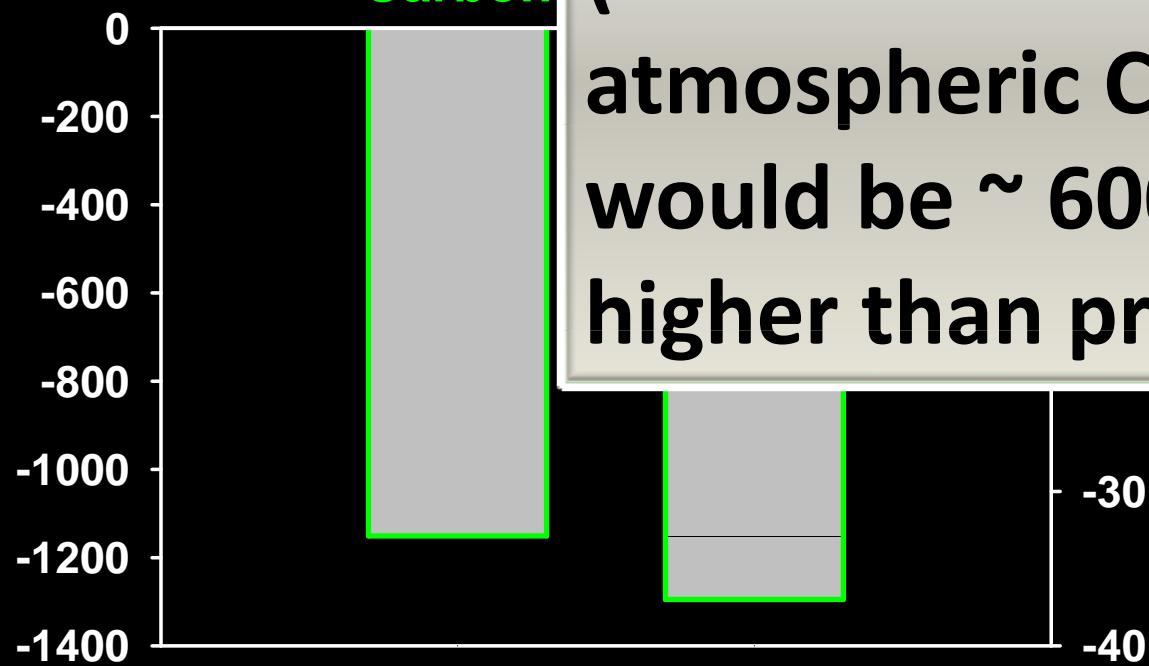
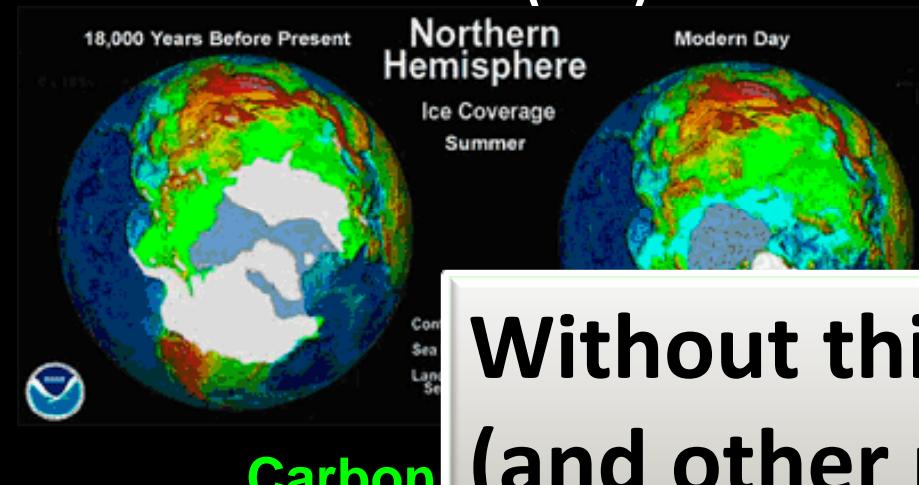




Lecture 2. Thoughts toward global nutrient cycling frontiers: Questions, techniques and opportunities

Benjamin Z. Houlton, Global Ecology and Biogeochemistry Lab, UC Davis
Web: www.houlton.lawr.ucdavis.edu; Email: bzhoulton@ucdavis.edu

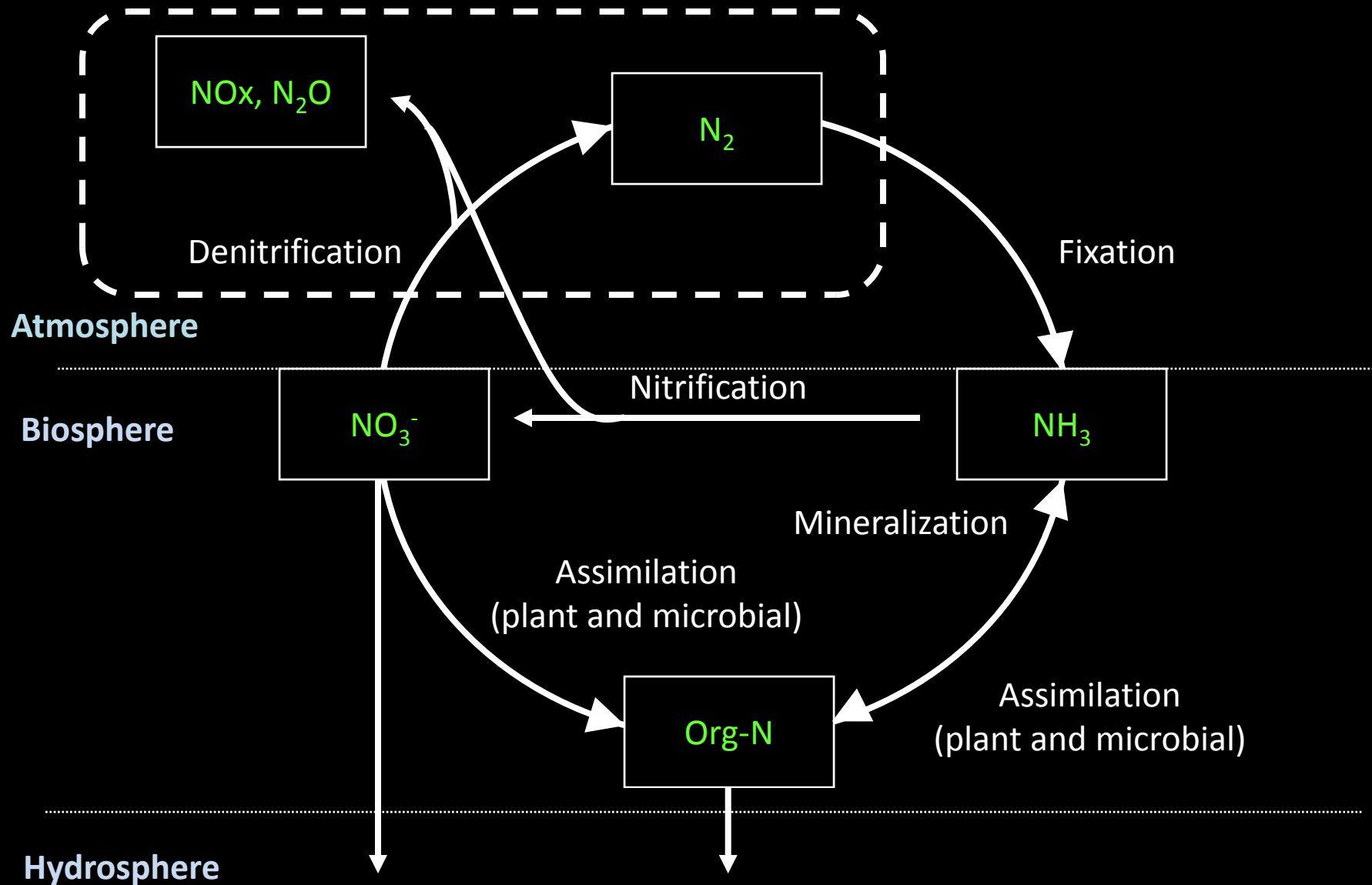
Carbon and nitrogen (and other nutrients) since LGM (Gt)

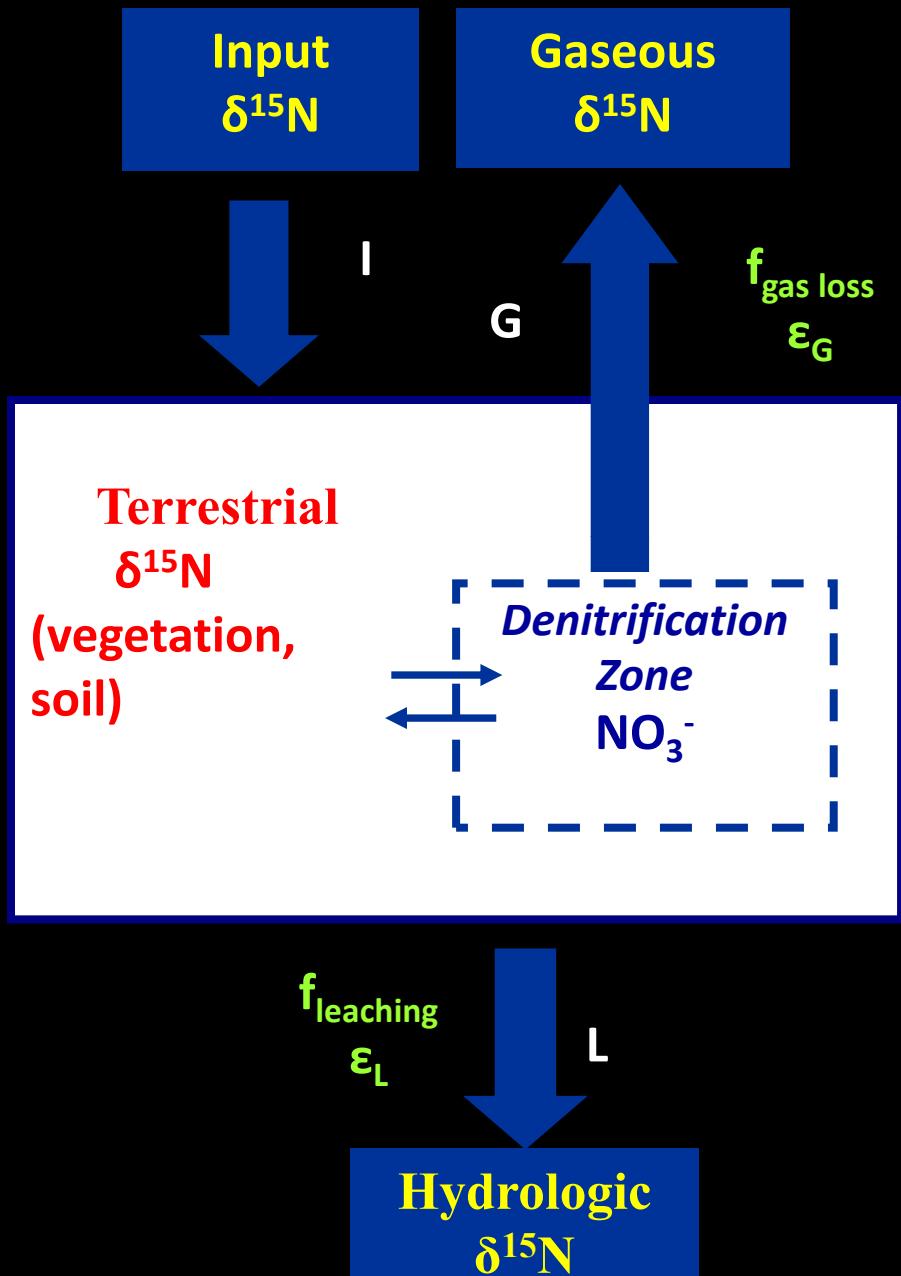


Without this added N (and other nutrients), atmospheric CO_2 would be ~ 600 ppm higher than present

1. Global hotspots of gaseous N losses in the terrestrial biosphere – **missing N sink?**
 - C,N, P modeling and natural isotopic constraints
2. N input via rock weathering – **missing N source?**
 - Cases study in natural forests and global calculations
3. Patterns of new production: N, P and satellites

Problem: N_2 production < 10^{-7} of atmospheric N_2 reservoir

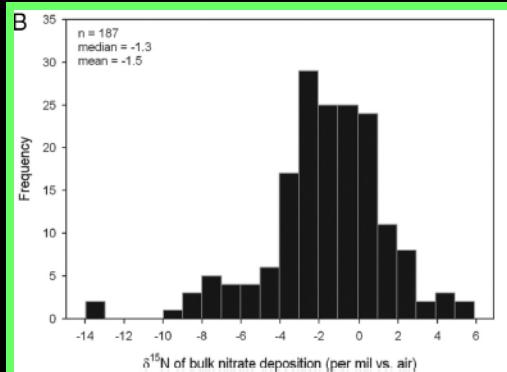




$$f_{\text{gas}} = \frac{\delta^{15}\text{N}_{TB} - \delta^{15}\text{N}_I + \epsilon_L}{\epsilon_L - \epsilon_G}$$

(Houlton et al., *PNAS*, 2006;
 Houlton et al., *PNAS*, 2007;
 Houlton and Bai, *PNAS*, 2009
 Bai and Houlton, *GBC*, 2009;
 Bai et al. *Biogeosciences*, 2012)

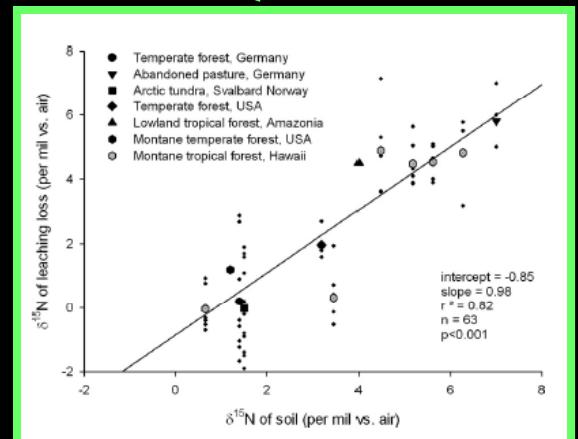
Global
data



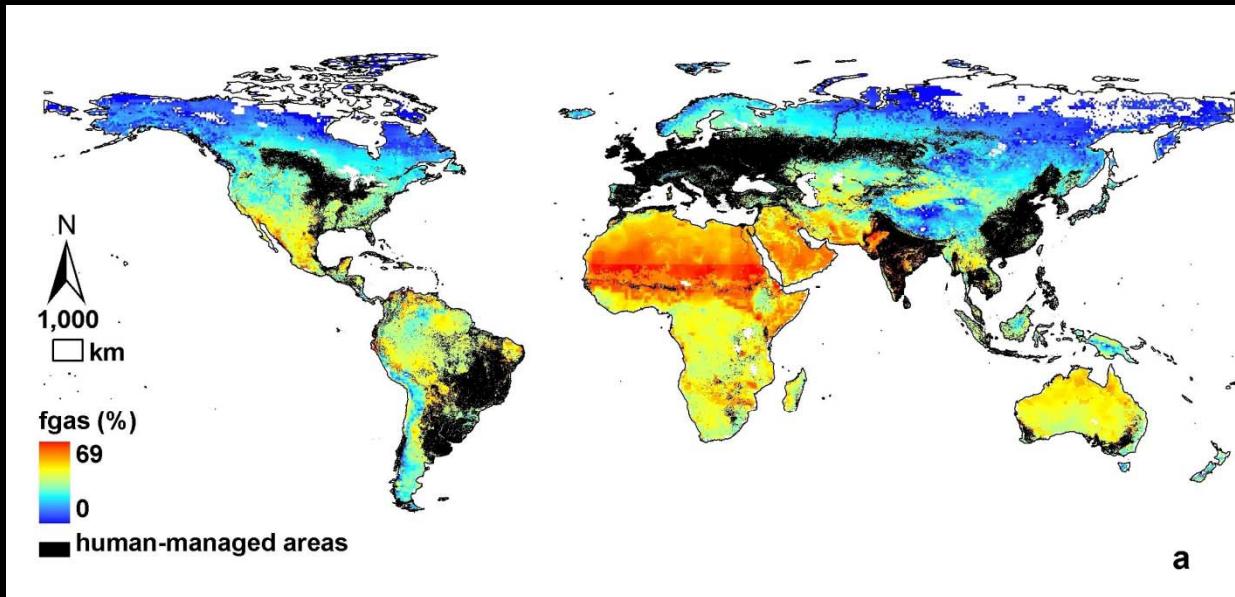
Observation
and
modeling

$$f_{\text{gas}} = \frac{\delta^{15}\text{N}_{\text{soil}} - \delta^{15}\text{N}_I - (\varepsilon_{\text{NH}3} - \varepsilon_L) \times f_{\text{NH}3} - \varepsilon_L}{\varepsilon_G - \varepsilon_L}$$

Lab and field
experiments



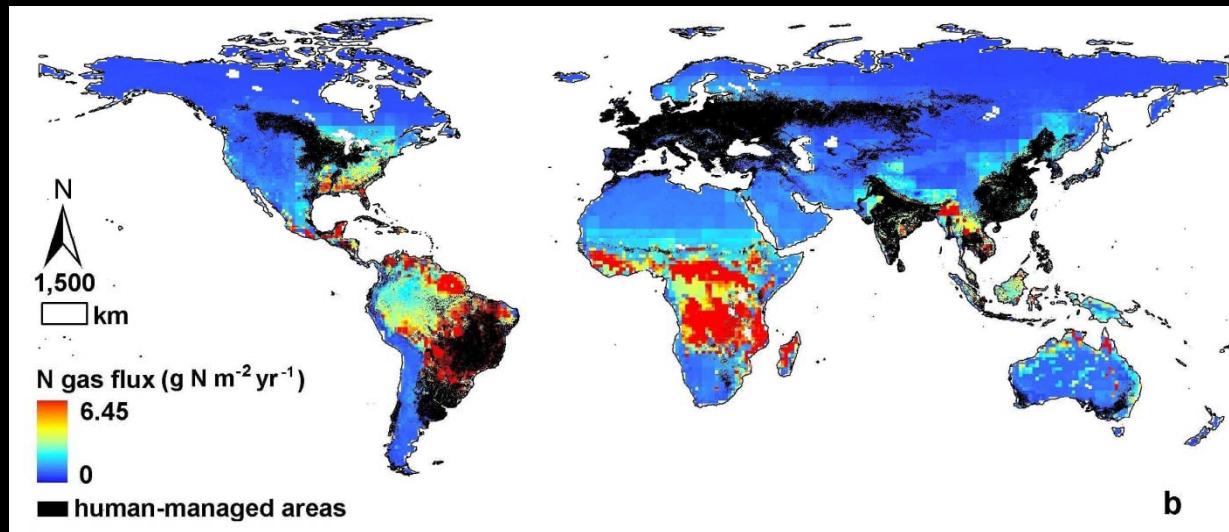
Denitrification gas loss fractions



a

$$f_{gas} = \frac{\delta^{15}N_{soil} - \delta^{15}N_I - (\varepsilon_{NH3} - \varepsilon_L) \times f_{NH3} - \varepsilon_L}{\varepsilon_G - \varepsilon_L}$$

Denitrification fluxes



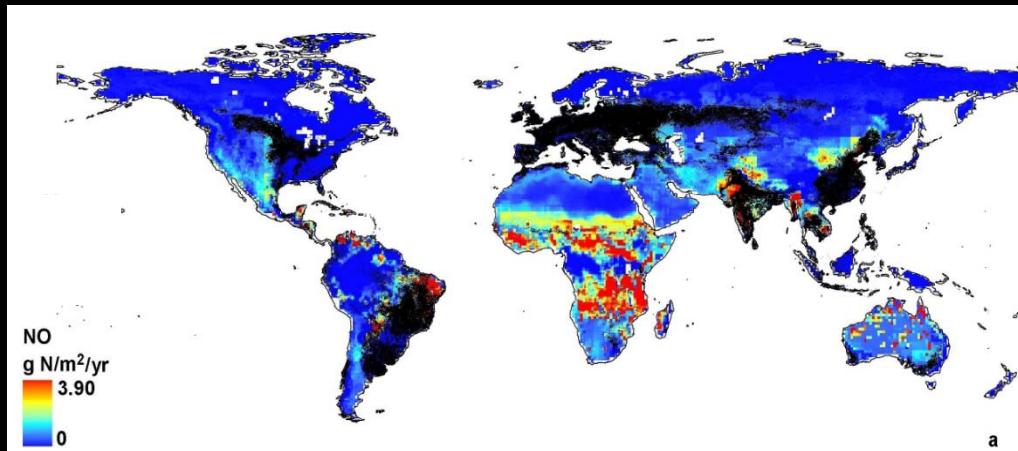
$$f_{gas} = \frac{\delta^{15}N_{soil} - \delta^{15}N_I - (\varepsilon_{NH_3} - \varepsilon_L) \times f_{NH_3} - \varepsilon_L}{\varepsilon_G - \varepsilon_L}$$

+

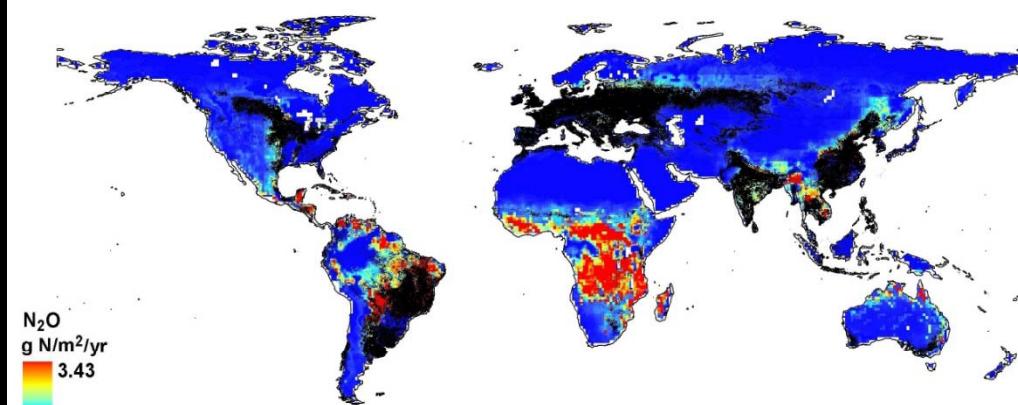
N deposition (Lelieveld and Dentener, *JGR*, 2000)

N fixation (Houlton et al., *Nature*, 2008)

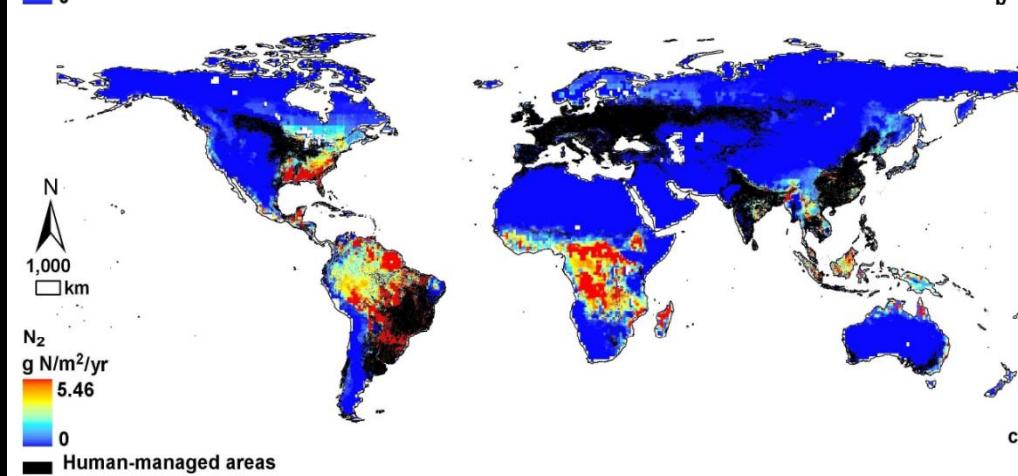
NO_x^*
11 – 20 Tg
*(cut in half with canopy)



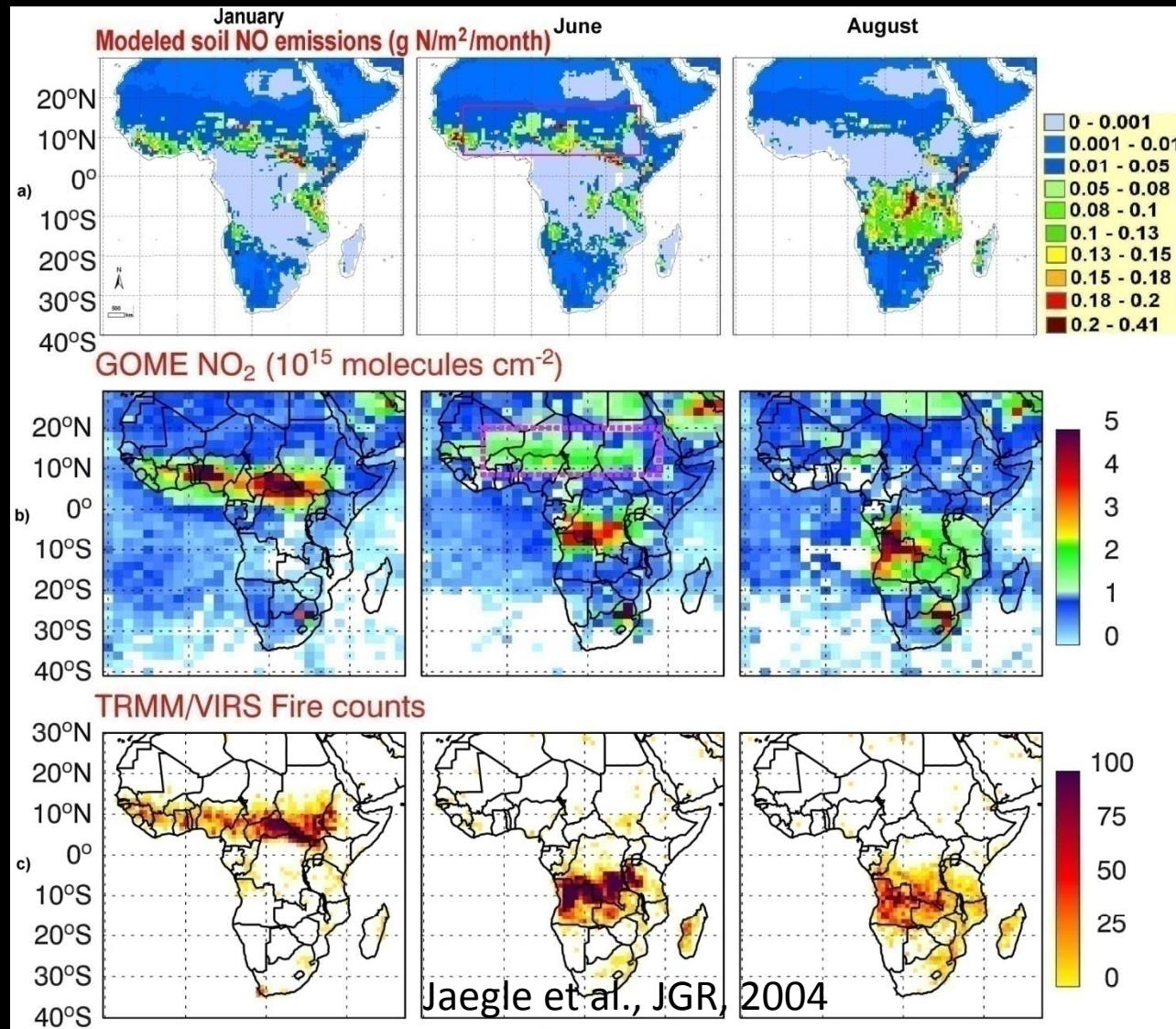
N_2O
7 – 13 Tg
(Ag. ~3)



N_2
15 – 27 Tg

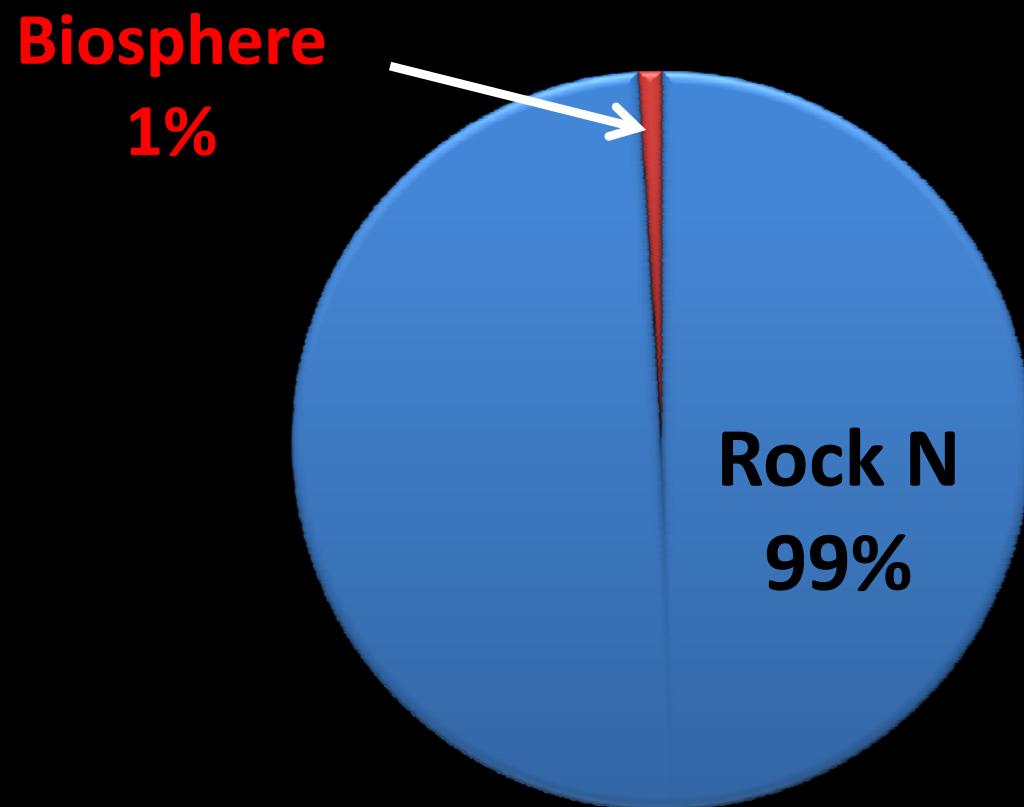


Example: isotopic modeling and satellites



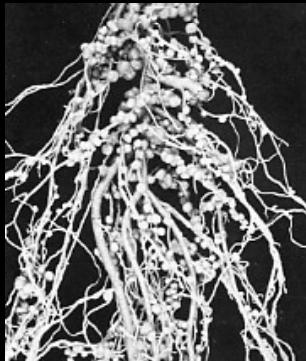
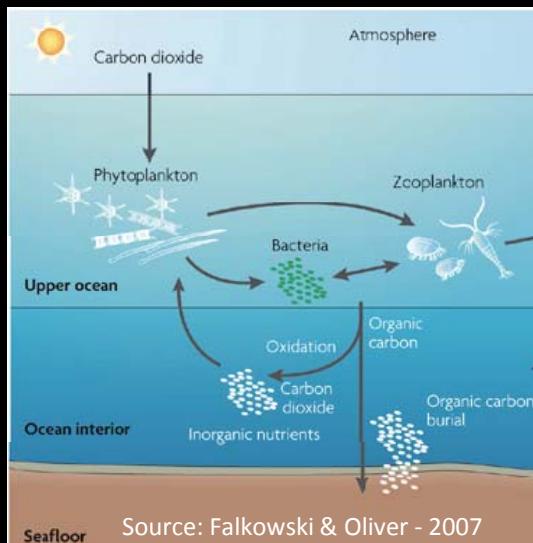
1. Global hotspots of gaseous N losses in the terrestrial biosphere – missing N sink?
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Global distribution of “fixed” N

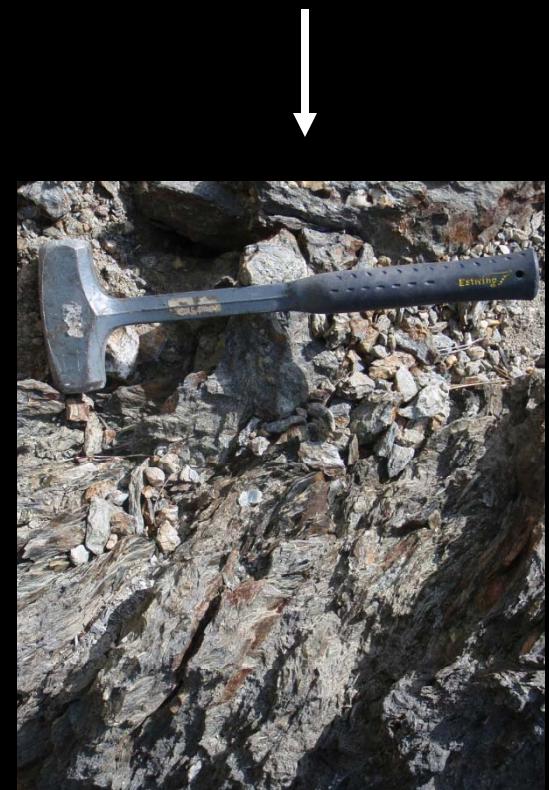


Origin and Genesis of Geologic N

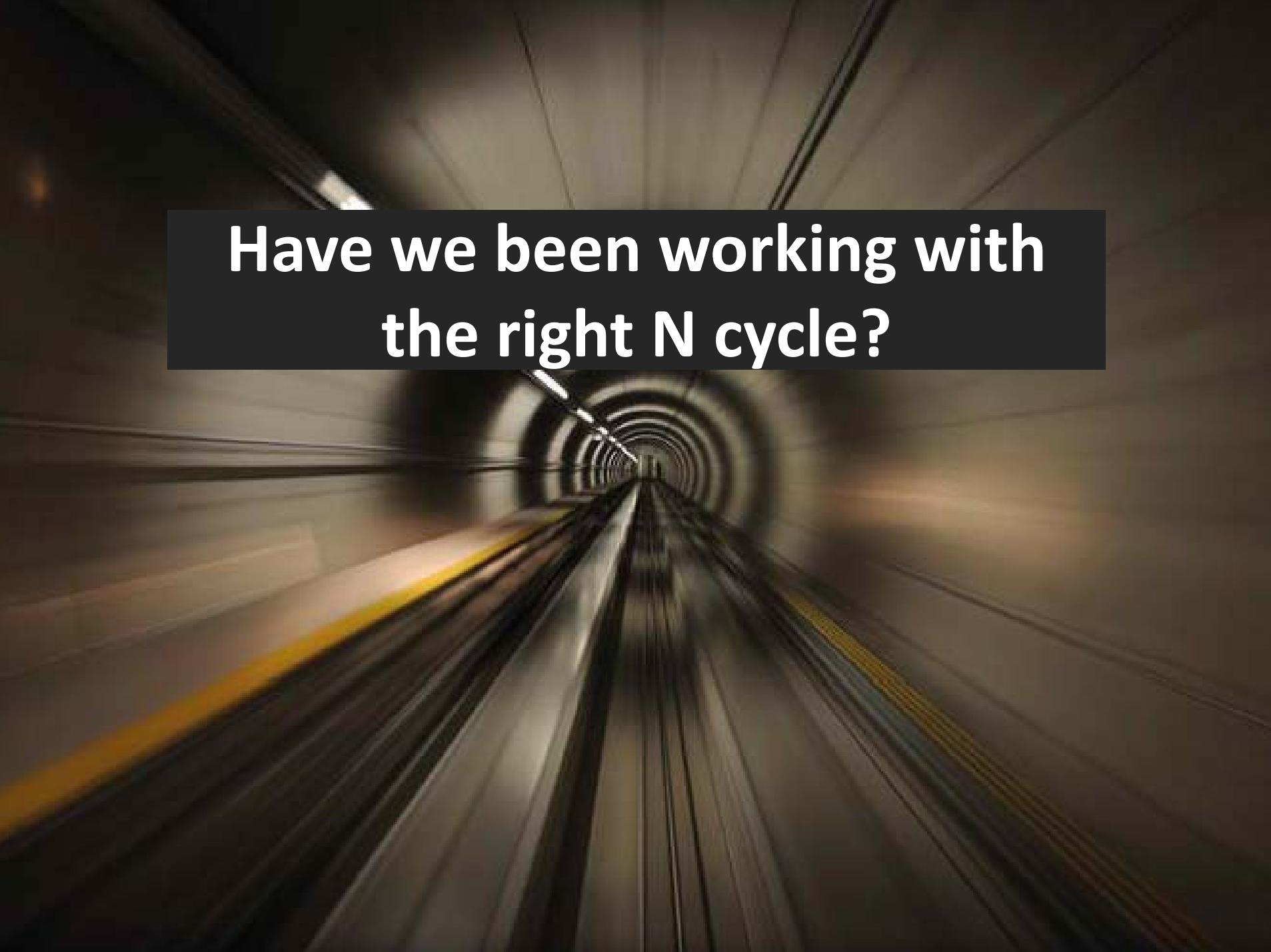
N fixation in the past → Diagenesis → Low-grade metamorphism



Organic N in shale and mudstones



Fixed NH_4^+ in silicate rocks



**Have we been working with
the right N cycle?**

Study Sites

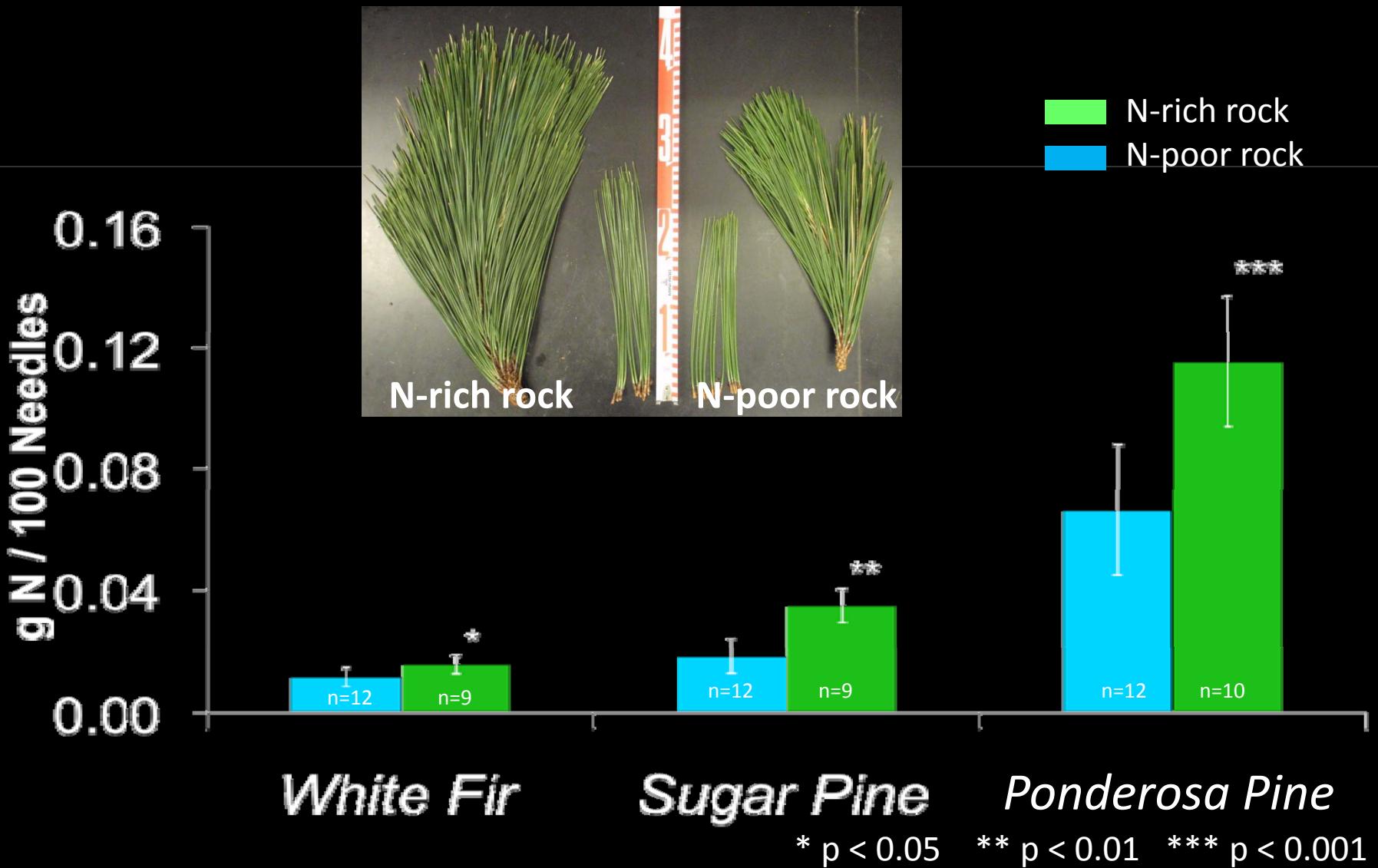


	<u>N-rich rock</u>	<u>N-poor rock</u>
Geology:	Mica Schist	Diorite
Elevation:	1750 m	1500 m
N deposit:	< 1kg ha ⁻¹	< 1kg ha ⁻¹
Aspect:	NE & N	N
Precip:	1520 mm	1400 mm
Temp:	9° c	10.5° c
Rock:	682 ppm	55 ppm

Species sampled:

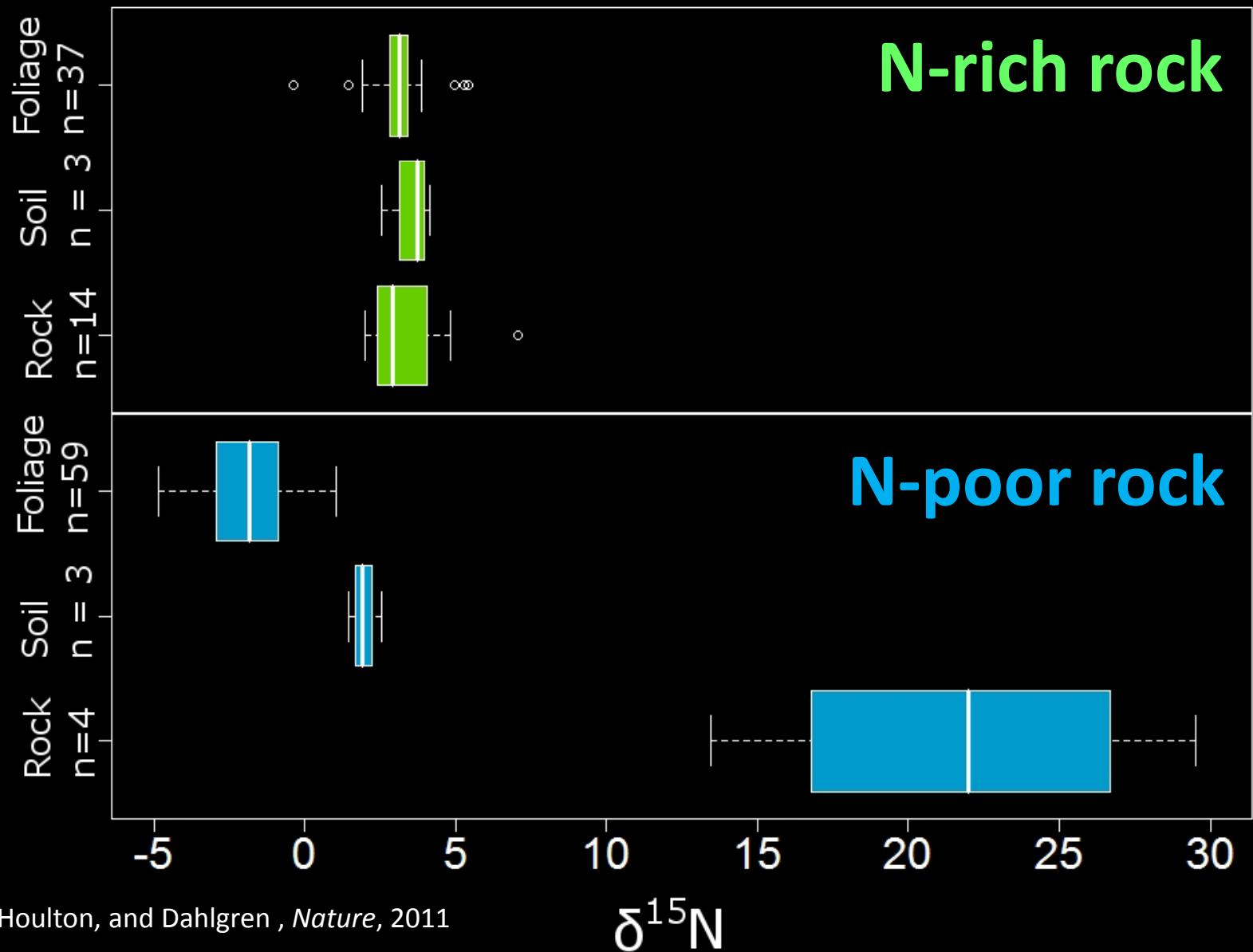
- Abies concolor* (White Fir)
- Pinus lambertiana* (Sugar Pine)
- Pinus ponderosa* (Ponderosa Pine)
- Calocedrus decurrens* (Incense Cedar)

Total N in Plant Foliage



Morford, Houlton, and Dahlgren, *Nature*, 2011

$\delta^{15}\text{N}$ of Plant – Soil – Rock System



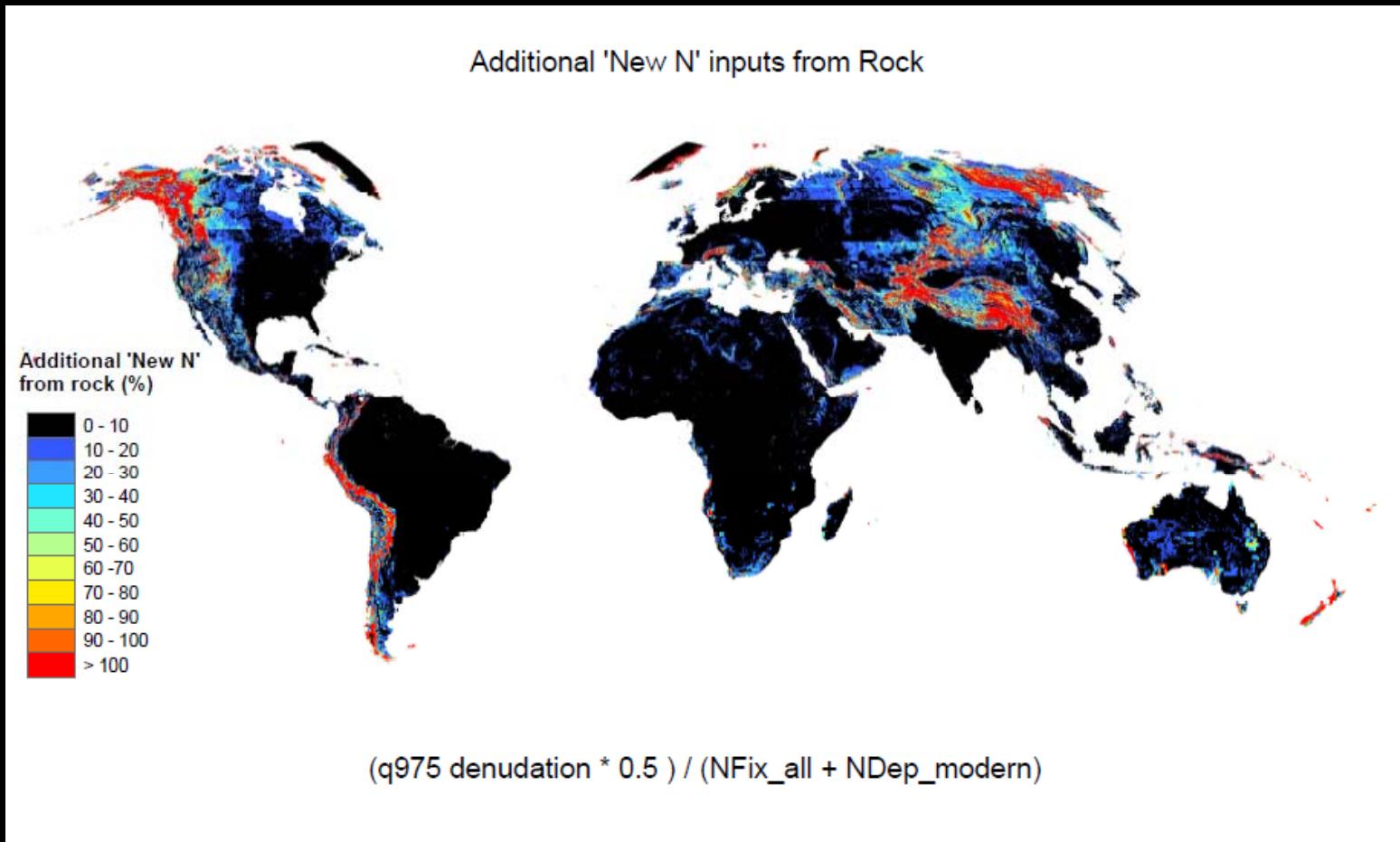
Estimated N Inputs

	Isotope balance	Isotope balance with fractionations	Uplift	Lab weathering
Fraction of total N inputs via rock weathering	1.22	0.30	0.47	0.64

3 to 10 Kg N/ha/yr, or ~ doubling of N inputs

Morford, Houlton, and Dahlgren , Nature, 2011

Global implications



~ 24 Tg N at the global Scale, cf. ~120 for natural N fixation, 5 for lightening

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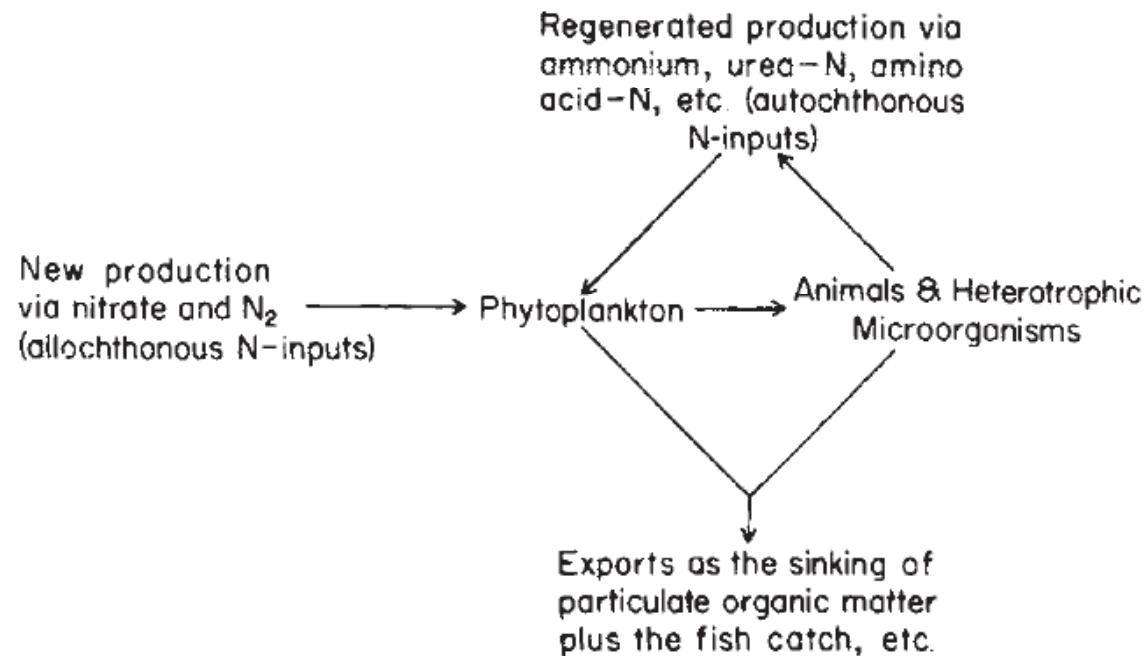
Particulate organic matter flux and planktonic new production in the deep ocean

Richard W. Eppley

Institute of Marine Resources, A-018, Scripps Institution of Oceanography, University of California, San Diego, La Jolla, California 92093

Bruce J. Peterson

Ecosystems Center, Marine Biological Laboratory, Woods Hole, Massachusetts 02543



$$f = \text{new production/total} = \\ \sim 18\% \text{ for new nitrogen}$$

Defining new vs. recycled production for the terrestrial biosphere

$$NPP_{total}(N, P) = NPP_{new}(N, P) + NPP_{recycle}(N, P)$$

$$f_{new}(N, P) = \frac{NPP_{new}(N, P)}{NPP_{total}}$$

$$f_{recycle}(N, P) = \frac{NPP_{recycle}(N, P)}{NPP_{total}}$$

$$NPP_{new}(N, P) = NPP_{total}(N, P) \times f_{new}(N, P)$$

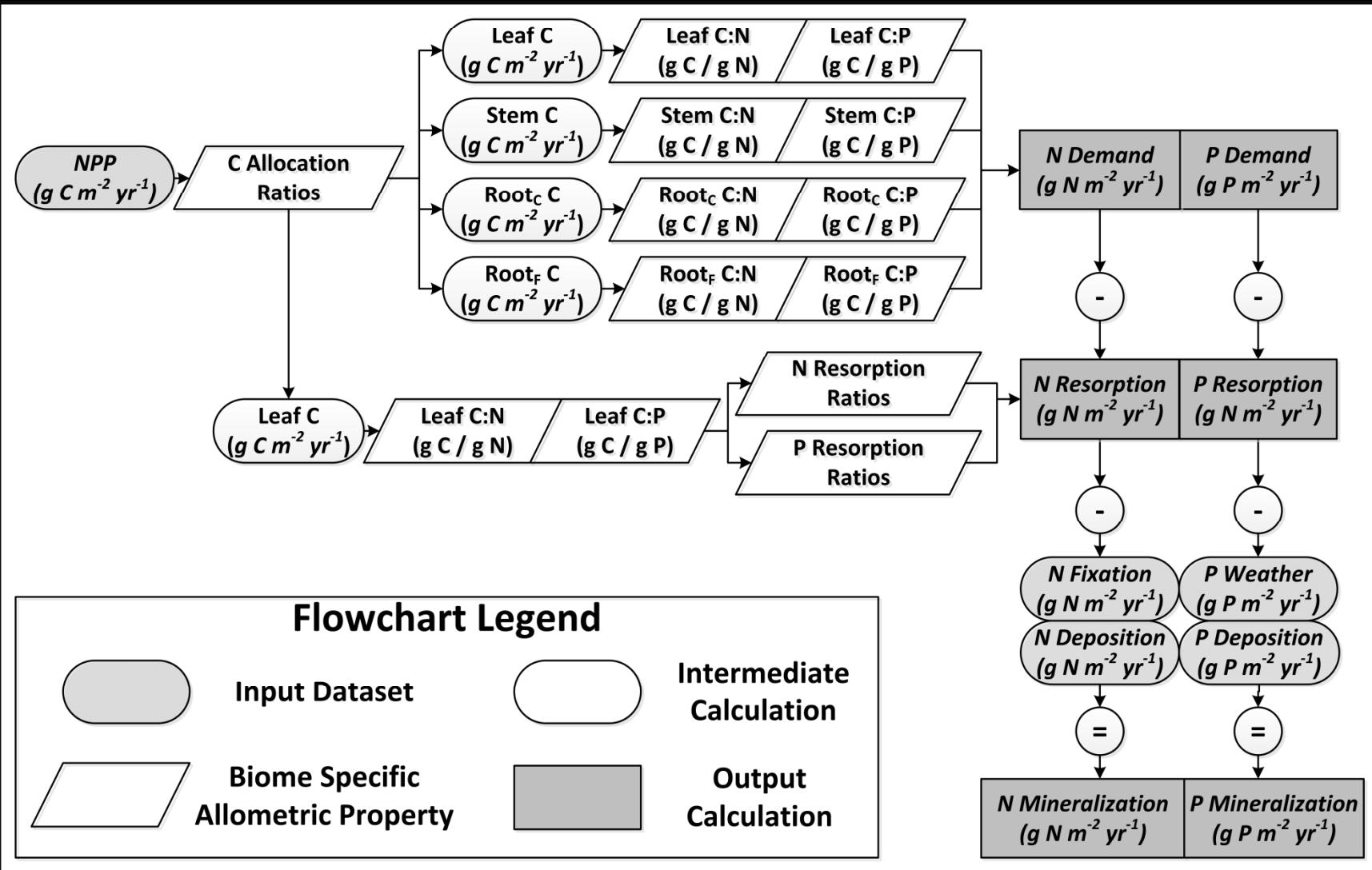
New production vs. recycled production

Total plant N demand = N Resorption + Net N Mineralization +
N Fixation (f plant) + N Deposition (f plant)

Total plant P demand = P Resorption + Net P Mineralization + P
Dust (f plant) + P Weathering (f plant)

New N Production = N Fixation (f plant) + N Deposition (f plant)
N Resorption + Net N Mineralization

New P Production = P Dust (f plant) + P Weathering (f plant)
P Resorption + Net P Mineralization



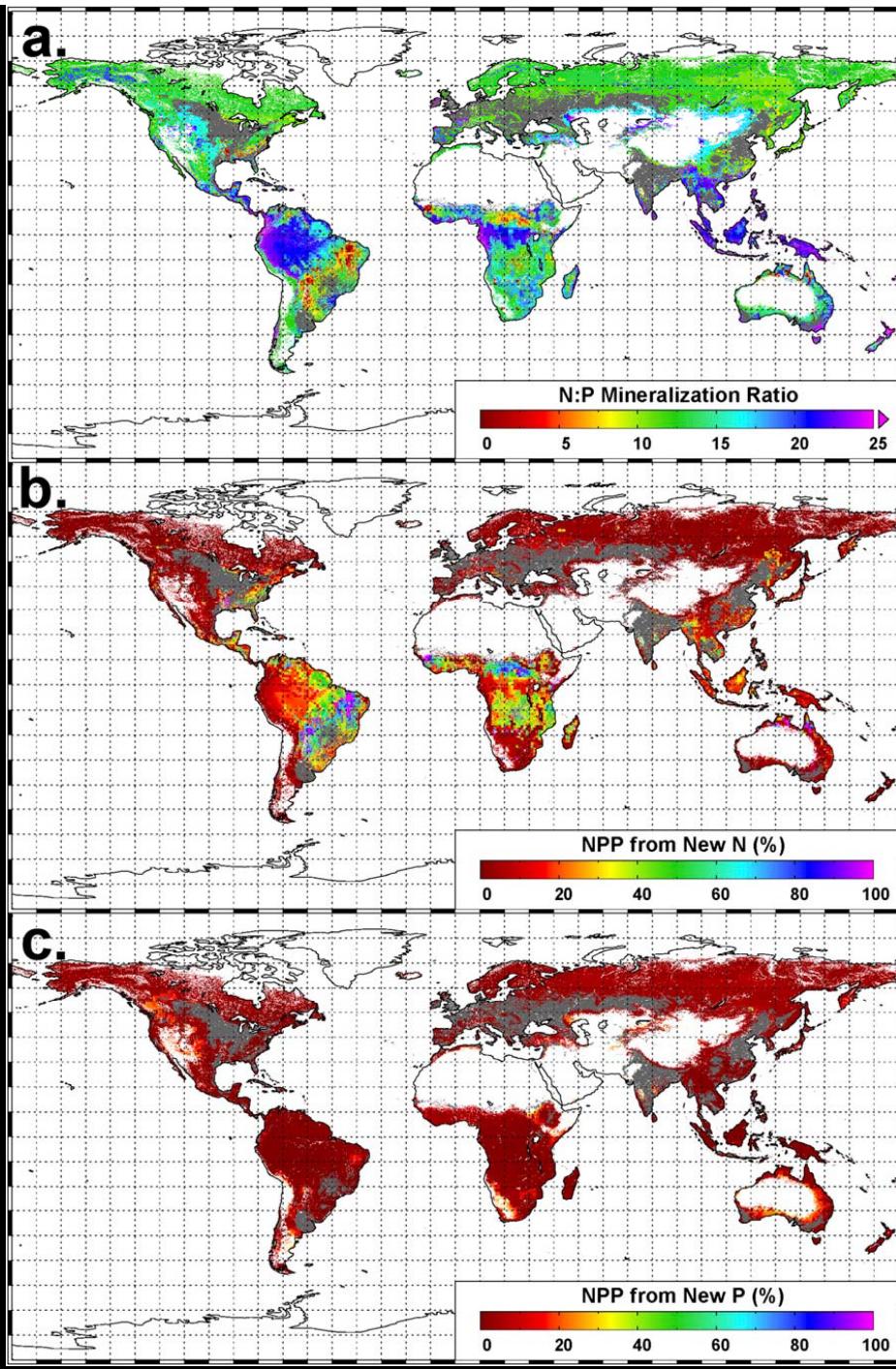
Flowchart Legend

Input Dataset

Intermediate Calculation

7 Biome Specific Allometric Property

Output Calculation

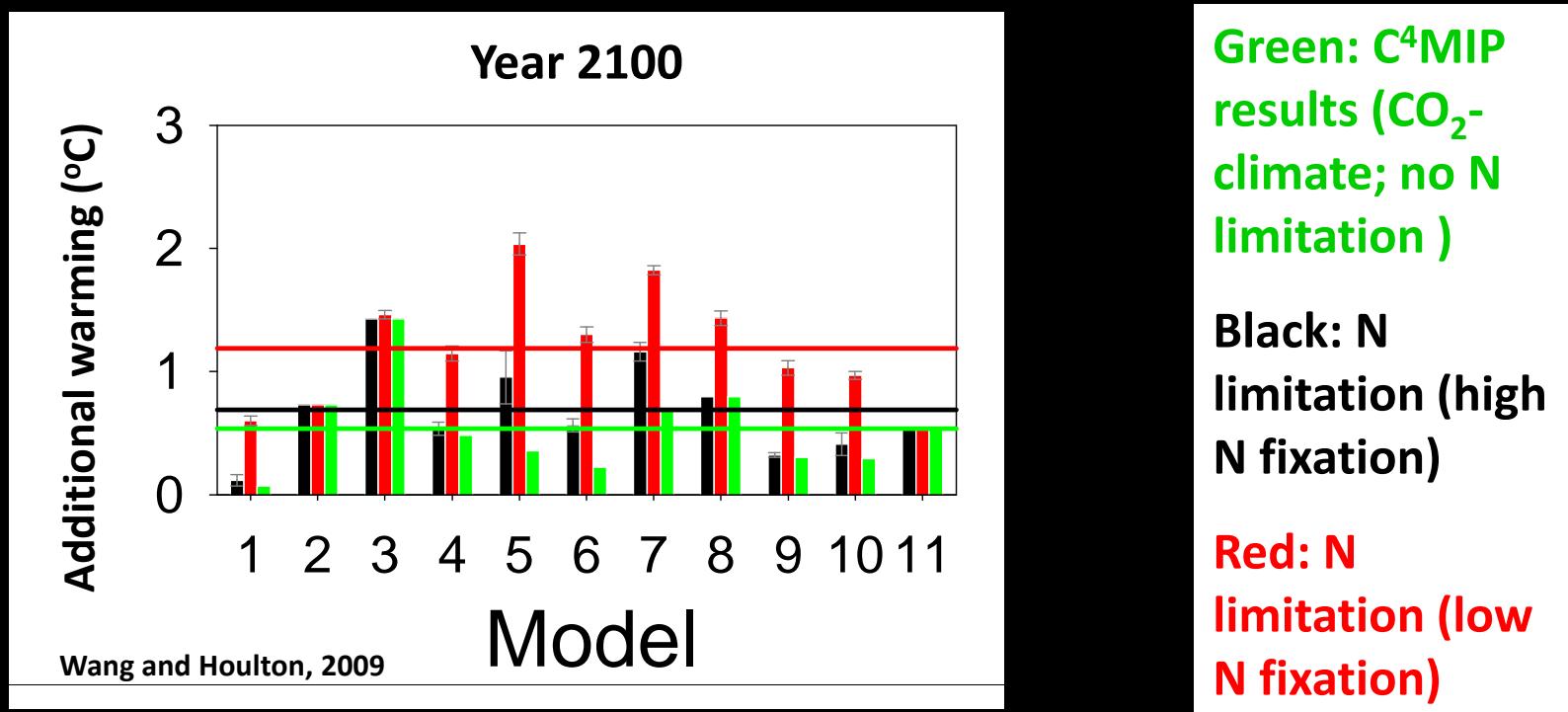


Global Implications of terrestrial f

- Terrestrial new production is 16% for nitrogen; <2 % for phosphorus
- Natural nitrogen fixation accounts for more than 10% of new nitrogen production; nitrogen deposition makes minor contribution (<1%)
- Tropical systems – particularly tropical forests – display highest capacity for new nitrogen production, perhaps C uptake in general (~18% vs. ~3% for extra-tropical systems)

Closing thoughts

We have much to learn about global nutrient balances...with significant implications



2100-1900: 0.54°C (C⁴MIP); 0.69 to 1.19°C (N limiting)

Acknowledgments

- Andrew W. Mellon Foundation
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- Kearney Foundation of Soil Science
- David and Lucile Packard Foundation
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Questions?