

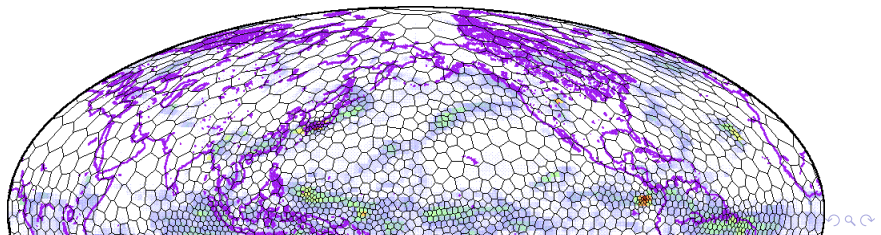
Curl-free pressure gradients over orography in a solution of the fully compressible Euler equations with long time-steps

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Engineering, Leeds University, UK²

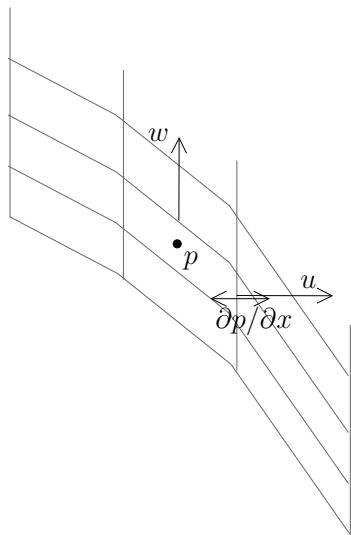
7 April 2014



Motivation

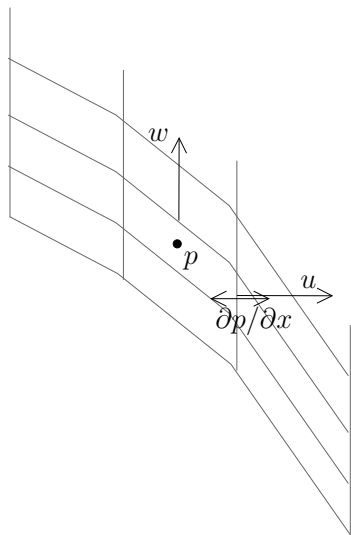
- ▶ Mimetic properties over orography
 - ▶ using mimetic horizontal discretisation in the vertical
- ▶ Long time-steps
 - ▶ suitable for massively parallel
 - ▶ suitable for unstructured grids
 - ▶ simpler than SISL (semi-implicit, semi-Lagrangian)

Representing Mountains



For numerous reasons in meteorology the cells should line up in columns

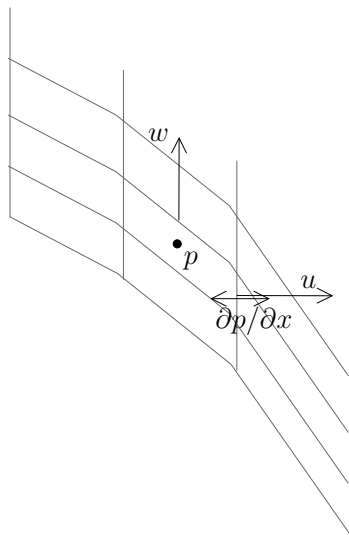
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\therefore the mesh is non-orthogonal over orography

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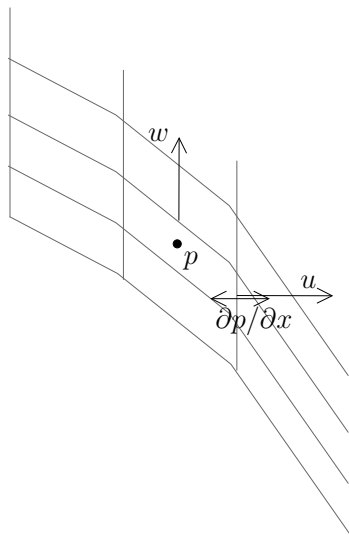


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Usual approach: orthogonal prognostic velocity variables u , v , w in horizontal and vertical directions

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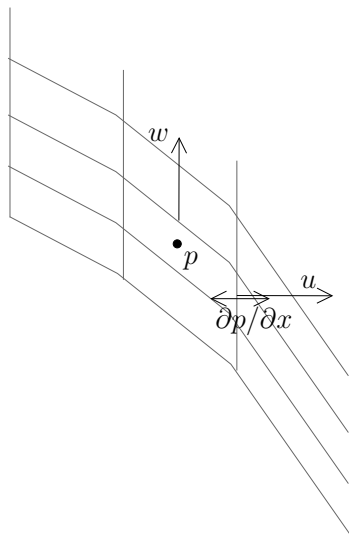
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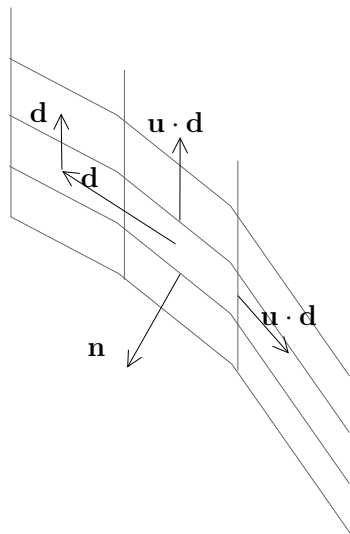
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\therefore find $\frac{\partial p}{\partial x}$ co-located with u without knowing p at this altitude (eg Klemp, Zangl)

\rightarrow pressure gradients not curl free

Alternative: non-orthogonal prognostic variables (covariant)



Following horizontal discretisations on non-orthogonal grids:

Prognostic variables: $\mathbf{u} \cdot \hat{\mathbf{d}}$

where $\mathbf{d}_f = \mathbf{x}_i - \mathbf{x}_j$

→ curl free pressure gradients

→ no spurious generation of vorticity

Non-orthogonal prognostic variables (covariant)

Need mass flux $u_n = \mathbf{u} \cdot \hat{\mathbf{n}}$
in continuity equation

Requires operator H from space of all
 u_d s to space of all u_n s.

$$u_n = H u_d$$

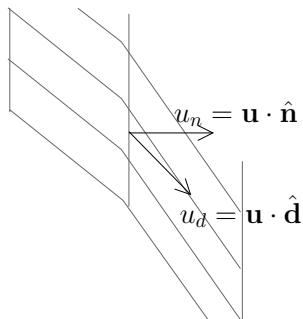
(following Thuburn, Dubos, Cotter, 2014)

- First reconstruct full velocity at face f
from $u_{d'}$ at the surrounding faces, f' :

$$\mathbf{u} = T_i^{-1} \sum_{f'} \mathbf{d}_{f'} A_{f'} u_{d'}$$

$$\text{where } T = \sum_{f'} \hat{\mathbf{d}}_{f'} \hat{\mathbf{d}}_{f'}^T A_{f'}$$

Least squares fit which reconstructs a
uniform velocity field
The resulting H is asymmetric which violates energy conservation



- Next take component in
direction \mathbf{n} and correct the
component in direction \mathbf{d} so
that the result is exact on
an orthogonal face:

$$u_n = \mathbf{u} \cdot \hat{\mathbf{n}} + \left(u_d - \mathbf{u} \cdot \hat{\mathbf{d}} \right) \left(\hat{\mathbf{n}} \cdot \hat{\mathbf{d}} \right)$$

Results

Resting stratified atmosphere over a steep mountain

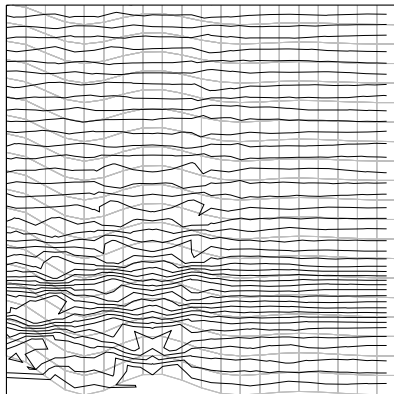
- ▶ should remain stationary
- ▶ potential temperature contours should remain horizontal

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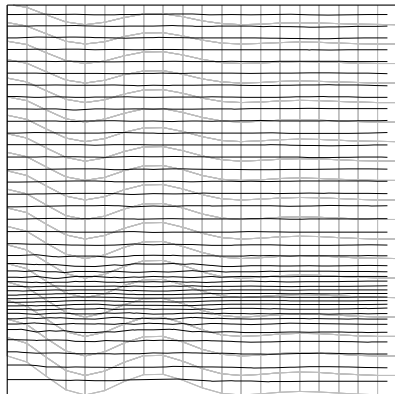
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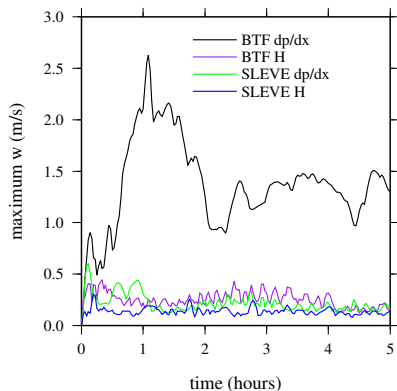
$\partial p/\partial x$ version, implicit gravity waves
 $\Delta t=100$ s. Maximum $N\Delta t=2$



H version, implicit gravity waves
 $\Delta t=100$ s. Maximum $N\Delta t=2$



Maximum Spurious Velocity



New models

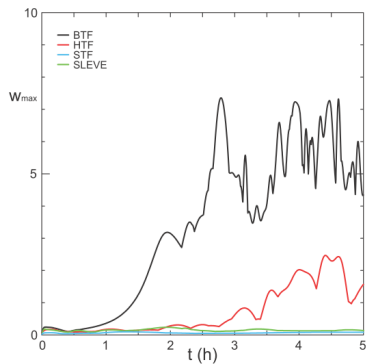
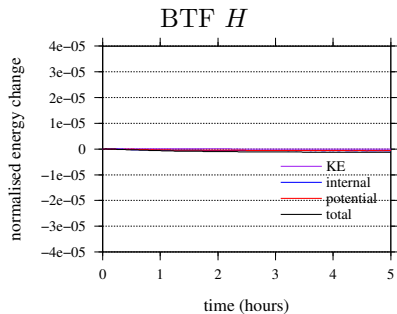
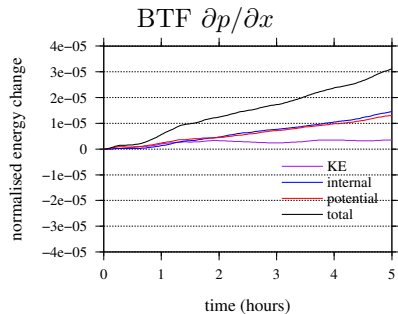


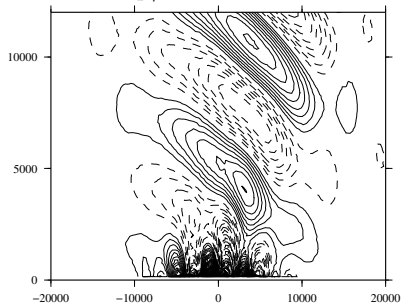
Figure 4 from Klemp [2011]

Energy Conservation



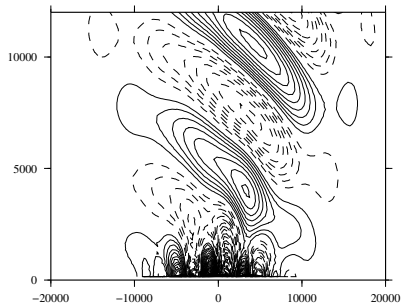
Gravity-waves over orography $\Delta t=40s$, $u\Delta t/\Delta x \approx 1$, w every .05m/s

$\partial p/\partial x$ version

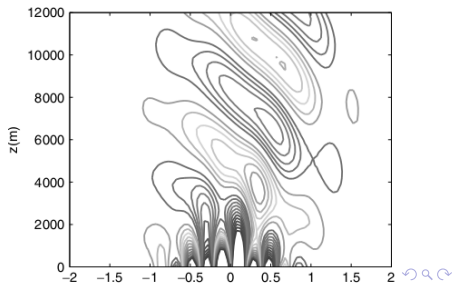
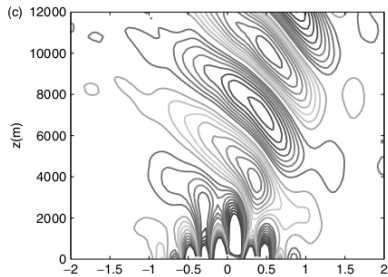


Met Office [Melvin et al., 2010]

H version



Linear analytic solution



Semi-implicit acoustic and gravity waves

For structured, lat-lon grid models, this is usually done by

- ▶ treating the z coordinate direction differently
- ▶ expressing variables as mean and perturbation quantities

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- ▶ expressing variables as mean and perturbation quantities

Neither of these are necessary:

Starting from the Euler equations:

$$\text{Momentum} \quad \partial \mathbf{u} / \partial t + \mathbf{u} \cdot \nabla \mathbf{u} + 2\Omega \times \mathbf{u} = \mathbf{g} - c_p \theta \nabla \Pi$$

$$\text{Continuity} \quad \partial \rho / \partial t + \mathbf{u} \cdot \nabla \rho + \rho \nabla \cdot \mathbf{u} = 0$$

$$\text{Potential temperature} \quad \partial \theta / \partial t + \mathbf{u} \cdot \nabla \theta = 0$$

$$\text{State} \quad \Pi^{\frac{1-\kappa}{\kappa}} = R\rho\theta/p_0$$

$$\text{where potential temperature} \quad \theta = T (p_0/p)^\kappa$$

$$\text{Exner function of pressure} \quad \Pi = (p/p_0)^\kappa$$

In order to treat acoustic and gravity wave implicitly, these must ALL be combined to form a linearised equation for Π

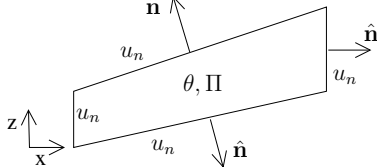
Semi-implicit acoustic and gravity waves

Prognostic variables:

$u_n = \mathbf{u} \cdot \hat{\mathbf{n}}$ velocity normal to each cell face

θ at cell centre

Π at cell centre



Overview:

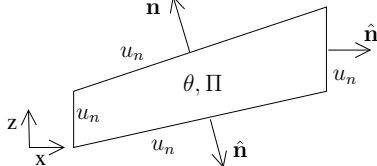
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- ▶ Substitute θ equation into momentum equation

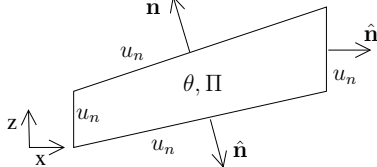
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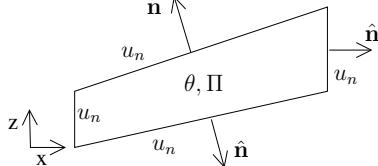
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Overview:

- ▶ Substitute θ equation into momentum equation
- ▶ Substitute u_n into continuity equation
- ▶ Use equation of state to replace ρ with Π in continuity equation:
→ Helmholtz equation for Π

Substitute θ into momentum equation

Rearrange θ equation to give θ^{n+1} in terms of u_n^{n+1} (1st order in time for brevity):

$$\partial\theta/\partial t + \mathbf{u} \cdot \nabla\theta = 0 \rightarrow \theta^{n+1} = \theta^n - \Delta t \mathbf{u} \cdot \nabla\theta^n$$

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where $\mathbf{u}^\perp = \mathbf{u} - (\mathbf{u} \cdot \hat{\mathbf{n}}) \hat{\mathbf{n}}$
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and write as:

$$u_n^{n+1} = G (u' - \Delta t c_p \theta' \nabla_n \Pi^{n+1})$$

Substitute u_n^{n+1} into the continuity equation to get ρ^{n+1} in term of Π^{n+1}

Final Construction of Helmholtz equation

$$u_n^{n+1} = G(u' - \Delta t c_p \theta' \nabla_n \Pi^{n+1})$$

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Use equation of state to replace ρ^{n+1} with Π^{n+1} :

$$\rho^{n+1} = \Psi \Pi^{n+1}$$

where $\Psi = (\rho^\ell)^{\frac{2\kappa-1}{\kappa-1}} (R\theta/p_0)^{\frac{\kappa}{\kappa-1}} \approx (p_0/R)^{0.4} \rho^{0.6}/\theta^{0.4}$

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Solve to find Π^{n+1} in terms of Π^n then back substitute to get ρ^{n+1} , u_n^{n+1} and θ^{n+1} .

This is VERY convergent and allows long time steps w.r.t. gravity and acoustic wave speeds ... but what about advection ...

Sub time-steps for advection

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Sub time-steps for advection

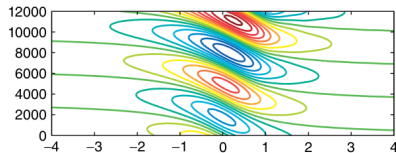
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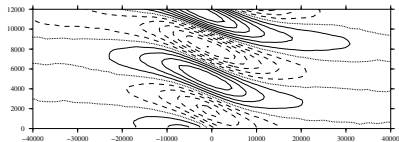
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- ▶ Linear stability analysis does not reveal any time-step restrictions (not shown)

Hydrostatic Mountain Waves (To test long time-steps)

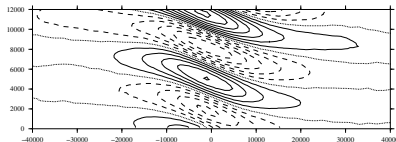
Linear analytic solution



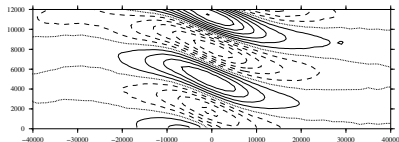
Implicit gw, $\Delta t = 20\text{s}$, $Co = 0.2$, $N\Delta t = 0.4$



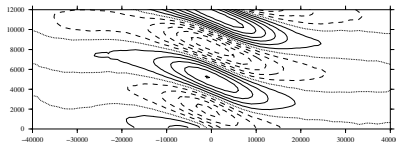
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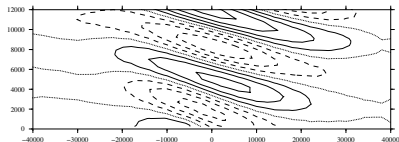
Implicit gw, $\Delta t = 100\text{s}$, $Co = 1$, $N\Delta t = 2$



Implicit gw, $\Delta t = 200\text{s}$, $Co = 2$, $N\Delta t = 4$



Resolution $\div 2$, $\Delta t = 500\text{s}$, $Co = 2.5$, $N\Delta t = 10$



Stable at long time-steps but accuracy is lost because θ advection is implicit rather than sub-stepped - needs sorting

Conclusions

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