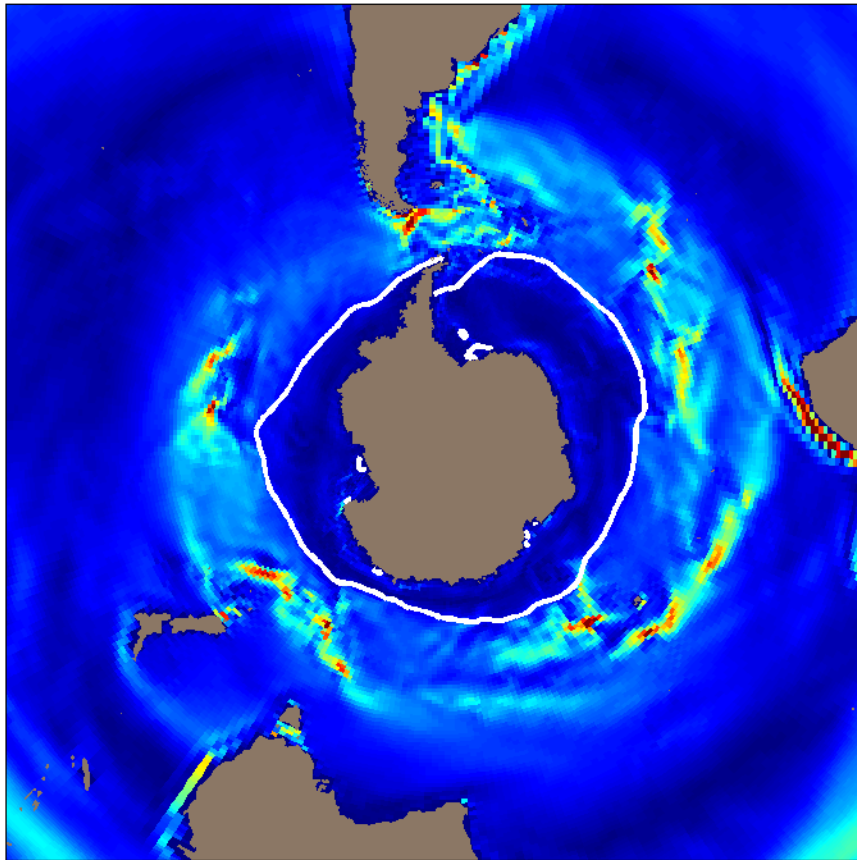
Ocean  
3.5°x3.5°  
x 5 levels

Diagnostic  
atmosphere

# *Simulated Surface Currents Around Antarctica*

1 degree Simulation

0.1 degree Simulation



0 10 20 30 40 50 60

Surface current speed in cm/s for October

## *The GM (1990) Parameterization*

$$\frac{\partial T}{\partial t} + (\underline{u} + \underline{u}^*) \cdot \underline{\nabla} T = \underline{\nabla}_\rho \cdot (\kappa \underline{\nabla}_\rho T)$$

$$w^* = -\underline{\nabla} \cdot (\kappa \underline{\nabla} \rho / \rho_z), \quad \underline{\nabla} \cdot \underline{u}^* = 0.$$

Assumes that eddies advect temperature and salinity and diffuse them along constant density surfaces. The advection represents the effects of unresolved baroclinic instability, because it ensures a global sink of mean potential energy.



$(u, v, w)$  is velocity in directions  $(x, y, z)$

continuity equation  $u_x + v_y + w_z = 0$ .



Integrate over depth  $\int u_x dz + \int v_y dz = 0$

form 2-D **barotropic streamfunction in**

**Sverdrups** ( $1 = 10^6 m^3/s$ ).



Integrate over  $x$   $\int v_y dx + \int w_z dx = 0$

form 2-D **overturning streamfunction.**



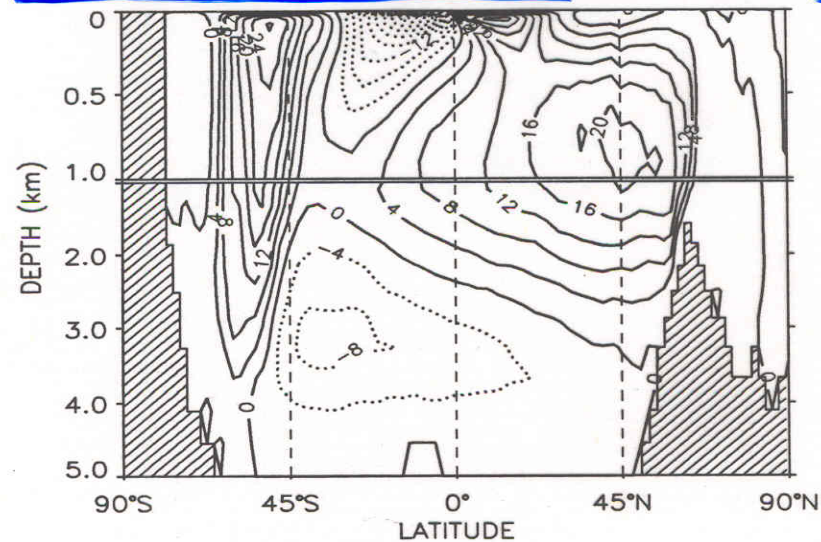
Original GM  
implementation  
in a very low  
resolution  
 $4^\circ \times 3^\circ \times 20$   
levels global  
ocean model;

Danabasoglu  
et al. (1993).

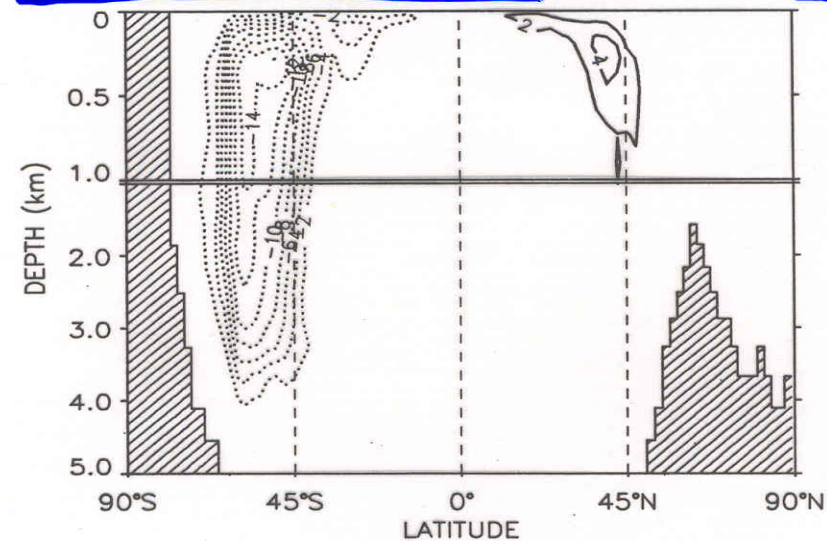
GLOBAL MERIDIONAL OVERTURNING STREAMFUNCTION (Sv)

(3x3) **CSM MODEL**

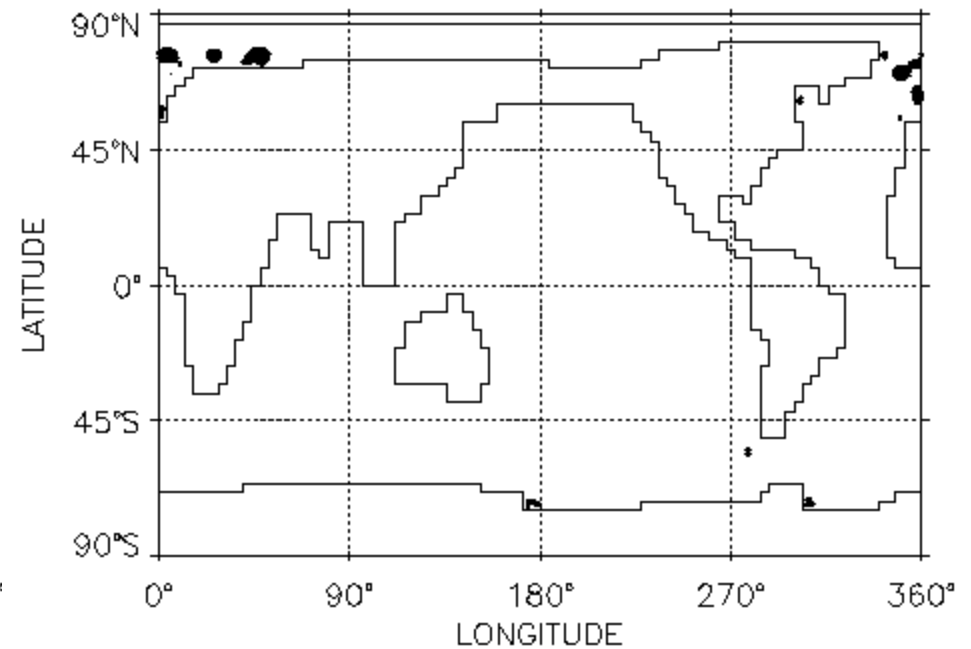
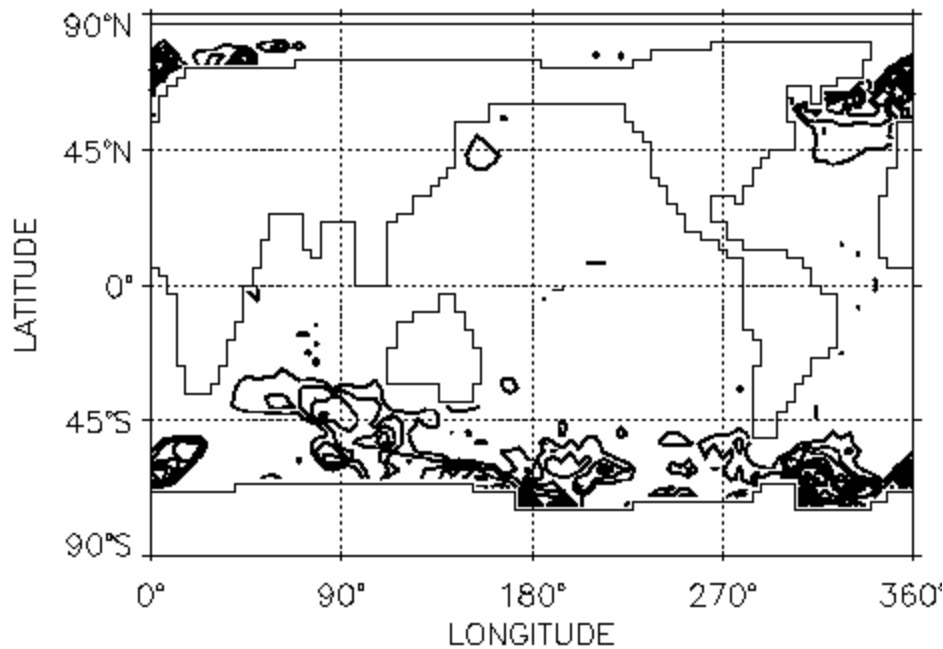
EULERIAN-MEAN VELOCITY (cont. int. = 4 Sv)



EDDY-INDUCED VELOCITY (cont. int. = 2 Sv)



# *Deep Water Formation 4° × 3°*



Horizontal Mixing

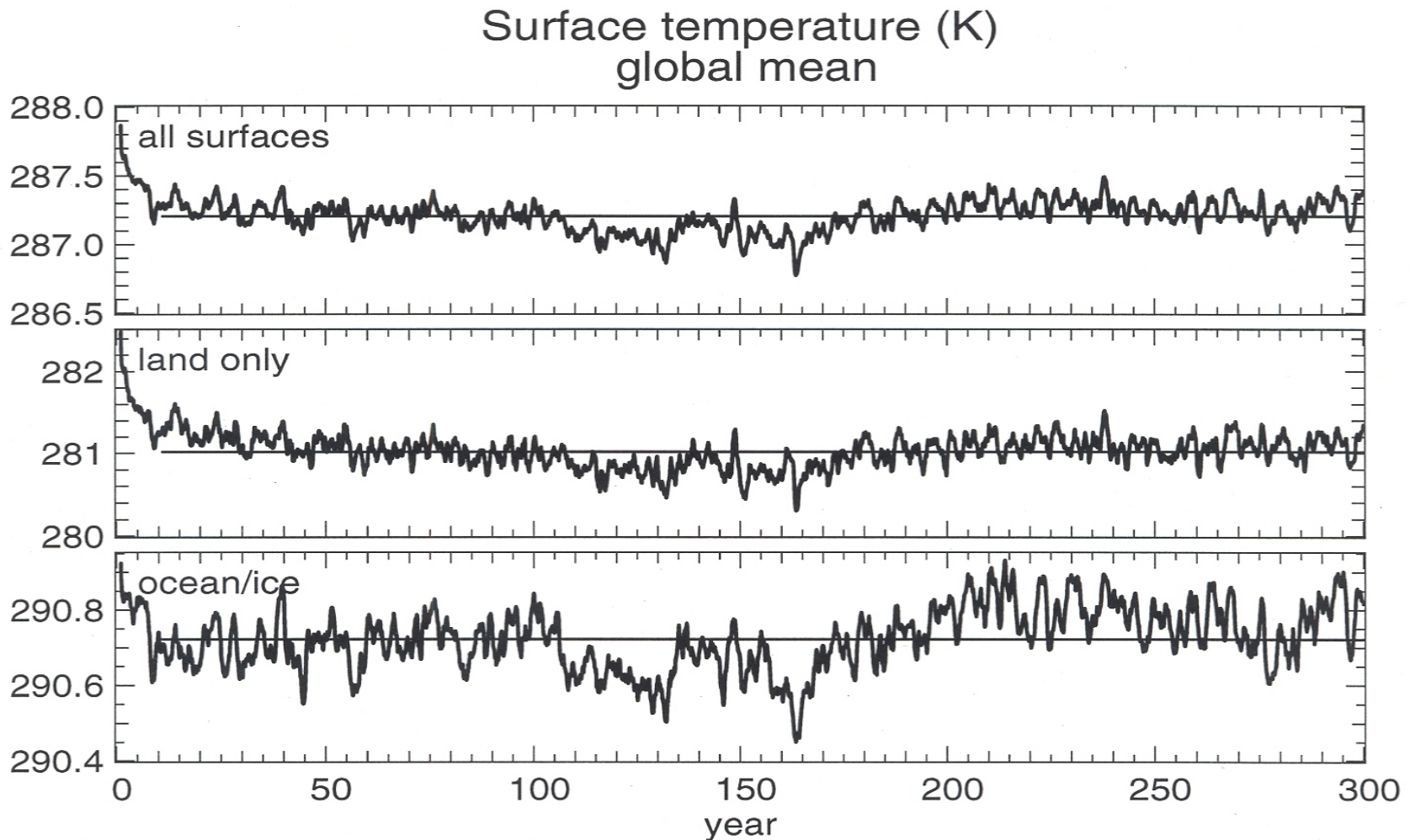
GM 1990

Danabasoglu et al. (1993).



# CSM 1 was the first climate model to produce a non-drifting control run without “flux corrections”

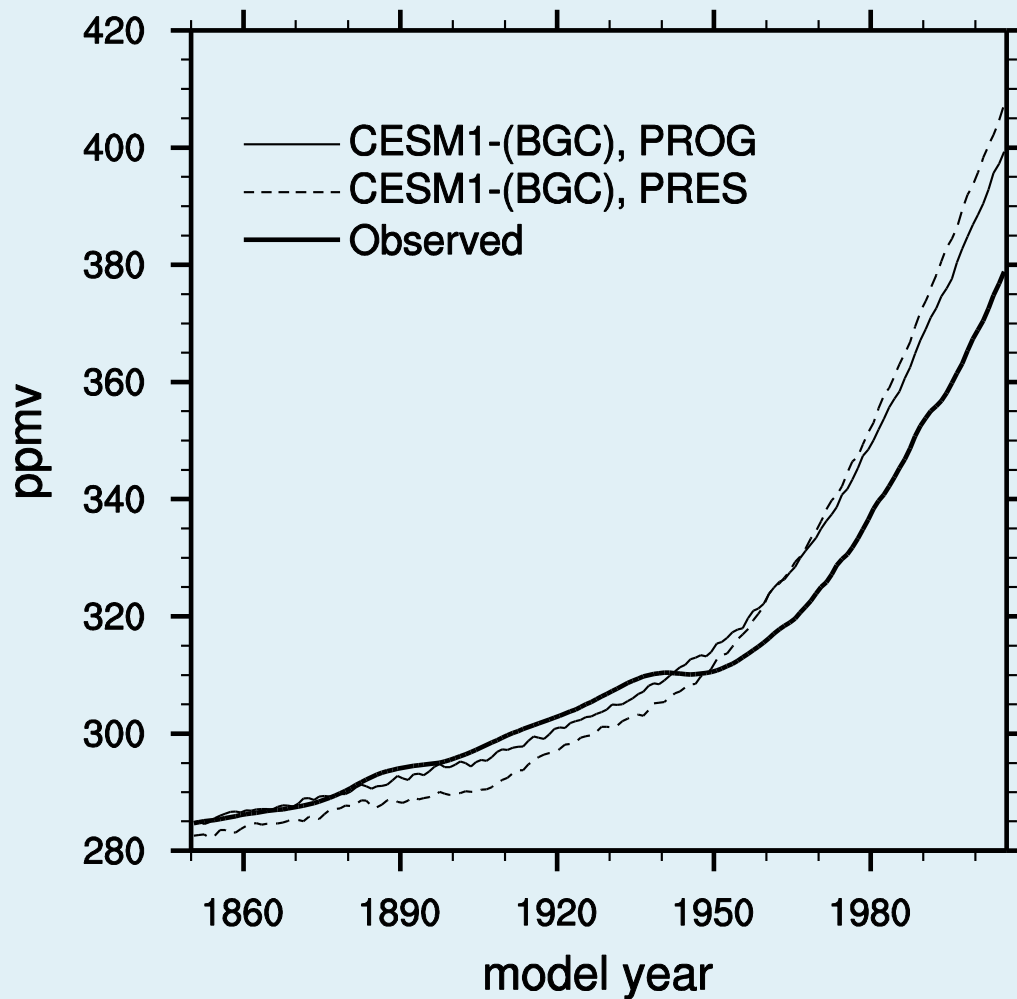
FIG 8



## ***Components added to CCSM to make it into an Earth System Model (CESM)***

- **Carbon cycle components in the land, ocean & atm.**
- **A component that uses predictive chemistry.**
- **A whole atmosphere version of the atm (WACCM) that reaches up into the stratosphere and beyond.**
- **A land ice component that models the Greenland ice sheet, and will model the Antarctic ice sheet.**

# CO<sub>2</sub> in 20<sup>th</sup> Century Experiments



Modeled increase of CO<sub>2</sub> over 1850-2005 too large:

Observations: 94 ppmv

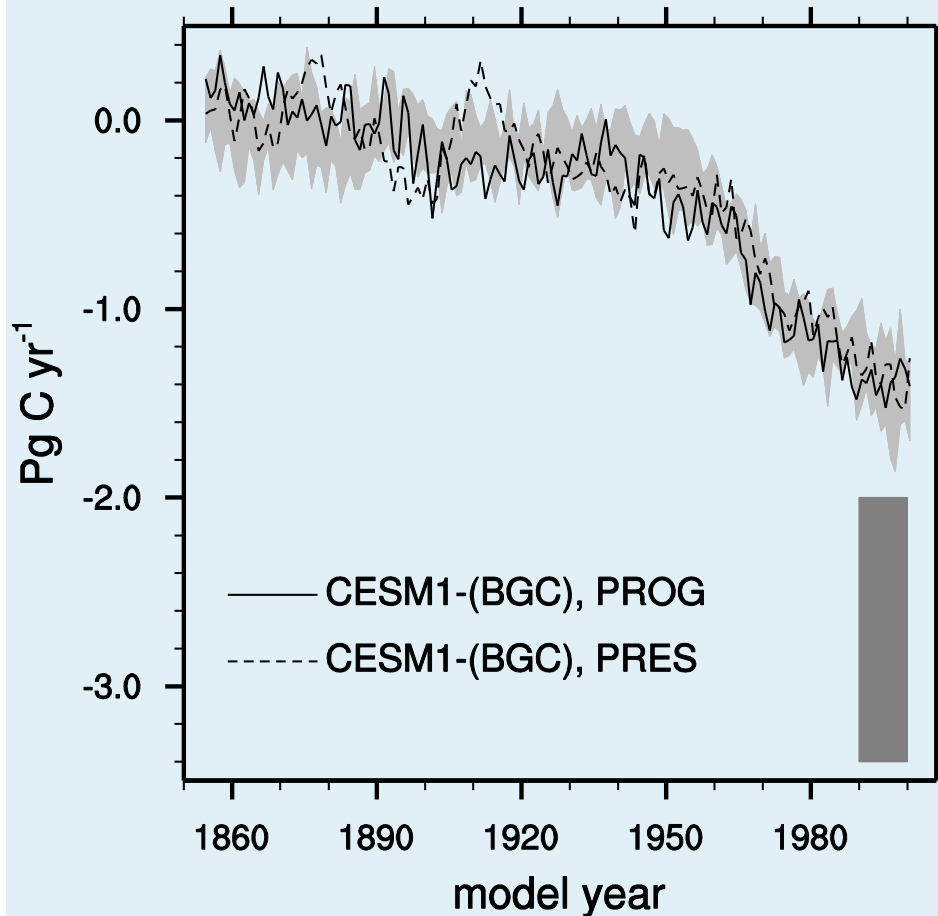
Prognostic CO<sub>2</sub>: 114 ppmv

Diagnostic CO<sub>2</sub>: 125 ppmv

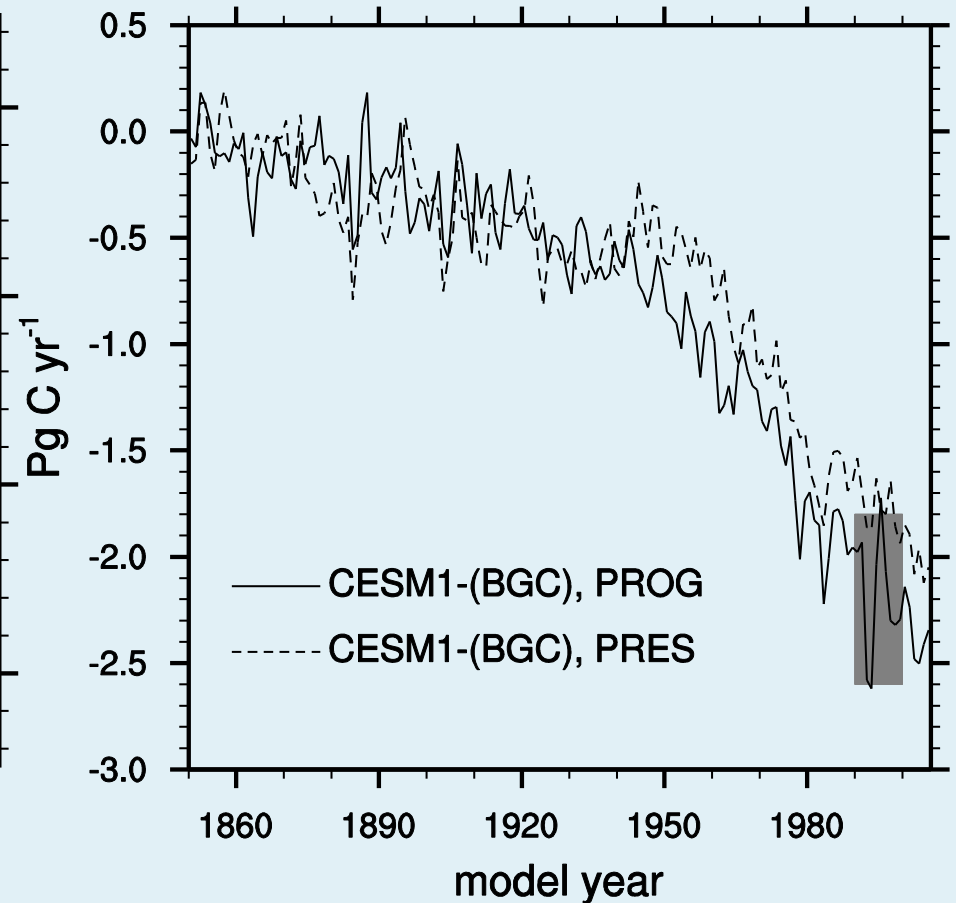
Lindsay et al. (2013)

# 20<sup>th</sup> Century CO<sub>2</sub> Sinks from Atmosphere

## Land Uptake



## Ocean Uptake



Lindsay et al. (2013)



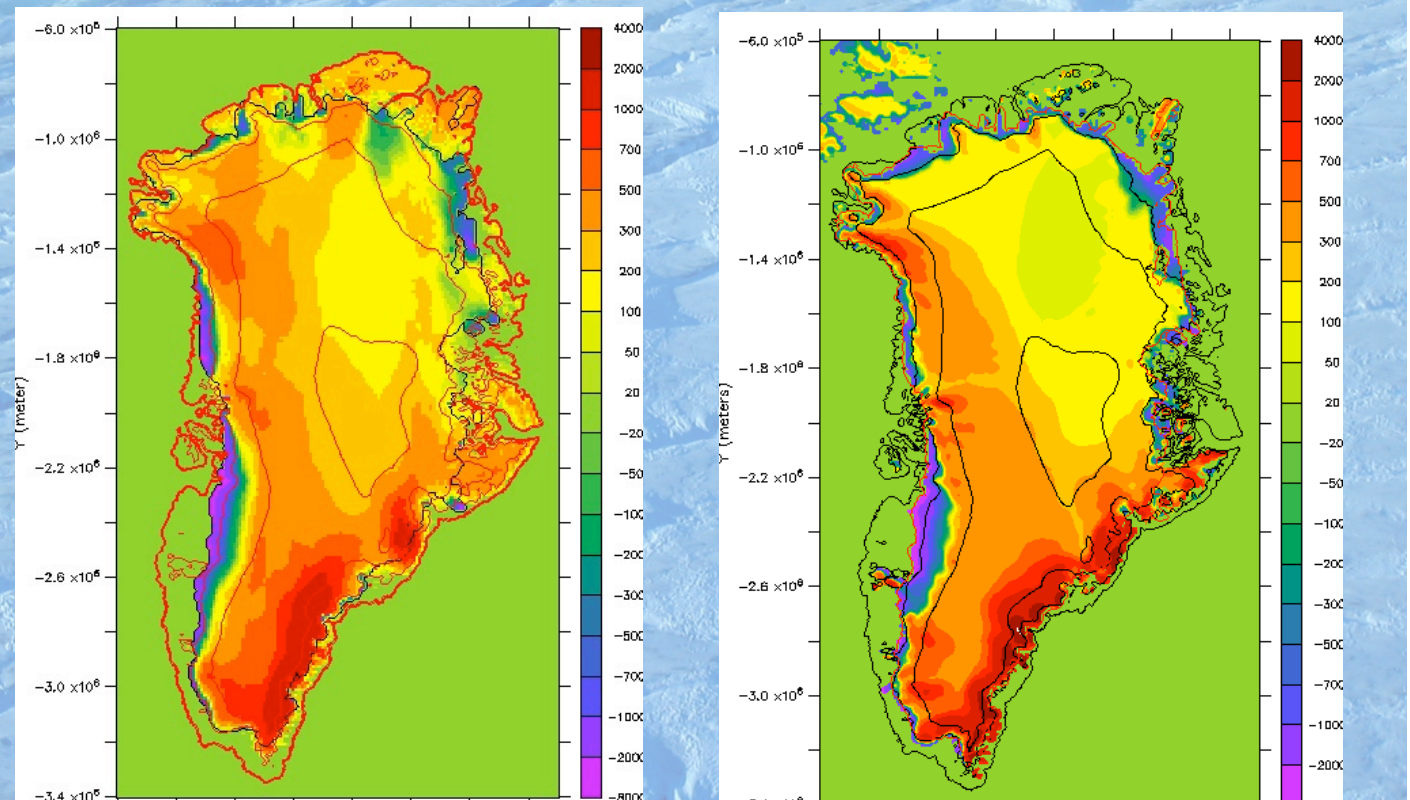
# A community ice sheet model in CESM

CESM1 includes Glimmer, the Community Ice Sheet Model & a new surface mass balance scheme for ice sheets in CLM.

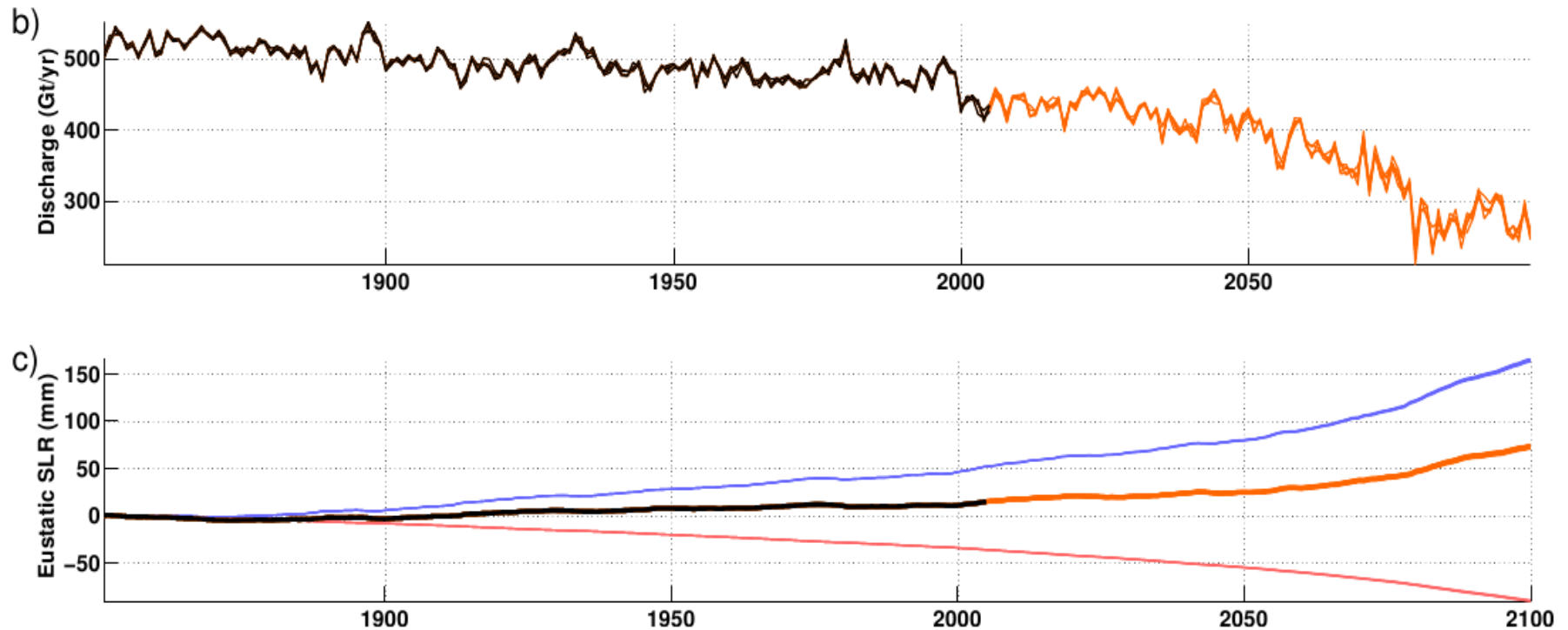
**Left:** Greenland SMB: CLM on  $1^\circ$  grid forced by CAM output, downscaled to 10 km **Right:** Greenland SMB from a high-resolution regional climate model (Ettema et al. 2009).

Red = net  
accumulation

Blue = net  
ablation



# Greenland ice sheet discharge and sea level rise over the 20<sup>th</sup> Century and 21<sup>st</sup> Century (RCP8.5)



c) Actual sea level rise (orange): Rise is 58mm over 2005 – 2100.  
Rise due to change in surface mass balance (blue): Fall due to  
change in Greenland ice dynamics (pink). Lipscomb et al. (2013).

# Climate Sensitivites

## Equilibrium

- Globally-averaged surface temperature increase due to a doubling of CO<sub>2</sub>.
- Calculated using a slab upper ocean model, so the integration comes into equilibrium in about 30 yrs.
- CCSM4 -- 3.20°C
- CESM1 -- 4.10°C

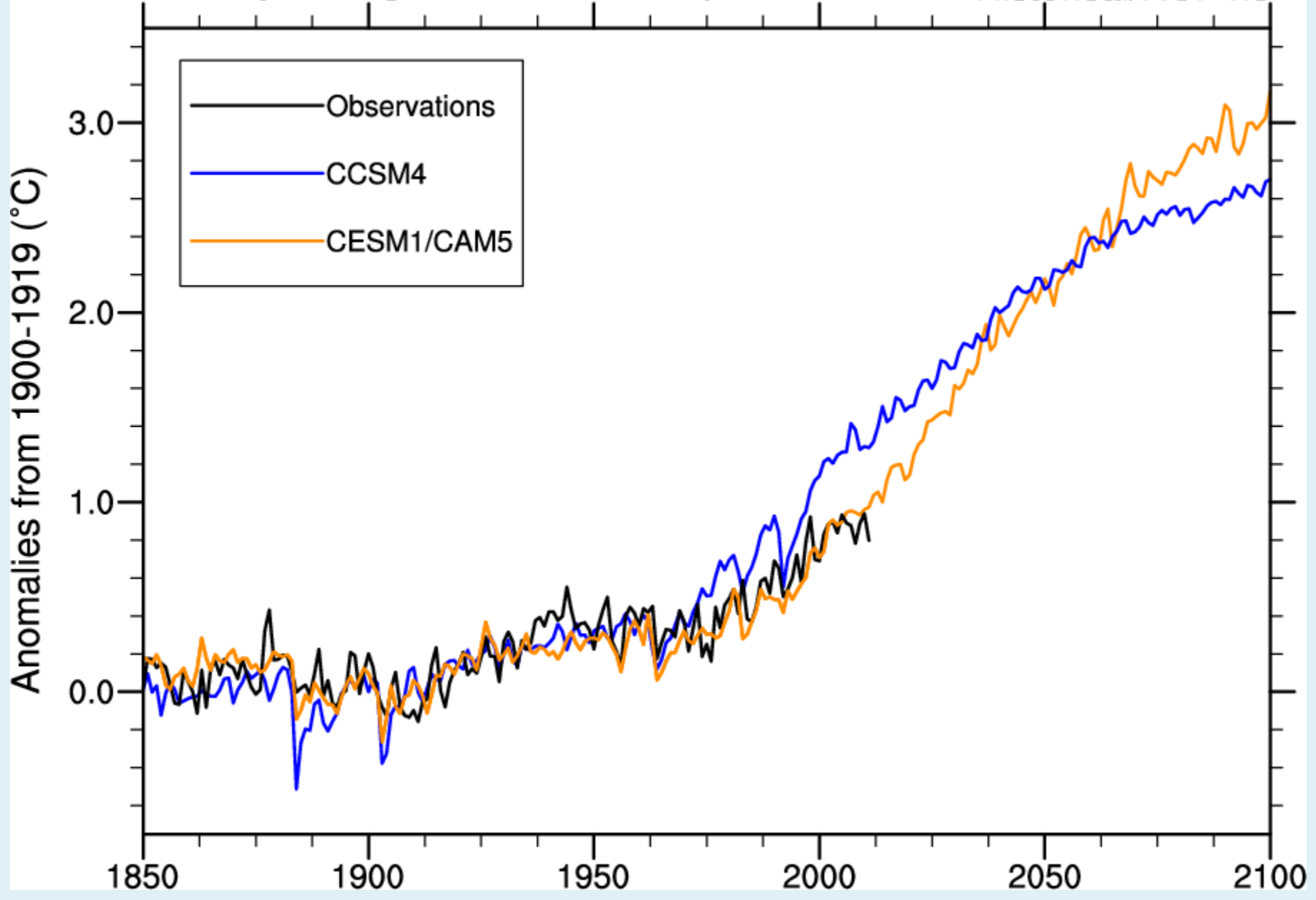
## Transient

- Global surface temperature increase at doubling of CO<sub>2</sub> in a run where it increases at 1% per year: Average over years 60 -- 80.
- Calculated using the full depth ocean component.
- CCSM4 -- 1.73°C
- CESM1 -- 2.33°C

CESM1 includes secondary effects due to aerosols

# Globally averaged surface air temperature

Historical/RCP4.5

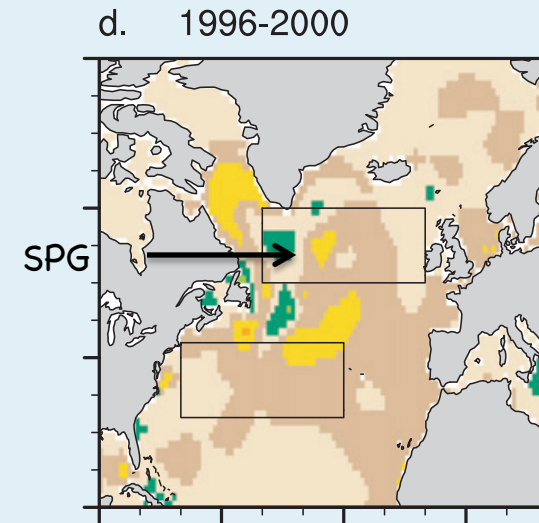
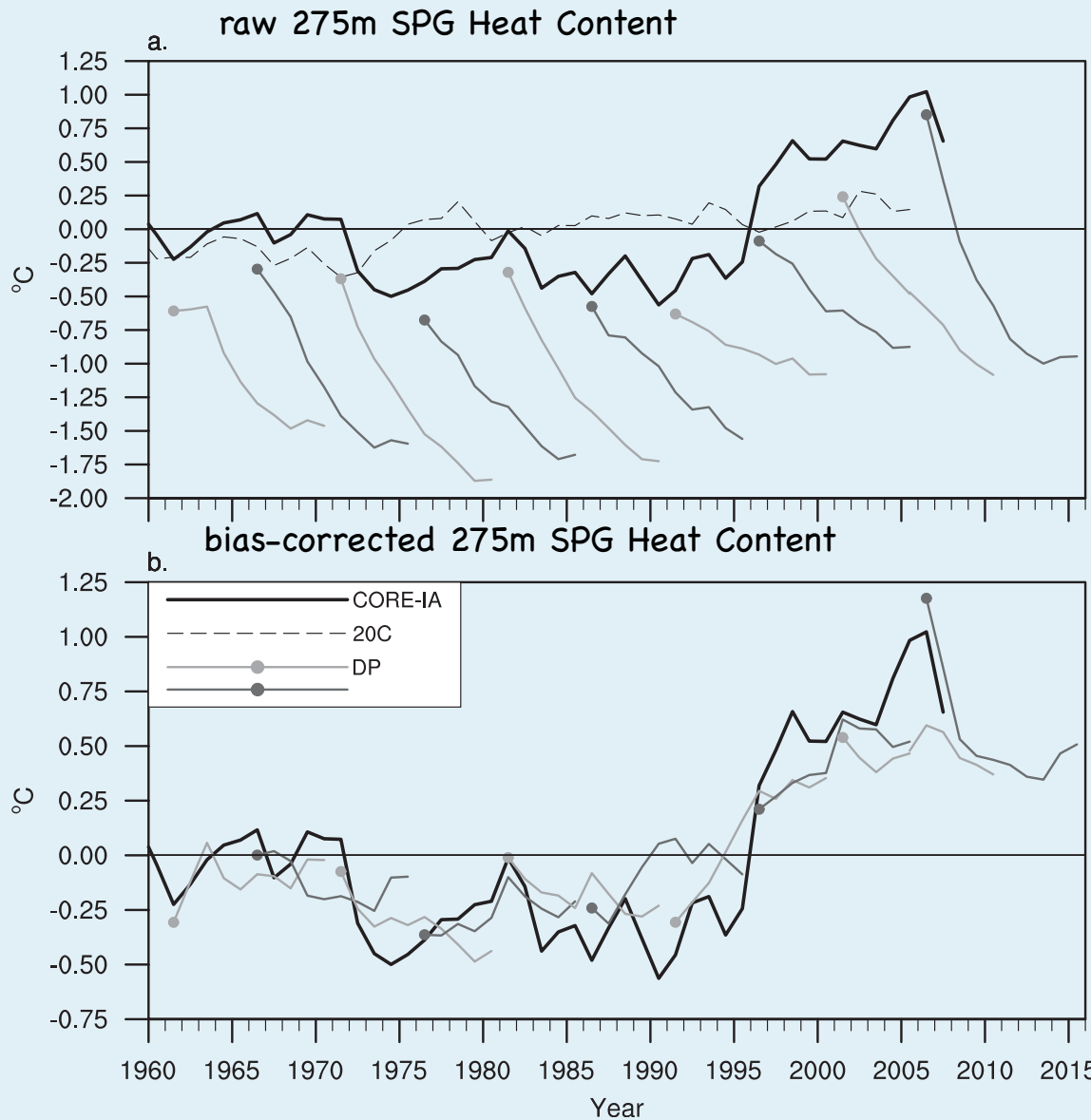




# Decadal Forecasts with the CESM

- Fully-coupled CESM1 simulations initialized from historical ocean and sea ice states for Jan 1<sup>st</sup> of 1961, 1966, ..., 2006
- Ocean & sea ice initial conditions from an ocean/sea-ice simulation of 1948–2007 forced by atmospheric observations.
- Full-field initialization (with bias correction)
- 10-member ensembles for each of 10 start dates
- Case study of the mid-1990s warming in the subpolar gyre of the North Atlantic Ocean.

# Decadal Prediction in North Atlantic Subpolar Gyre



After bias correction, the CESM has significant skill in predicting SPG heat content & SST up to a decade in advance.

Yeager et al. (2012)

# ***Conclusions***

- **The first coupled climate models were run in the early 1970s, and used very coarse resolution.**
- **These models had to use flux corrections in order to maintain the present climate in coupled runs.**
- **The CSM1 (run in 1996) was the first to maintain the present ocean temperatures in a coupled run.**
- **Earth System Models have interactive components for the carbon cycle, and possibly others, such as Chemistry, high-top Atmosphere and Land Ice.**