

Modeling Arctic terrestrial processes and feedbacks

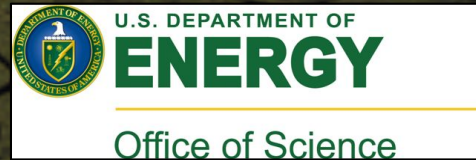
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Sean Swenson, Andrew Slater,
Chad Thackeray, Chris Fletcher,
Merritt Turetsky, Christina
Schaedel**

**National Center for Atmospheric
Research, Boulder, CO**

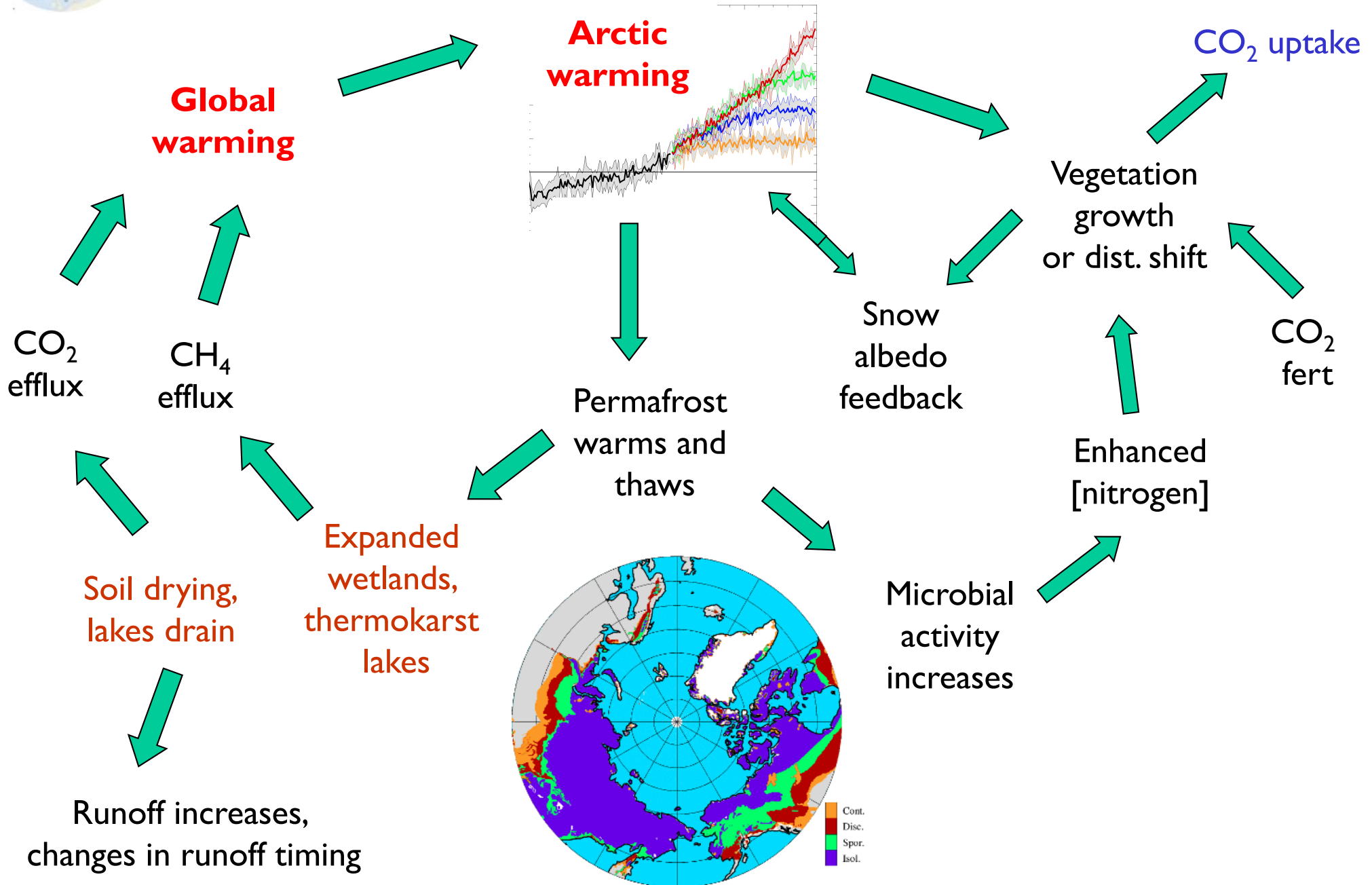


NCAR is sponsored by the National Science Foundation



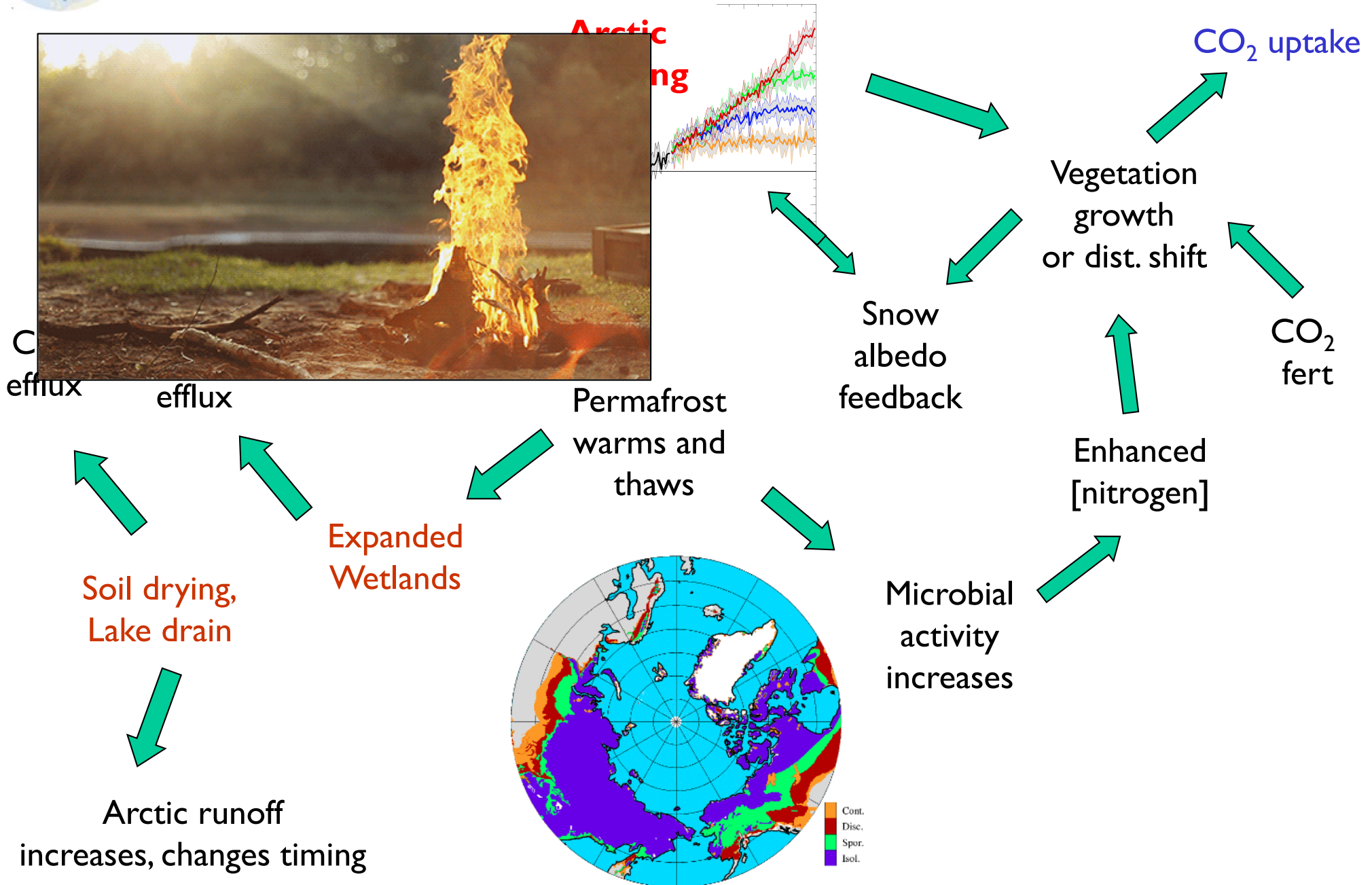


Arctic terrestrial climate-change feedbacks





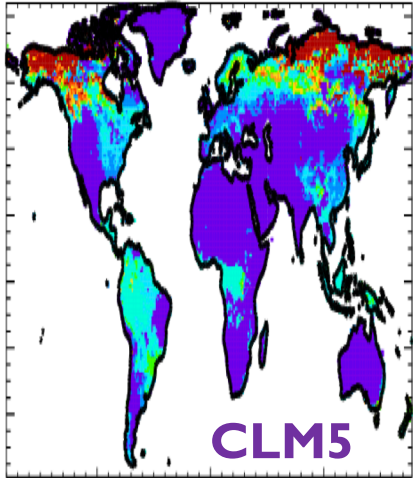
Arctic terrestrial climate-change feedbacks





Permafrost climate-carbon feedback

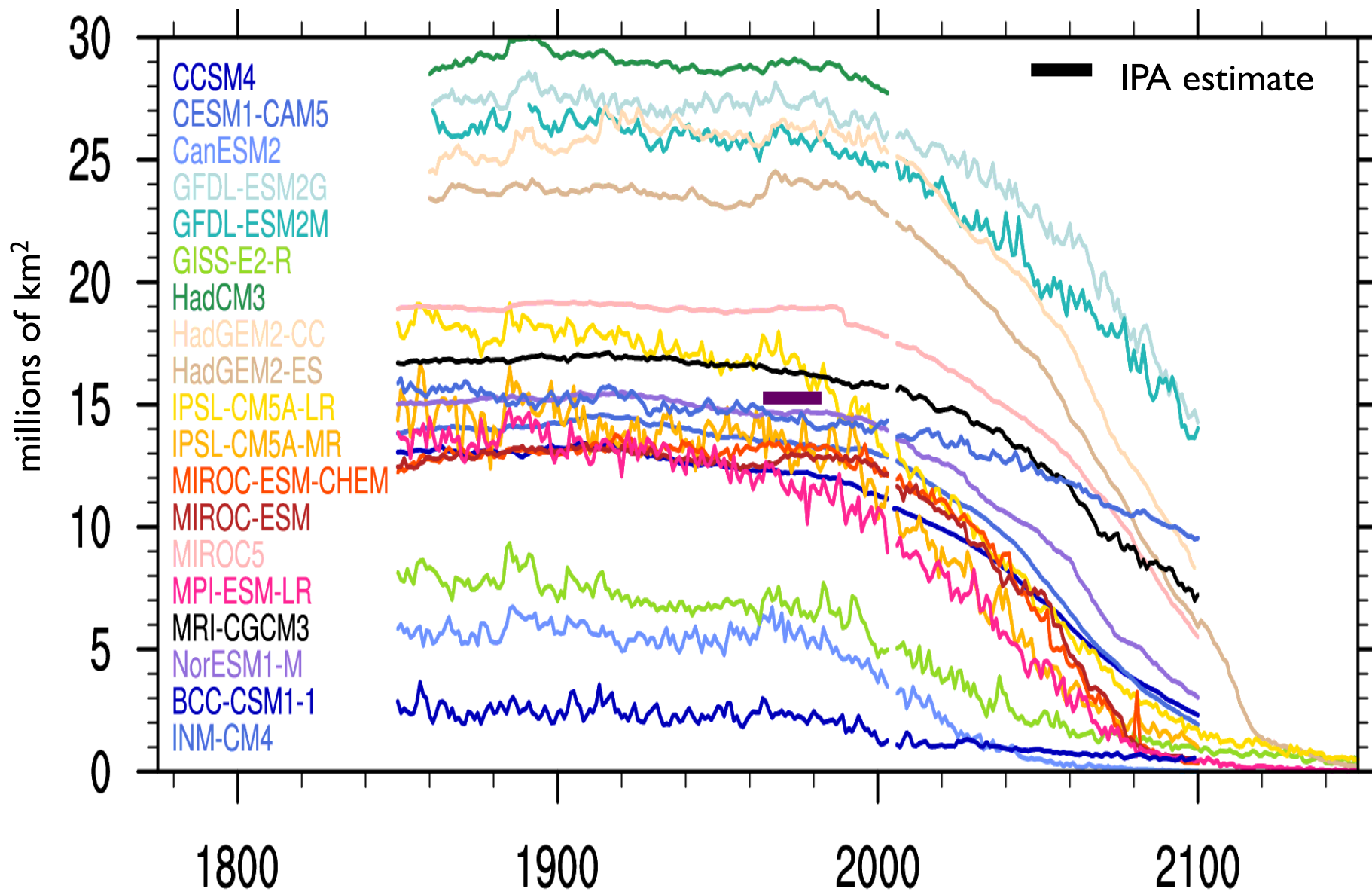
Ecosystem Carbon



- >1700 PgC stored in permafrost soils
- Substantial permafrost thaw projected, especially at high emission scenarios
- Permafrost climate-carbon feedback not represented in CMIP5 models

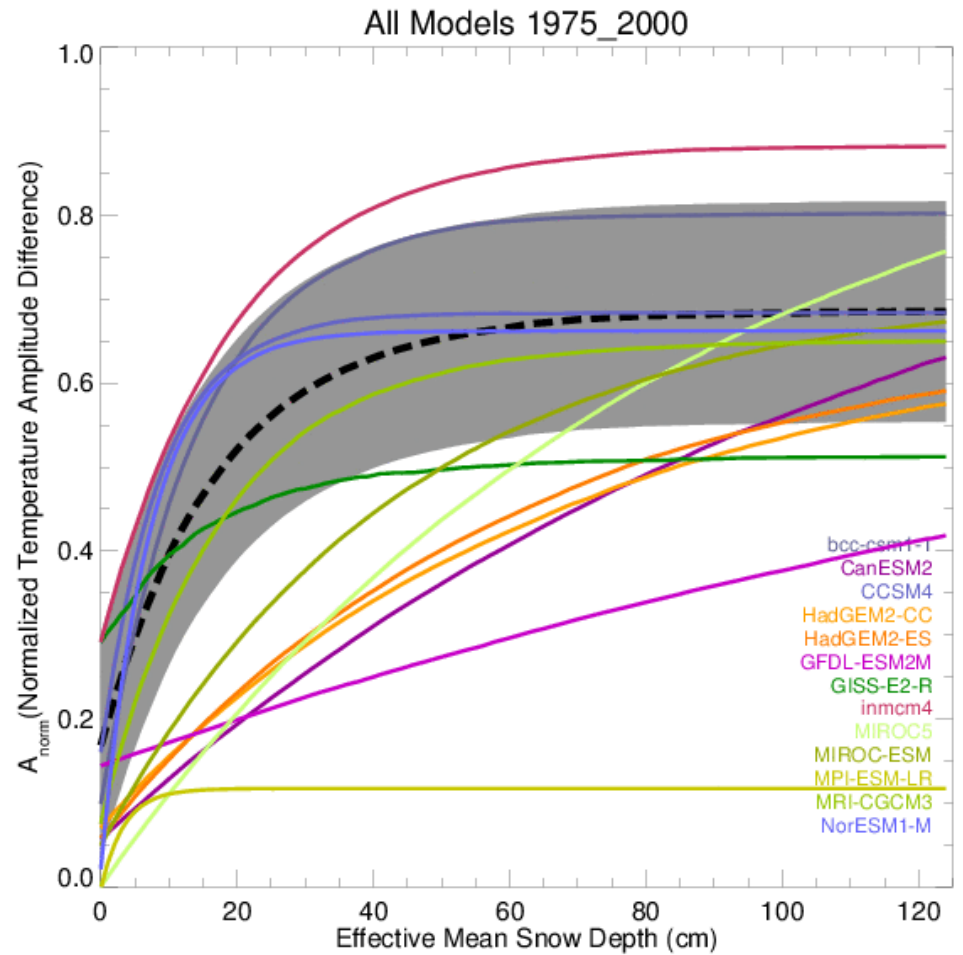
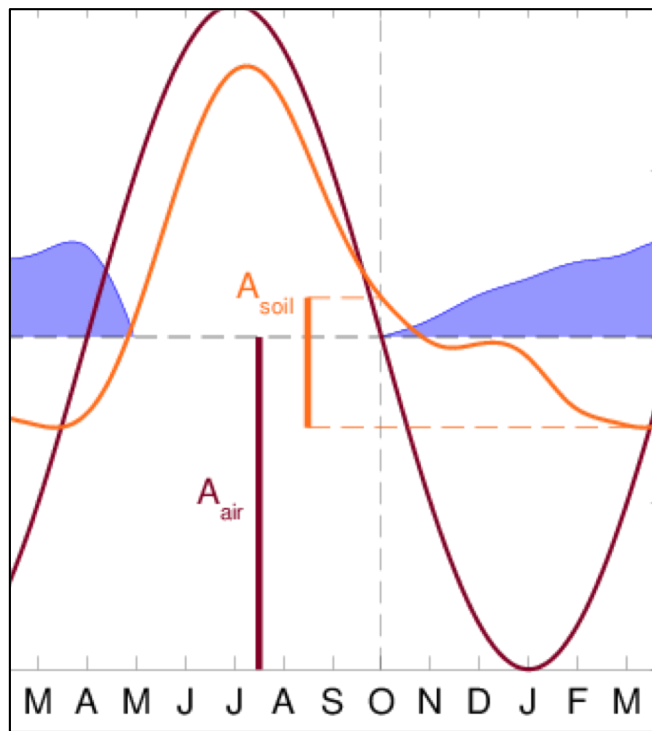


CMIP5 Models: Near-surface permafrost extent (RCP 8.5)





A snow heat transfer metric

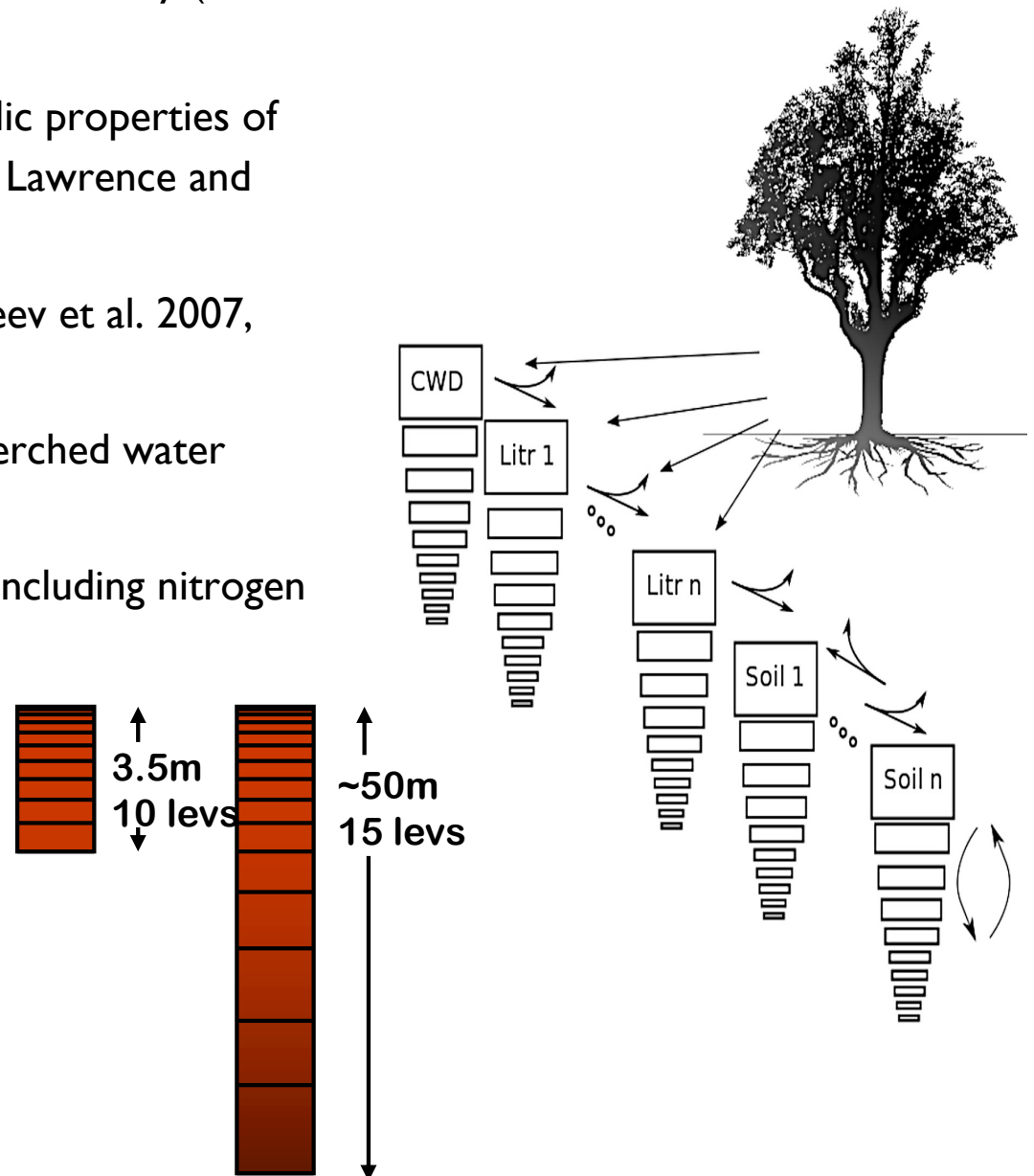


Many models do not correctly represent snow insulation



Key land model features for permafrost simulations

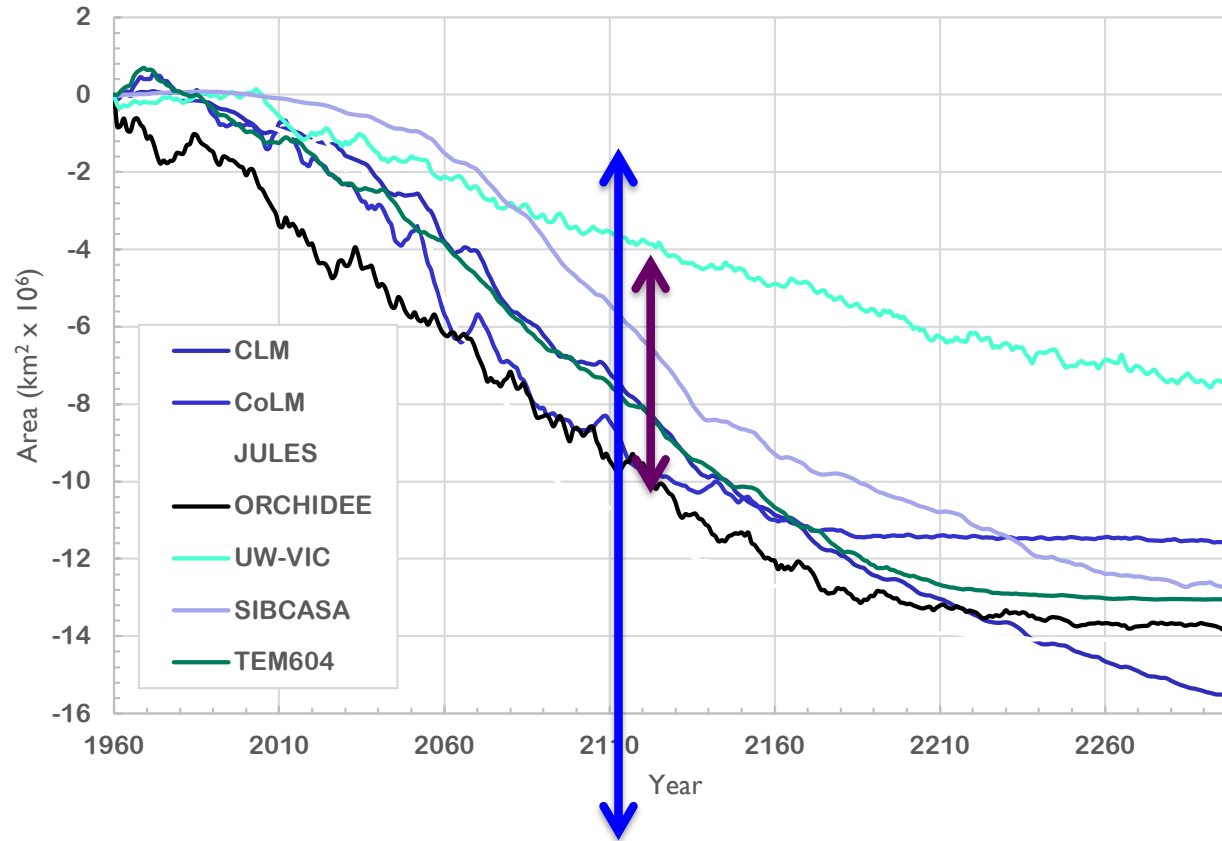
- Snow model that treats snow insulation reasonably (Koven et al. 2013)
- Explicit treatment of thermal and hydraulic properties of soil organic matter (Nicolsky et al. 2007, Lawrence and Slater, 2008)
- Deep ground column ~50m depth (Alexeev et al. 2007, Lawrence et al., 2008)
- Cold region hydrology, ice impedance, perched water table (Swenson et al. 2012)
- Vertically-resolved soil biogeochemistry including nitrogen (Koven et al. 2014)
- CH₄ emissions (Riley et al., 2013)
- Soil excess ice (Lee et al. 2015)





PCN: "Permafrost-enabled Model intercomparison"

Permafrost Area Loss 1960-2299 (RCP 8.5)

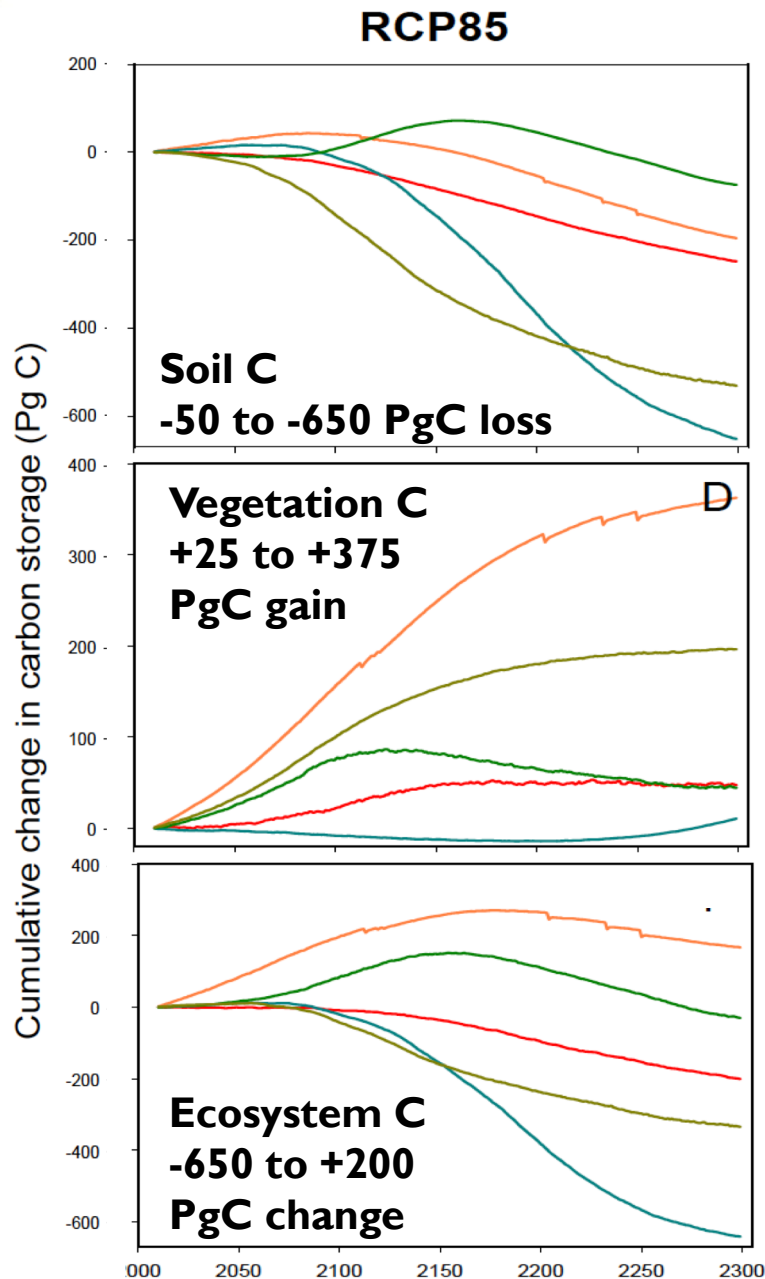


PCN 4 -10 million km²
CMIP5 1-18 million km²



PCN: “Permafrost Model intercomparison”

Diverse permafrost C predictions



Needs for permafrost-carbon feedback modeling

- Standardize structural representation of permafrost and carbon
- Develop data sets and methodologies to benchmark models
- Utilize models to assess sensitivities to processes
- Assess and represent C impact of permafrost thermokarst responses to warming (simple model estimates suggest +50% amplification of permafrost climate-carbon feedback)

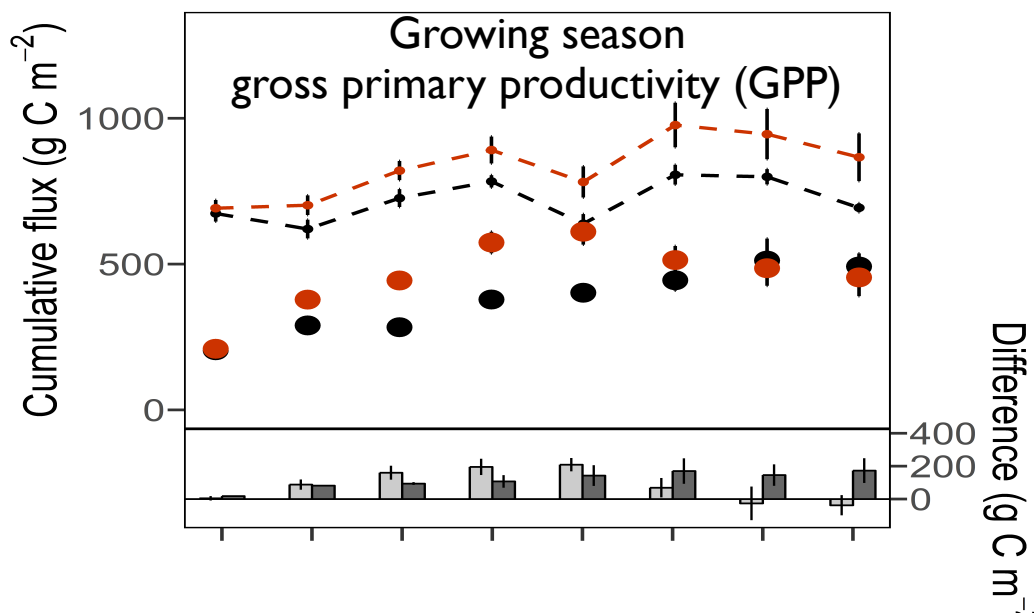
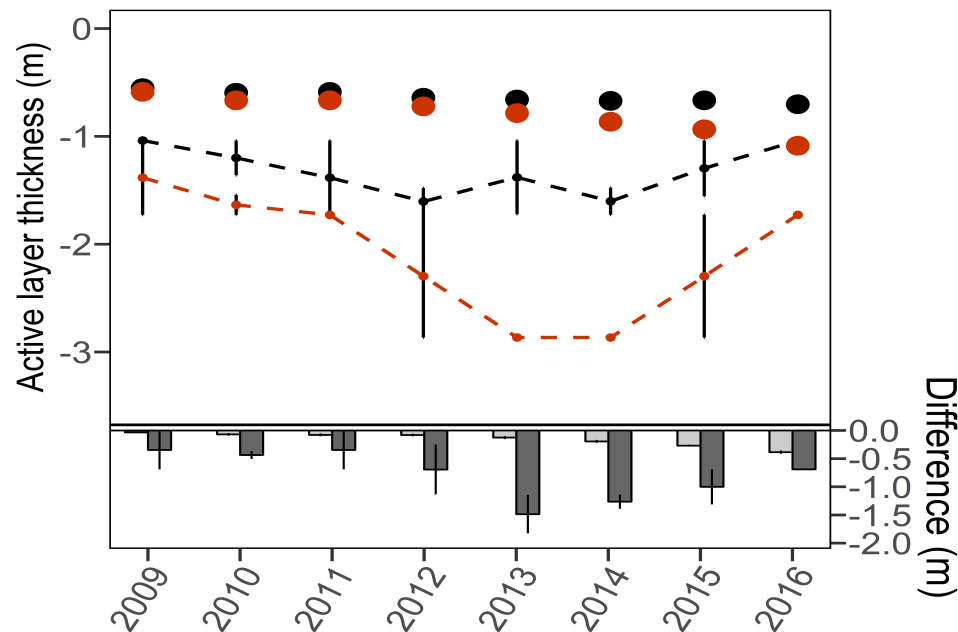


Benchmarking models against field experiments

Artificial warming
Snow fence experiment



- | | |
|--------------|---------------|
| Field | Model |
| ● Control | — Control |
| ● Warming | - - - Warming |

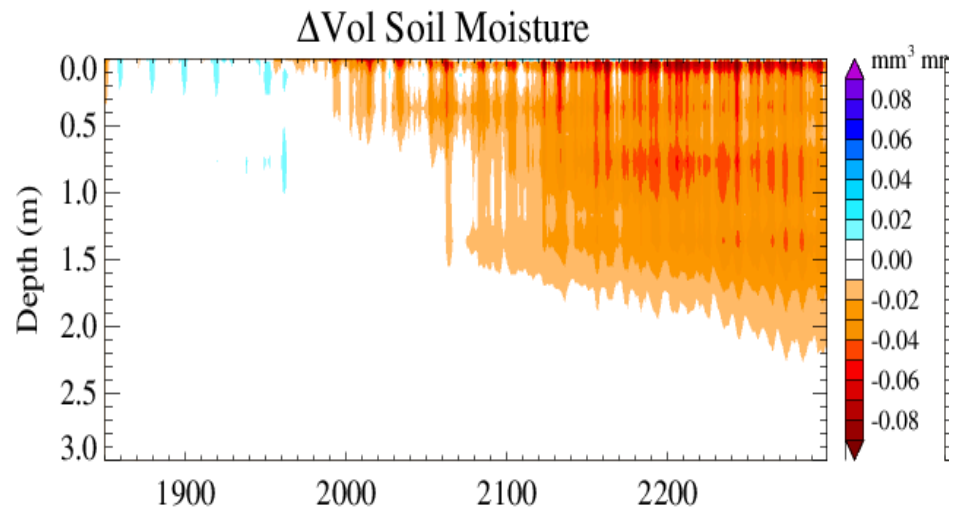
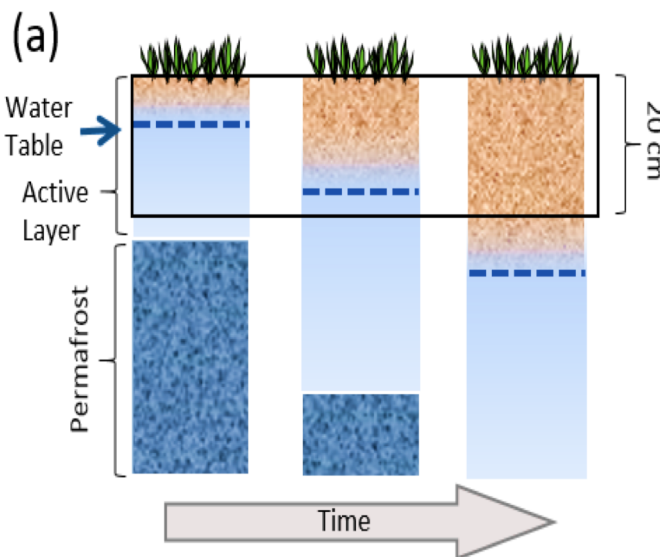
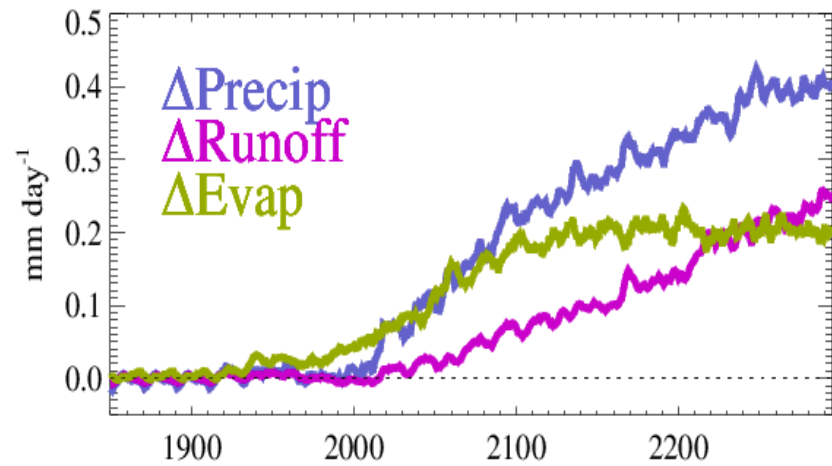
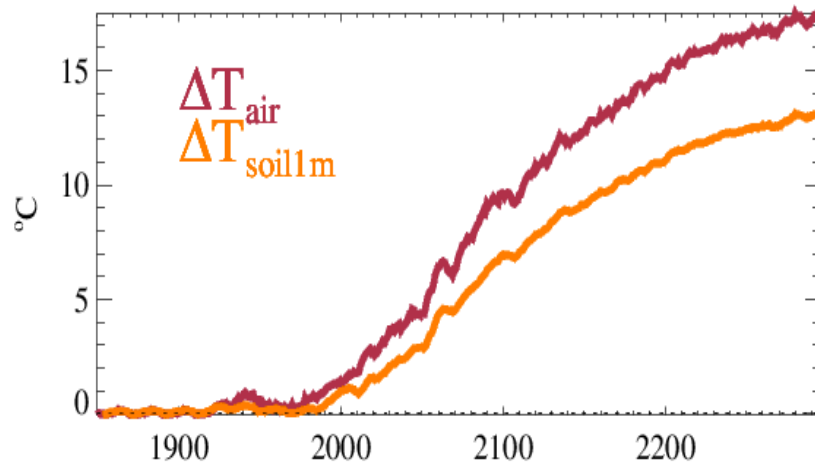




Using models to assess sources of uncertainty

Example: Uncertainty related to soil moisture projections

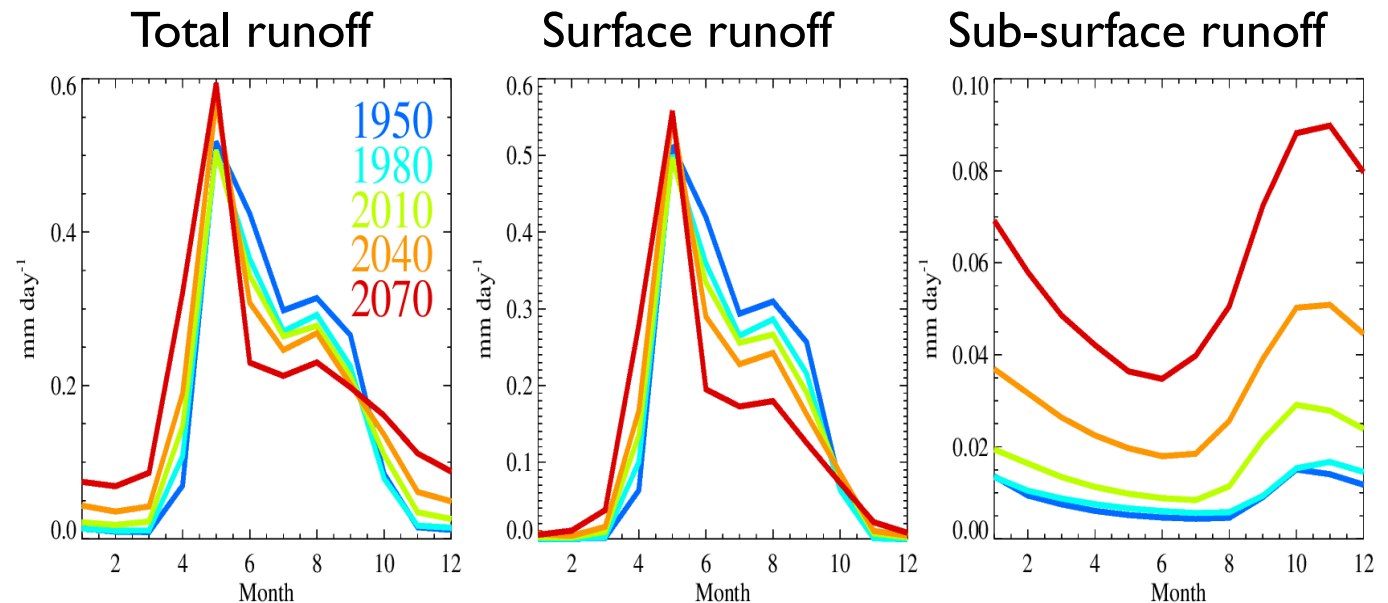
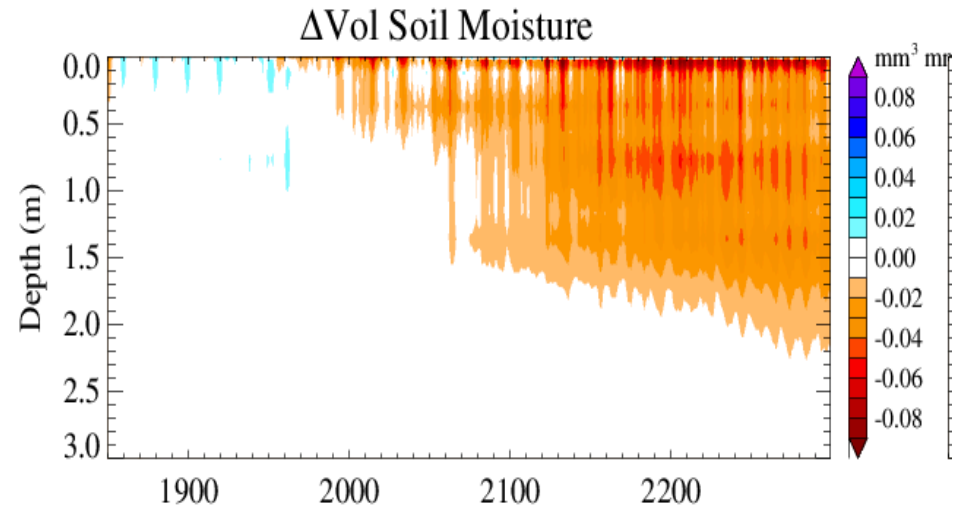
CESM Projections of temperature and water balance for permafrost domain (RCP8.5)





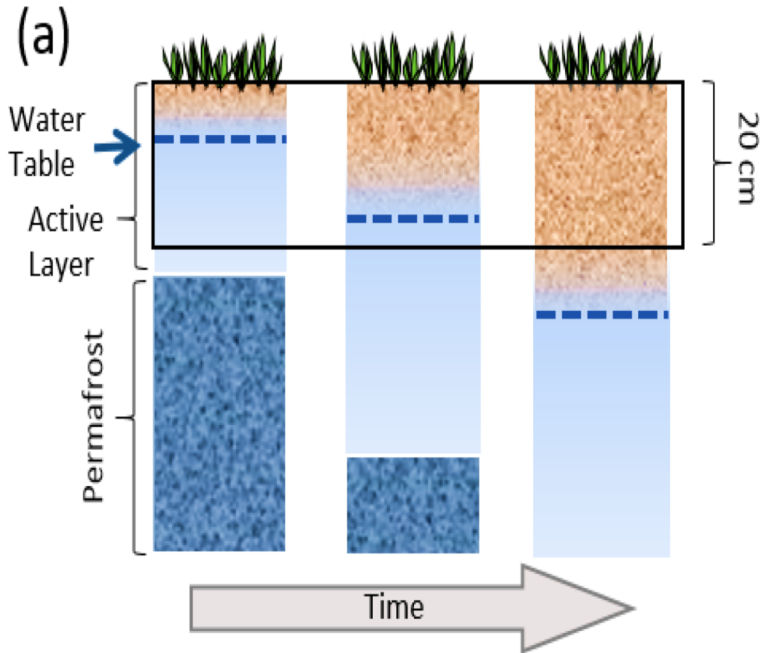
Permafrost-thaw driven transitions in runoff characteristics

- After permafrost thaw, transition to higher proportion baseflow
- Consistent with 'observations' and other hydrologic models (Walvoord and Striegl, 2007, Bense et al. 2009, Walvoord et al. 2012)
- High divergence in SM and runoff projections in PCN models (Andresen et al., in prep)

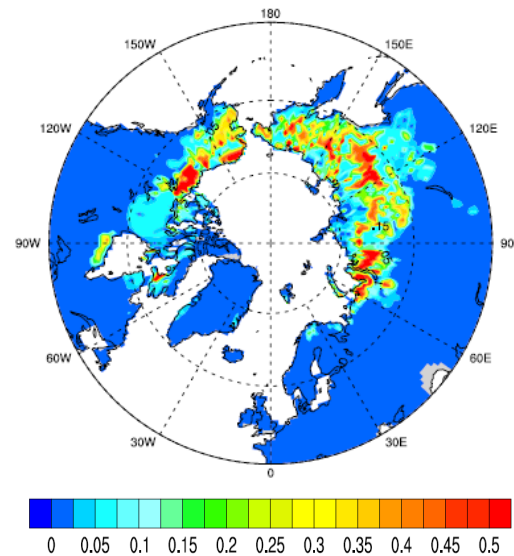




Active layer deepening and soil subsidence



CLM projection of subsidence by 2100

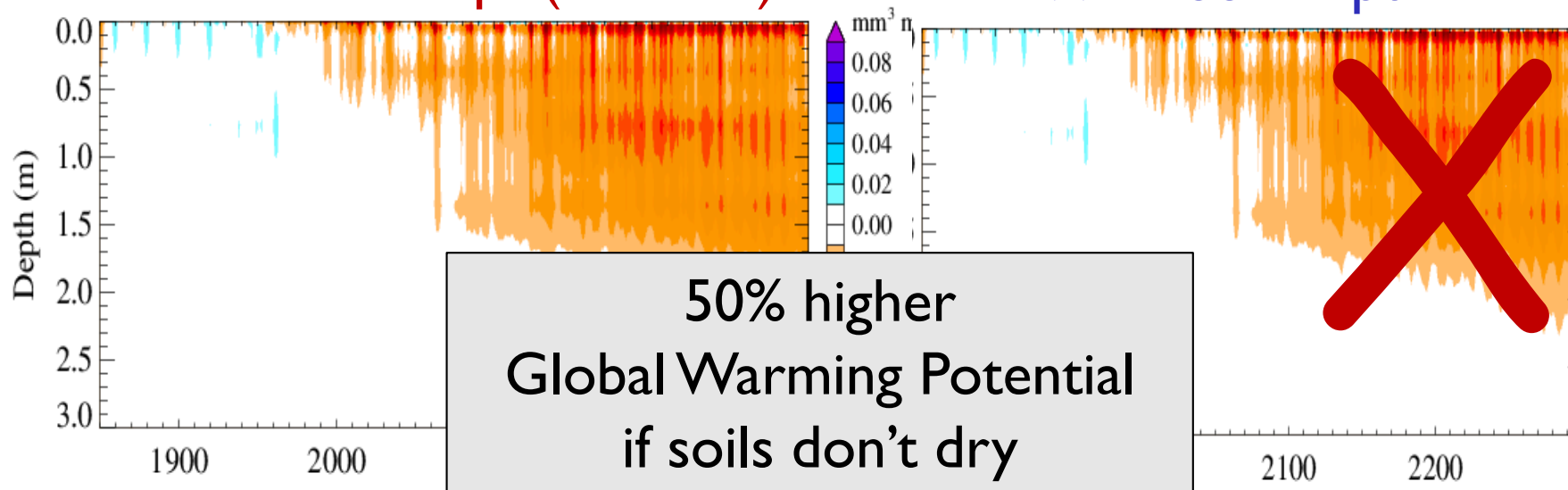




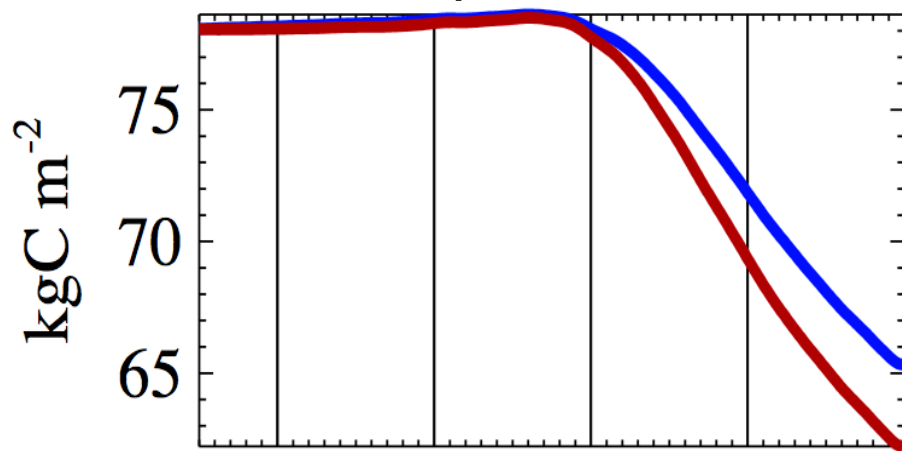
Permafrost carbon-climate feedback with and without soil drying

DRY Soil Expt (Control)

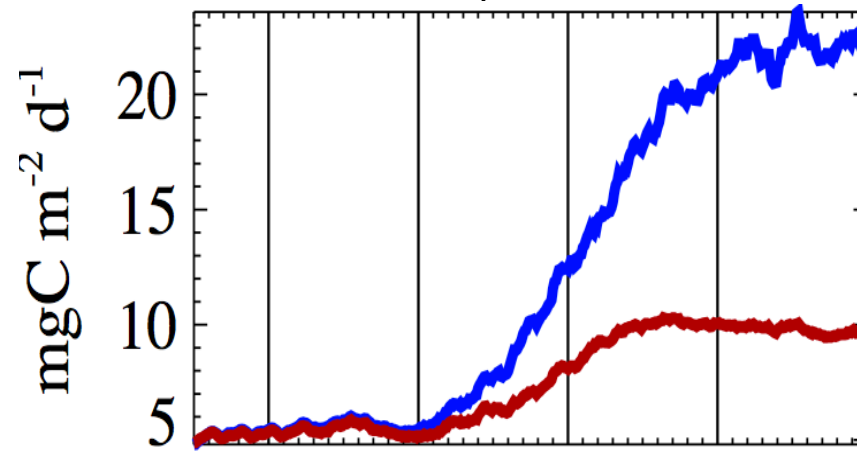
WET Soil Expt



Ecosystem Carbon



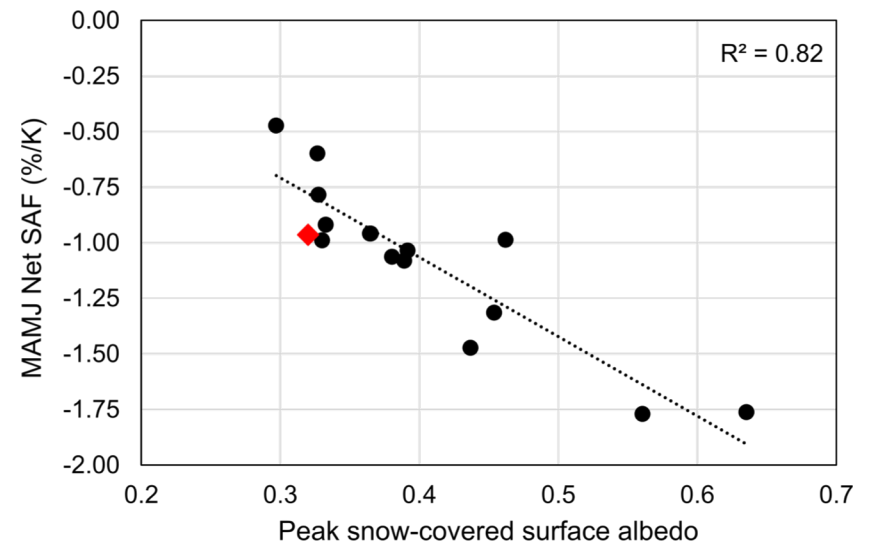
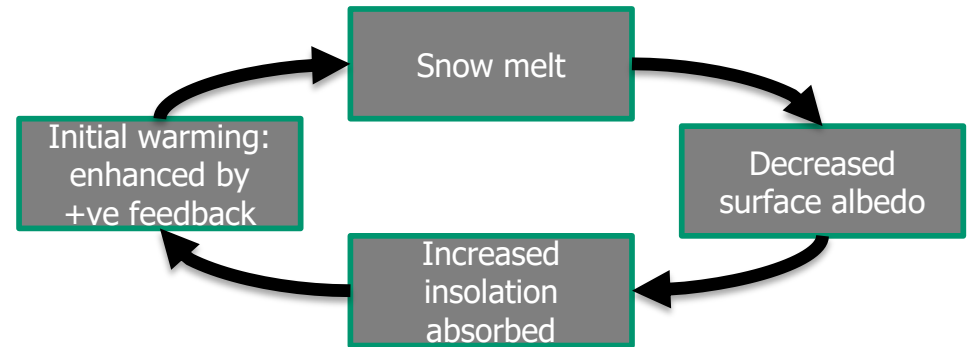
CH₄ emissions





Snow Albedo Feedback (SAF)

- SAF is a positive feedback climate mechanism and important driver of regional climate change
- Models exhibit large variability SAF
- Intermodel spread in SAF explains 40-50% of the CMIP5 variability in projected spring NH land warming.
- Much of the spread in SAF can be explained by differences in simulated maximum snow-covered surface albedo and the timing of the spring albedo transition



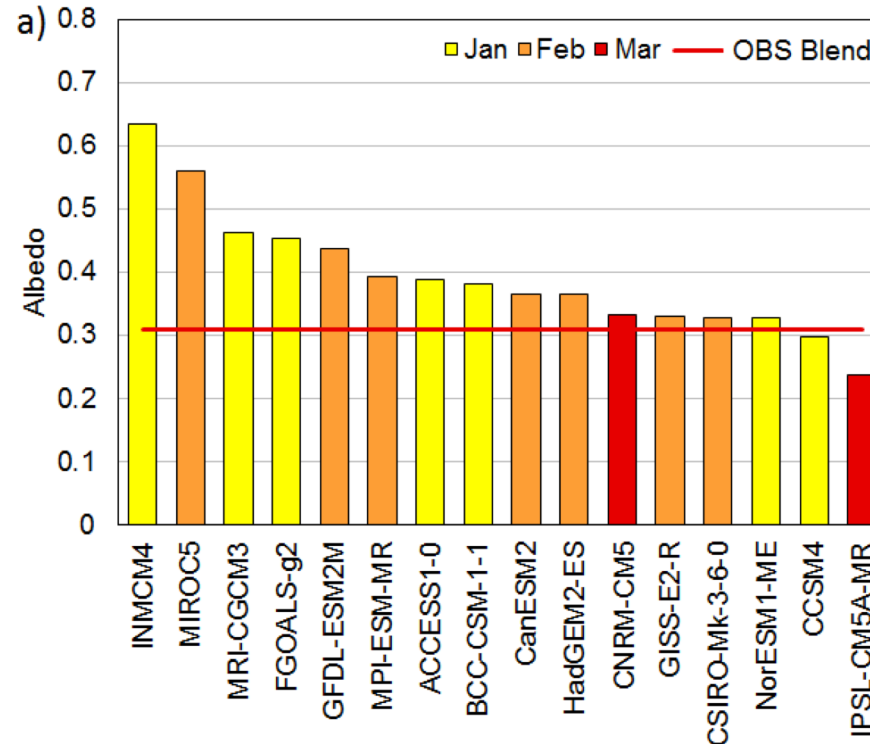
Relationship between peak snow-covered surface albedo and spring SAF from models (black) and OBS (red) across the boreal forest.



Large biases in snow-covered surface albedo

Max surface albedo

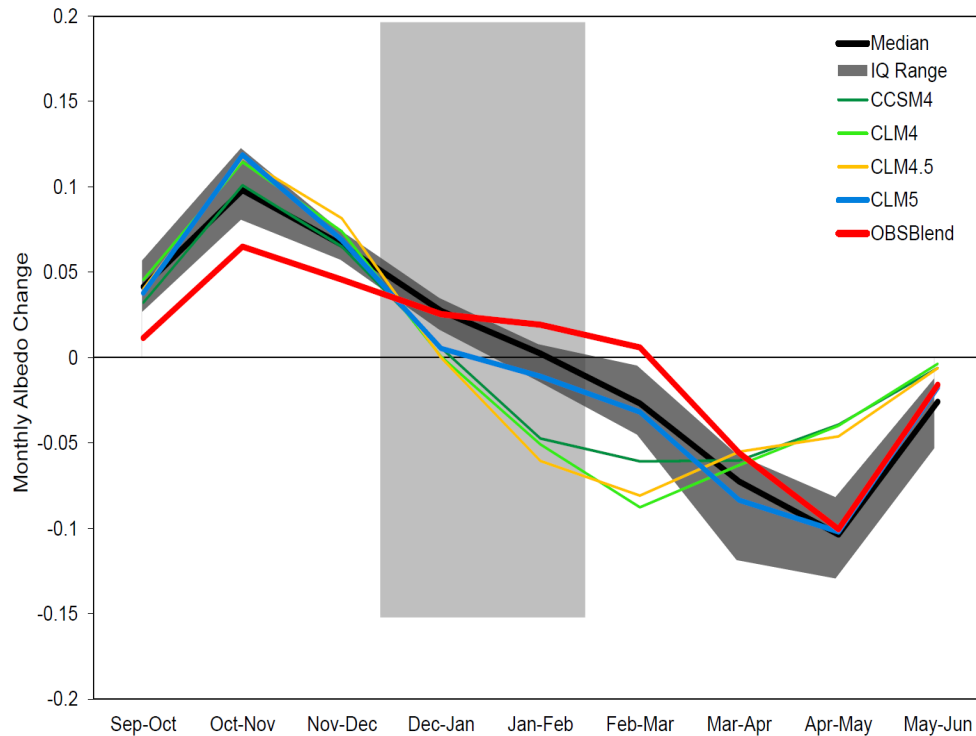
Boreal Forest



- Many climate models struggle to capture **timing** and/or **magnitude** of seasonal changes in albedo over boreal forest and Arctic tundra regions
- CCSM4: albedo decreases too early → weak SAF.



Reduction of SAF bias in CMIP6?



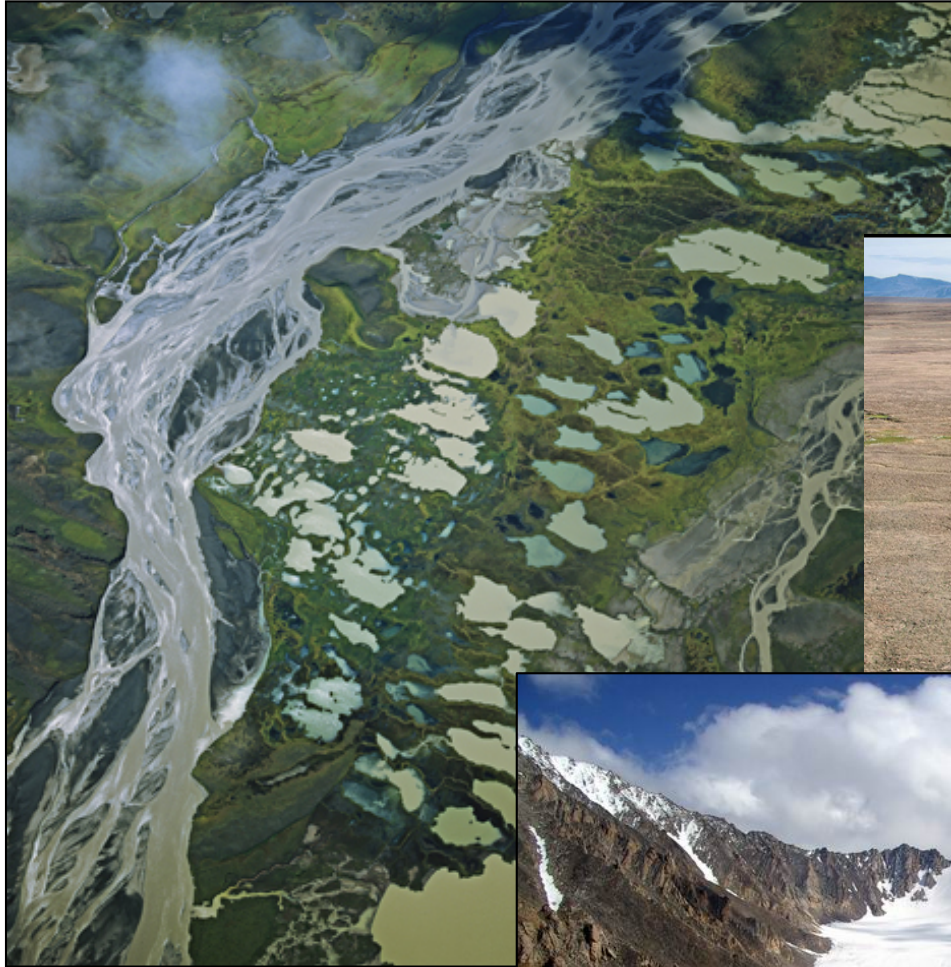
Monthly climatological albedo change across the boreal forest. The light gray box shows when observational uncertainty is largest.

Model	Boreal Spring SAF (%/K)
CCSM4	-0.60
CLM4	-0.64
CLM4.5	-0.68
CLM5	-0.83
MODIS	-0.87

- New canopy snow storage and unloading scheme reduced bias in seasonality of snow-covered surface albedo and thus, SAF
- Cautious expectation for reduced bias in SAF in CMIP6 models
- Snow-MIP to address snow-climate interactions



Challenge of heterogeneity

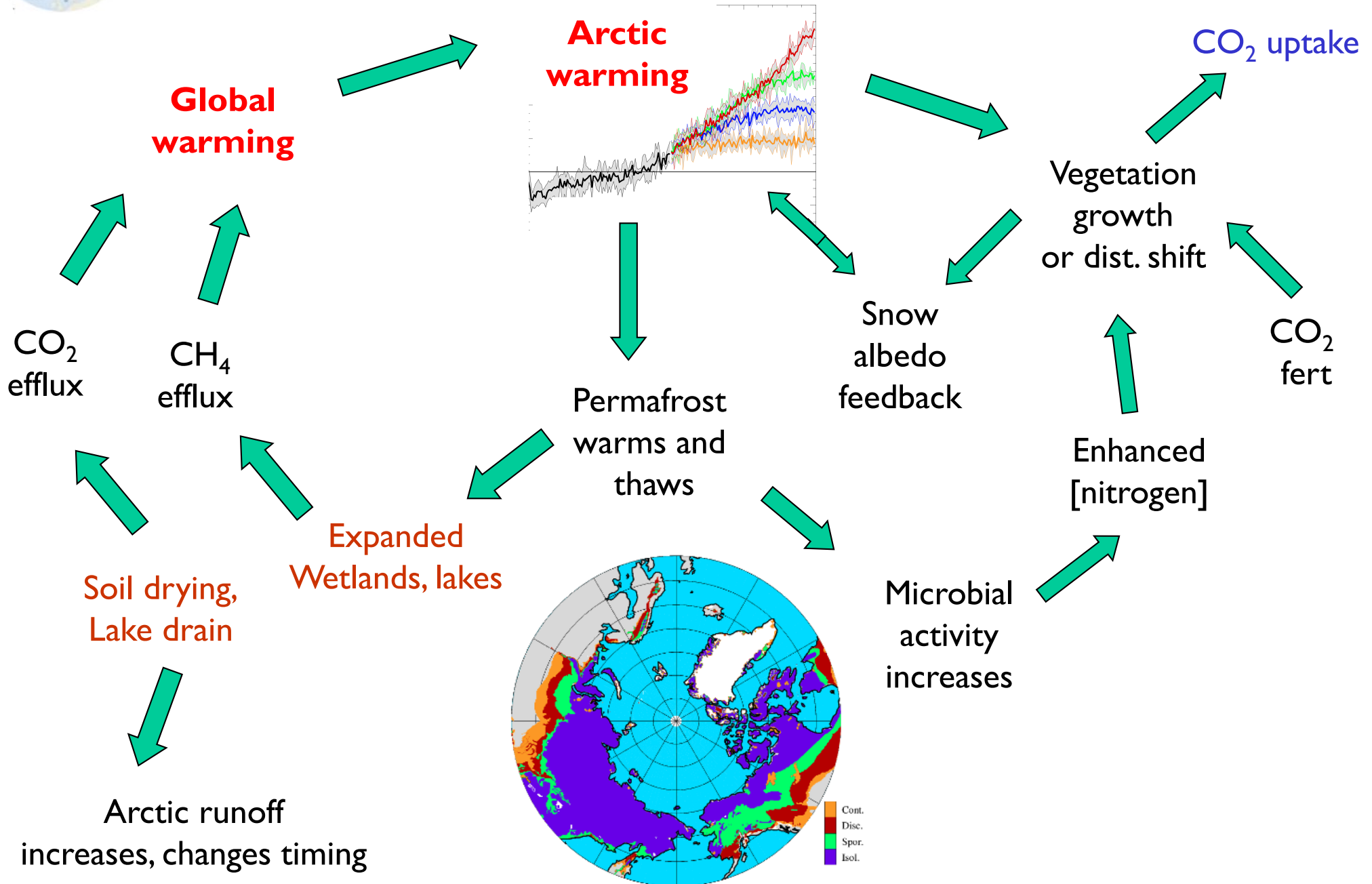


How much do unrepresented heterogeneous land responses to environmental change affect the strength of the overall feedbacks?





Potential Arctic terrestrial climate-change feedbacks





Permafrost in CMIP6 models?

Process/Model	CESM	GFDL	UKESM	MPI-ESM	IPSL	NorESM	EC-Earth
Permafrost physics	on	on	on	offline	on	on	on
Permafrost C	on	?	no	offline	offline	on	offline
CH₄ emissions	on	?	on	offline	offline	on	offline
CN interaction	on	on	on	on	on	on	on



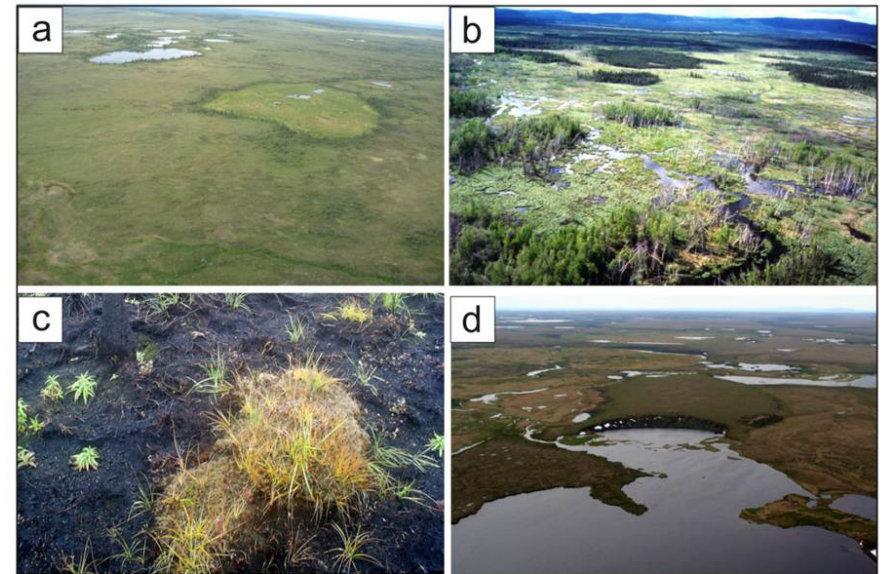
The challenge of heterogeneity

Example: Impact of thermokarst processes on permafrost C dynamics

Contrary to ‘top-down’ thaw, thermokarst processes can tap into deep permafrost C, resulting in rapid C release

Estimating magnitude of C loss due to ‘thermokarst’ response to warming

- (1) Define areas vulnerable to thermokarst processes
- (2) Document current extent of “thermokarst” features
- (3) Analyze recent trends in thermokarst processes
- (4) Assess impacts of thermokarst processes on landscape transitions and C dynamics
- (5) Initial assessment suggests that thermokarst could amplify permafrost climate-carbon feedback by 50%**



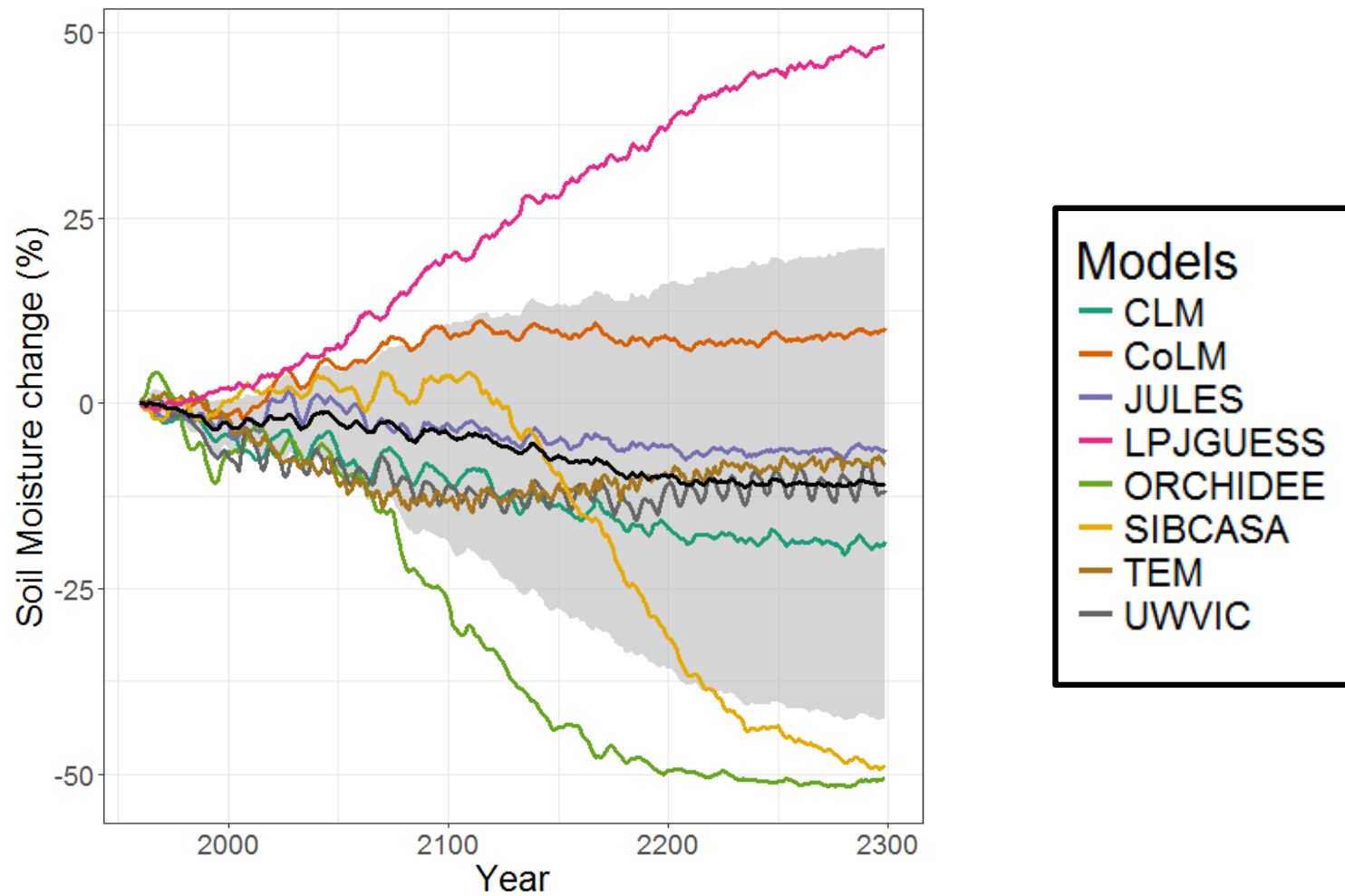
Thermokarst is subsidence of the surface that is caused by the melting of ground ice leading to fens/bogs, thermokarst lakes, thaw slumps, etc



- SAF spread was not reduced from CMIP3 to CMIP5 despite considerable land model development - largely due to shortcomings from two models.
 - The largest SAF biases arise because of structural errors relating to the distribution/type of vegetation or the parameterization of surface albedo (i.e. vegetation masking of surface) rather than parametric errors.
- Preliminary signs from ongoing model development are positive and suggest a likely improvement in SAF among most existing models.
- However, failure to update structural errors in a couple of models will likely limit the amount of reduction in SAF spread across the CMIP6 models. This drawback may further be exacerbated by the participation of a considerable amount of new modeling centers in CMIP6.
- Therefore, the extensive land model development undergone in many modeling centers may not achieve a great reduction in SAF spread across the CMIP6 models. To this cause, concerted efforts by the whole community are needed (e.g., ESM-SnowMIP).



High uncertainty in permafrost-domain soil moisture projections in PCN models

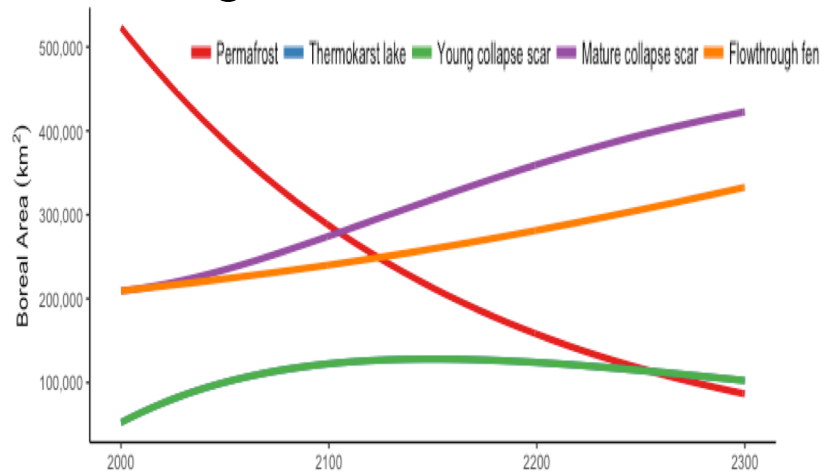




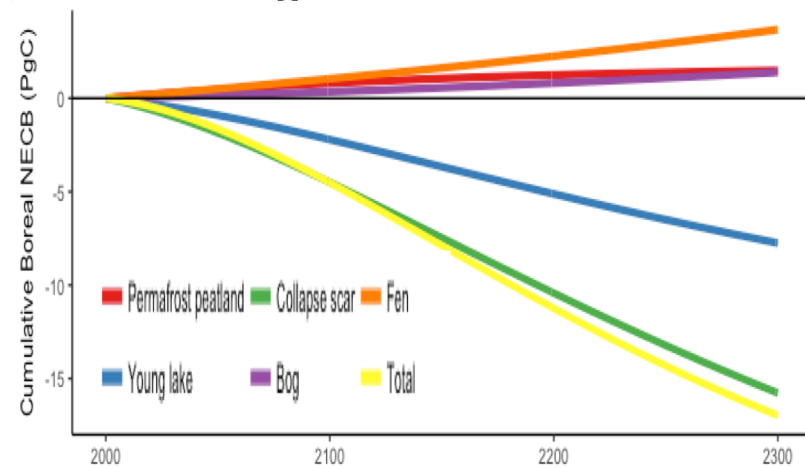
Thermokarst “state-and-transition” conceptual modeling

Ex. for Lowland Organic Terrain (Wetlands)

Change in area of successional state



Change in C balance



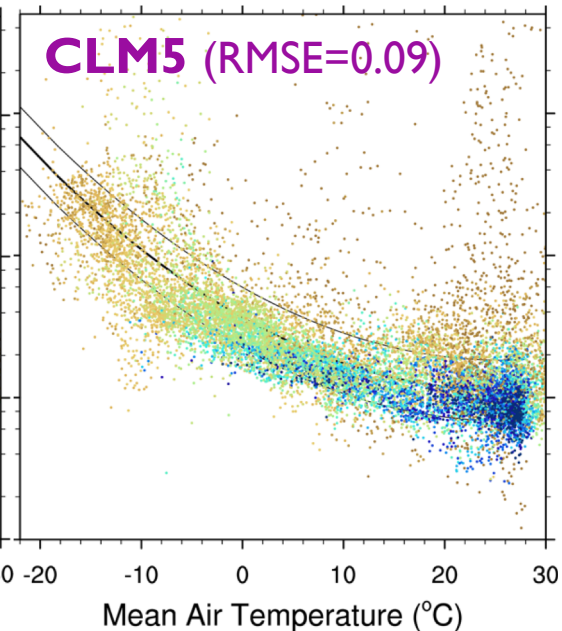
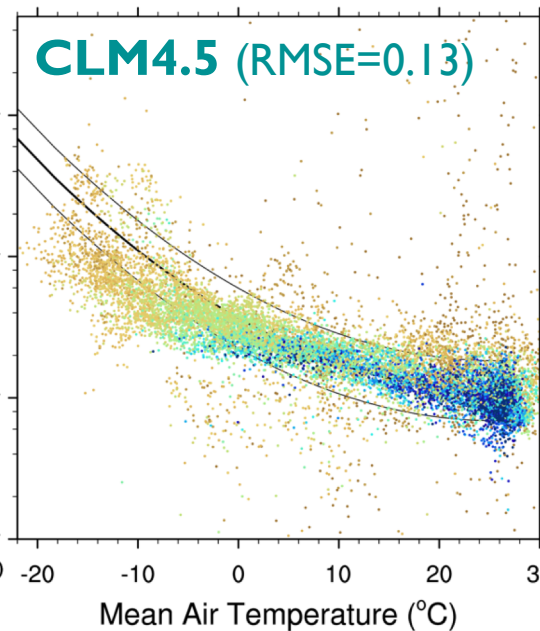
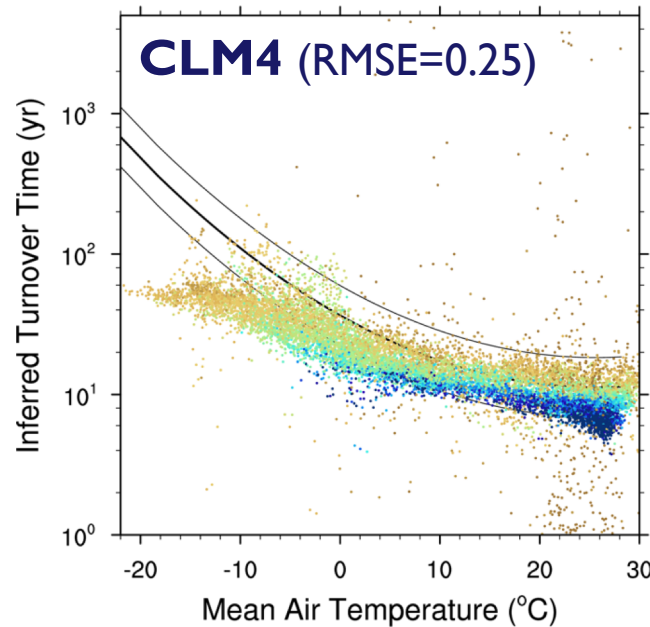
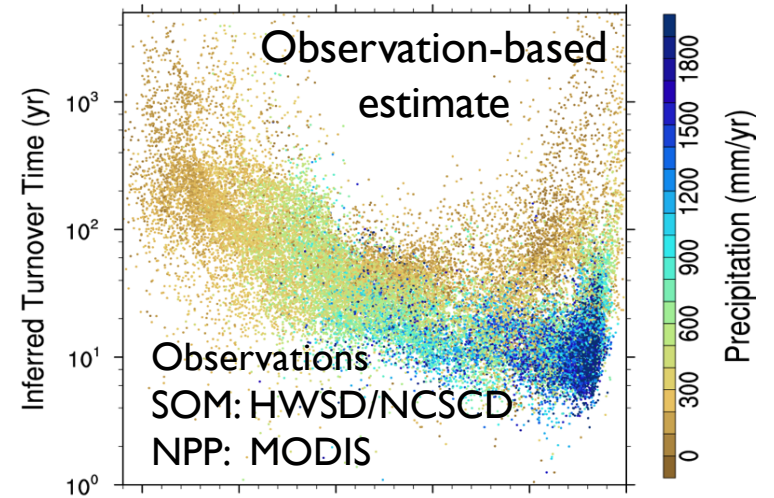
- Summed for all thermokarst processes, Global Warming Potential due to thermokarst ~50% of that due to ALT deepening
- Feedbacks under warming climate not captured by state-and-transition approach
- Challenge: integrate thermokarst parameterizations into ecosystem models



Metric for soil carbon turnover timescale

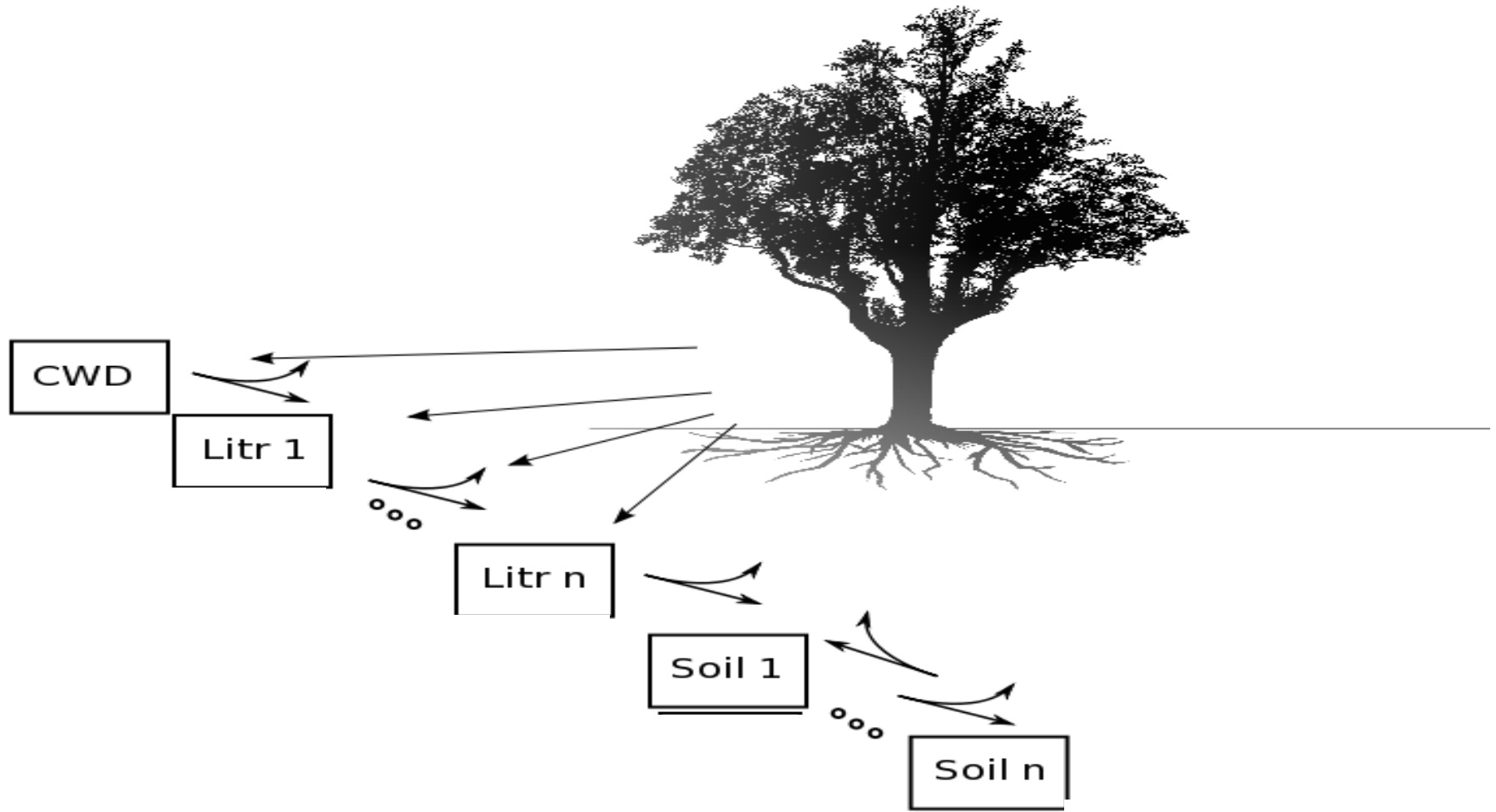
Inferred soil carbon turnover timescale

$$\tau = \frac{\text{carbon stocks (SOM)} (gC)}{\text{carbon inputs (NPP)} (\frac{gC}{s})}$$

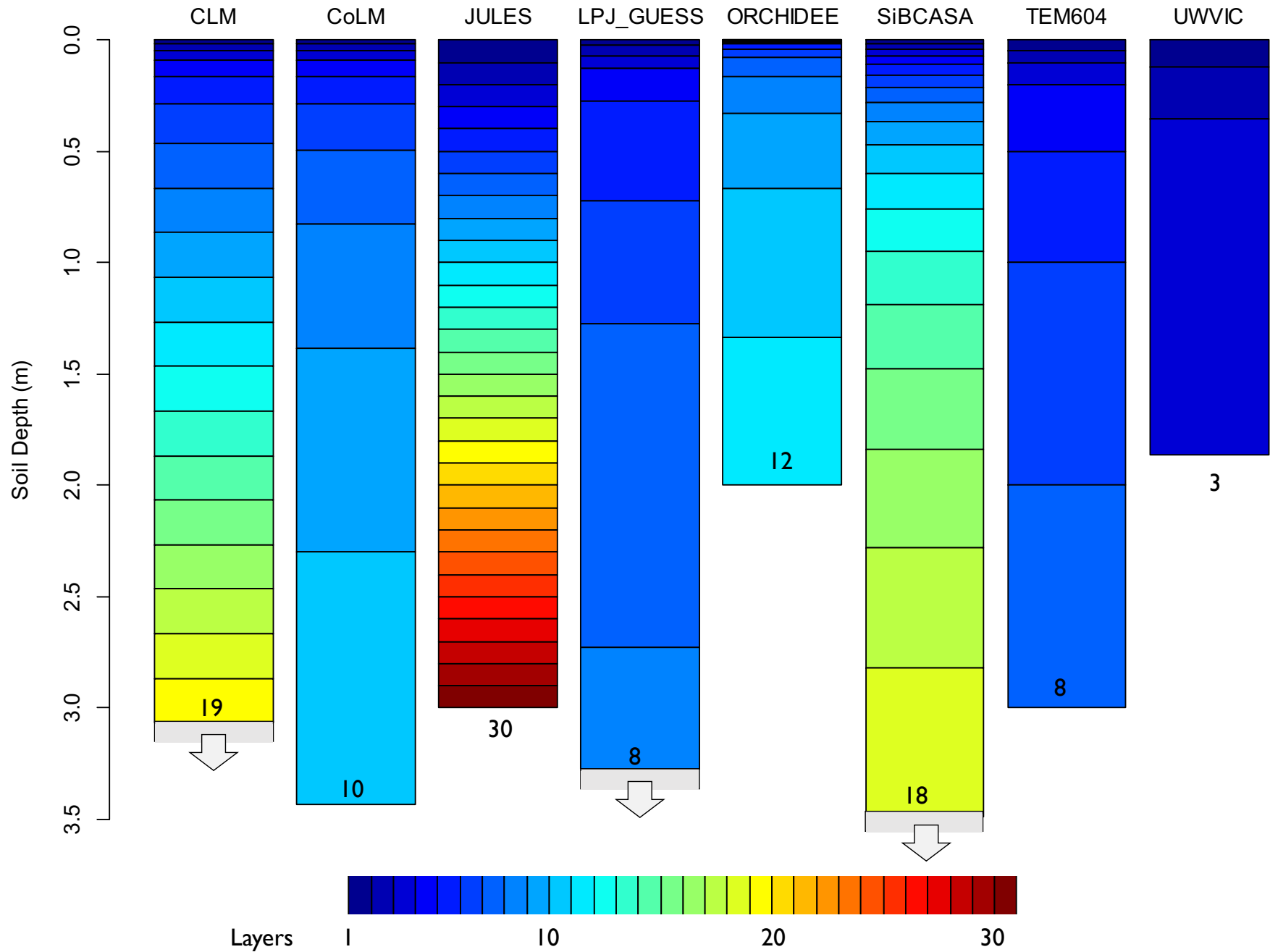




Vertically-resolved soil biogeochemistry



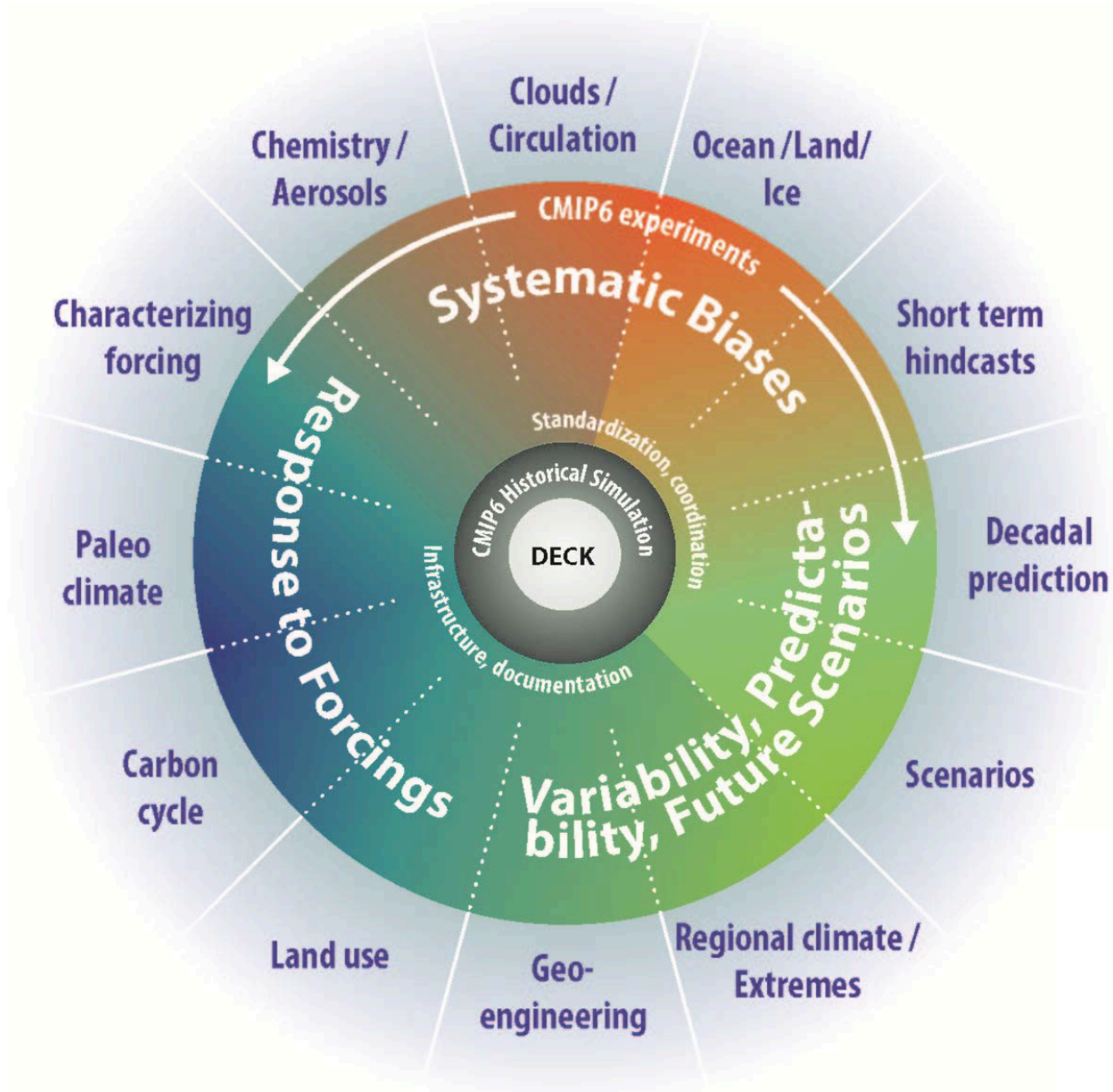
Diverse soil column configuration among land models



Increased focus on terrestrial processes in CMIP6

Coordinated activities to assess land role in climate and climate change

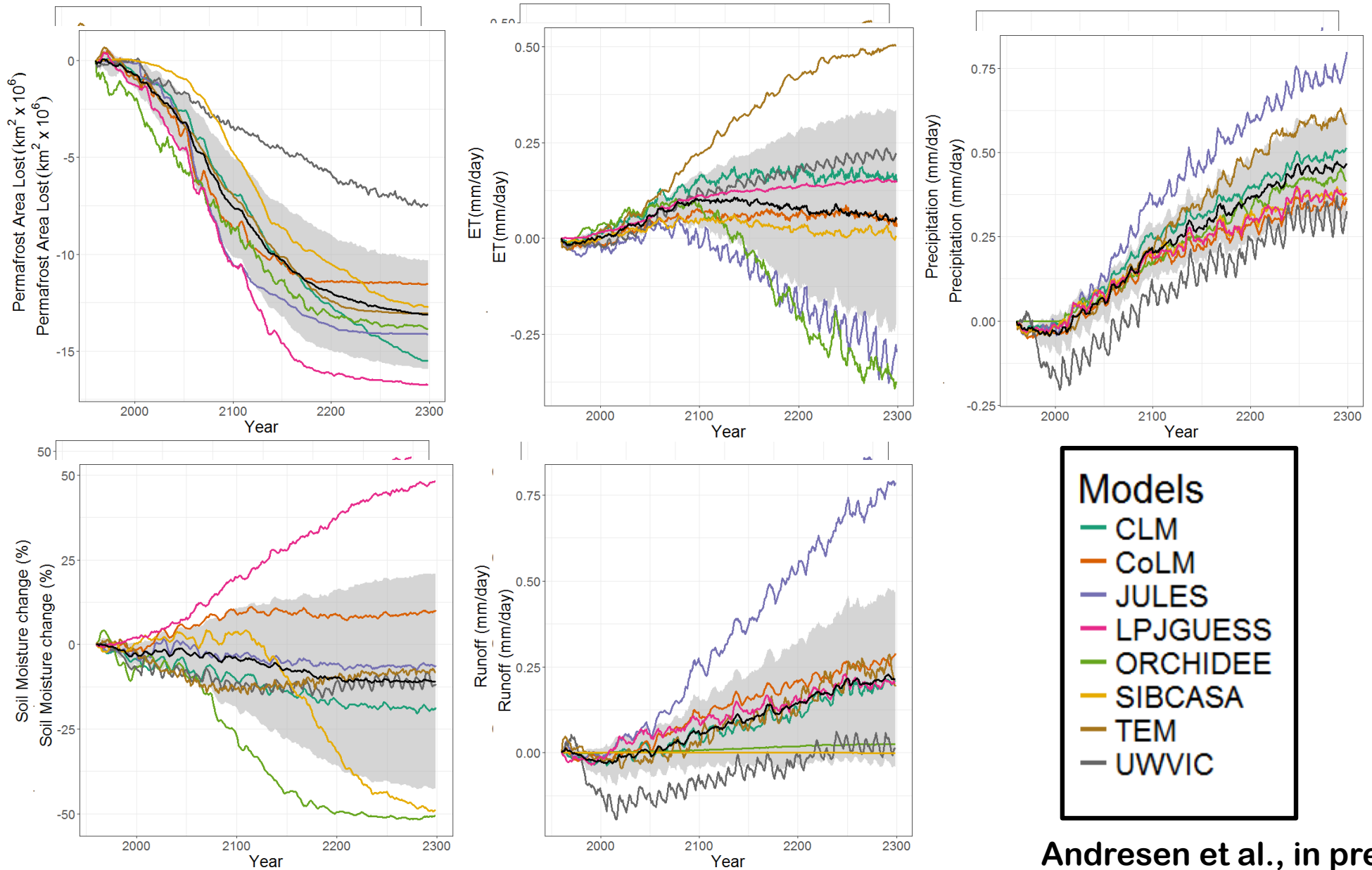
- **Land-only simulations forced with obs historical climate and common future**, land-systematic biases
- **Land Use = LUMIP**
land use forcing on climate, biogeophysics and biogeochemistry with policy relevance
- **Land = LS3MIP**
biogeophys feedbacks including soil moisture and snow feedbacks
- **Carbon Cycle = C4MIP**
land biogeochemical feedbacks on climate, **emissions-driven SSP5-8.5 21st and Extension to 2300**



Updated from Meehl et al., EOS, 2014



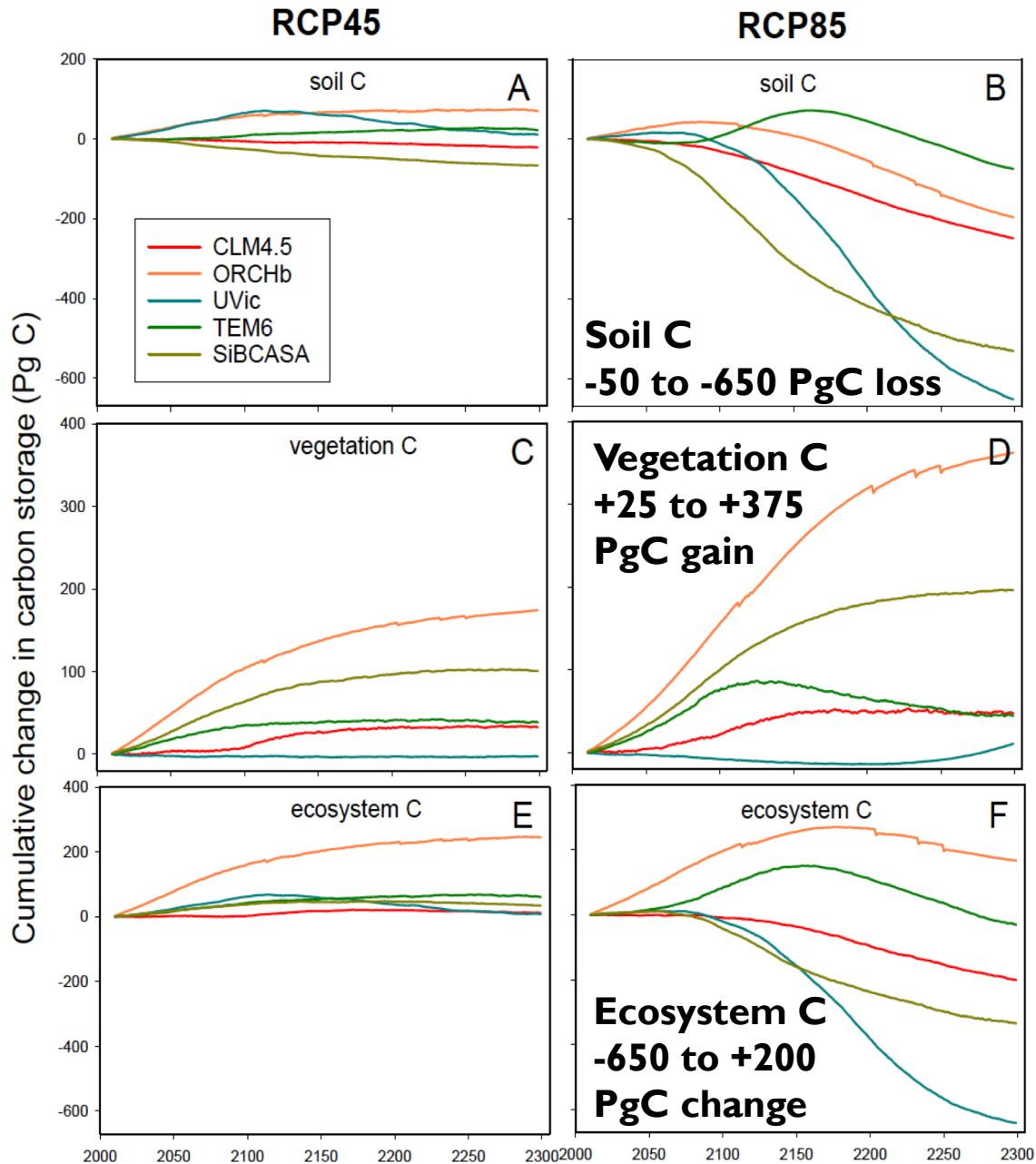
High uncertainty in permafrost-domain soil moisture projections





PCN: “Permafrost Model intercomparison”

Diverse permafrost loss predictions



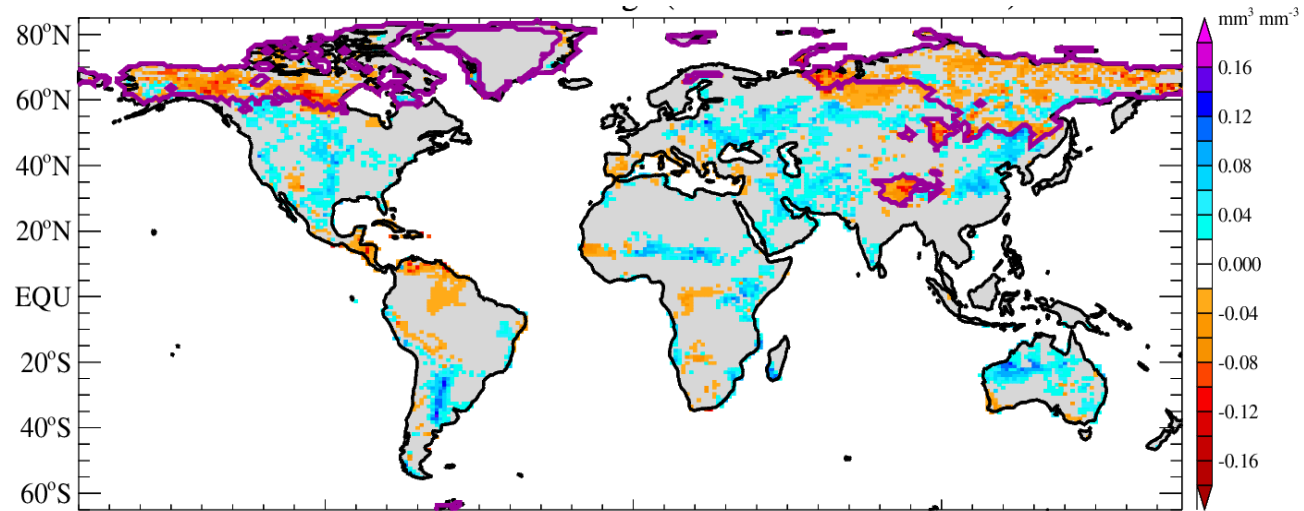
Needs for permafrost-carbon feedback modeling

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- Develop data sets and methodologies to benchmark models
- Utilize models to assess sensitivities to processes
- Assess and represent C impact of permafrost thermokarst responses to warming

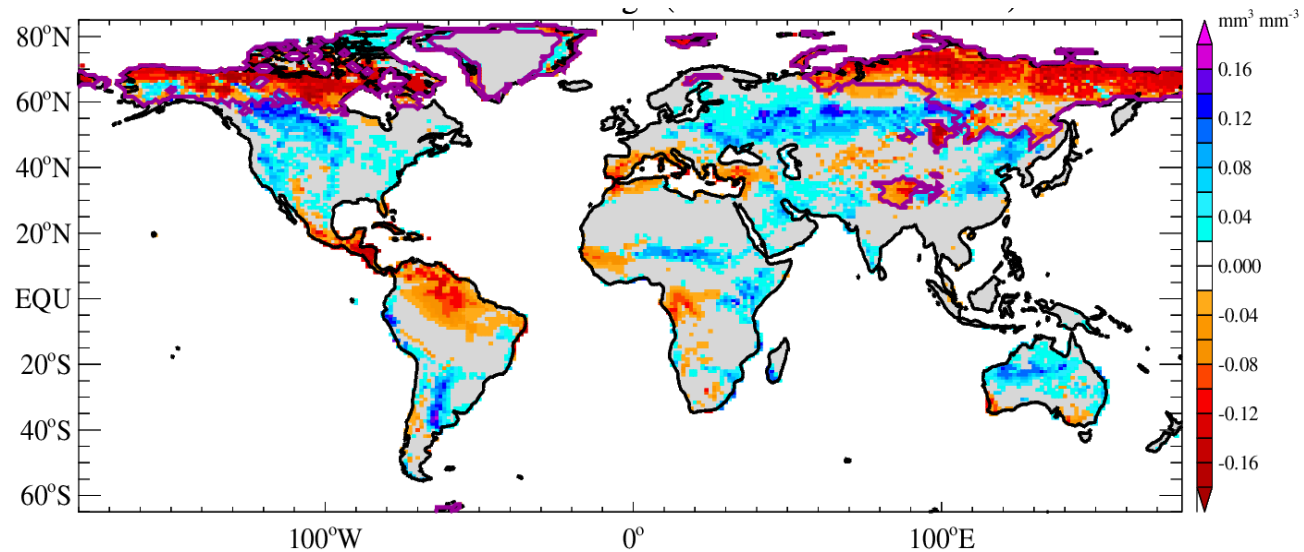


Projected soil moisture change (RCP8.5) CLM4.5

Column soil moisture change by 2100



Column soil moisture change by 2300





CLM representation of permafrost hydrology

