



Hydrology in the Community Land Model

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Terrestrial Sciences Section





What is Hydrology?

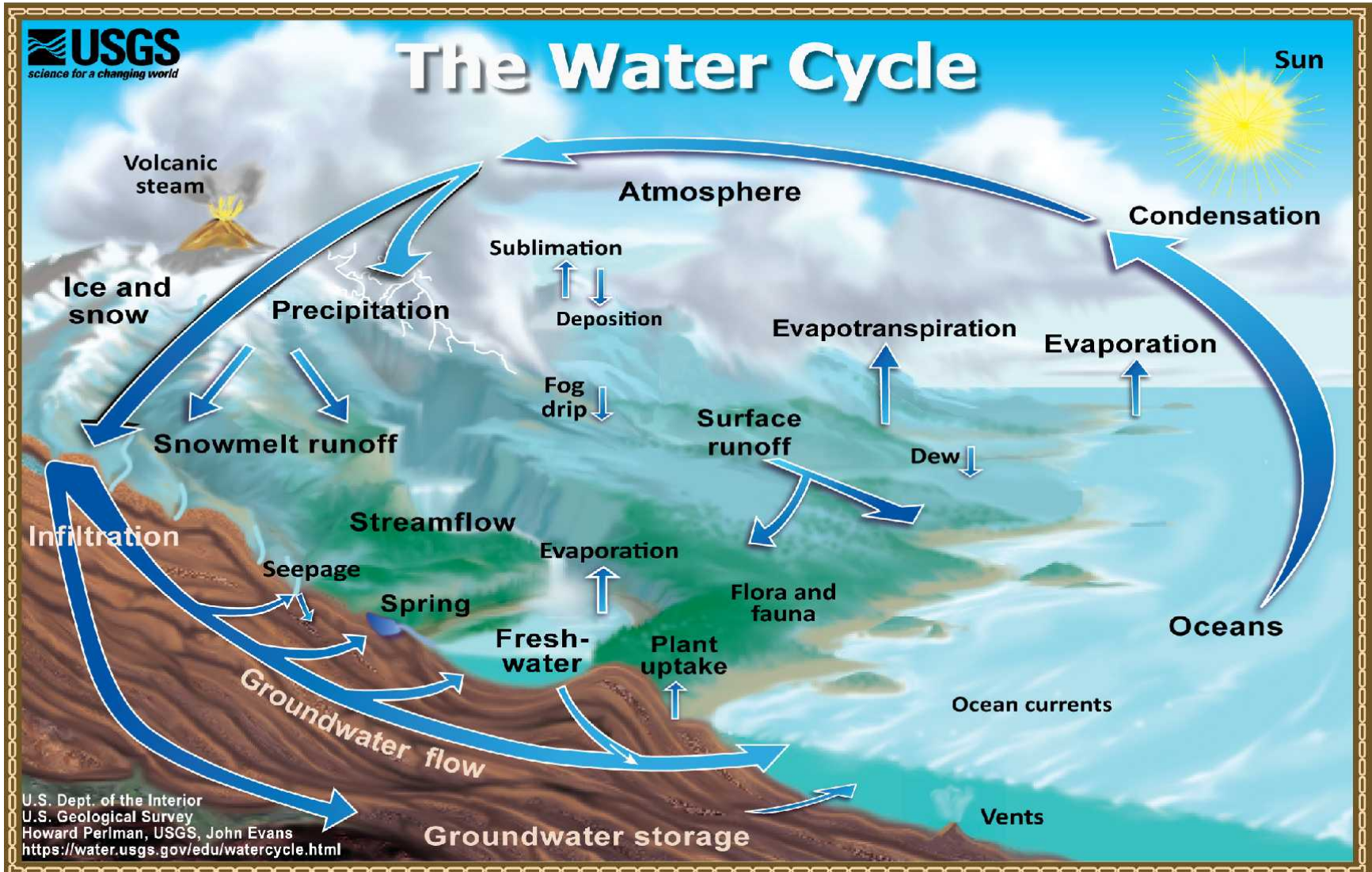
a) the study of water



What is Hydrology?

a) the study of water

b) the science that encompasses the occurrence, distribution, movement and properties of the waters of the Earth and their relationship with the environment within each phase of the hydrologic cycle. *





What is Hydrology in CTSM?

- a) the fluxes of water into and out of the land
- b) the redistribution of water within the land

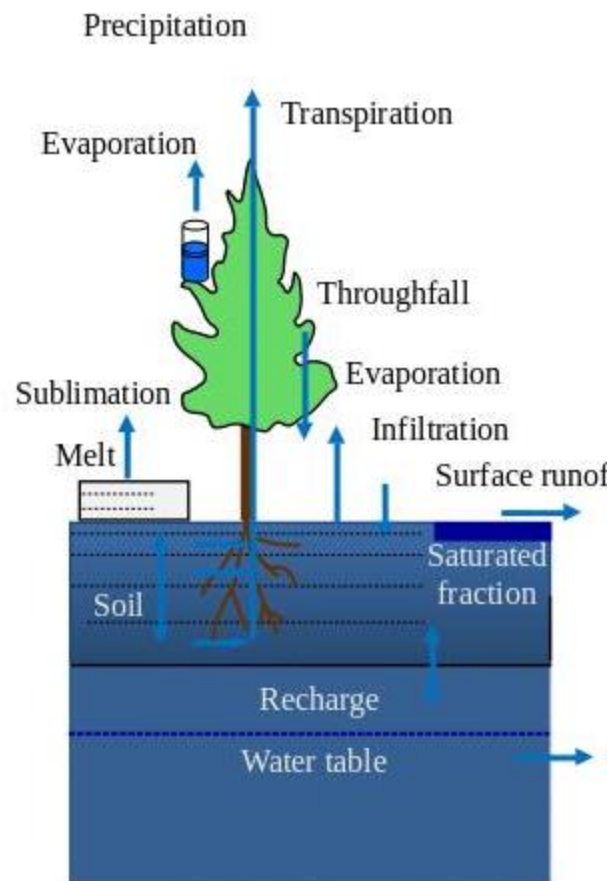


What is Hydrology in CTSM?

a) the fluxes of water into and out of the land

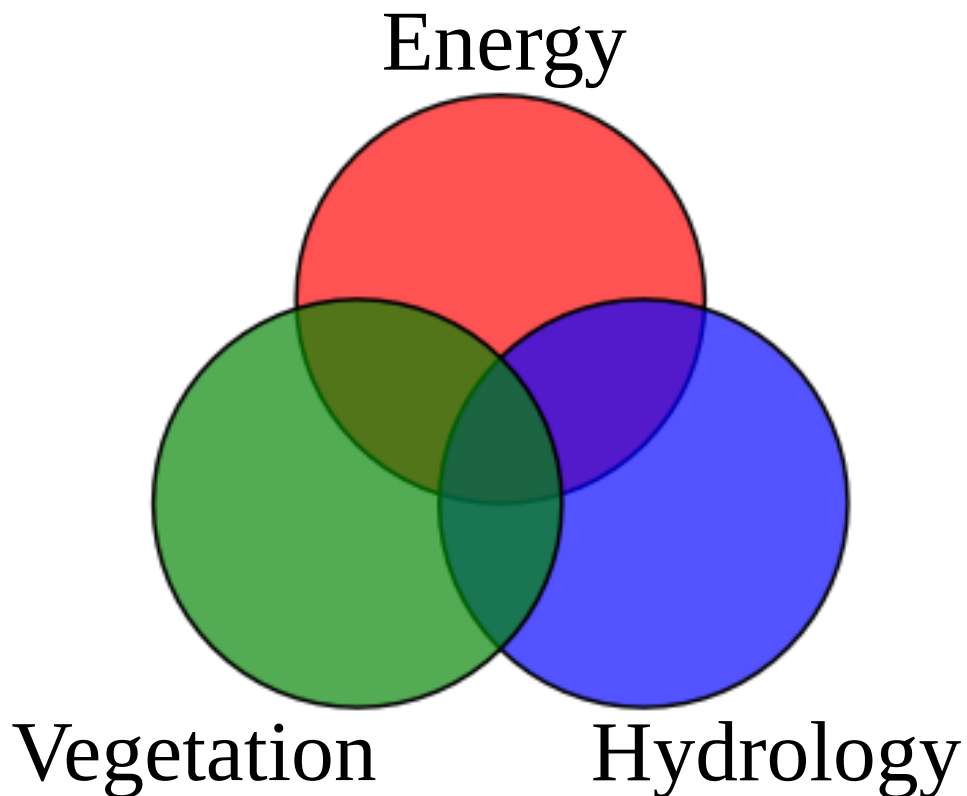
b) the redistribution of water within the land

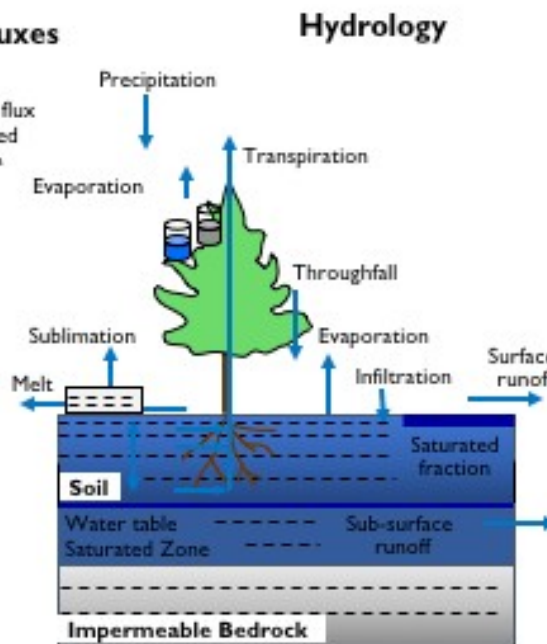
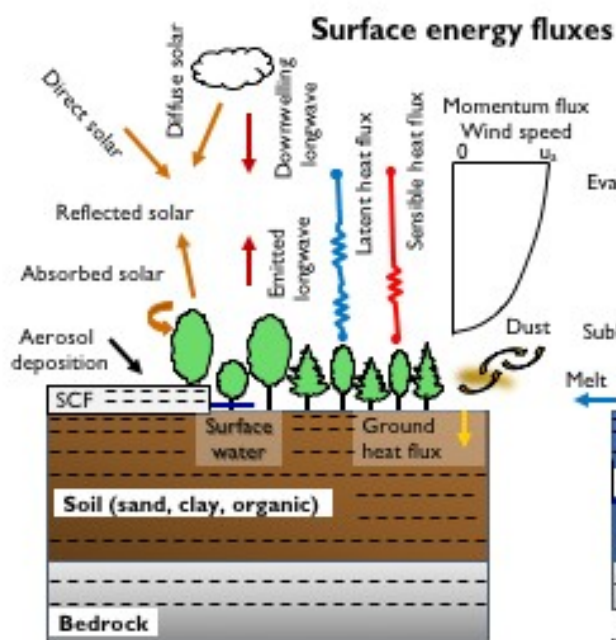
Hydrology





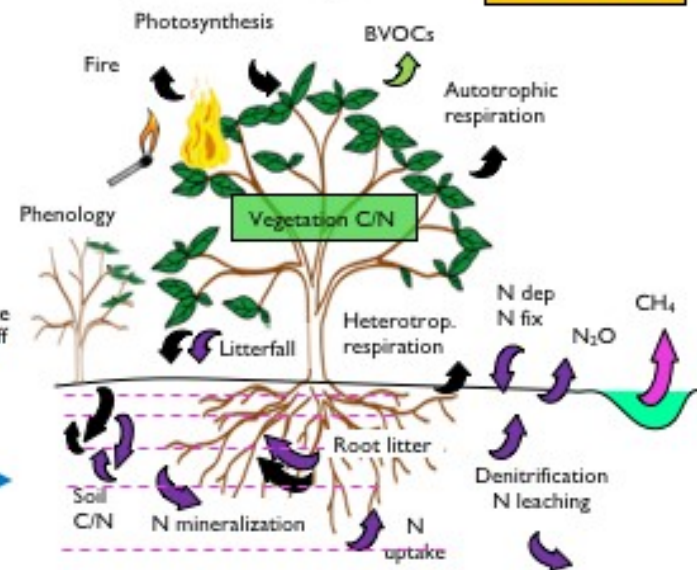
The movement of **water** is inextricably linked to the flow of **energy** and the life cycle of **vegetation**





Biogeochemical cycles

CLM5.0



From: *Technical Description of version 5.0 of the Community Land Model (CLM)*



The Water Balance

$$P = E + R + \Delta S$$

P = Precipitation

E = Evapotranspiration

R = Runoff

S = Storage



Hydrologic models have different objectives

Flood Forecasting \Rightarrow **R**

NWP, Climate Prediction \Rightarrow **E**

Drought Monitoring, Groundwater \Rightarrow **S**



Different objectives lead to different model structures

1-D \Rightarrow Darcy Flow (Infiltration/Recharge)

2-D \Rightarrow River Routing

3-D \Rightarrow Saturated Flow (Groundwater)



CLM is tasked with simulating *all* of these phenomena...

...therefore, *trade-offs* will be made.



CLM Water Balance Operations

Precipitation

⇒ Partitioning between rain and snow, or between stratiform and convective

⇒ Canopy interception, storage, and throughfall



CLM Water Balance Operations

Evaporation

⇒ Evaporation from Soil / Canopy / Snow /
Surface Water

⇒ Transpiration from vegetation



CLM Water Balance Operations

Runoff

⇒ Surface Runoff (Infiltration and/or Saturation Excess)

⇒ Subsurface Runoff (Baseflow)

⇒ River Routing



CLM Water Balance Operations

Storage

- ⇒ Soil Moisture
- ⇒ Groundwater and water table depth
- ⇒ Perched water table
- ⇒ Canopy water
- ⇒ Surface water
- ⇒ Snow

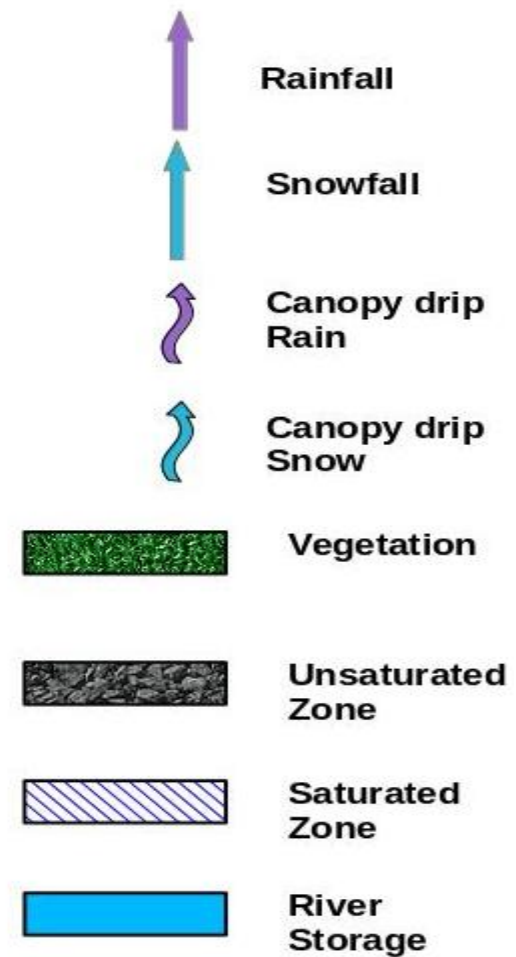
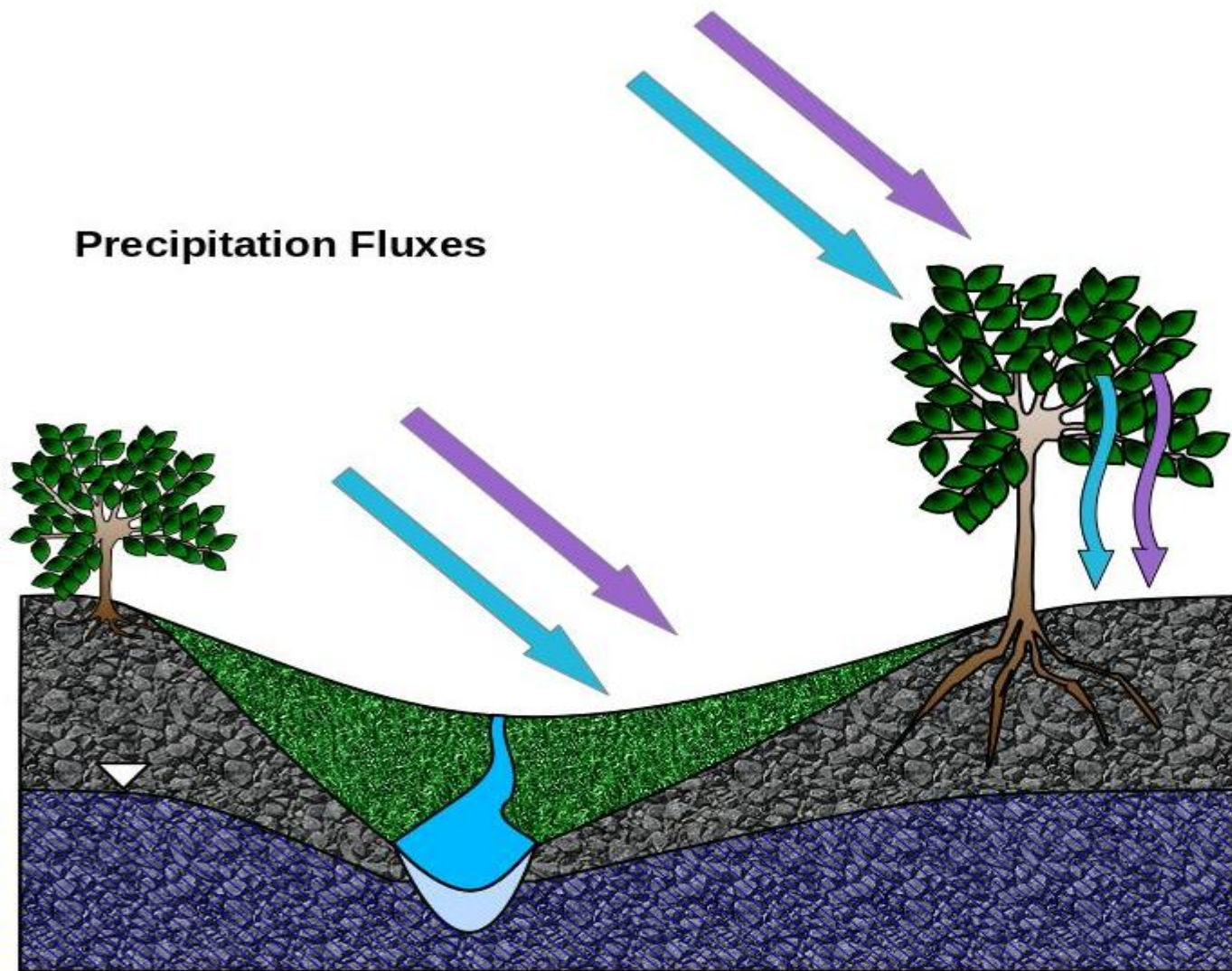


Canopy Hydrology

- Section 2.7.1 of the CLM Tech Note
- Interception / throughfall
- Leaf water storage and wetted fraction
- Evaporation from leaf surfaces

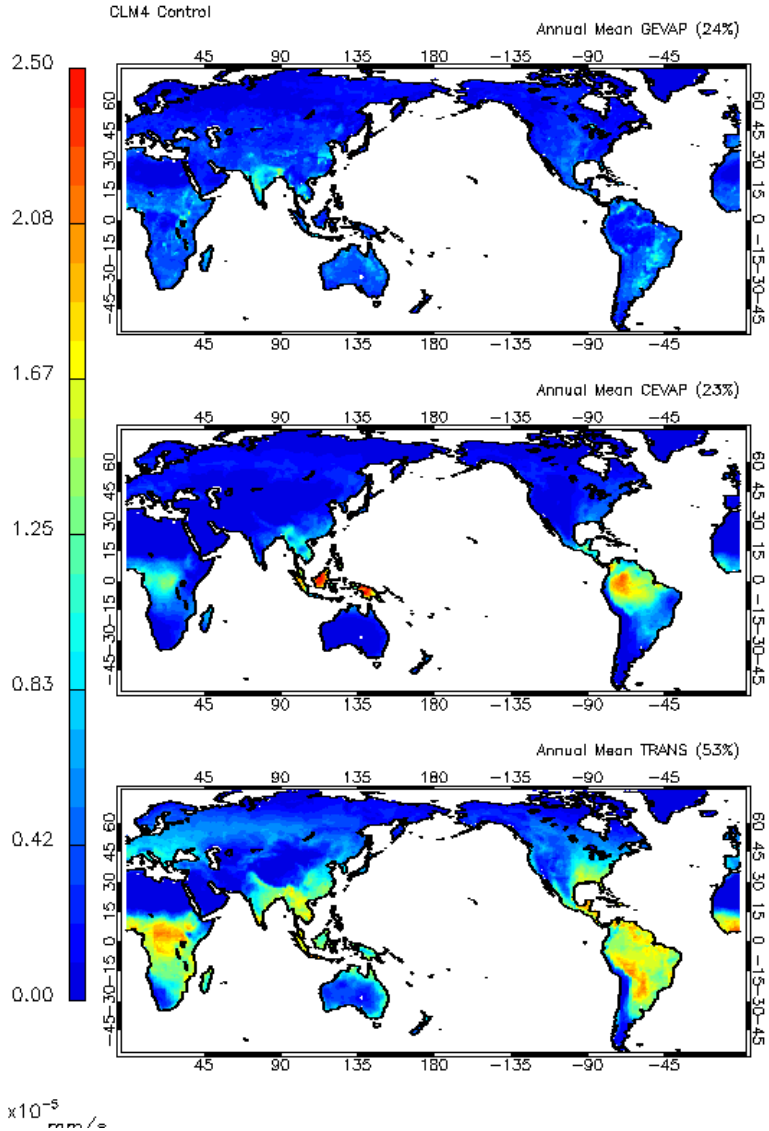


Precipitation Fluxes





Canopy Hydrology And Evapotranspiration Partitioning



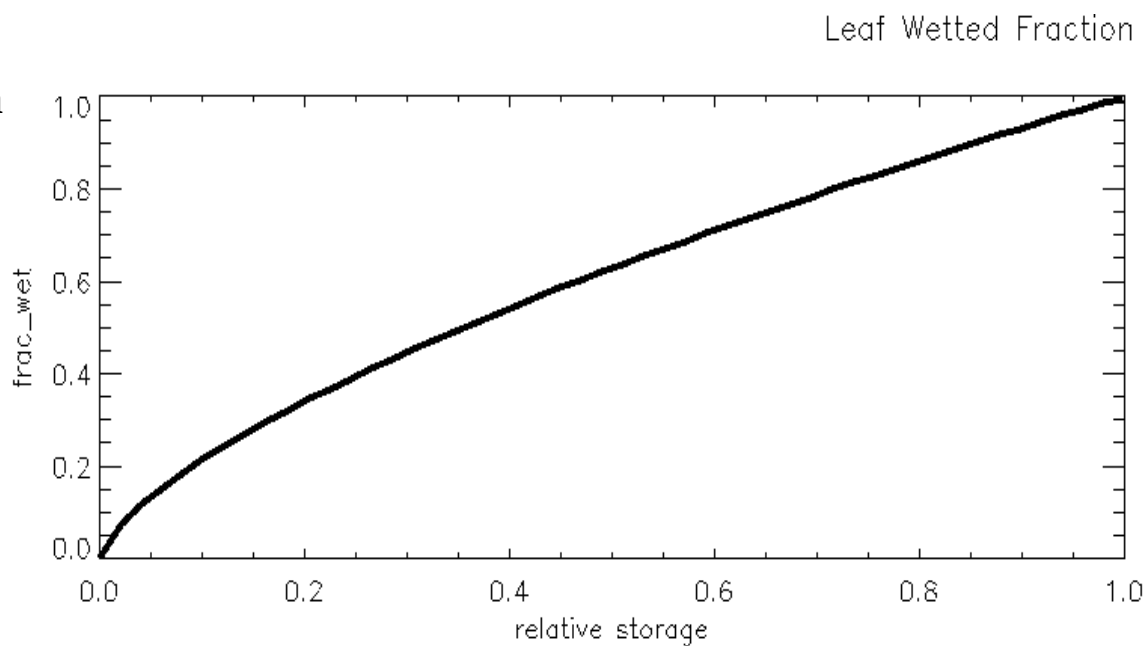
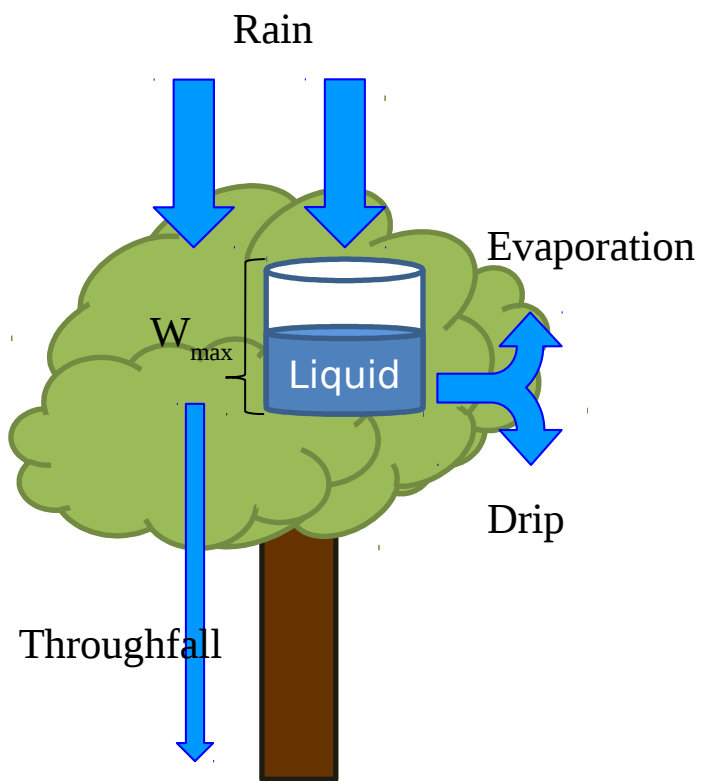
**Ground Evaporation:
24%**

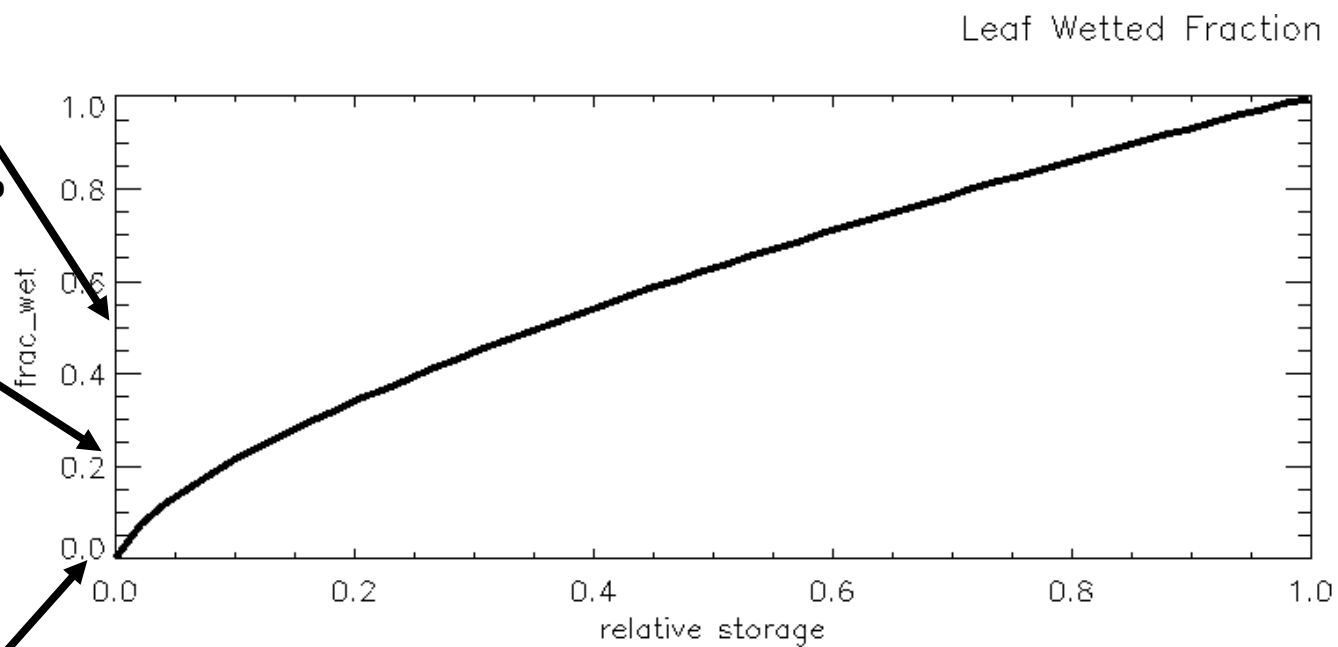
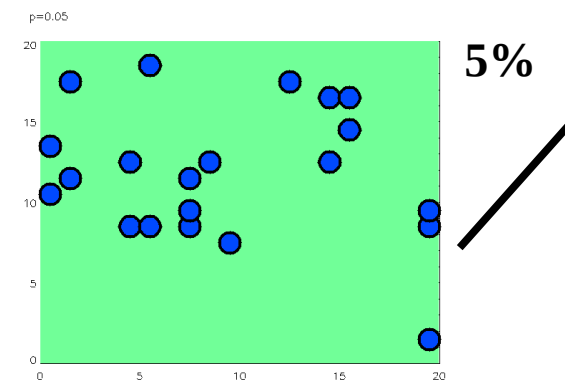
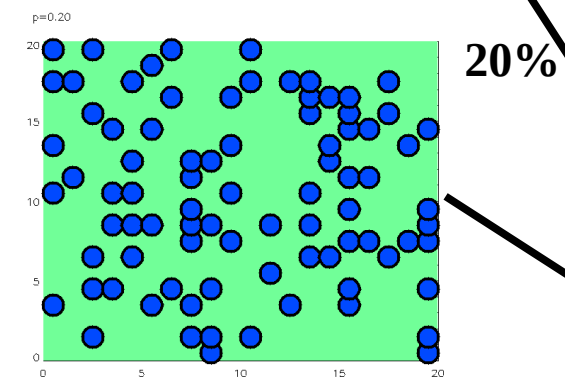
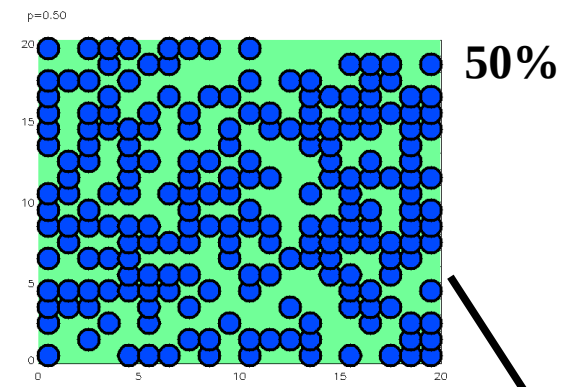
**Canopy Evaporation:
23%**

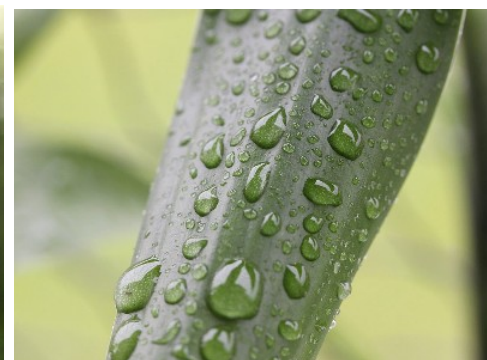
**Transpiration:
53%**



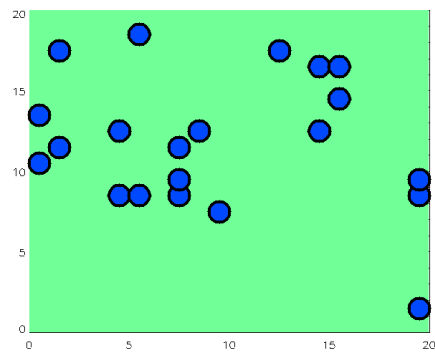
Leaf Wetted Area



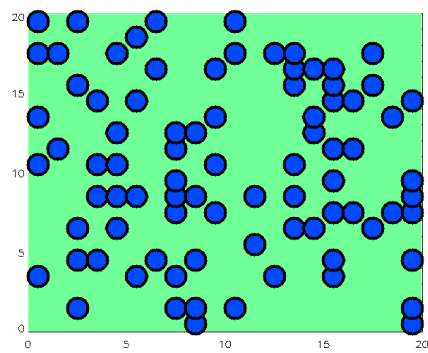




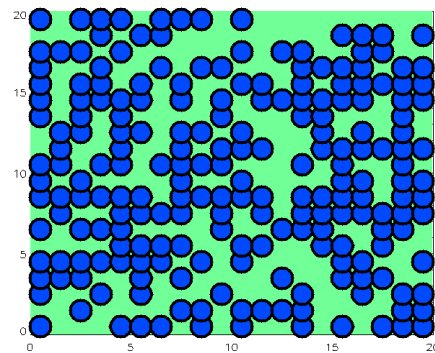
$p=0.05$



$p=0.20$



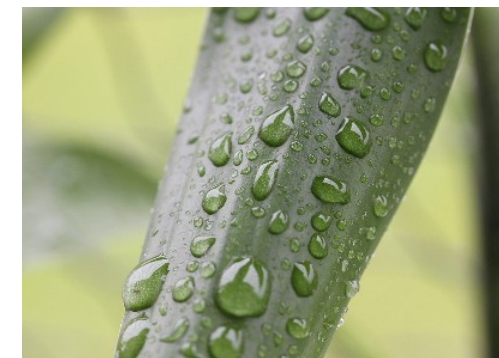
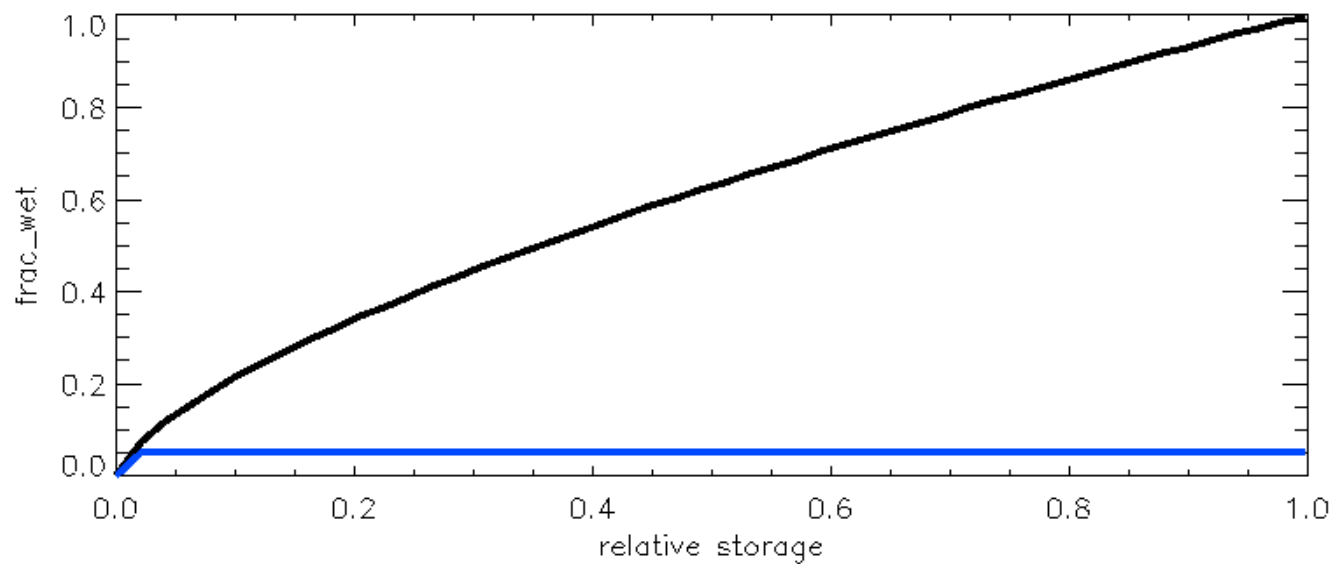
$p=0.50$





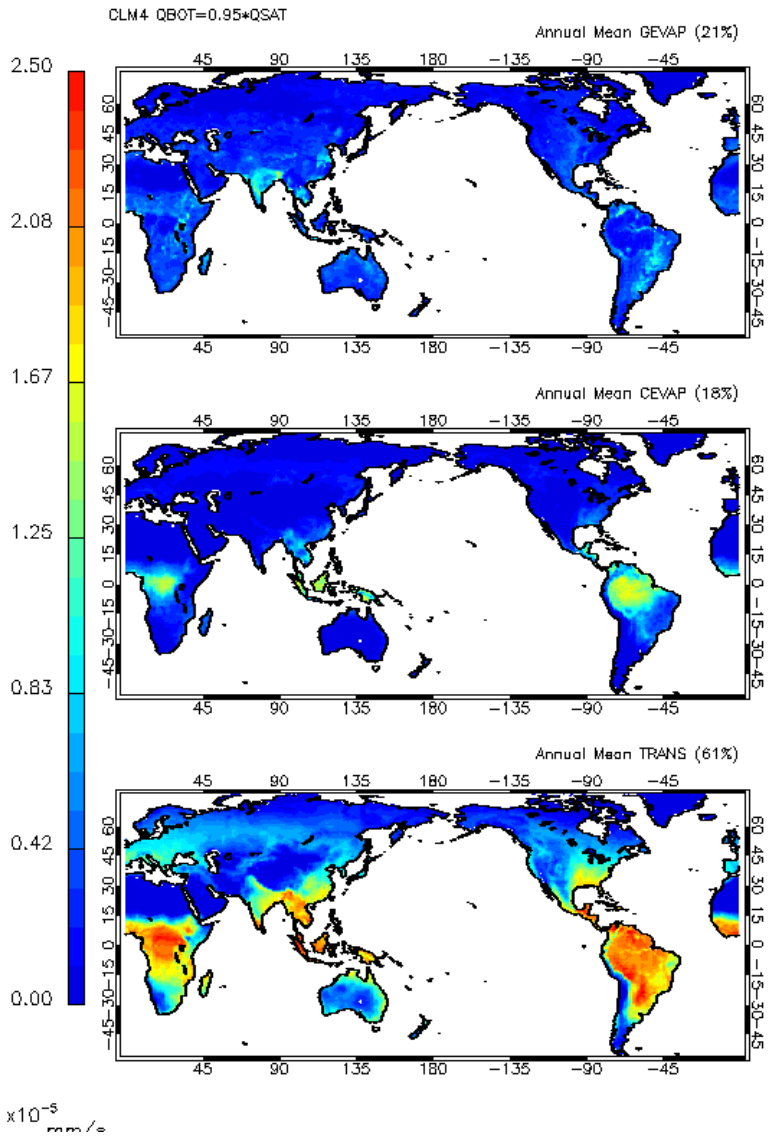
Leaf Wetted Area

Leaf Wetted Fraction





Canopy Hydrology And Evapotranspiration Partitioning



**Ground Evaporation:
21%**

**Canopy Evaporation:
18%**

**Transpiration:
61%**

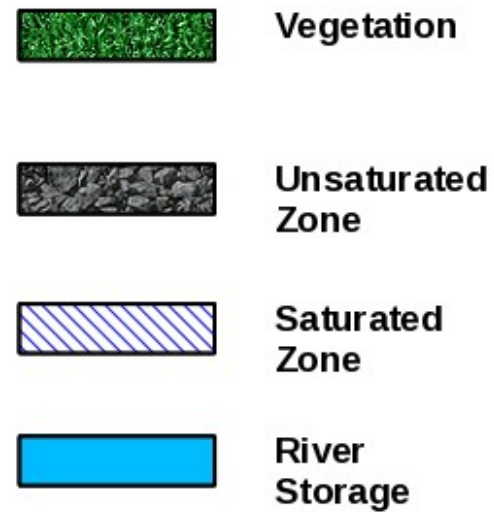
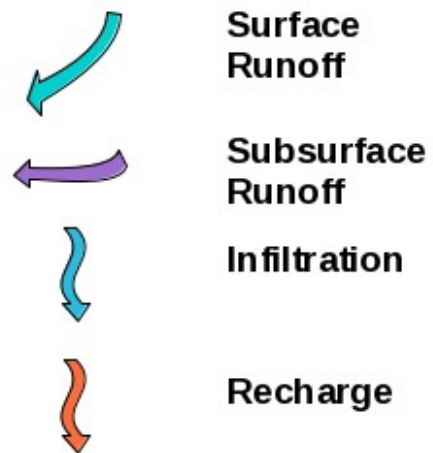
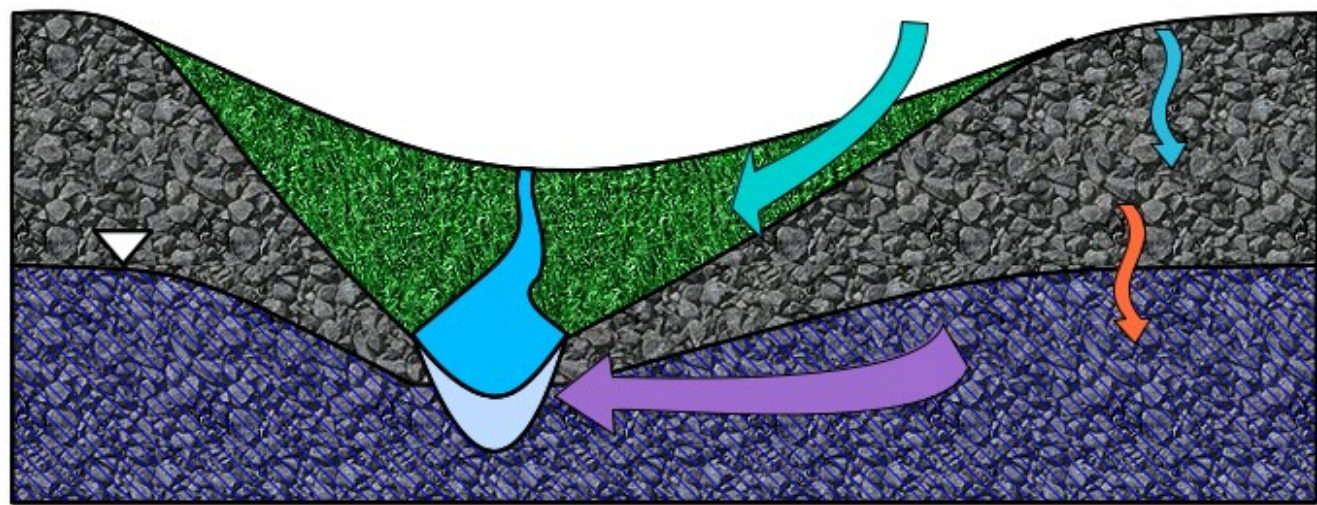


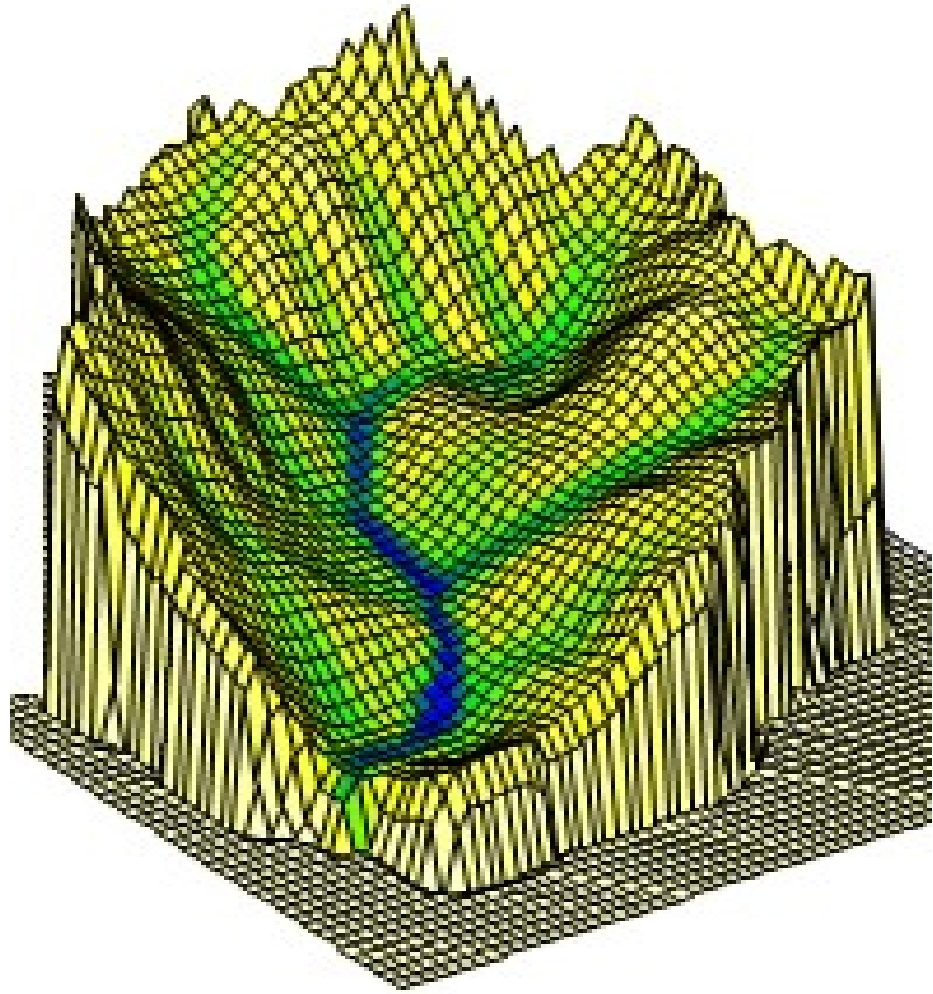
Runoff and Surface Water Processes

- Section 2.7.2 of the CLM Tech Note
- Surface runoff
- Saturated area
- Subsurface runoff
- River routing
- Surface water (wetlands)



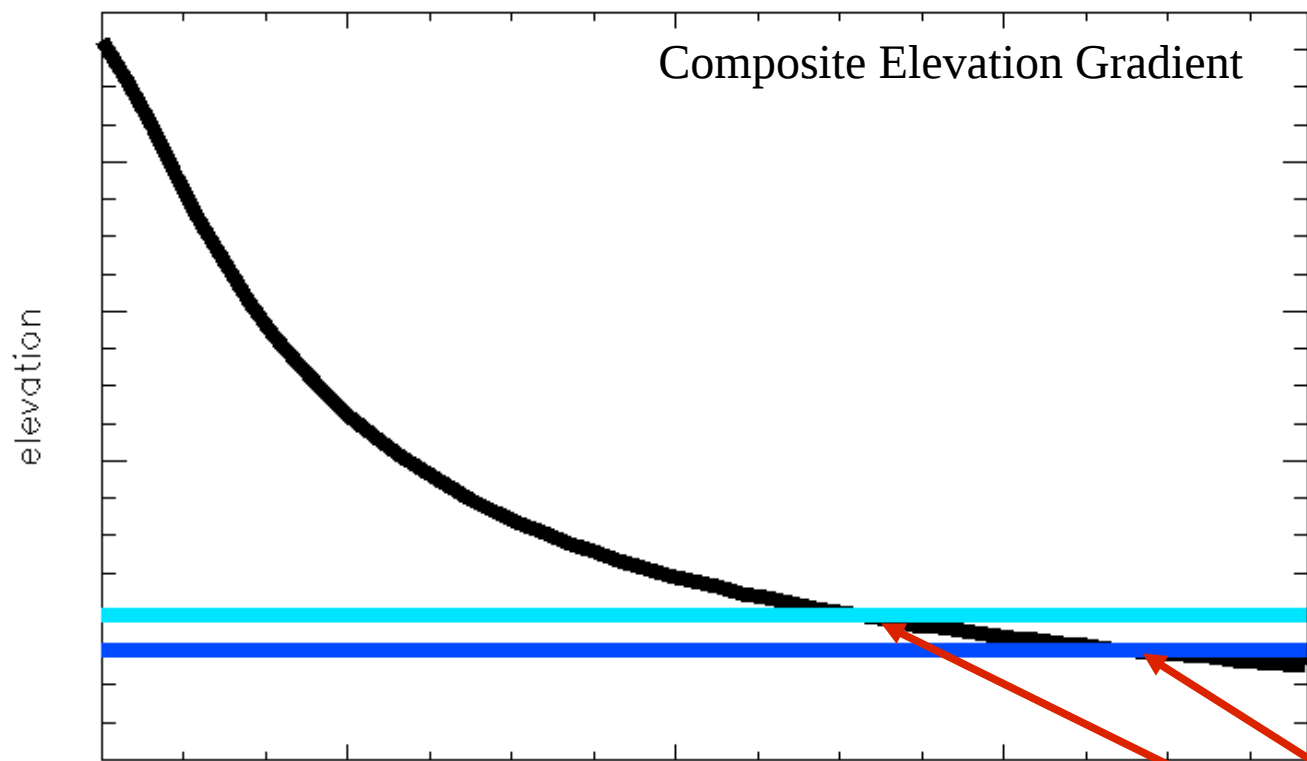
Runoff Generation and Infiltration





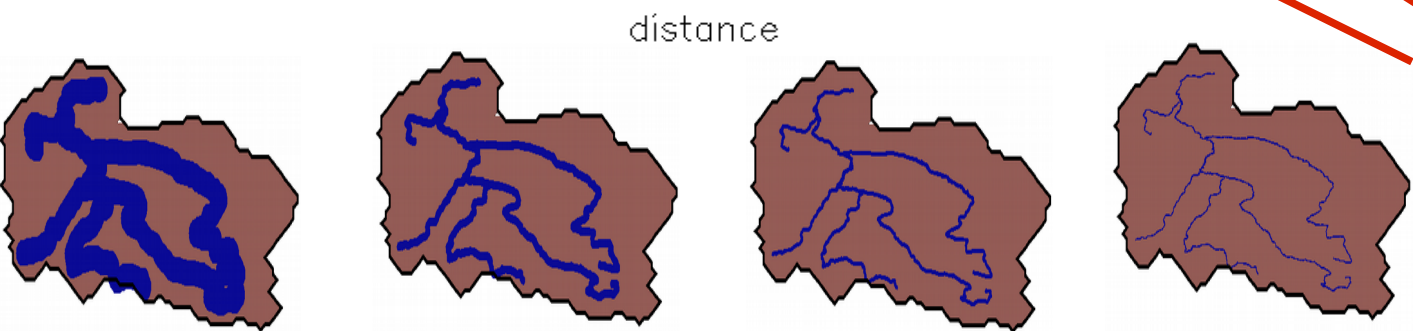
The water table determines the fraction of the area that is saturated

Saturated areas produce surface runoff

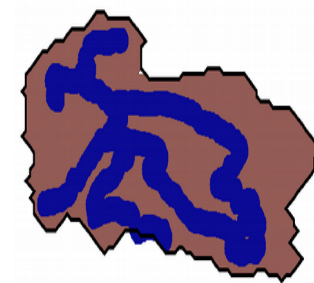
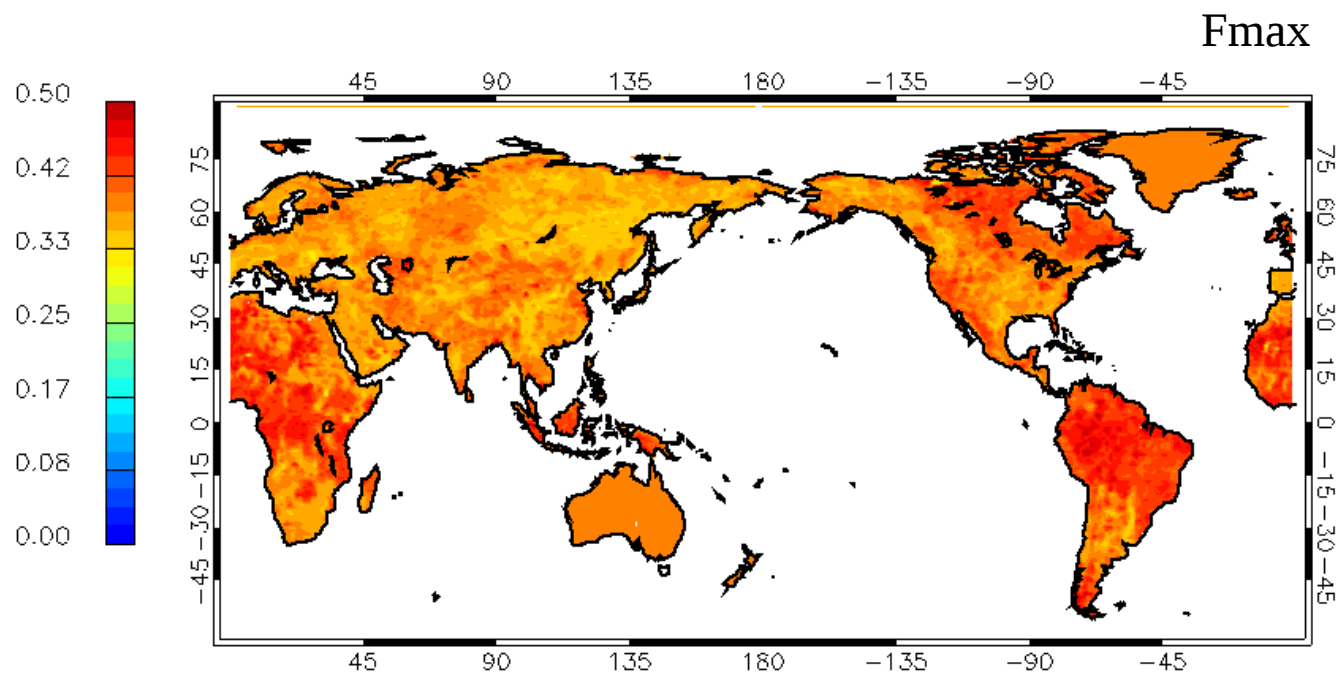


Saturated fraction = 0.36

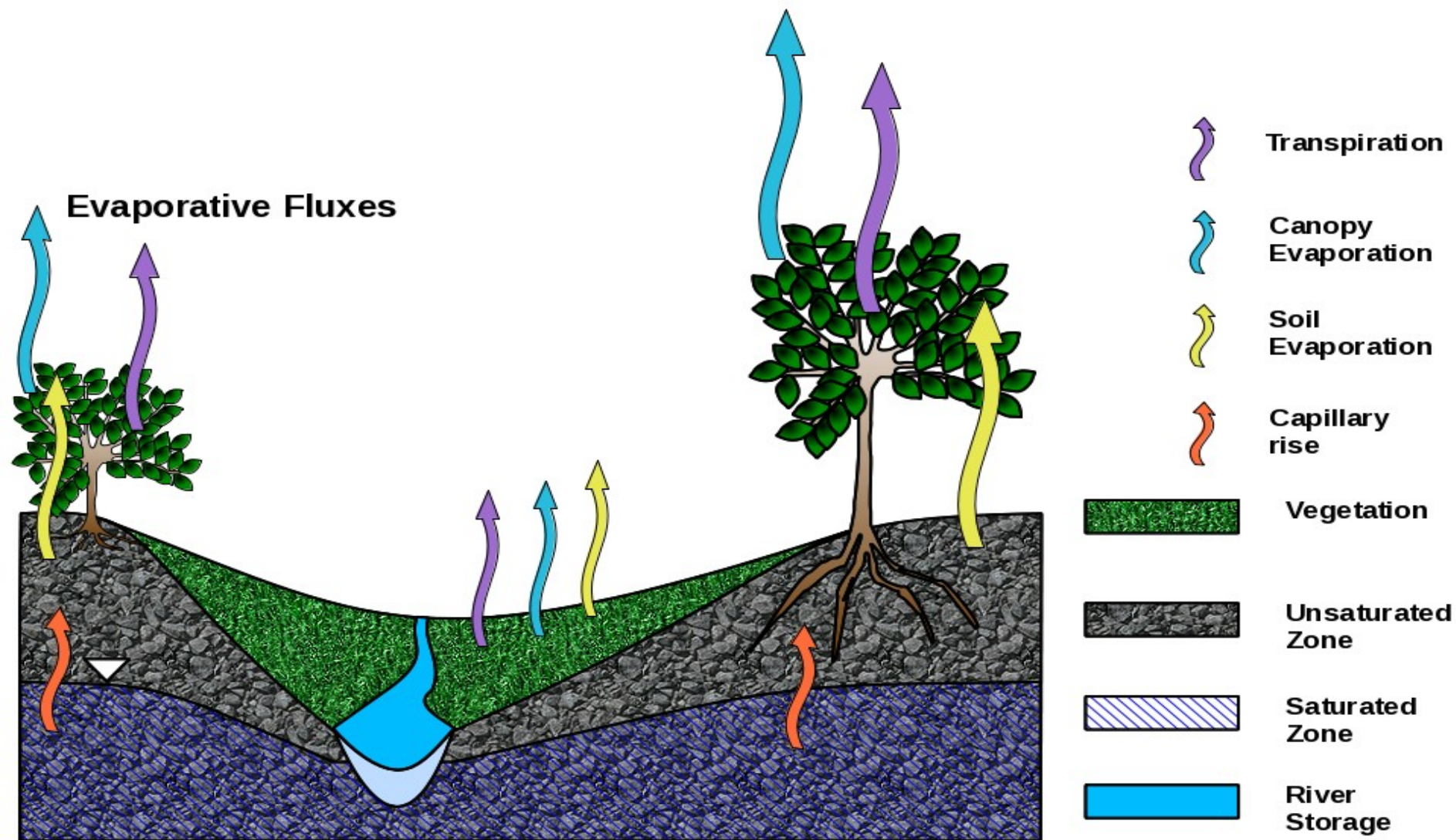
Saturated fraction = 0.14



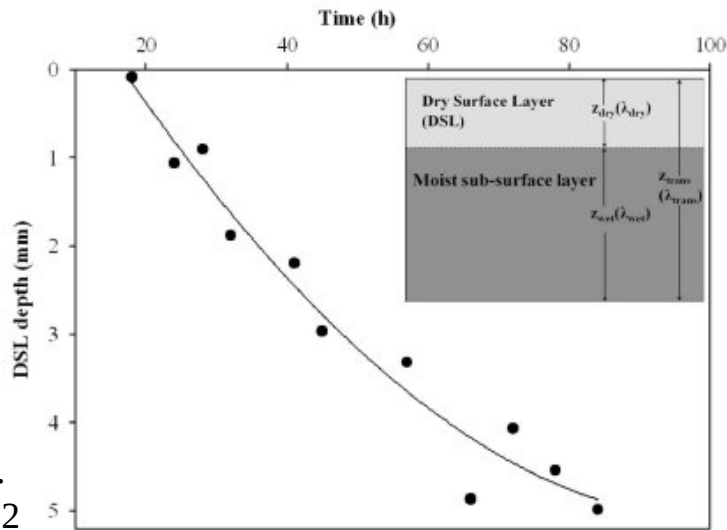
Point at which water table intersects surface determines saturated fraction



Fmax variable on surface data file determines maximum saturated fraction



Dry Surface Layer (DSL)



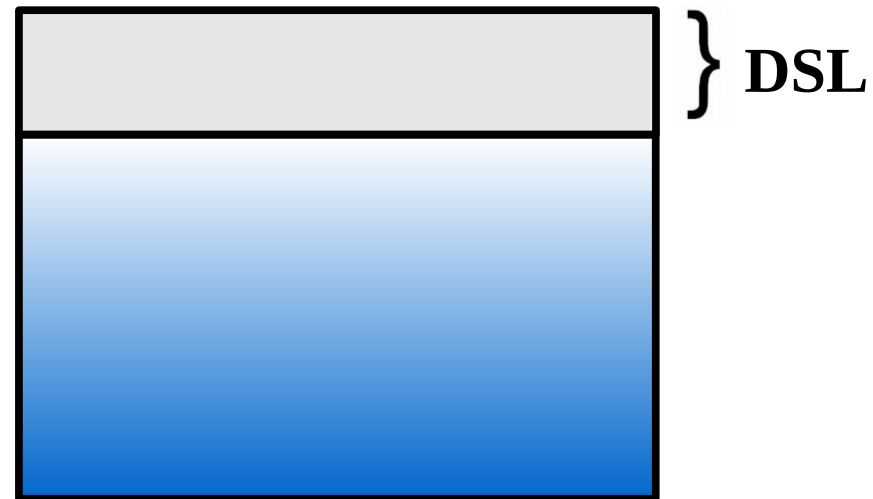
Deol et al.
WRR,2012

Figure 9. Estimated thickness of the dry surface layer (DSL) during transient evaporation. The line indicates the approximate trend. Inset: Conceptual structure of near-surface transition layer used for estimating thickness of the dry surface layer (DSL) with thickness (z) and thermal conductivity (λ).



Vapor Diffusion

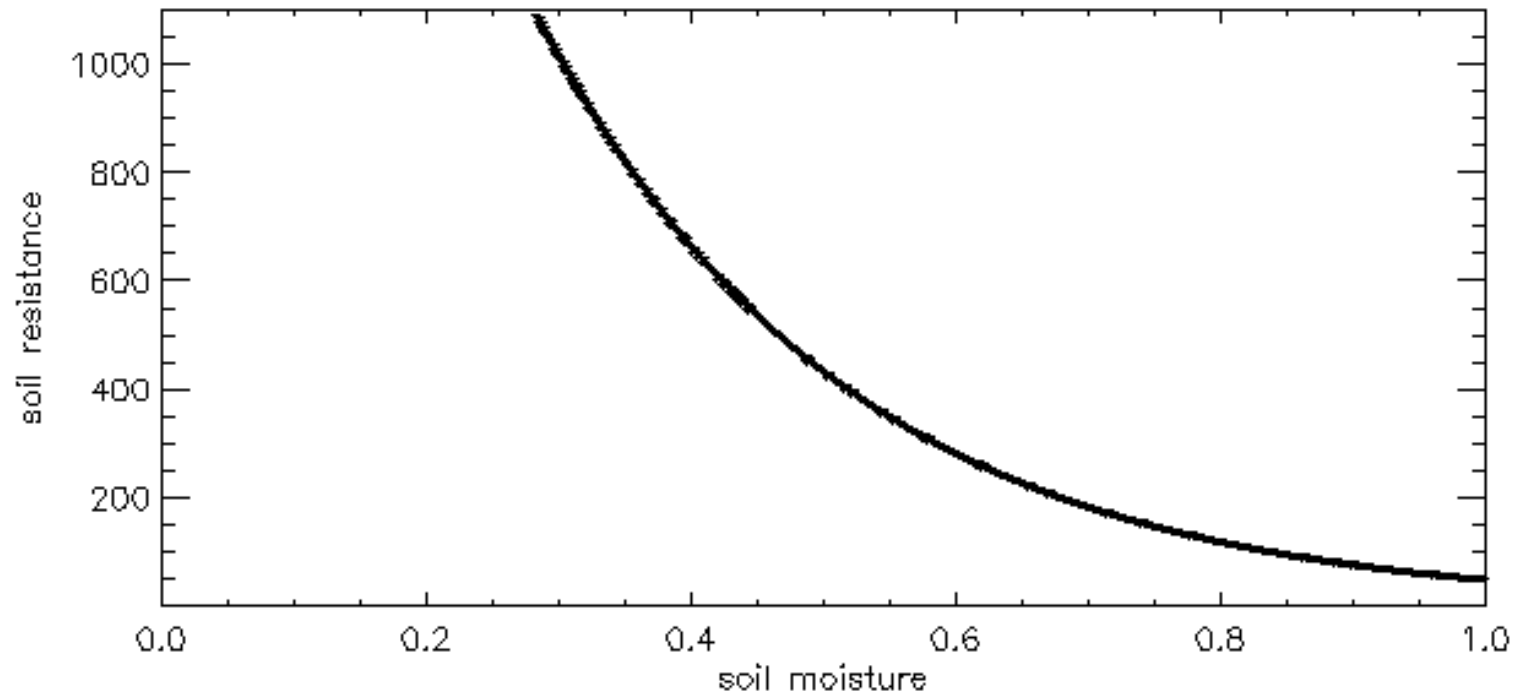
Liquid Diffusion



Vapor Diffusion < Liquid Diffusion



Soil Evaporative Resistance

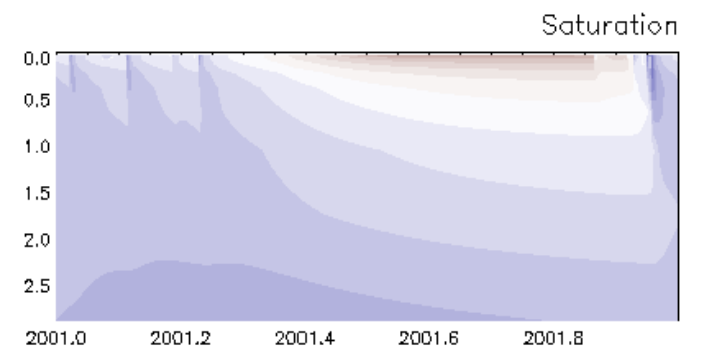
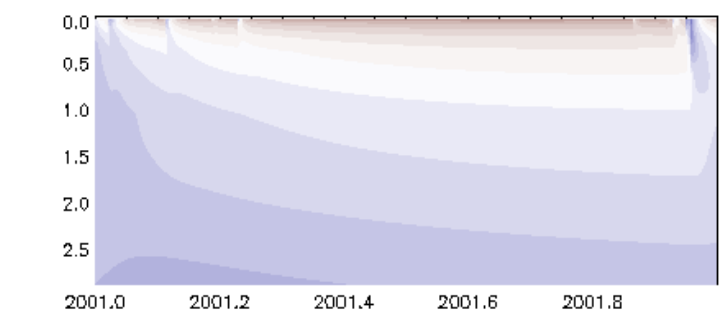
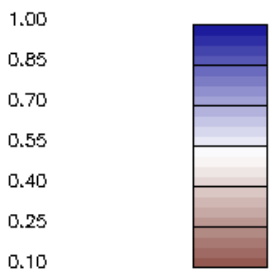
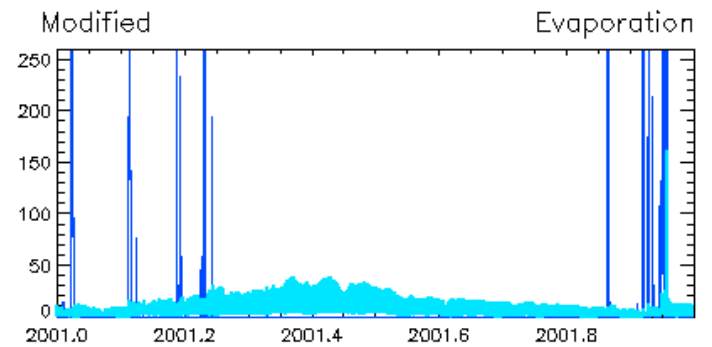
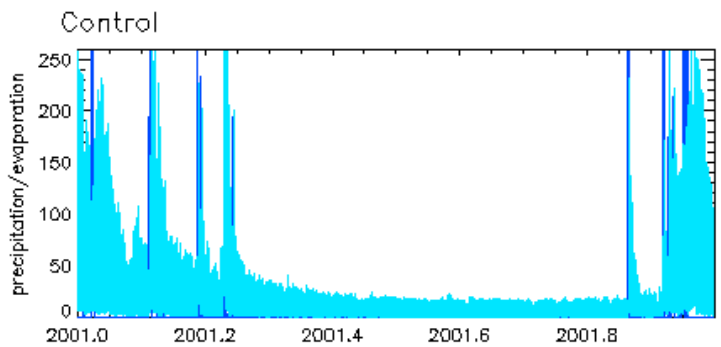
Surface Soil Moisture vs Soil Resistance



Section 2.5.2 of CLM Tech Note



Precipitation 
 Evaporation 



Little infiltration –
 precipitation rapidly
 returned to atmosphere

Larger infiltration events
 during winter – evaporation
 peaks during summer



Soil Moisture Redistribution

- Section 2.7.3 of the CLM Tech Note
- Moisture form of Richards equation with adaptive sub-stepping
- Water moves due to gravity and gradients in soil matric potential



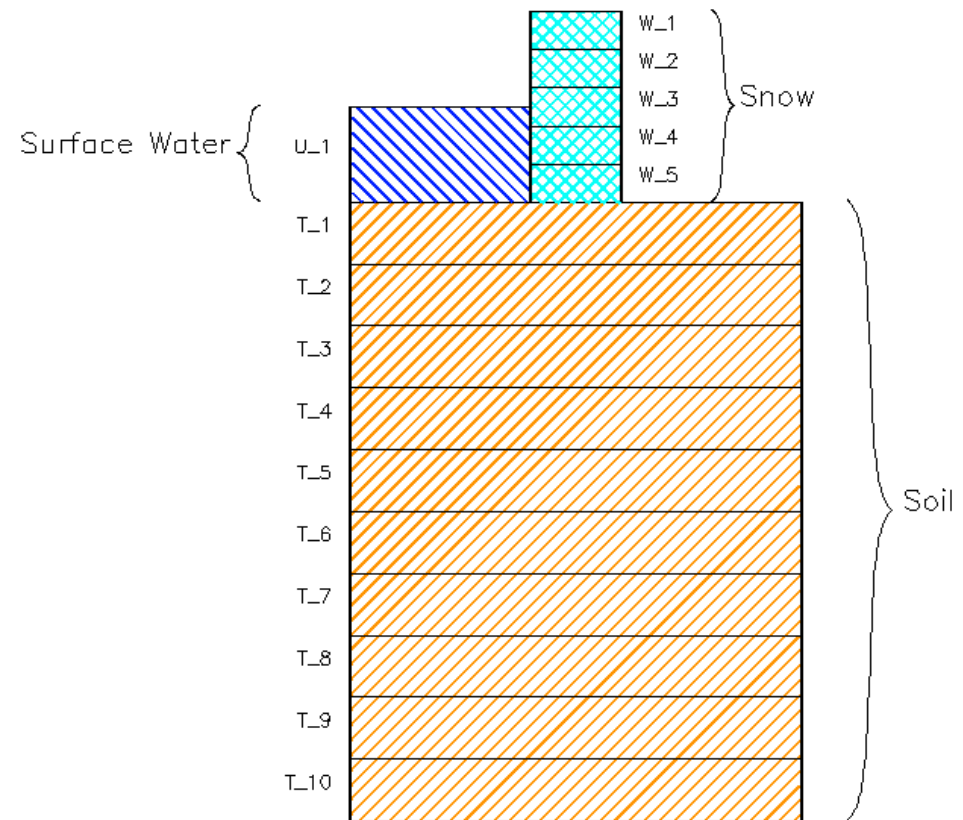
Soil model

Treats processes such as:

- Soil moisture redistribution
 - Infiltration
 - Darcy flow
 - Recharge
- Soil moisture phase change
- Soil temperature redistribution

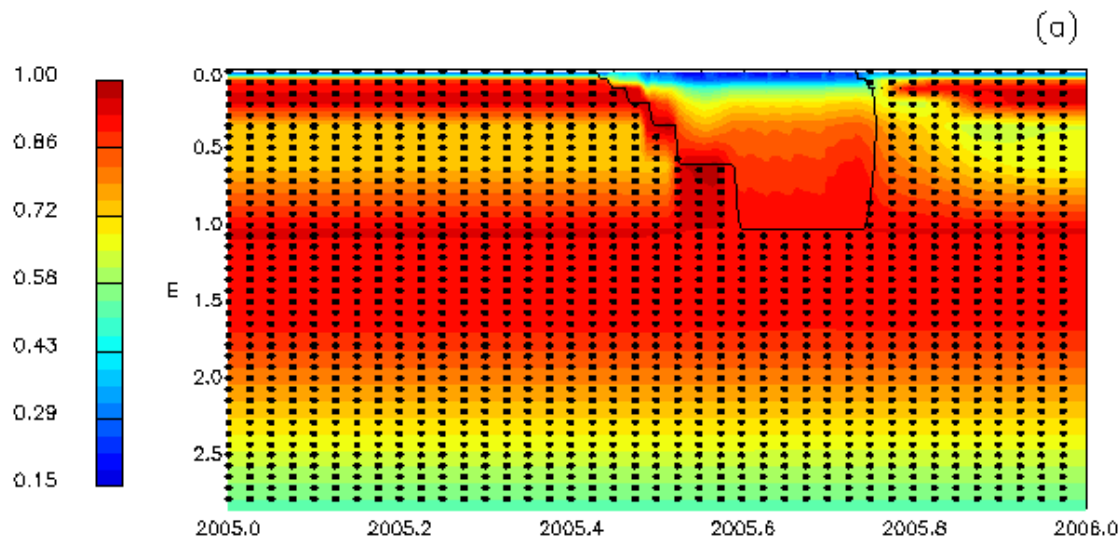
Default structure has 20 layers of variable thickness, spanning about 8 meters depth

- Thermal calculations use additional deep layers

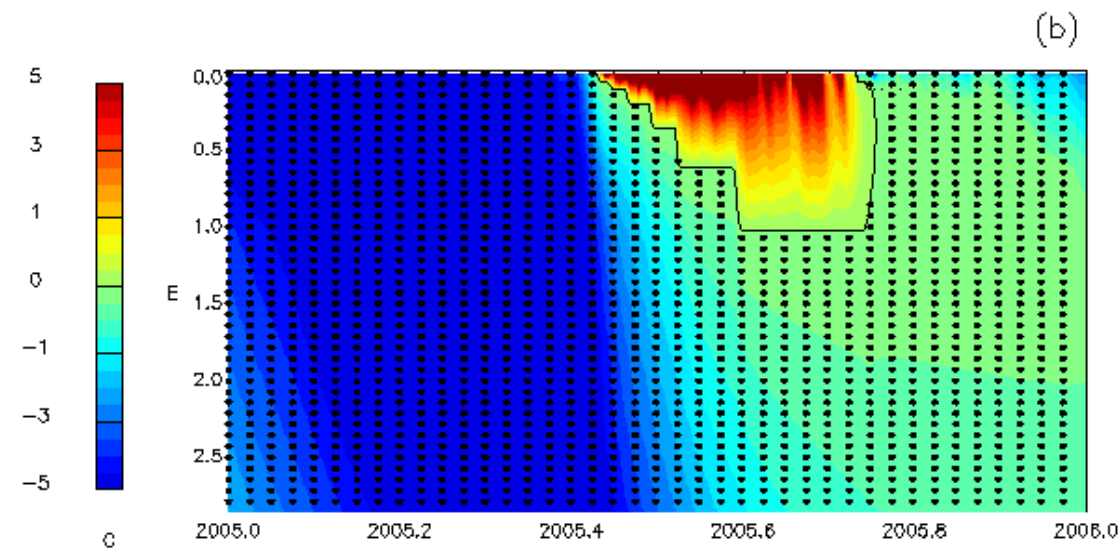




a) Soil moisture (% saturation)



b) Soil temperature ($^{\circ}\text{C}$)



Stippling indicates frozen soil



Percent sand

50.00

41.67

33.33

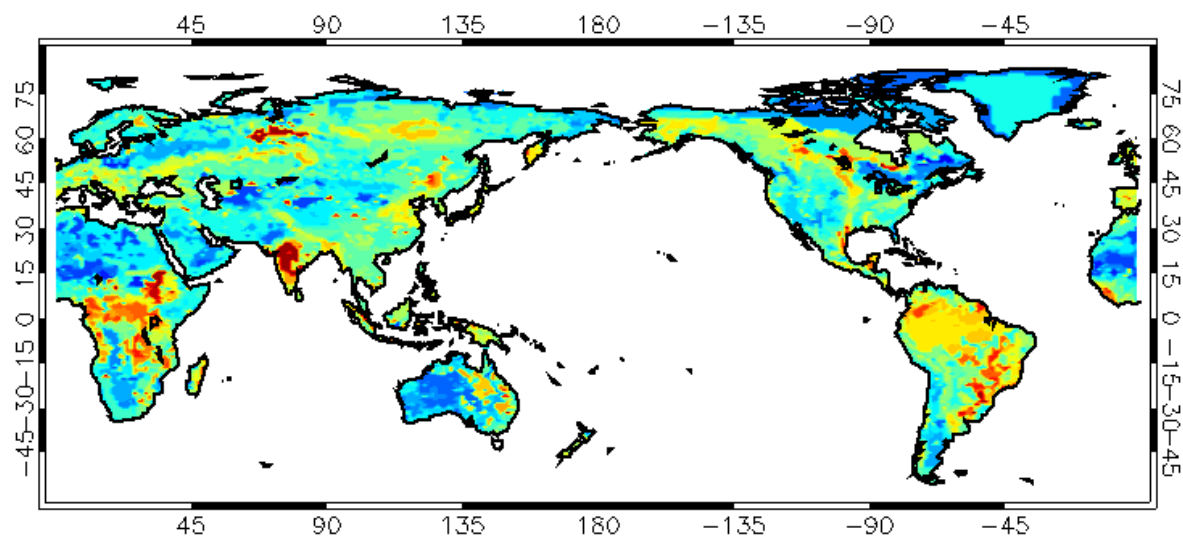
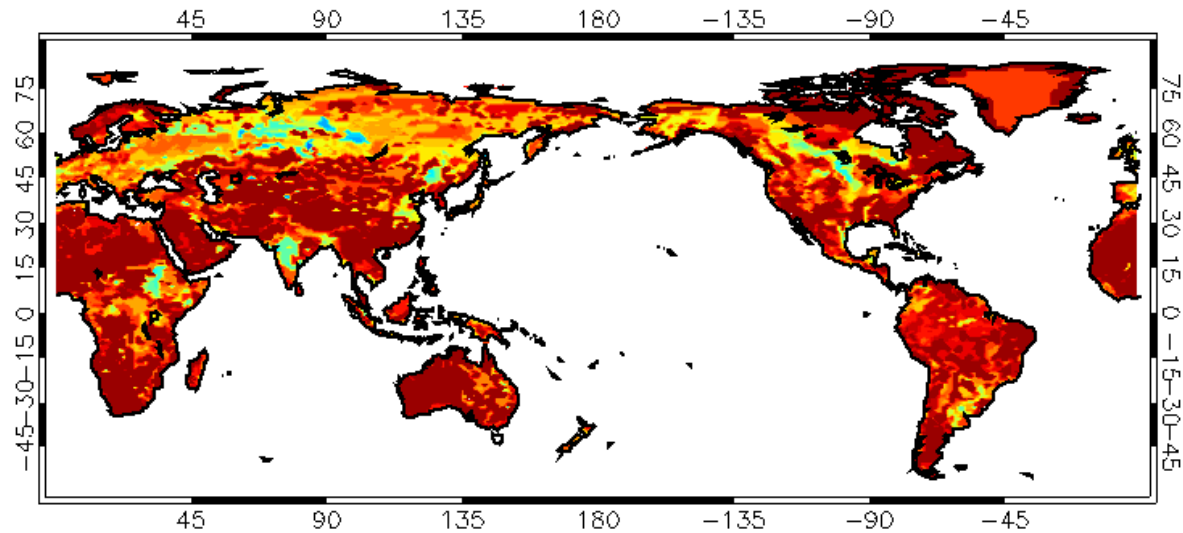
25.00

Percent clay

16.67

8.33

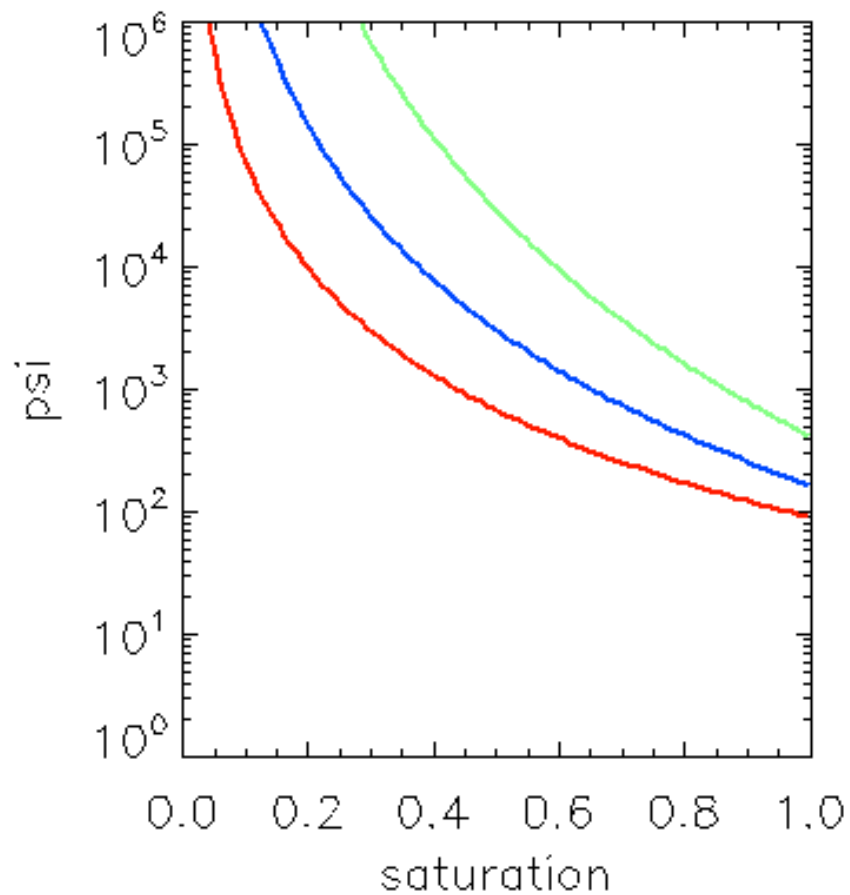
0.00



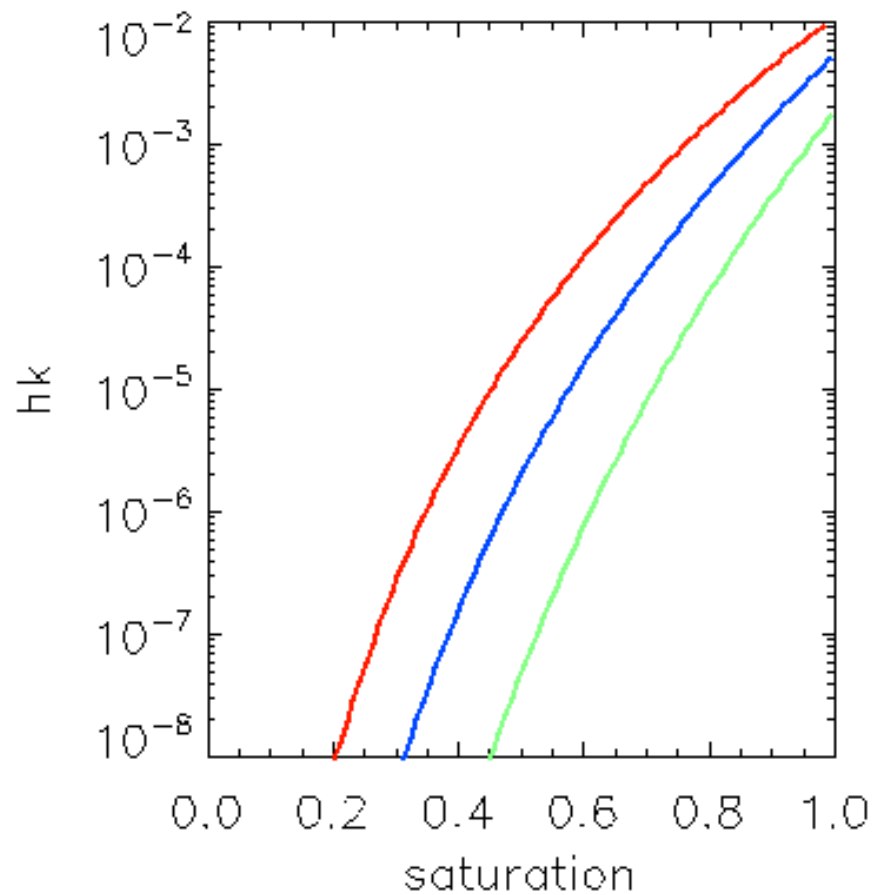


Soil Hydraulic Properties

Soil Matric Potential



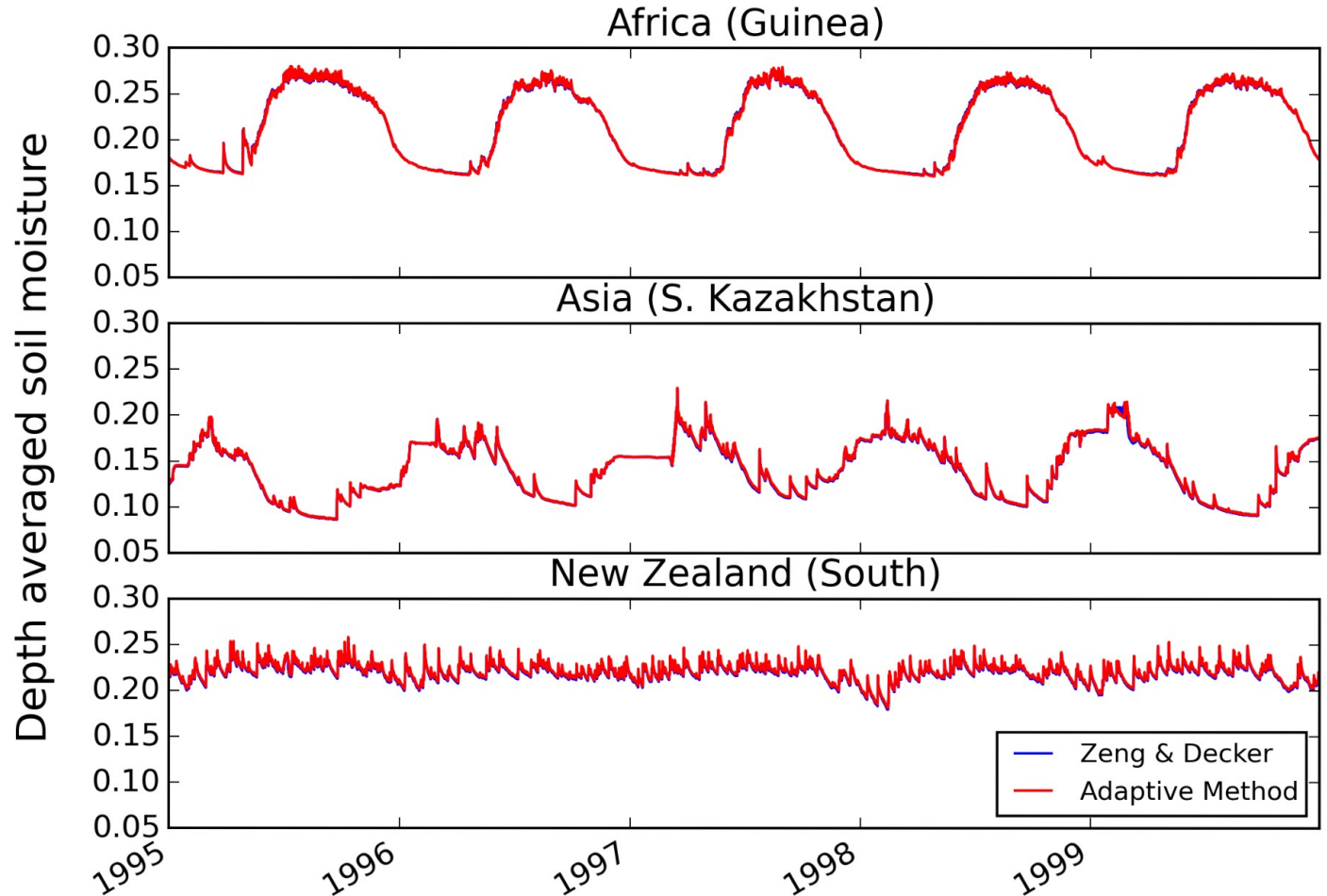
Soil Hydraulic Conductivity



Adaptive time stepping method for soil water distribution

Variable time step length depending on solution error

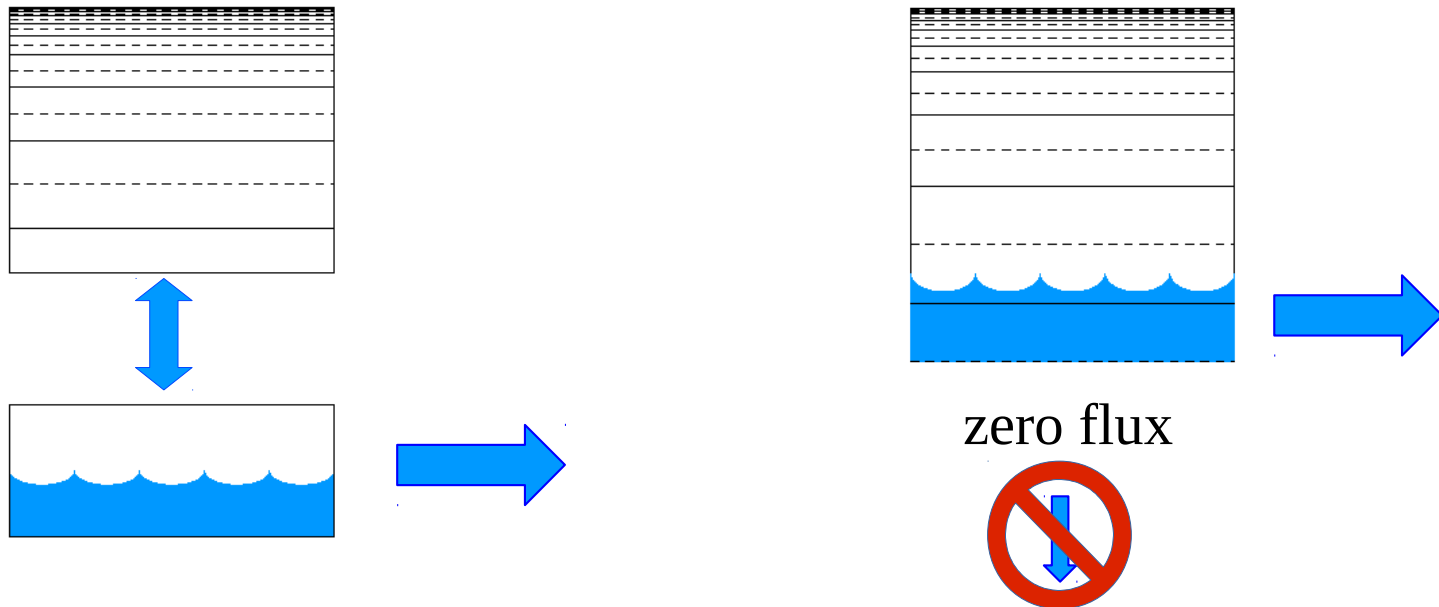
Eliminates numerical instabilities in Richards equation solution





Groundwater and Water Table Dynamics

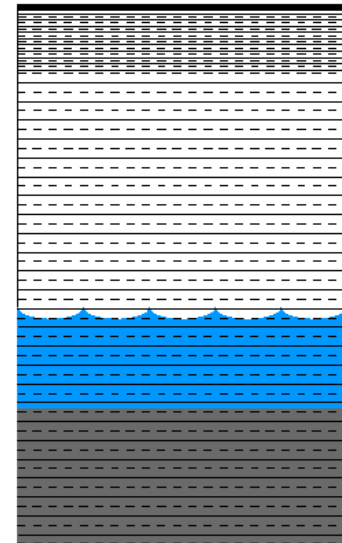
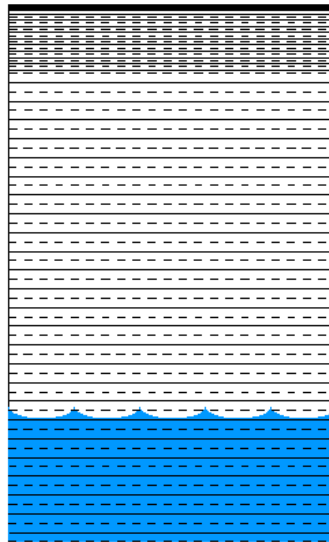
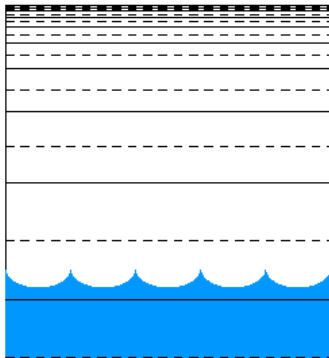
- bulk aquifer layer
- bedrock (zero vertical flux) lower boundary
- subsurface discharge depends on water table depth



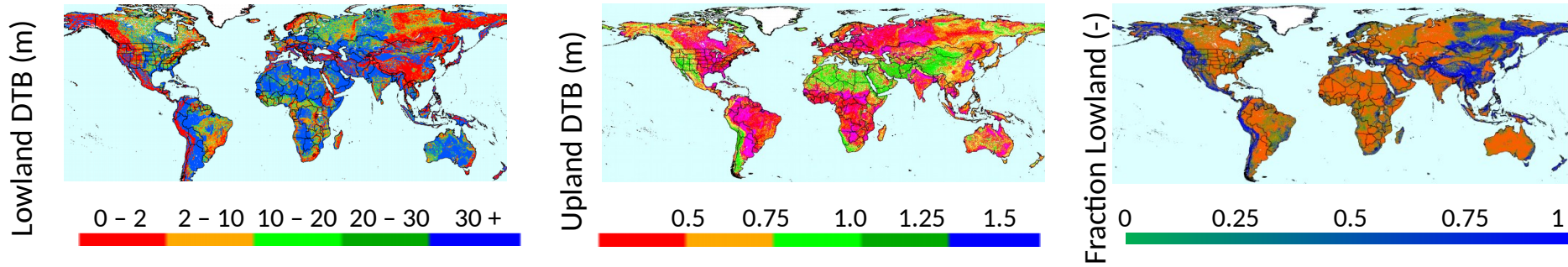


Soil Depth

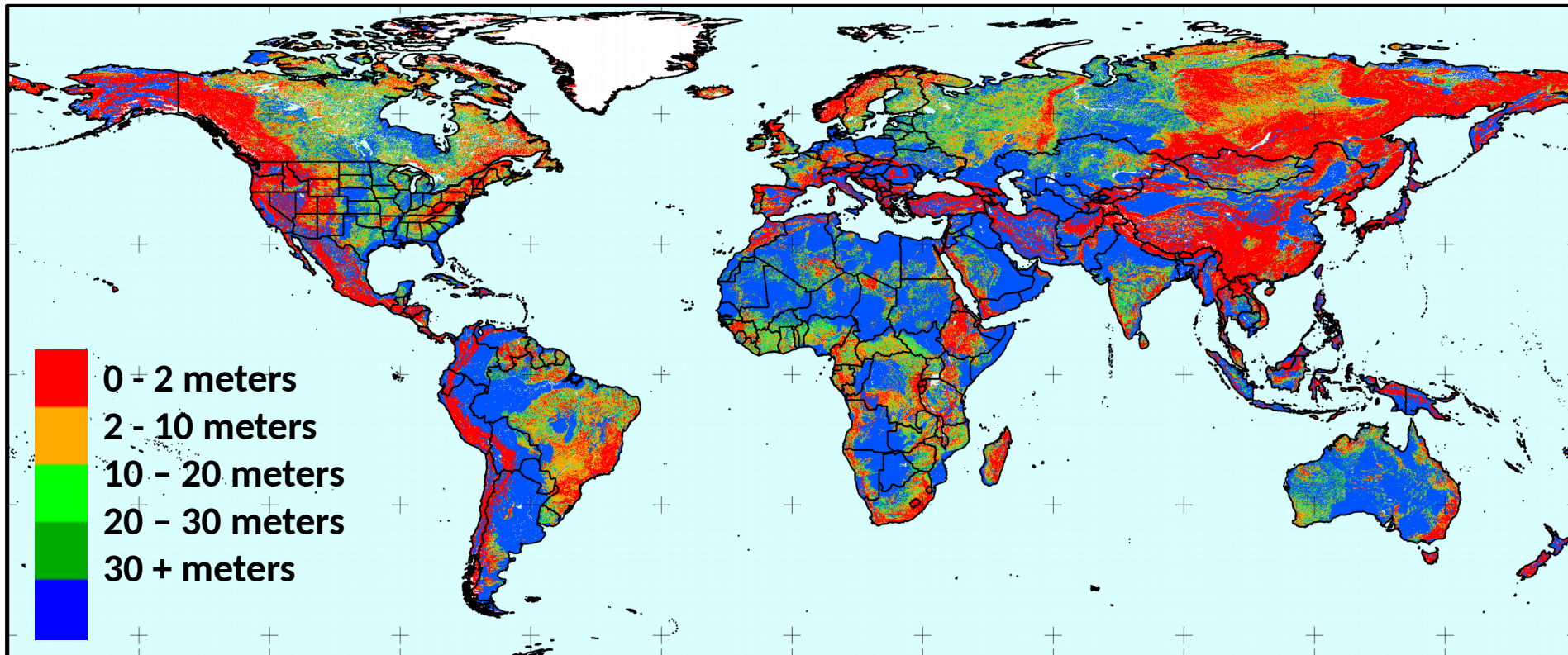
- deep soil / variable soil depth
- high vertical resolution soil



PRELIMINARY GLOBAL MAP OF DTB ESTIMATES



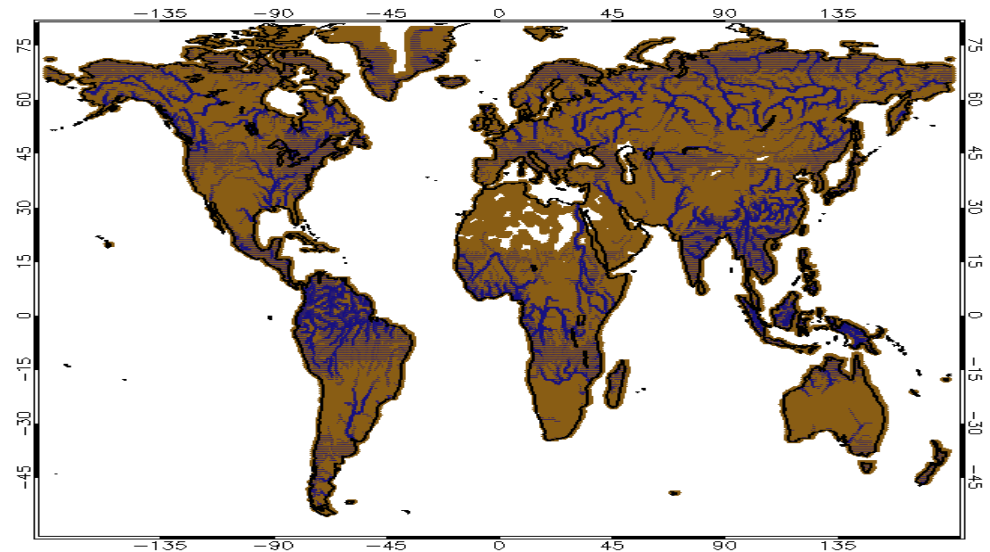
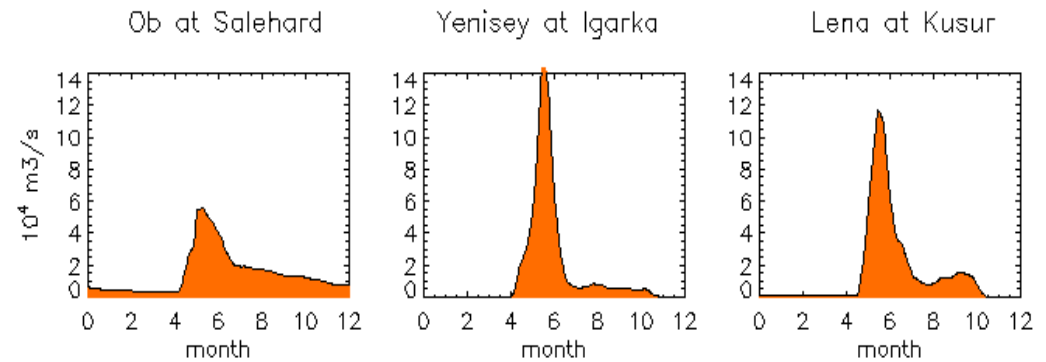
Overall Depth to Bedrock (~1 km resolution)





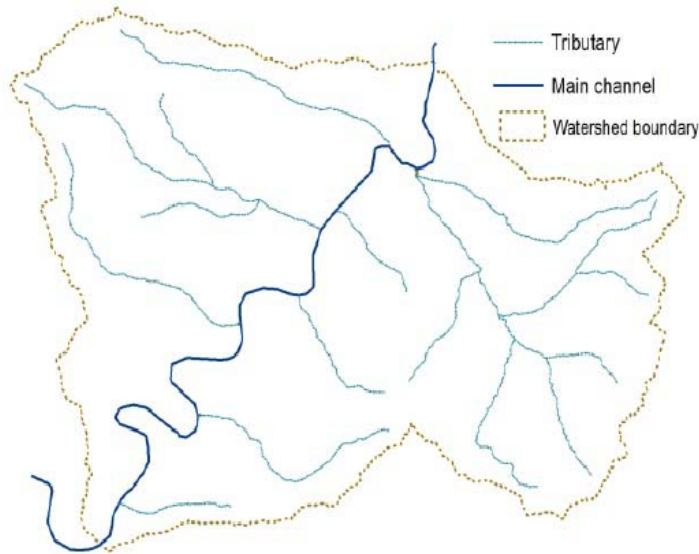
River model

- Routes runoff to the oceans
- Flow directions are obtained from an input dataset
- Calculates water volume and discharge

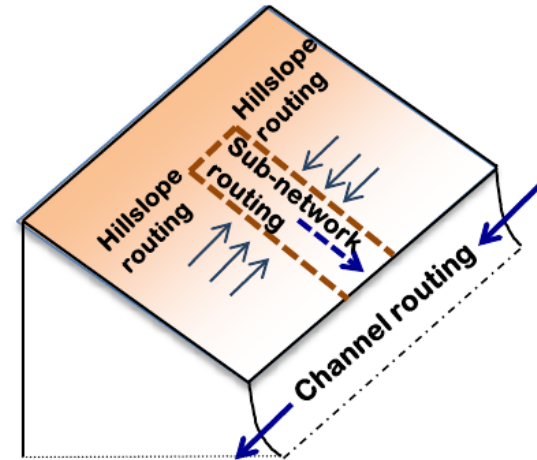


Model for Scale Adaptive River Transport

Real river network



Conceptualized network



- ▶ Hillslope routing accounts for event dynamics and impacts of overland flow on soil erosion, nutrient loading, etc.
- ▶ Sub-network routing: scale adaptive across different resolutions to reduce scale dependence
- ▶ Main channel routing: explicit estimation of in-stream status (velocity, water depth, etc).

(Li et al., JHM, 2013)



Model Validation Tools

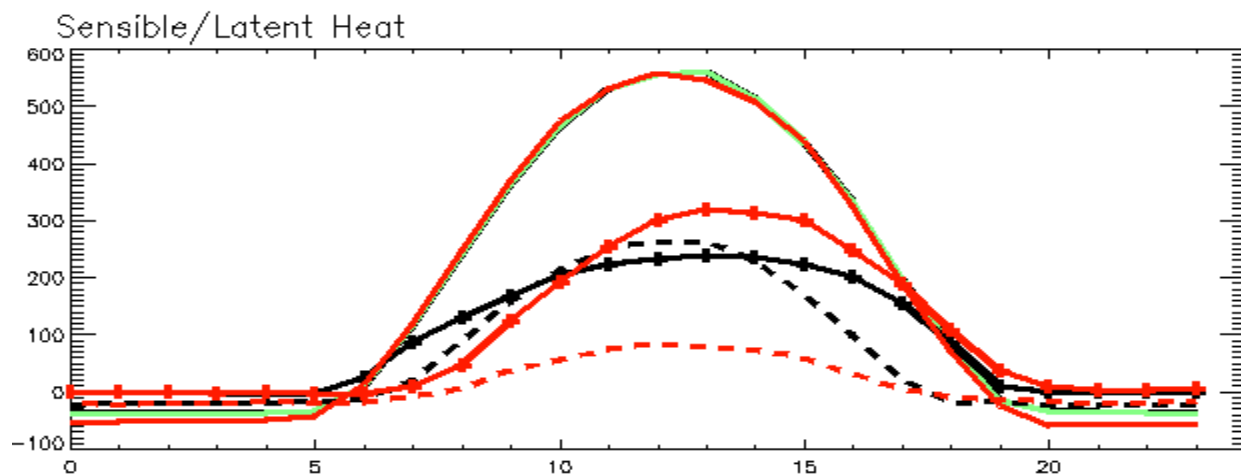
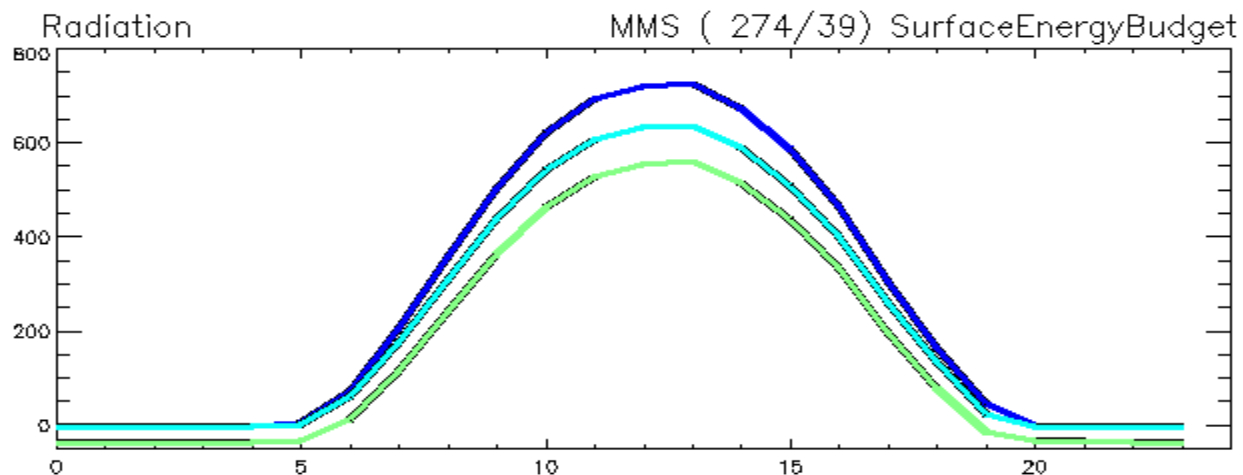
Ideally, should be:

- Global
- Directly comparable to modeled process/state/flux
- Same spatial / temporal scale
- High accuracy
- Long record

In reality, no datasets meeting these criteria exist...



Flux Towers

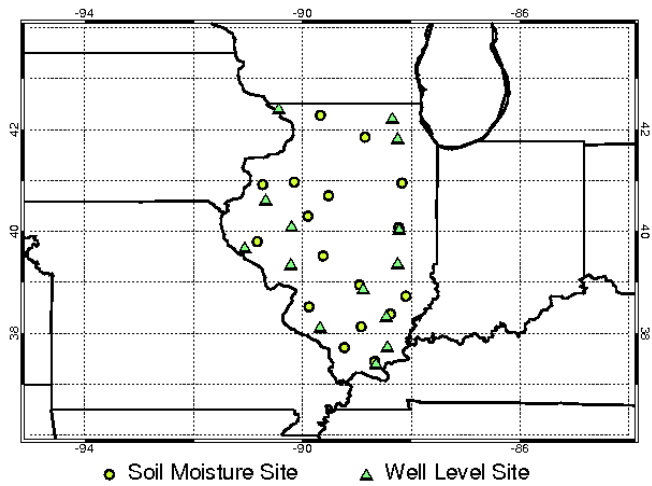


FSDS	—	Tower LH	—
FSA	—	Tower SH	- - -
RNET	—	CLM LH	—
		CLM SH	- - -

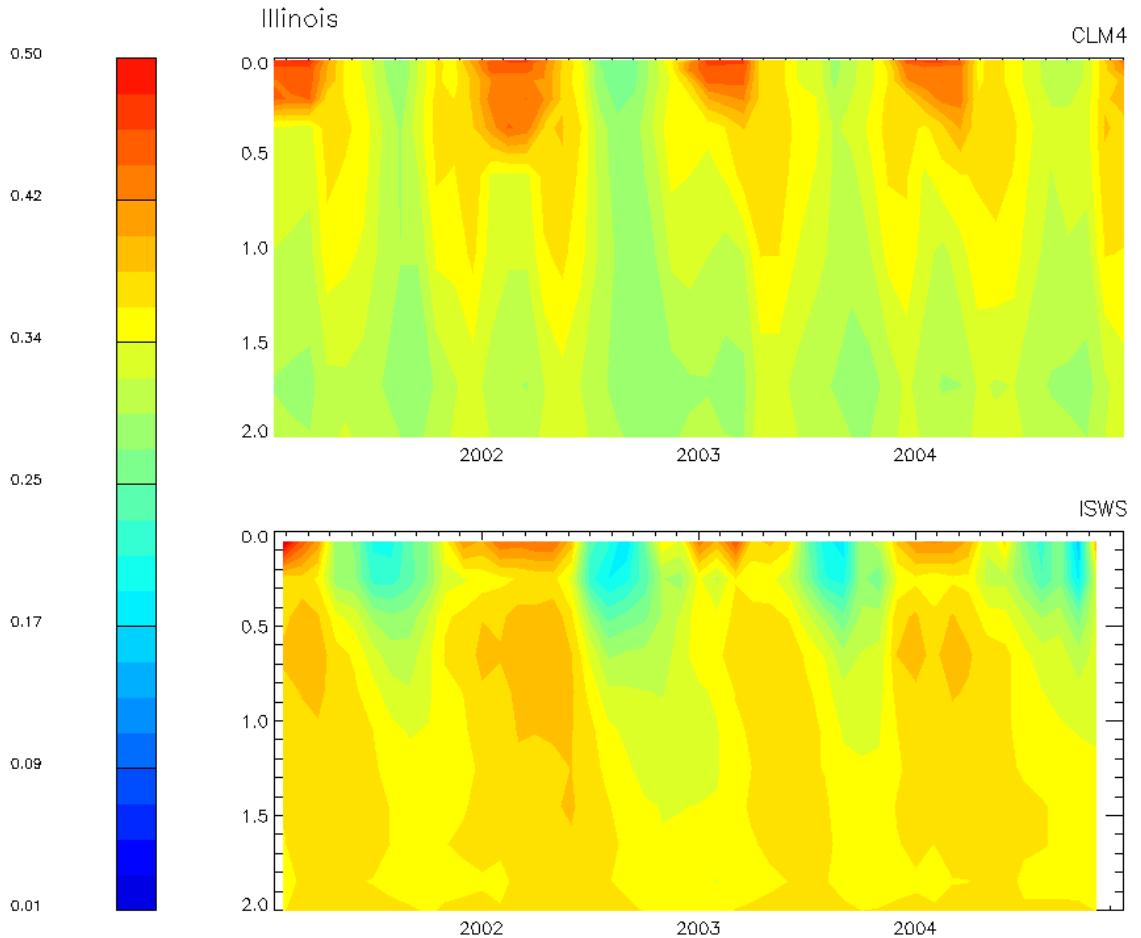


Soil Moisture Networks

Illinois Observations Locations

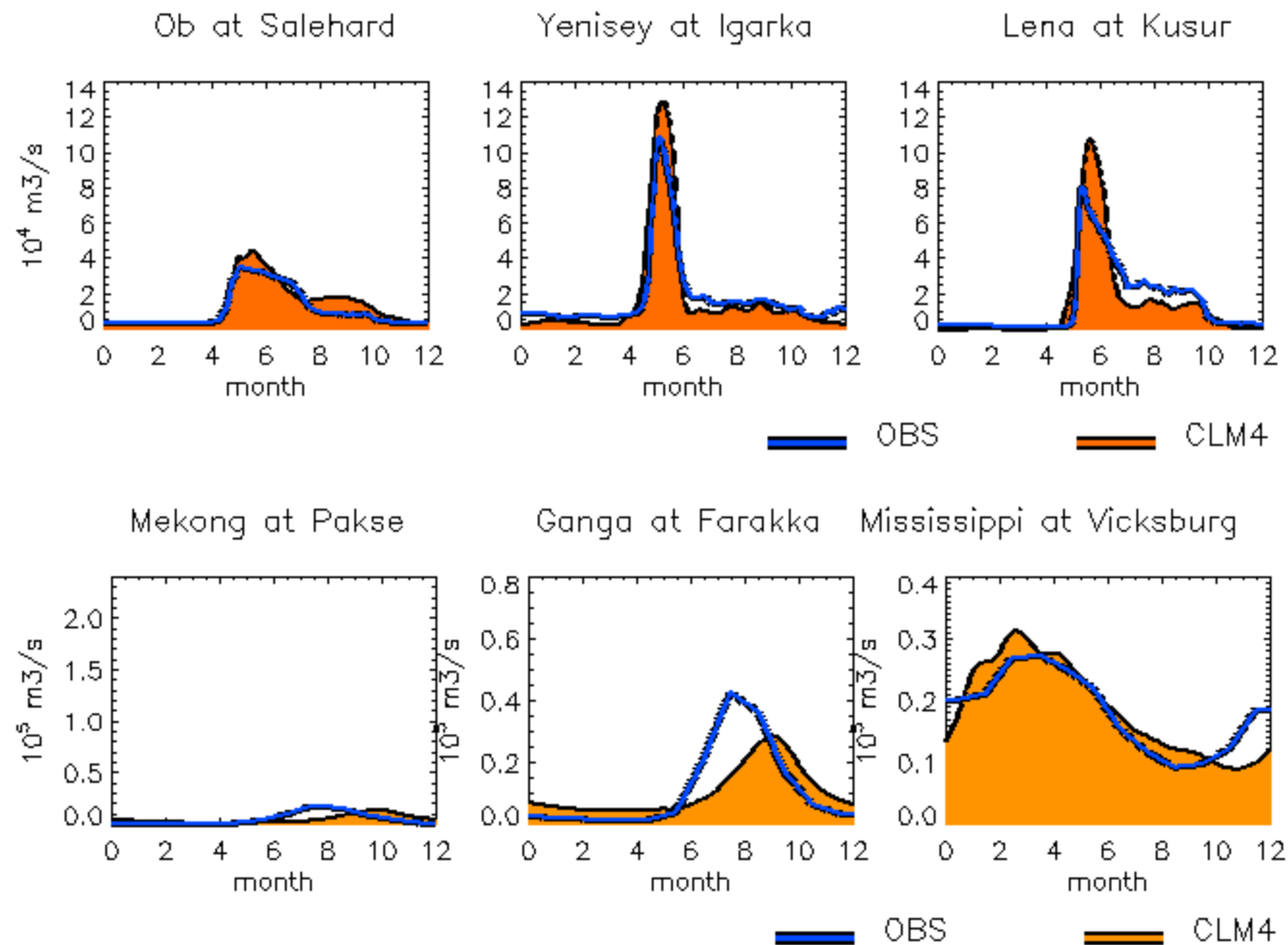


Top panel: CLM soil moisture
Bottom: Observed soil moisture





River Discharge



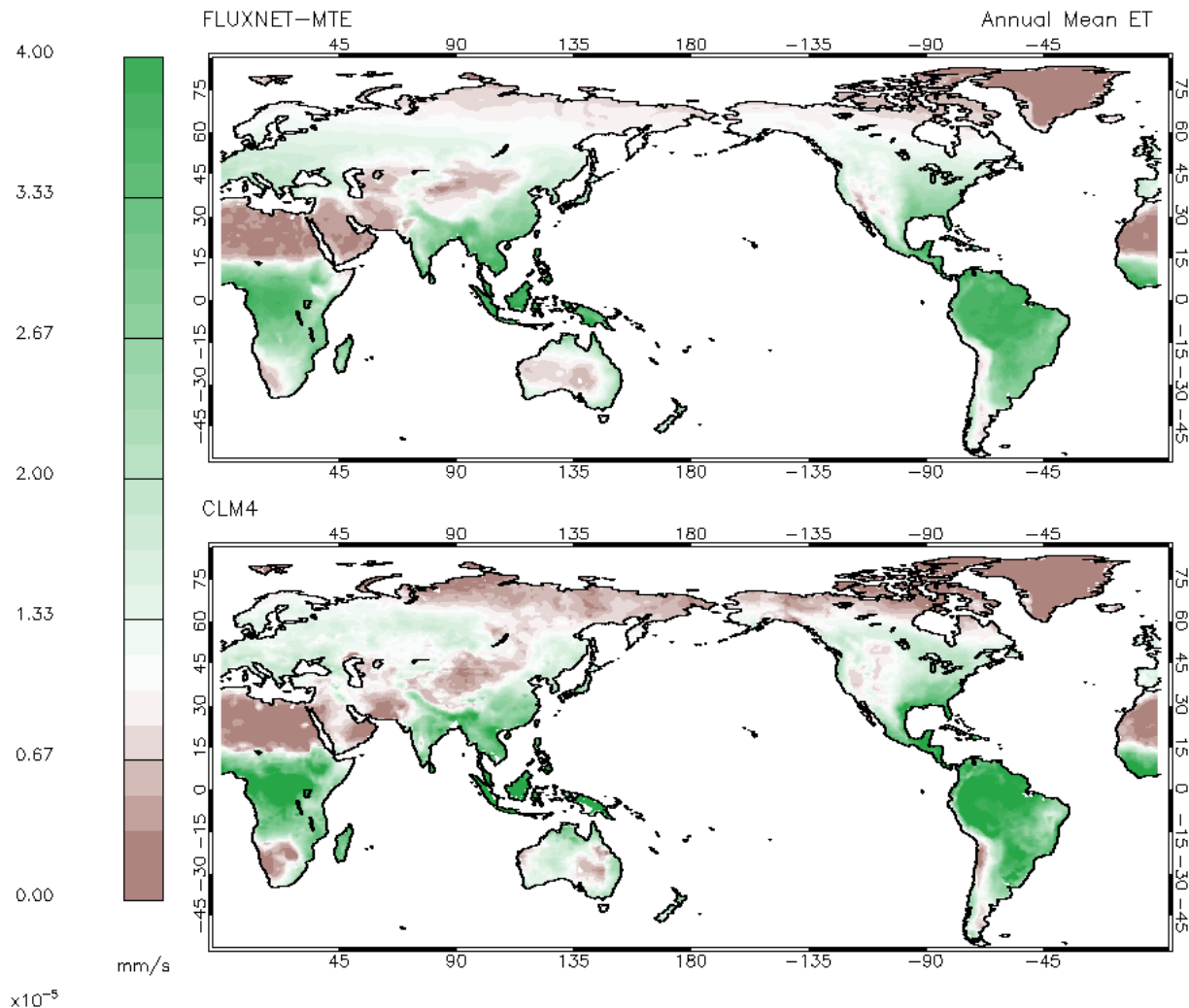


Gridded Model-Data Synthesis Products

Examples: FLUXNET-MTE, FLUXCOM

Top panel: FLUXNET-MTE
Bottom: CLM

Annual Mean Evapotranspiration

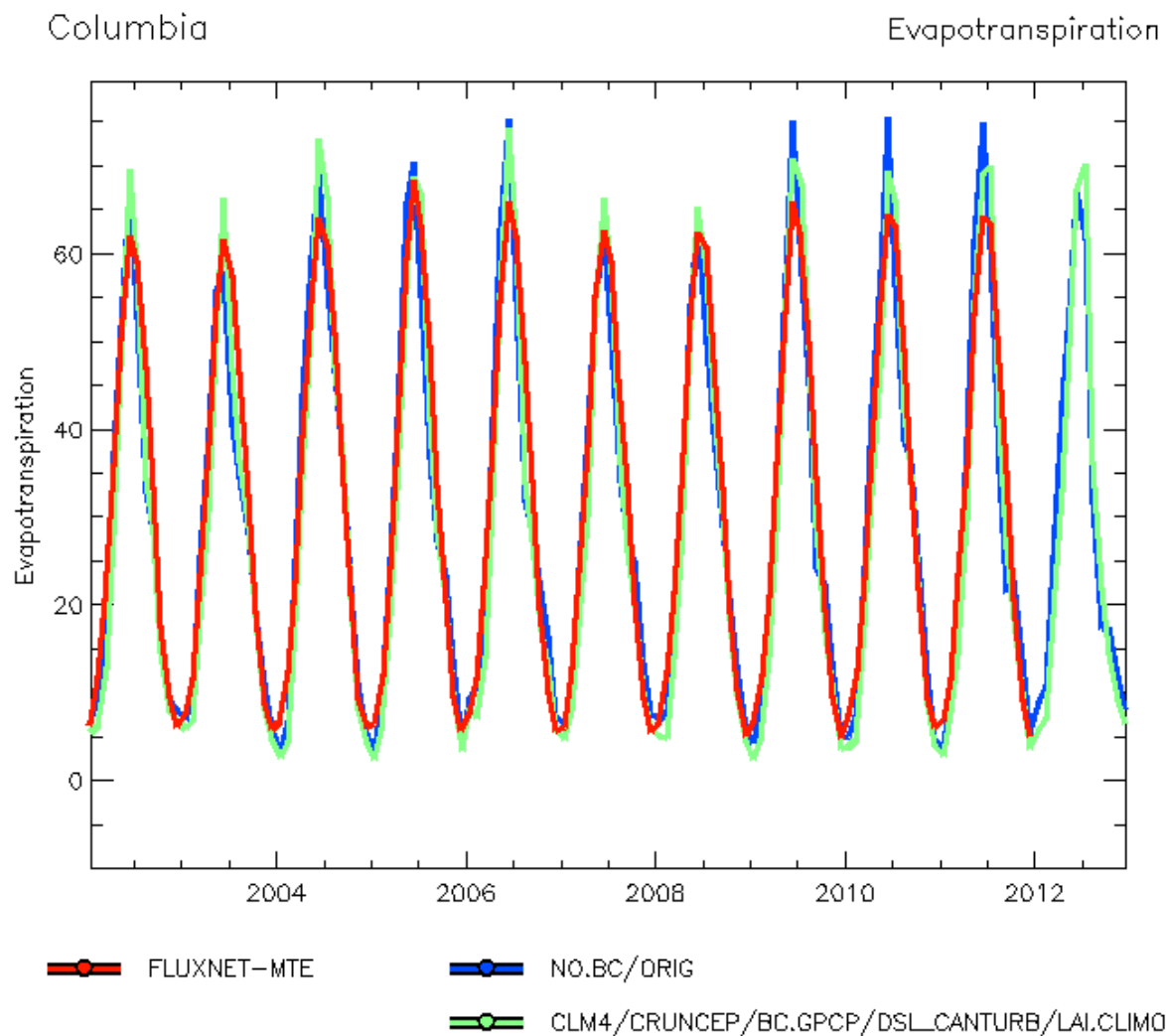




FLUXNET-MTE

Columbia River Basin Evapotranspiration

Red: FLUXNET-MTE
Blue/Green: CLM



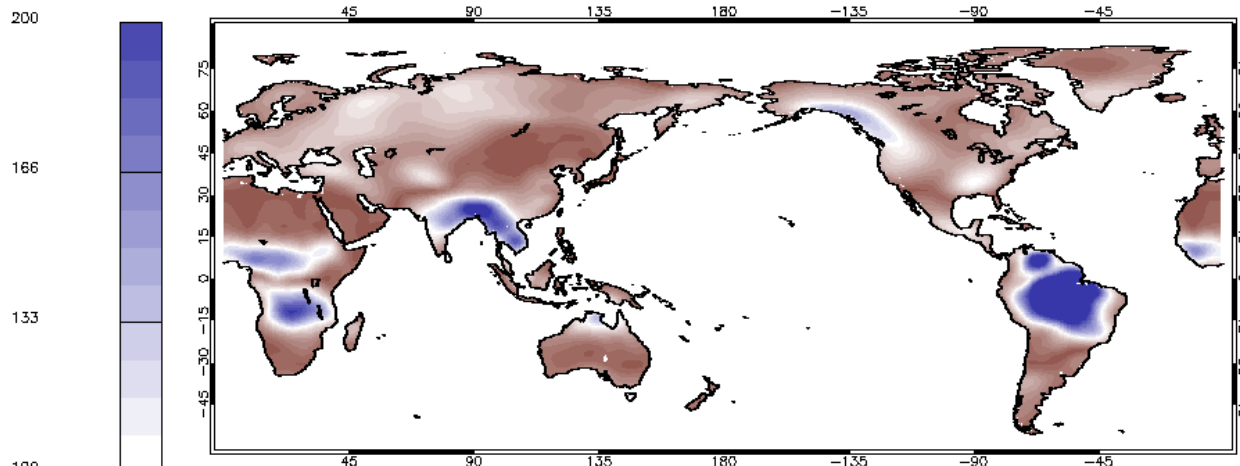


GRACE Total Water Storage

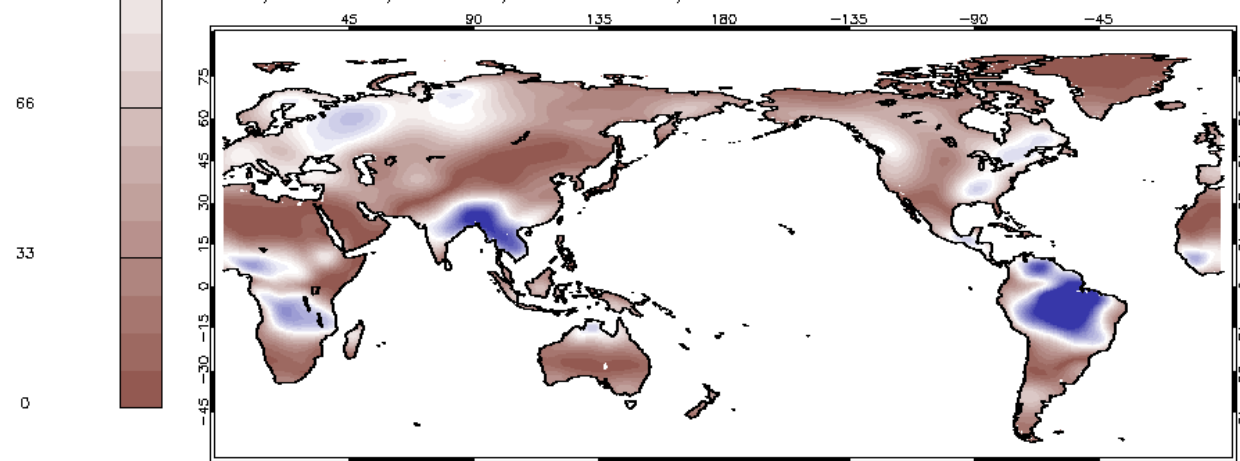
Mean Annual Amplitude of Total Water Storage

GRACE CSR RL05

Mean Annual Amplitude TWS



CLM4/CRUNCEP/BC.GPCP/DSL_CANTURB/LAI.CLIMO



mm

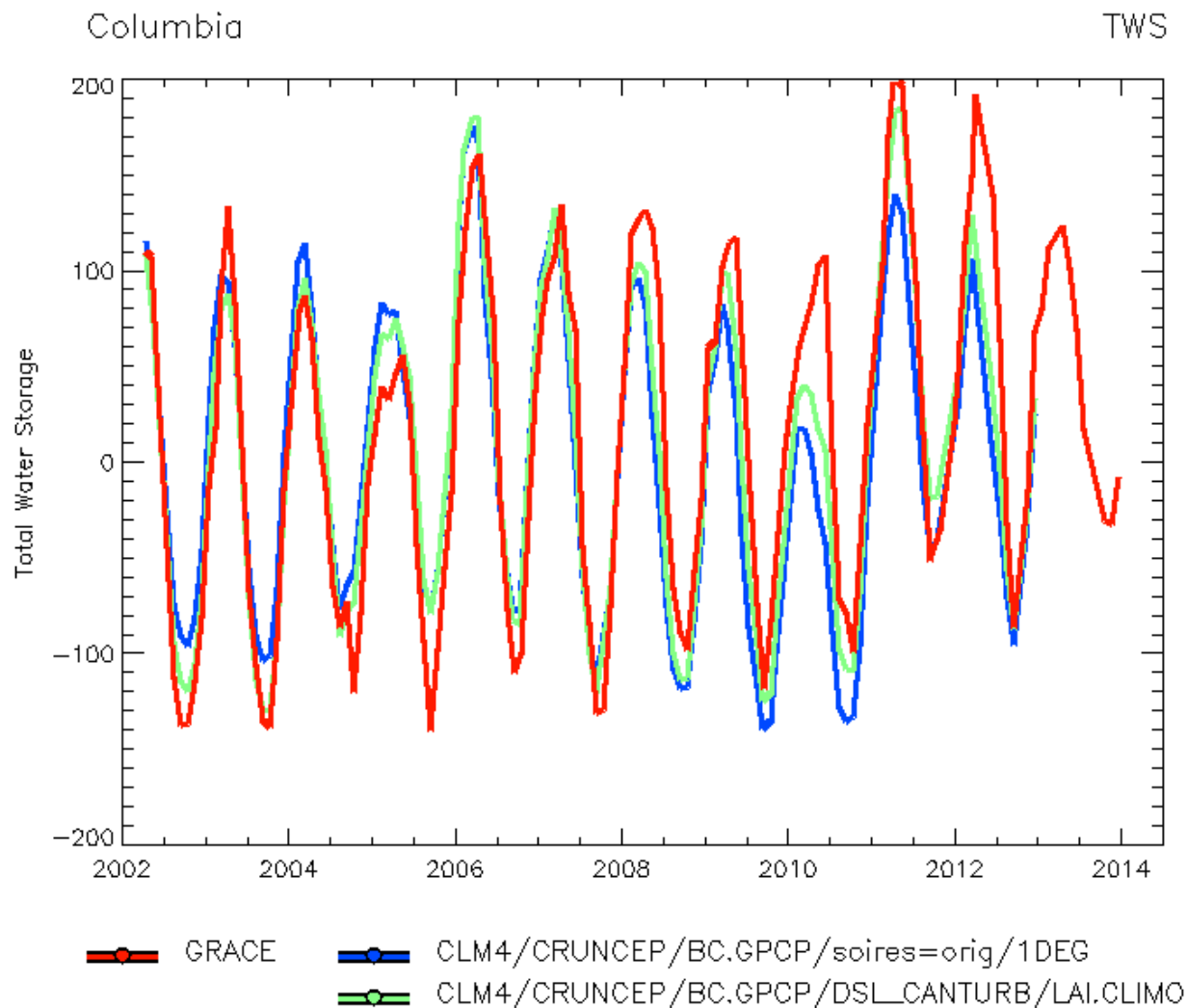
Top panel: GRACE
Bottom: CLM



GRACE Total Water Storage

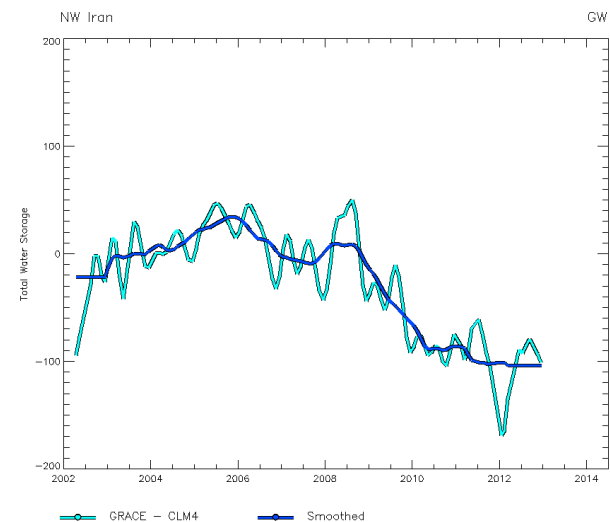
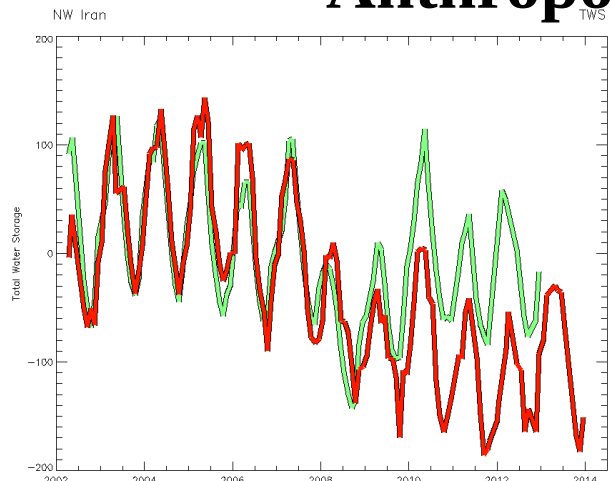
Columbia River Basin
Total Water Storage

Red: GRACE
Blue/Green: CLM



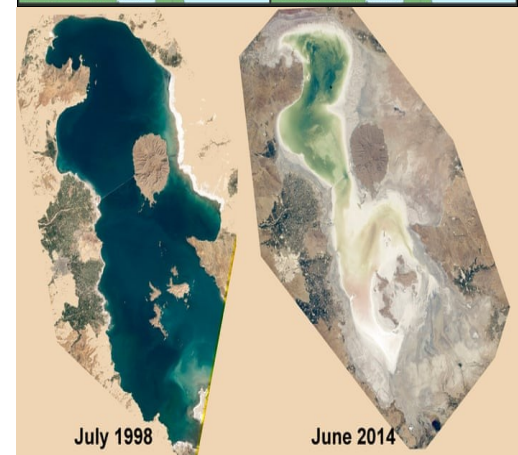
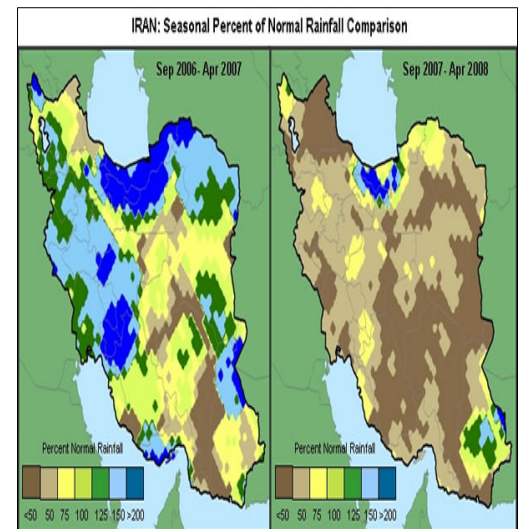


CLM Application Example: Anthropogenic Groundwater Withdrawal



Human-induced groundwater changes can be estimated by removing the CLM estimate of TWS from the GRACE estimate of TWS

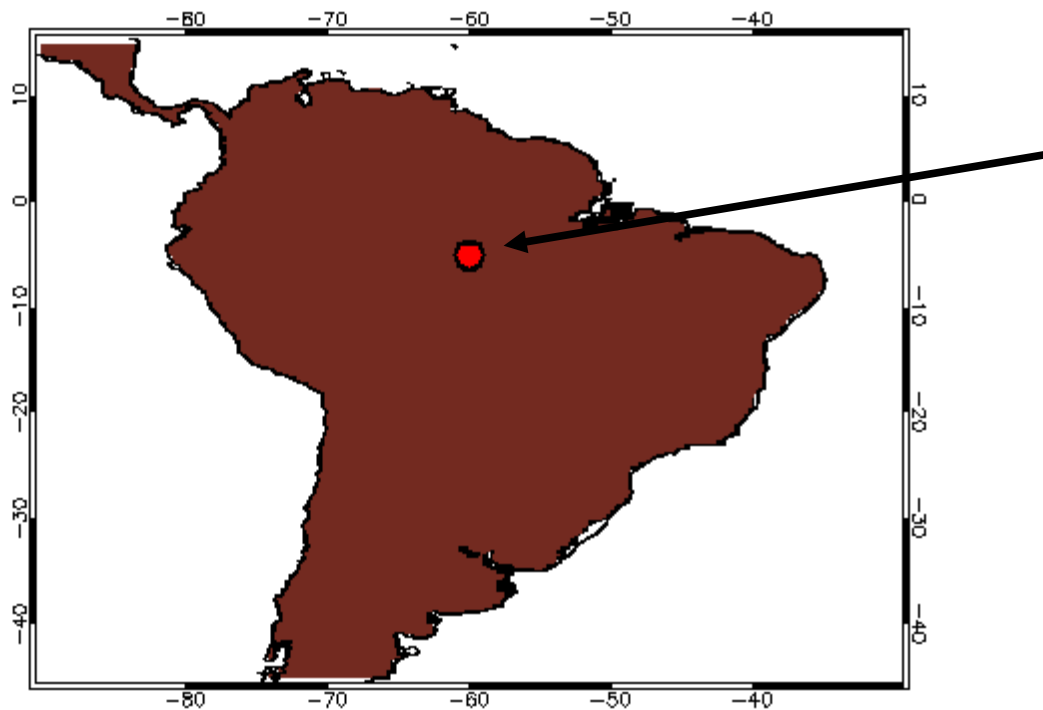
- GRACE TWS
- CLM TWS
- Groundwater





Example I

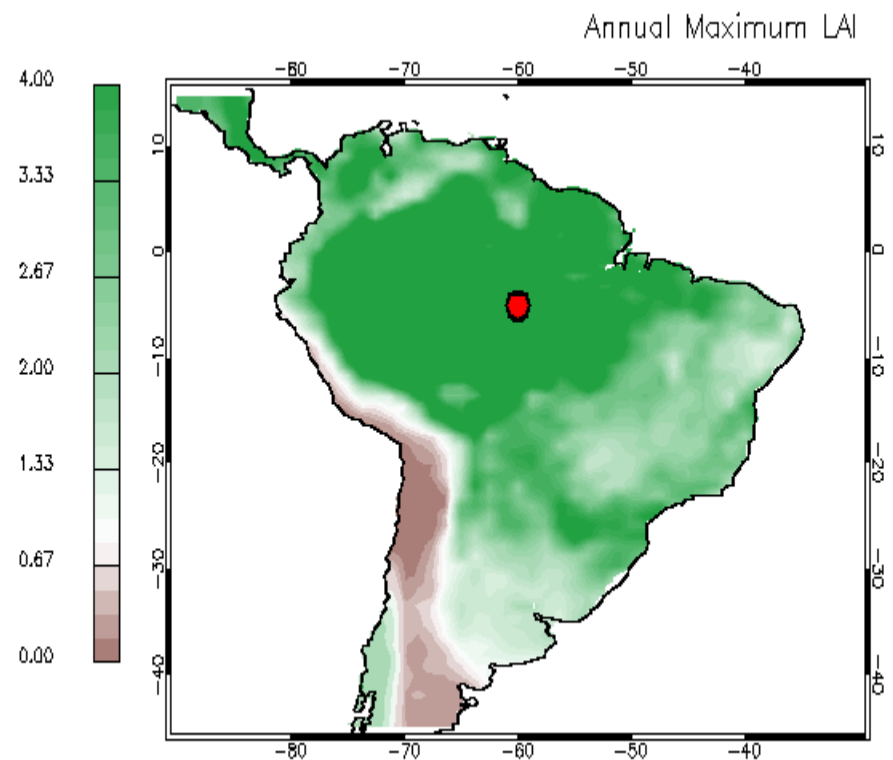
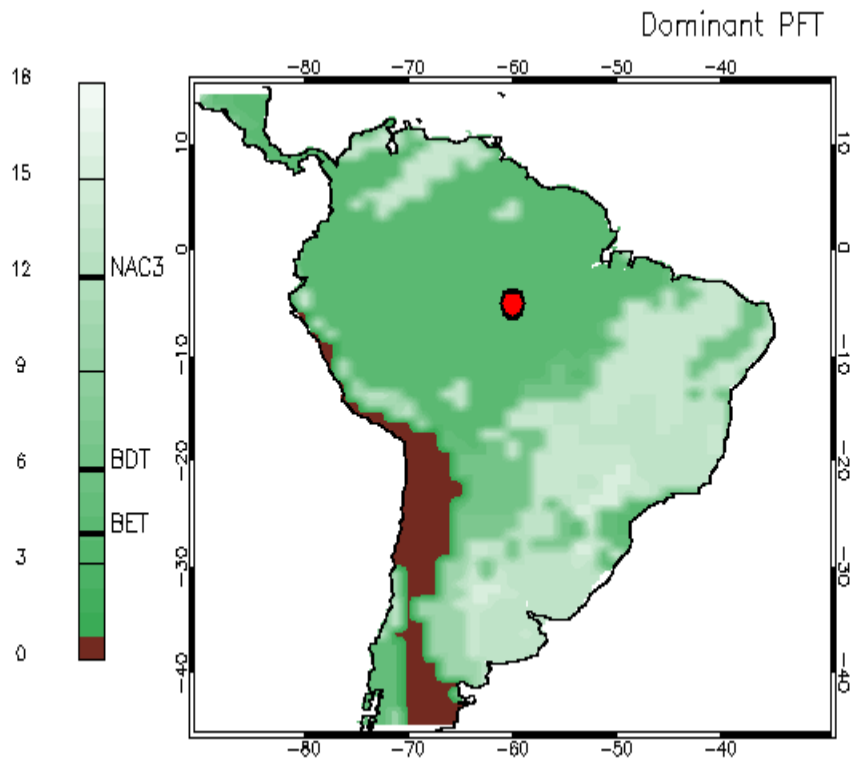
Effects of Parameter Change



60W / 5S

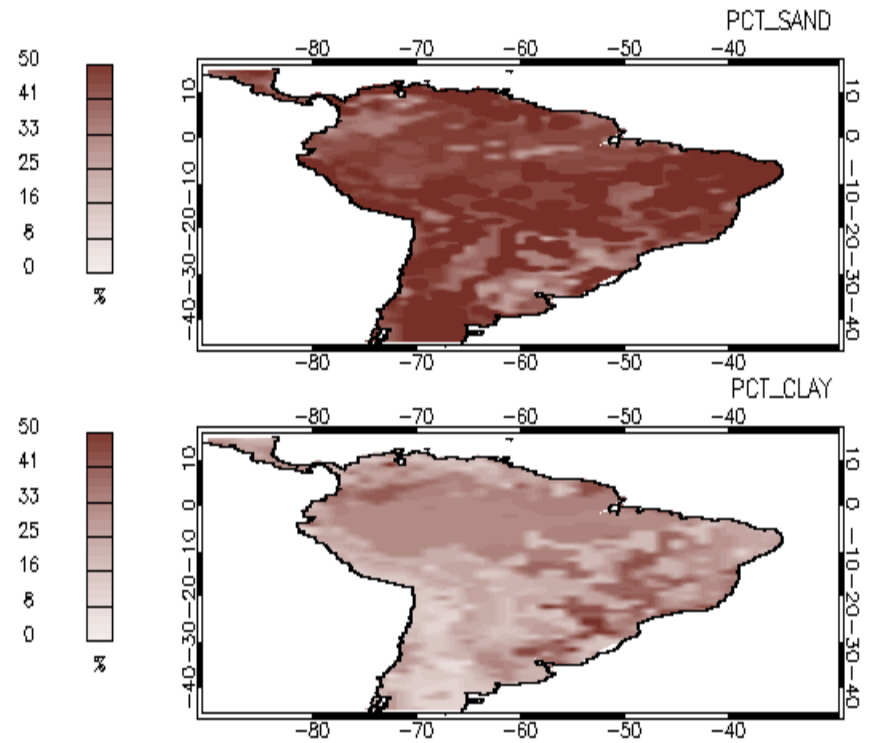
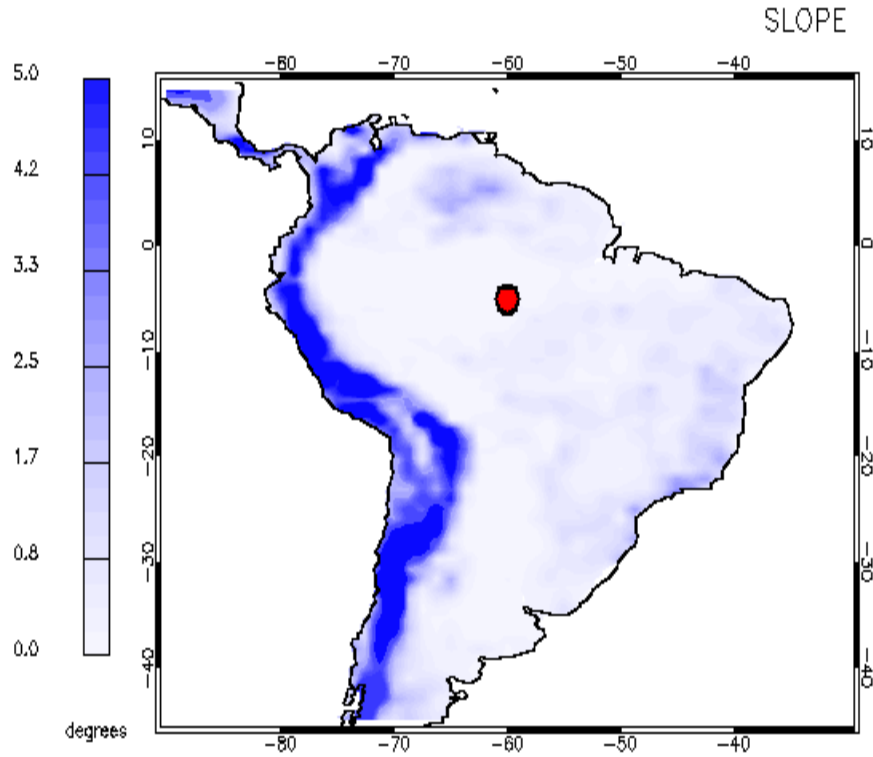


Hydrologically Relevant Surface Data





Hydrologically Relevant Surface Data

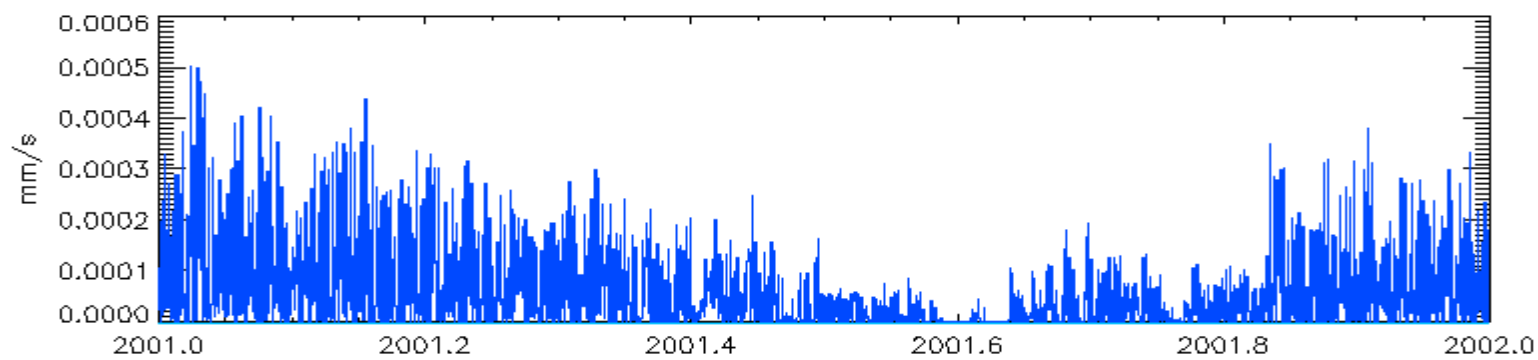




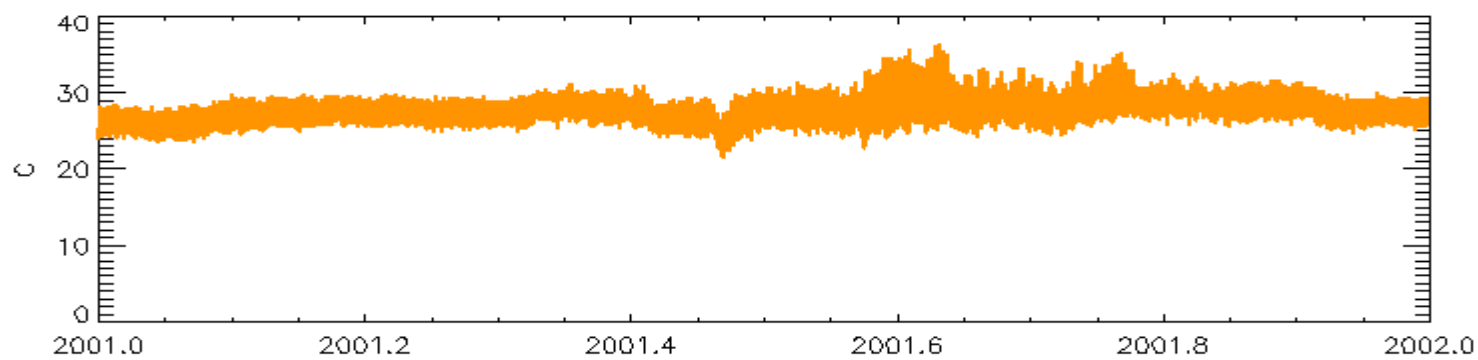
Time Series

lon:300.0/lat:-5.2

Precipitation



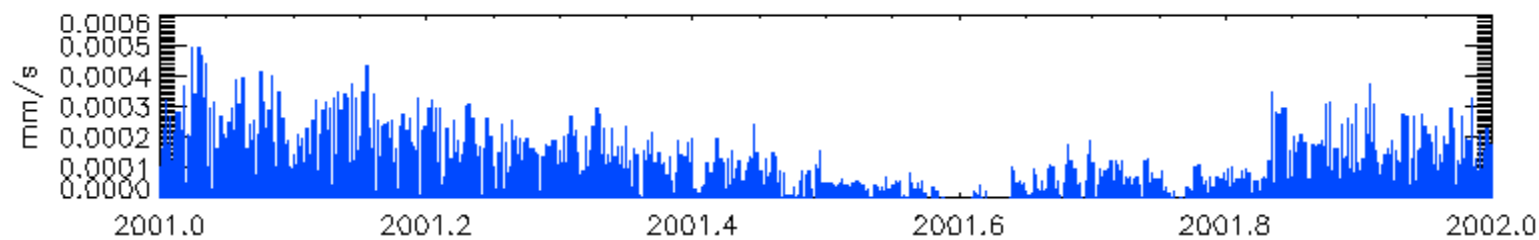
Air Temperature



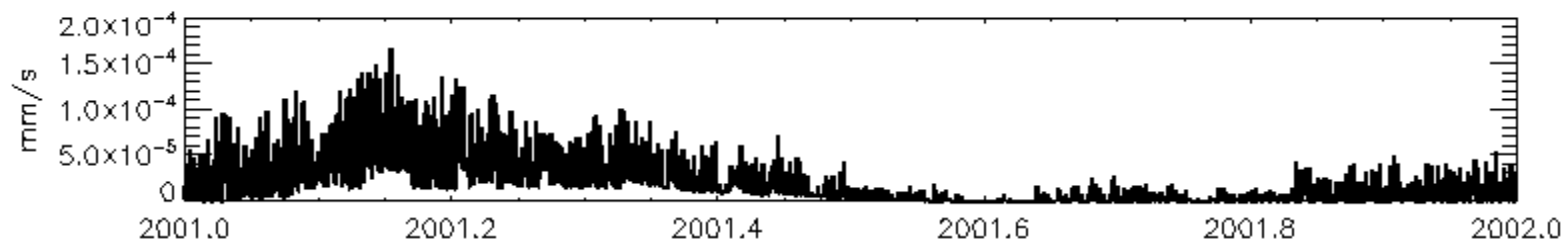


lon:300.0/lat:-5.2

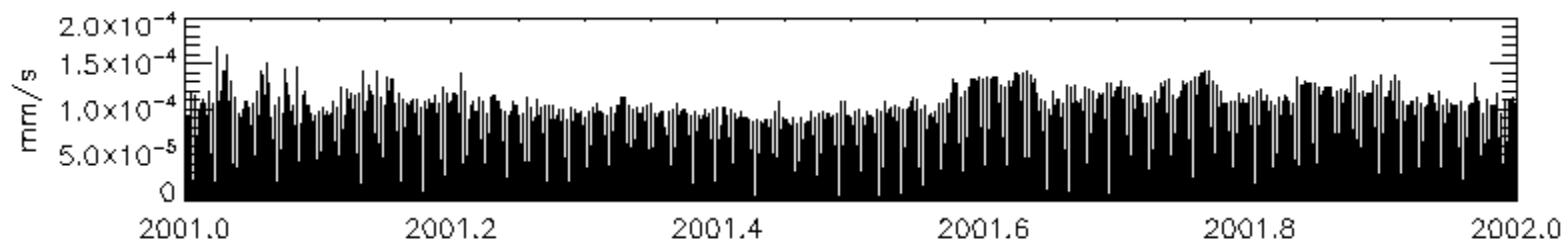
Precipitation



Runoff



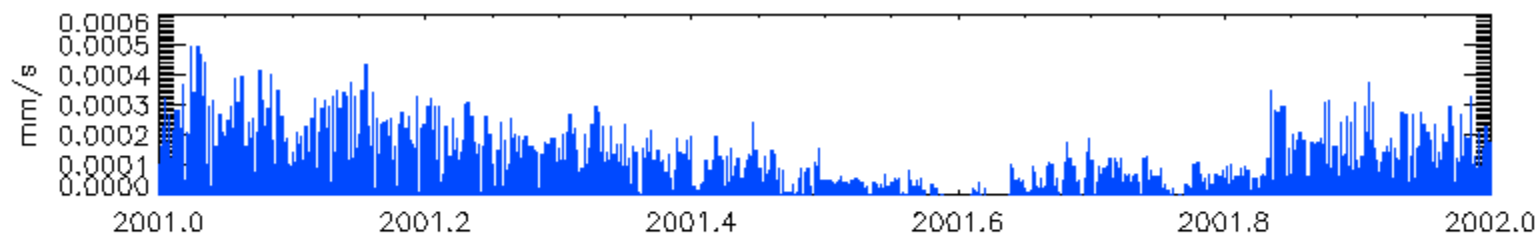
Evapotranspiration



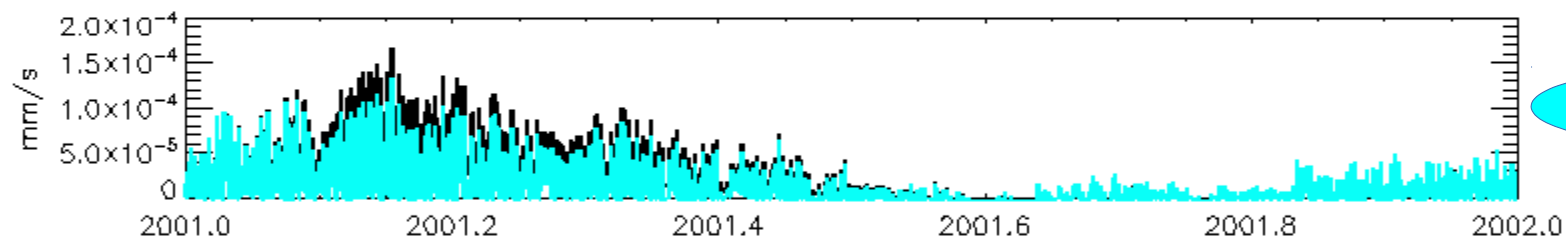


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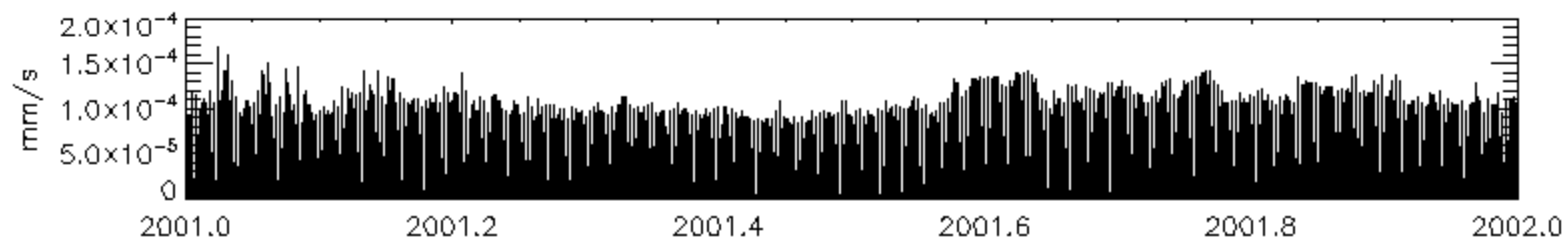
Precipitation



Runoff



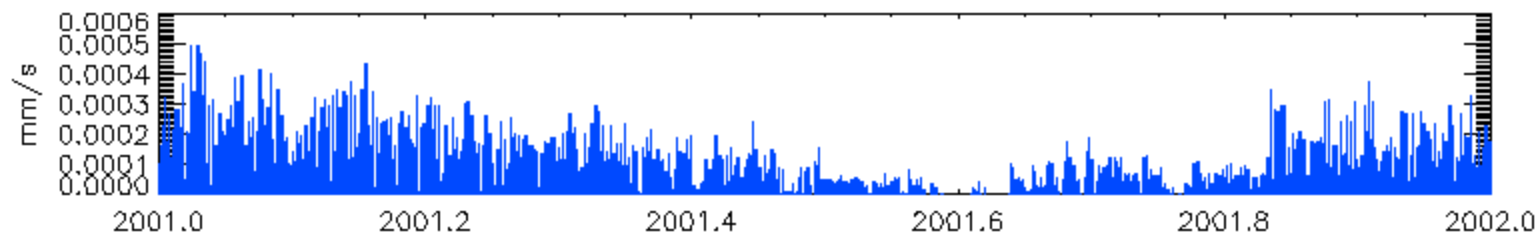
Evapotranspiration



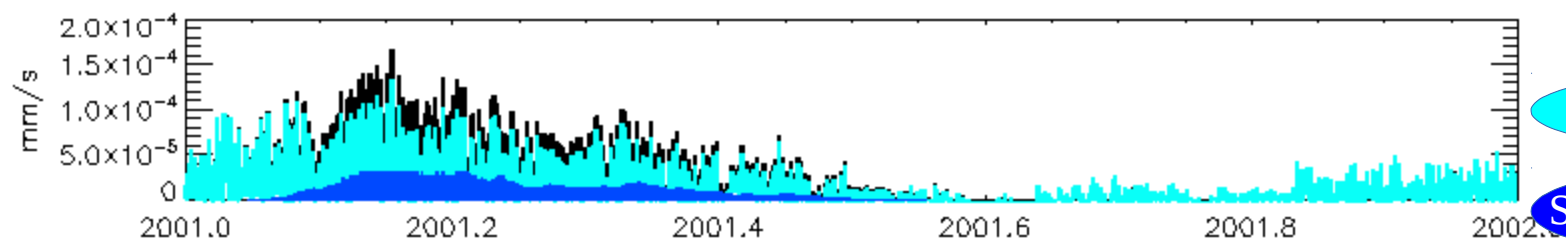


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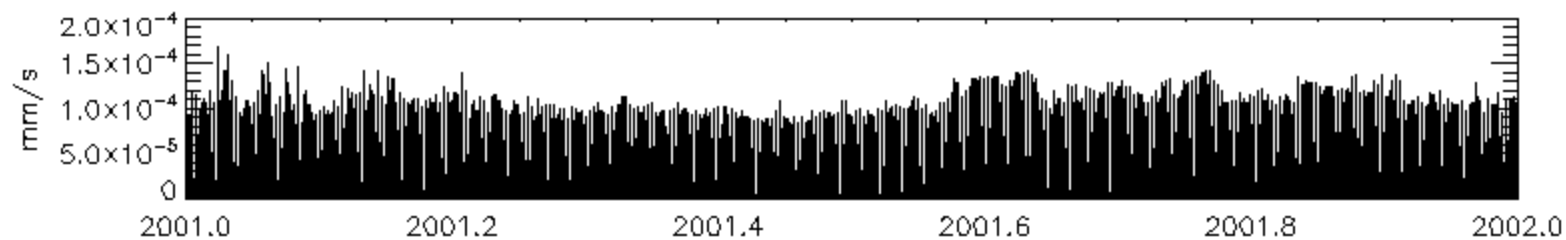
Precipitation



Runoff



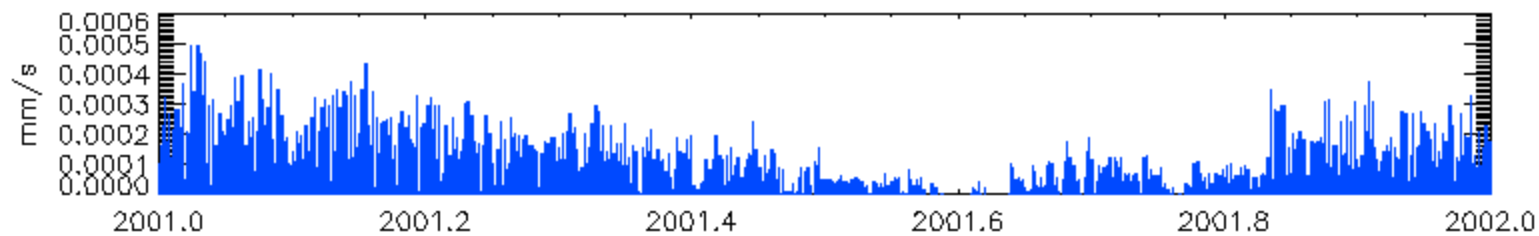
Evapotranspiration



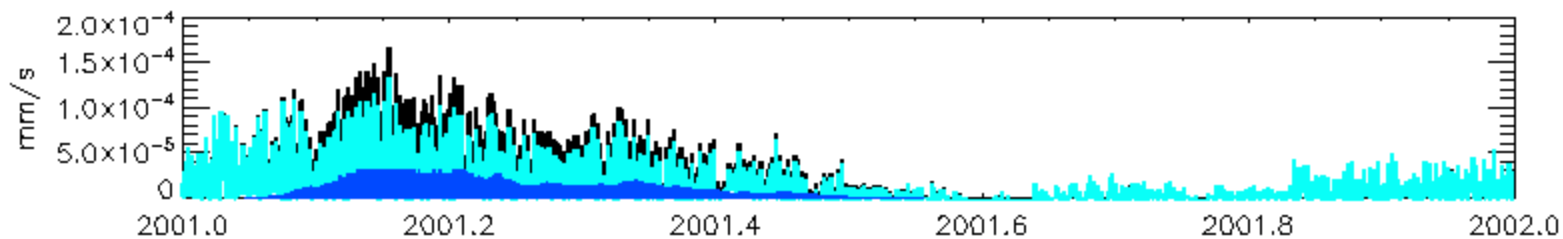


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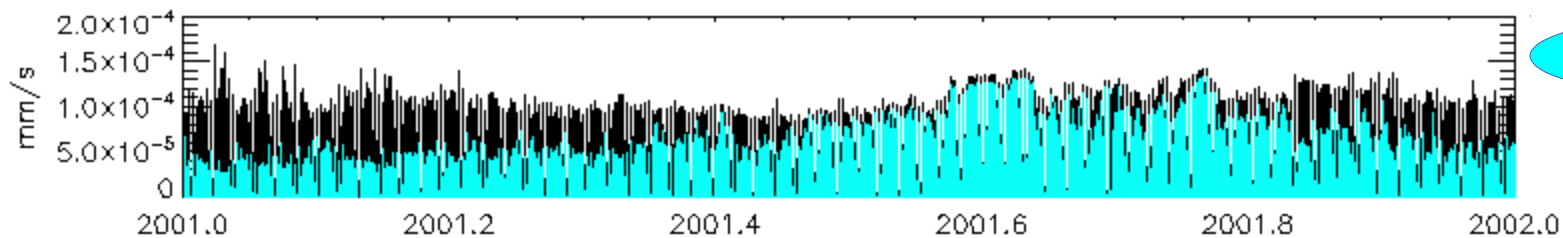
Precipitation



Runoff



Evapotranspiration



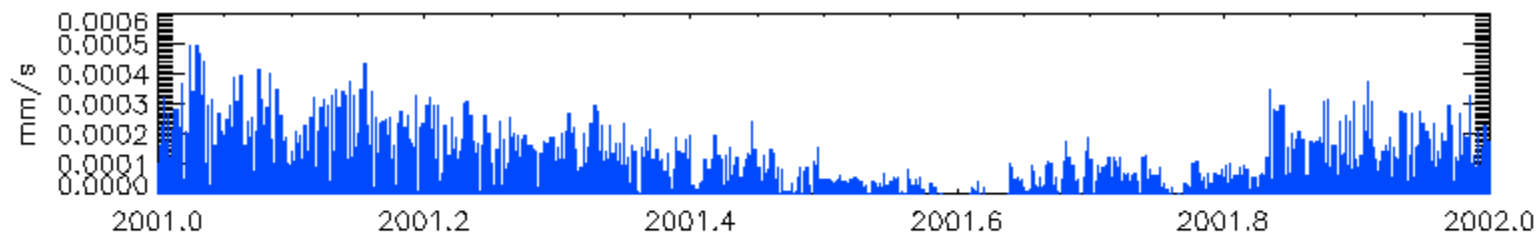
Total ET

Transpiration

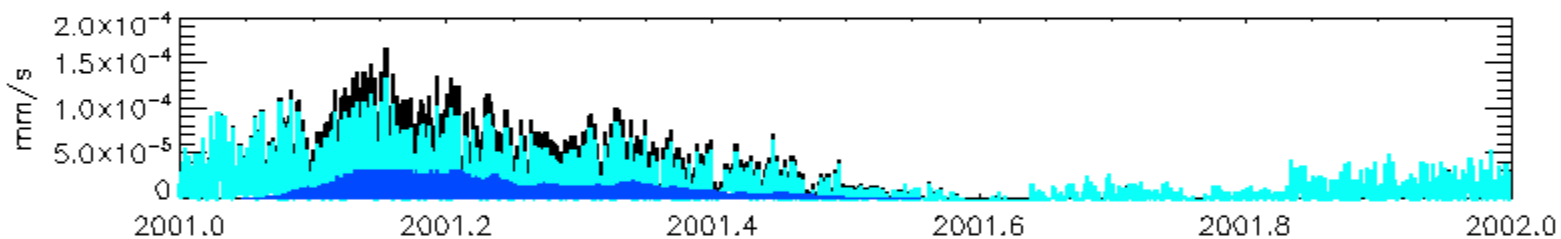


lon:300.0/lat:-5.2

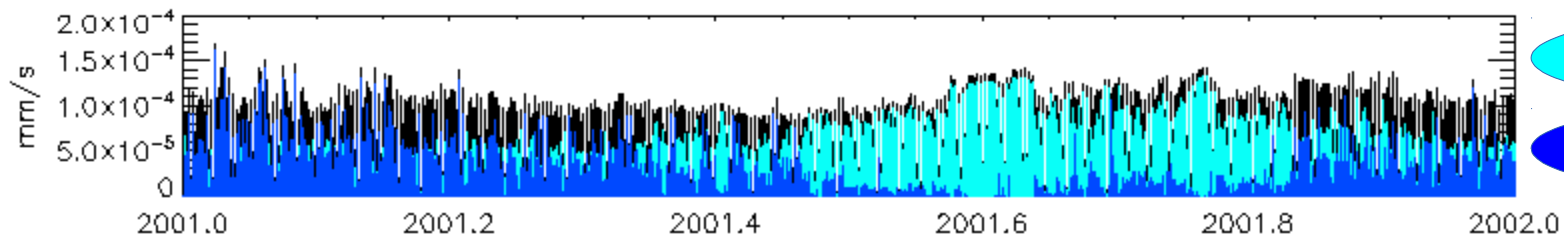
Precipitation



Runoff



Evapotranspiration



Total ET

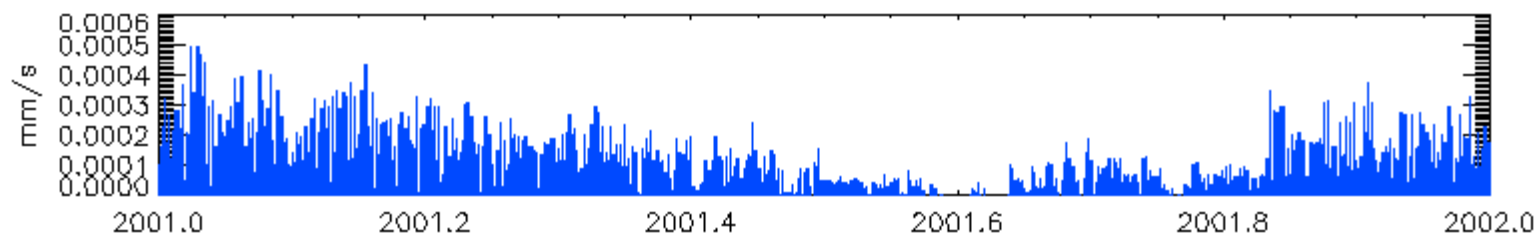
Transpiration

Canopy Evap

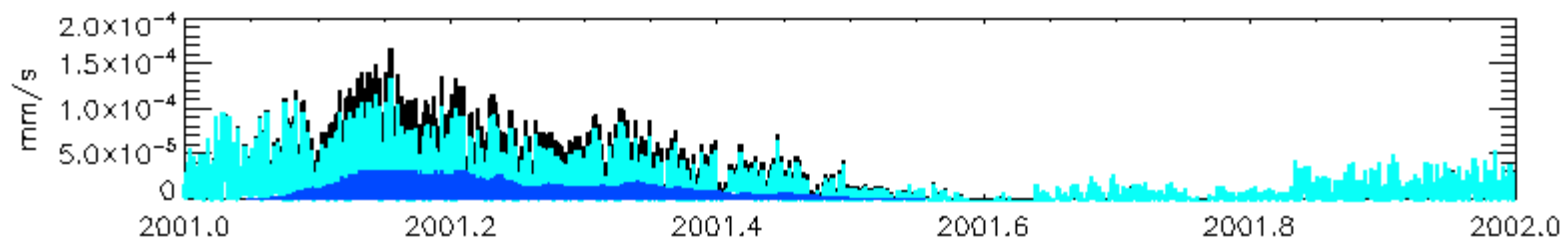


lon:300.0/lat:-5.2

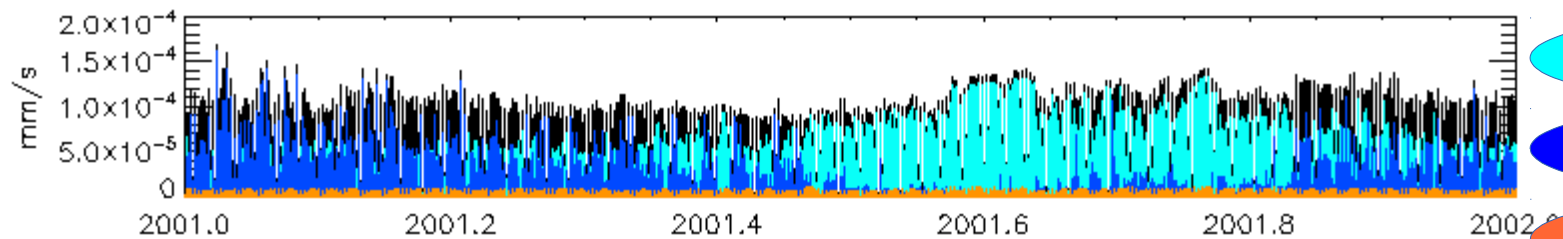
Precipitation



Runoff



Evapotranspiration



Total ET

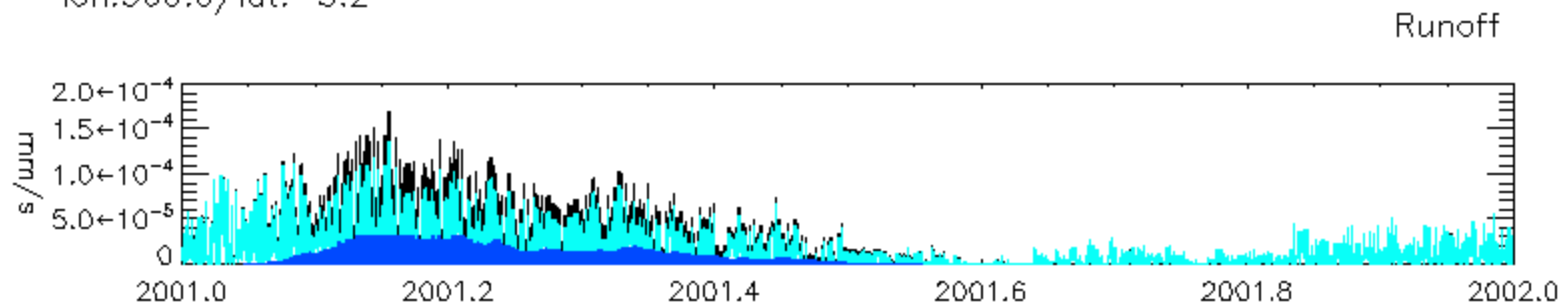
Transpiration

Canopy Evap

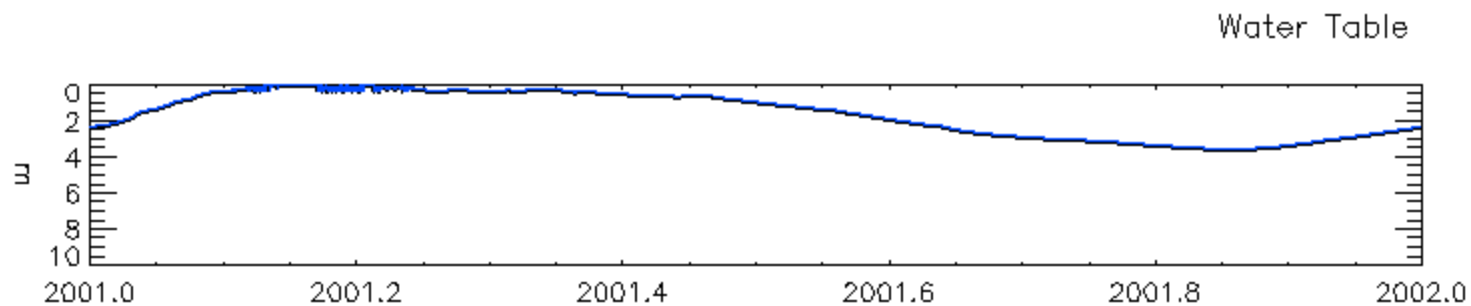
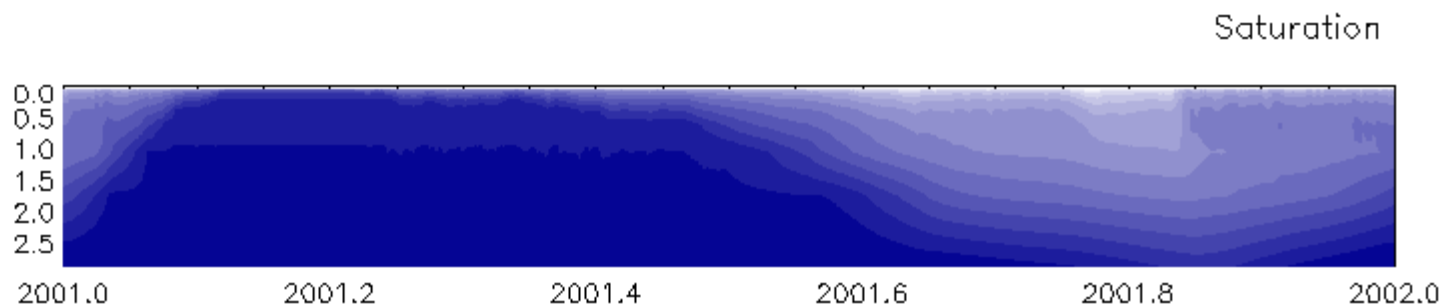
Soil Evap



lon:300.0/lat:-5.2



1.00
0.85
0.70
0.55
0.40
0.25
0.10





Example: Effects of Modifying the Water Table

$$\Delta ZWT = Q_{\text{drainage}} - Q_{\text{recharge}}$$

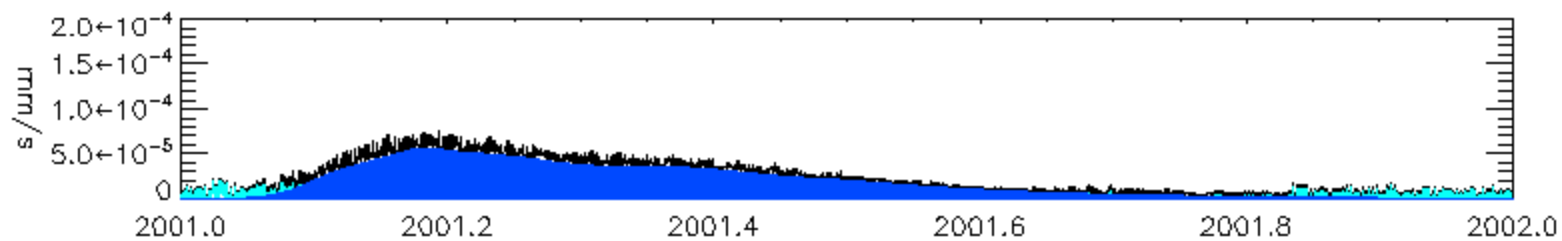
$$Q_{\text{drainage}} = A \exp(-f z)$$

$$Q_{\text{surface}} = F \exp(-g z) P_{\text{throughfall}}$$

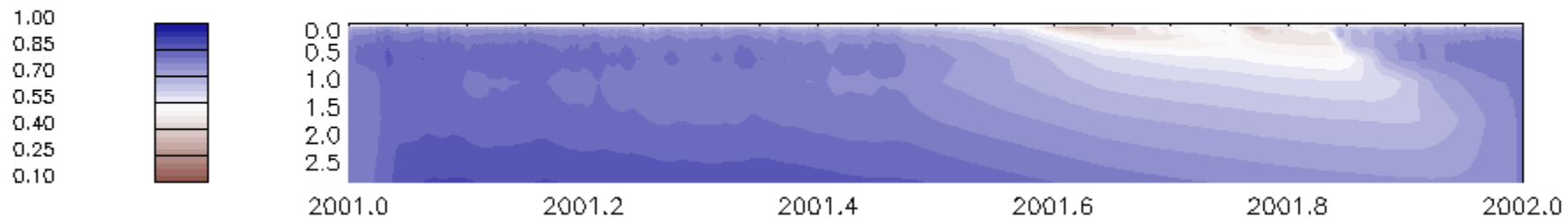


lon:300.0/lat:-5.2

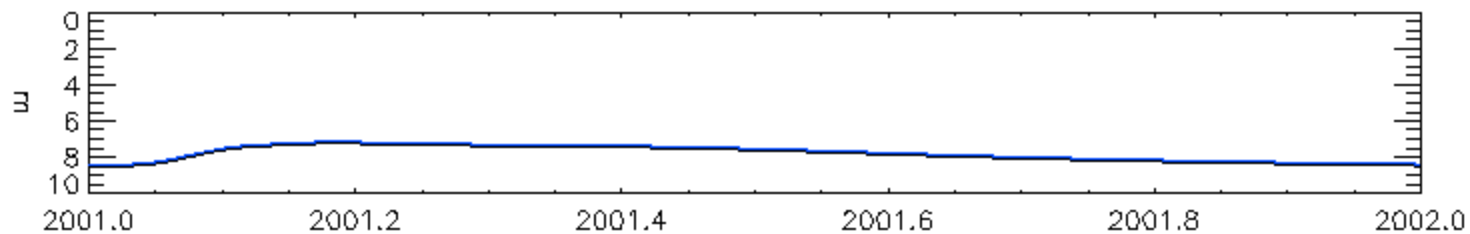
Runoff



Saturation



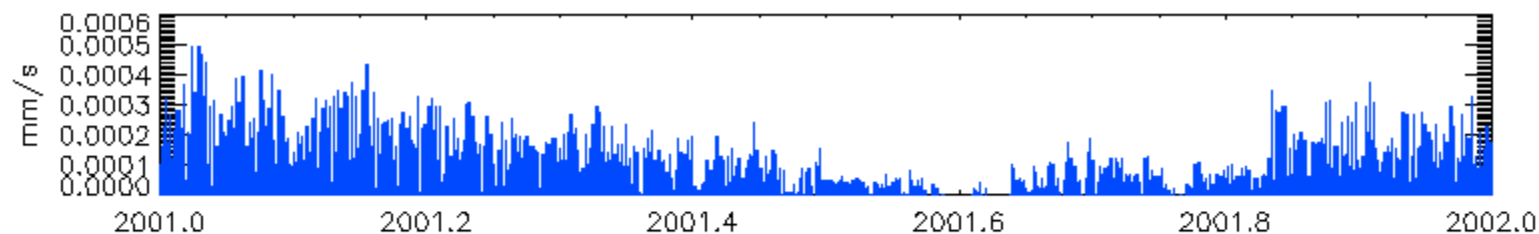
Water Table



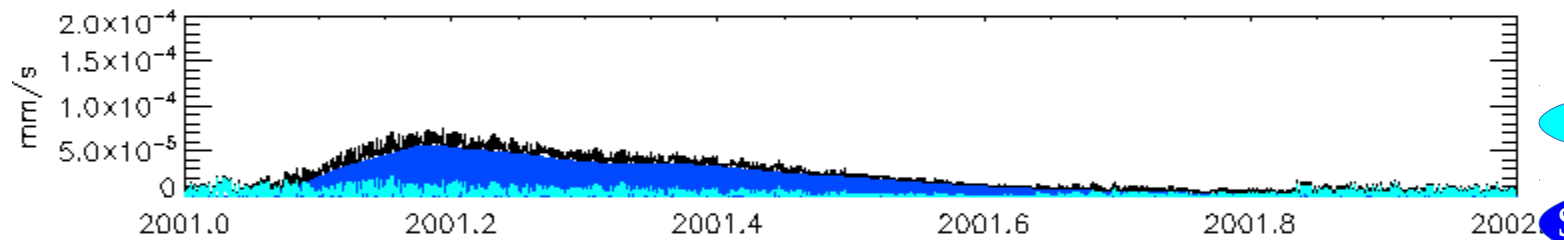


lon:300.0/lat:-5.2

Precipitation



Runoff

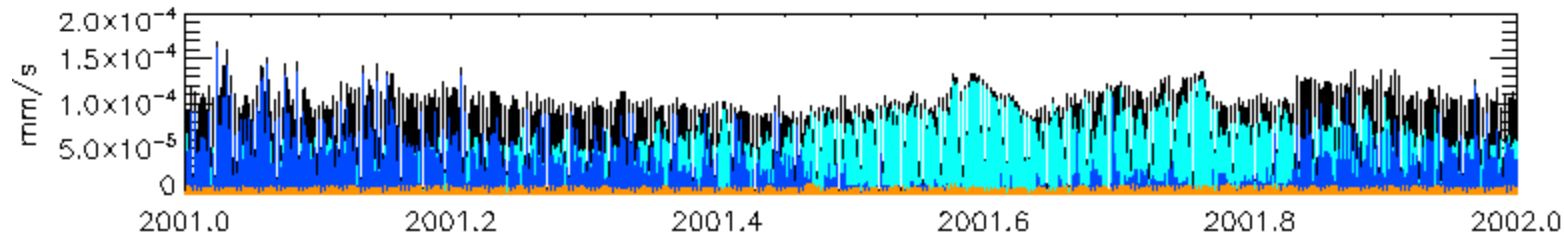


Total Runoff

Surface Runoff

Subsurface Runoff

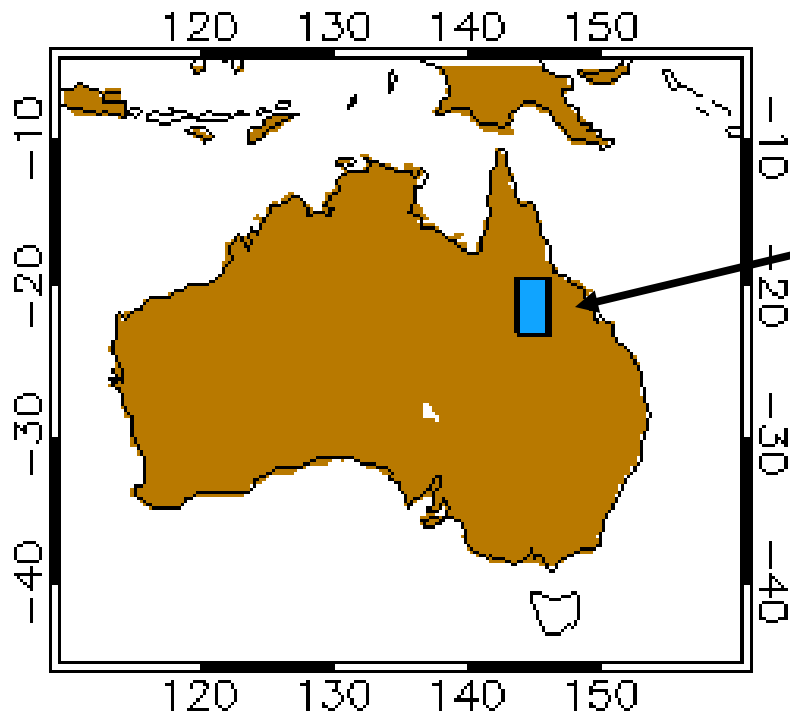
Evapotranspiration





Example II

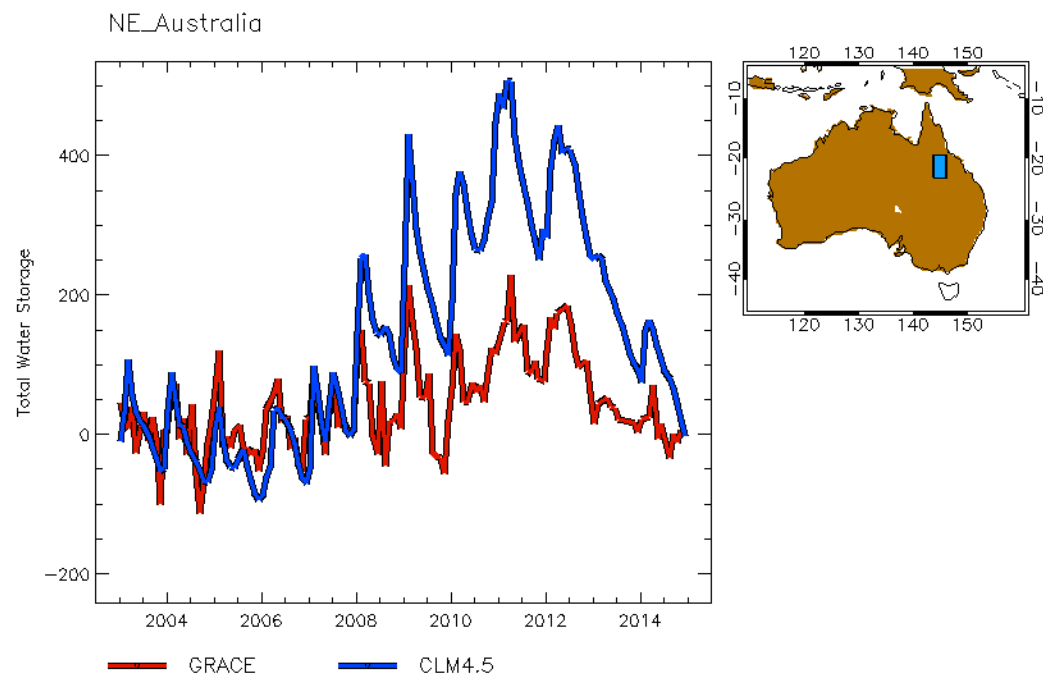
Model Structural Change



NE Australia

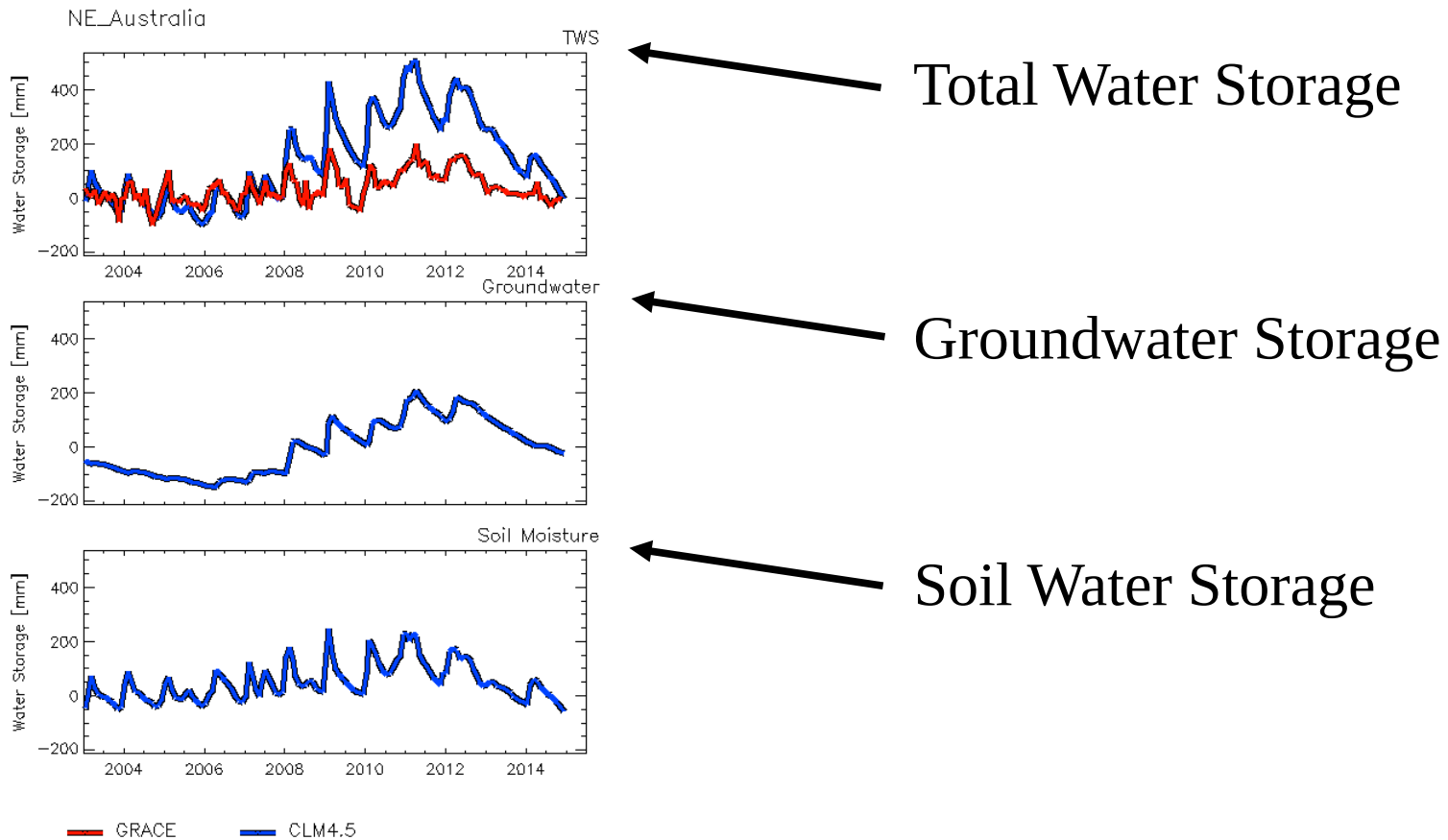


GRACE Water Storage Comparison



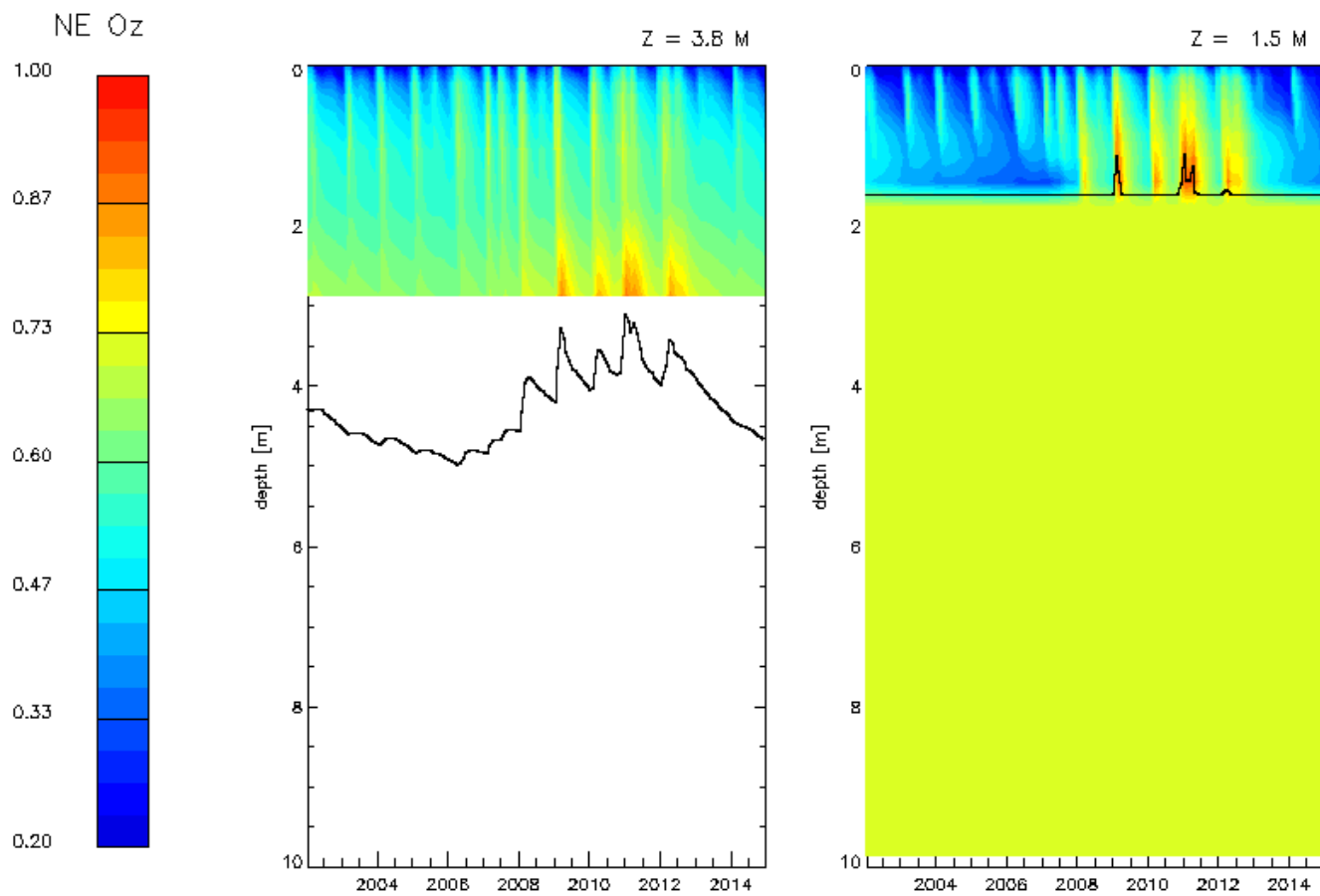


GRACE Water Storage Comparison



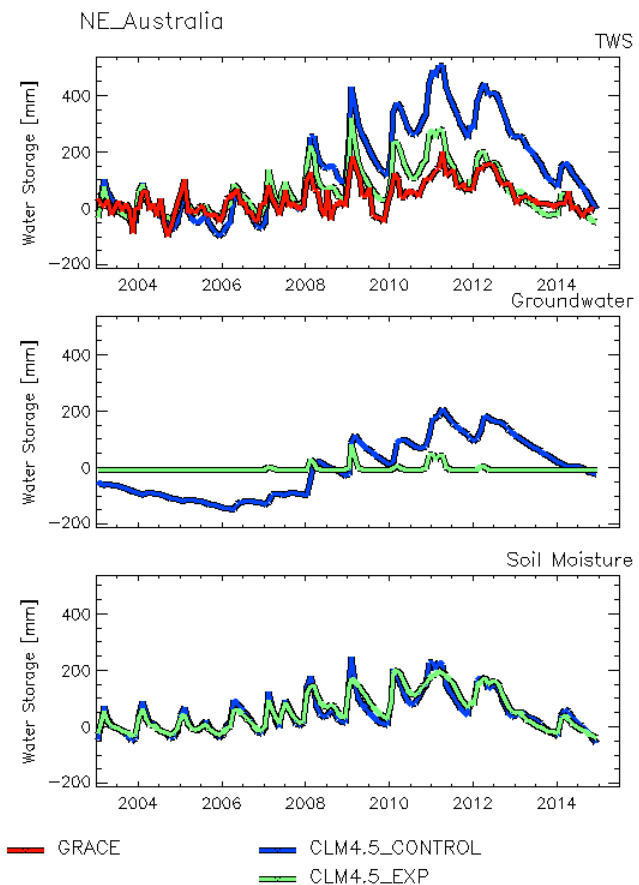
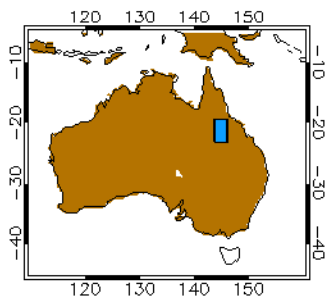
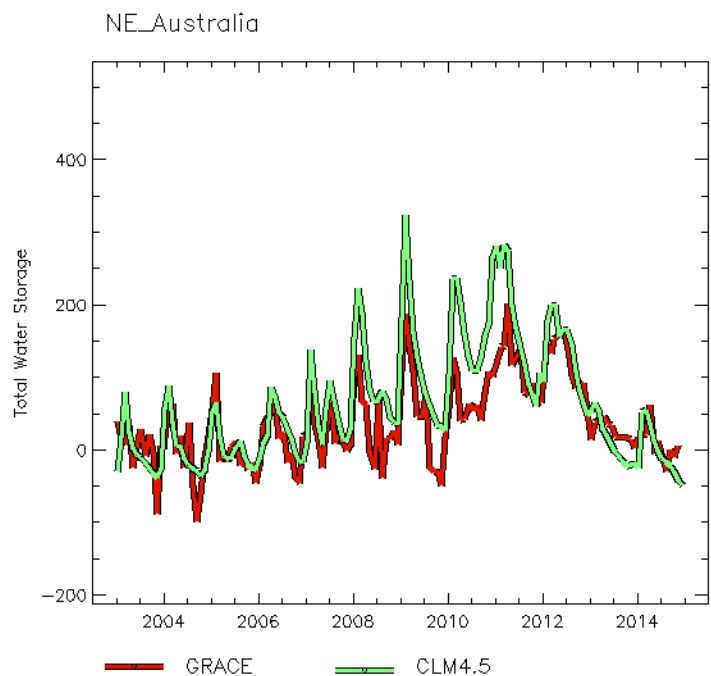


Spatially Variable Soil Depth





GRACE Water Storage Comparison





Current and Future Challenges

- Subgrid heterogeneity and covariance of vegetation, soil moisture, surface water and snow
- Within-canopy turbulent fluxes
- Human management and withdrawals
- Groundwater dynamics
- Dynamic lakes
- Hydrological response to land cover change





Questions?