

**Multiscale modeling of clouds  
and process studies  
using CRM, CSRM, and MMF frameworks**

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**Colorado  
State  
University**

*Knowledge to Go Places*

## Center for Multi-scale Modeling of Atmospheric Processes

[cmmap.colostate.edu](http://cmmap.colostate.edu)

Director: **David Randall**

**UCLA**

**SCRIPPS** INSTITUTE OF OCEANOGRAPHY

**Hampton U**  
"Our Home by the Sea"



the City College of New York

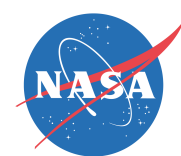


**CCSR**  
UNIV. TOKYO



**UCSD**

**UCAR**



UNIVERSITY OF MARYLAND



THE UNIVERSITY OF UTAH

Pacific Northwest National Laboratory  
Operated by Battelle for the U.S. Department of Energy



UNIVERSITY OF WASHINGTON

**IBM**

**AVML**



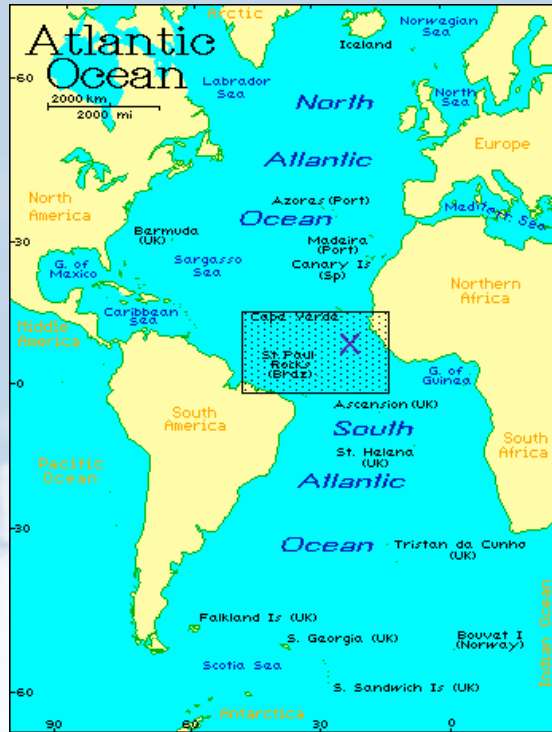
Poudre School District



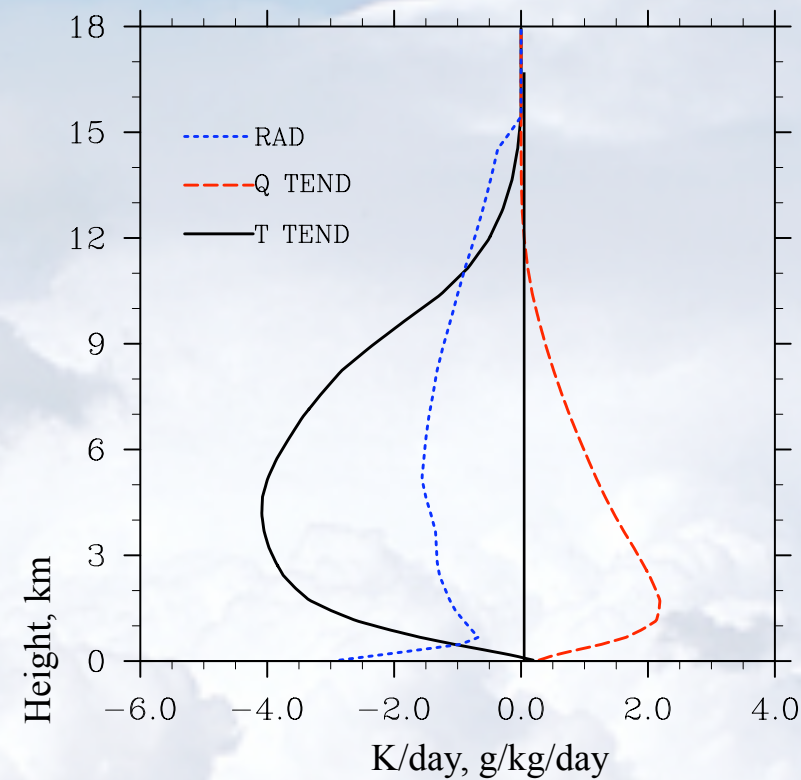
# 'Benchmark' LES of organization of convection on mesoscale

## GATE Phase III Mean Conditions

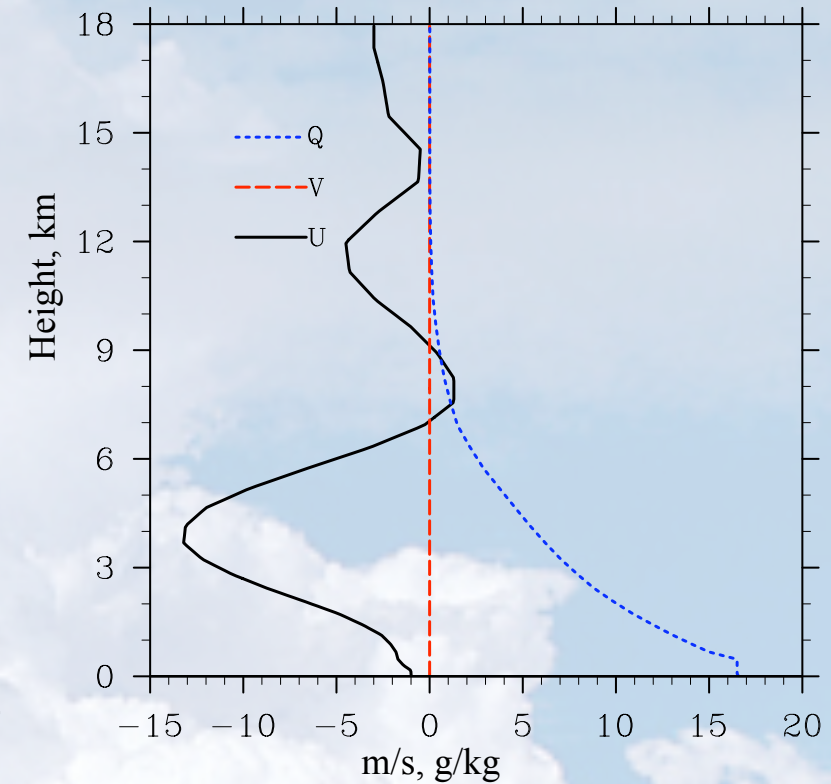
30 August - 19 September, 1974



### Large-Scale Forcing



### Initial Profiles

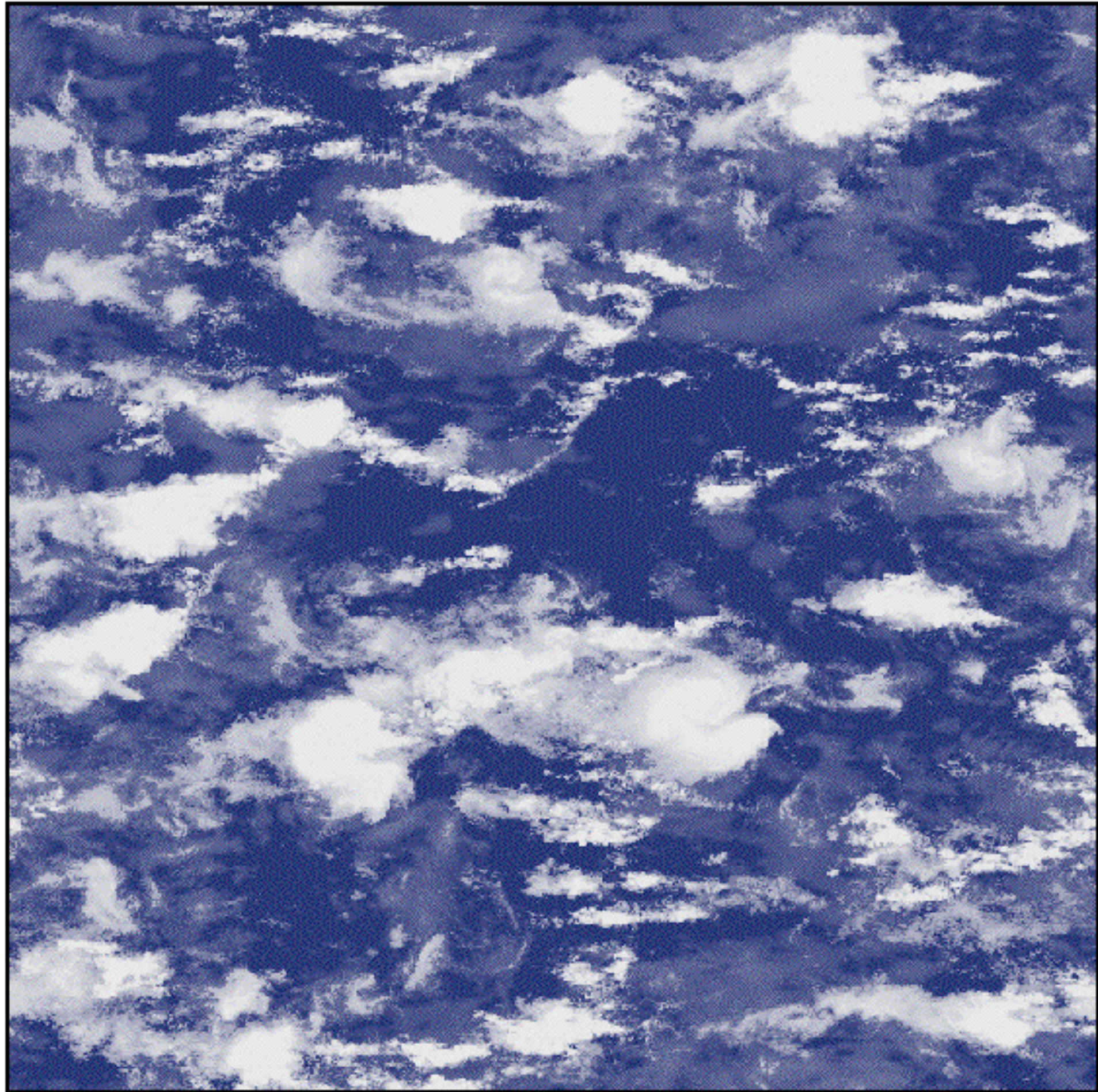


- **System for Atmospheric Modeling, SAM 6.7:**

- **Grid: 2048 x 2048 x 256, or 205 x 205 x 27 km<sup>3</sup>;**
- **Horizontal res. 100m; periodical lateral boundaries;**
- **Vert. res. 50m below 1km, 50-100m @1-5km; 100m @5-18km; 100-300m above;**
- **Time step 2 sec;**
- **Initialization: Random small temperature noise at the lowest grid levels;**
- **Durarian of run: 24 hours**
- **2048 processors; IBM BlueGene BG/L @BNL;**



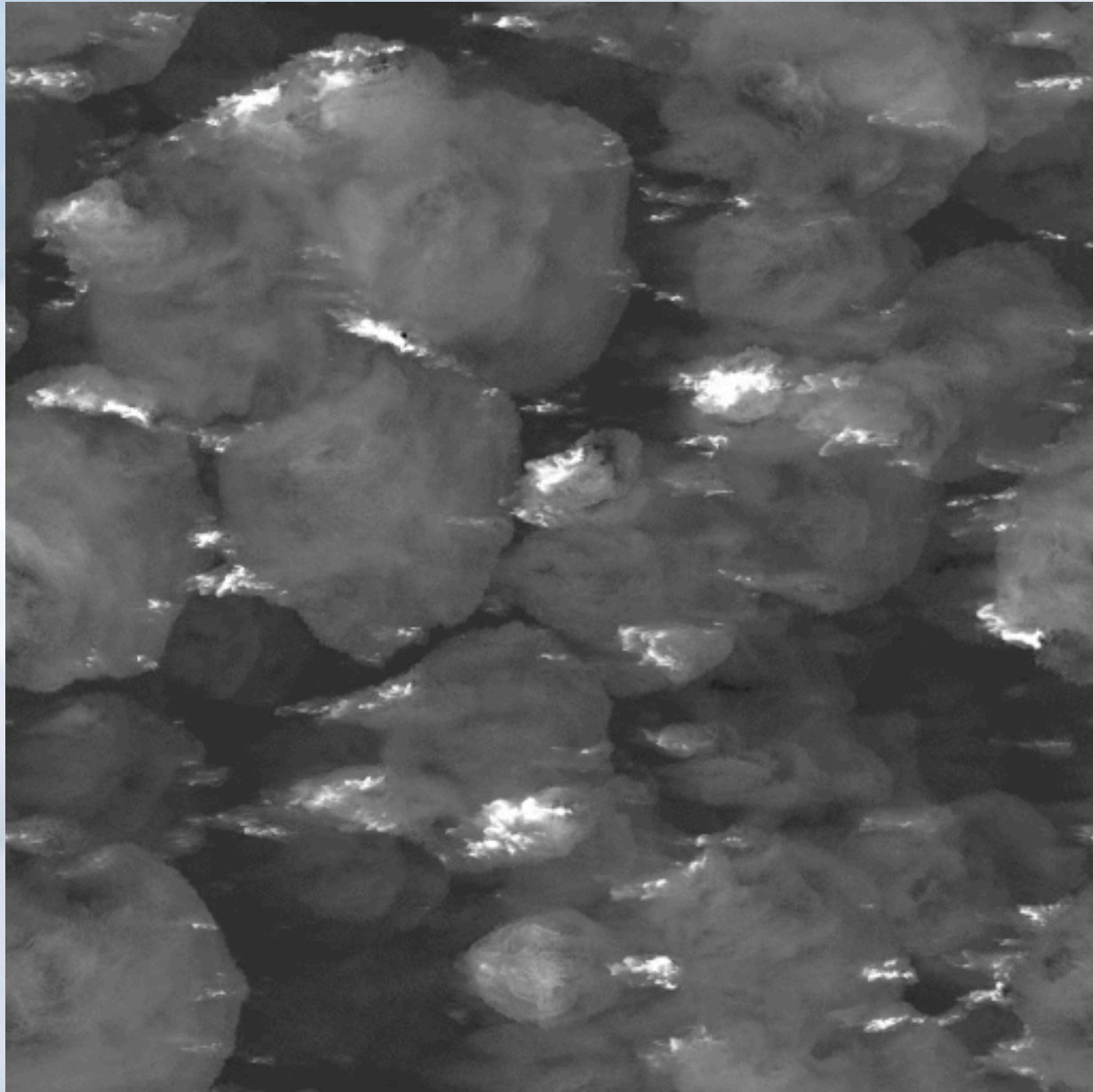
205 km



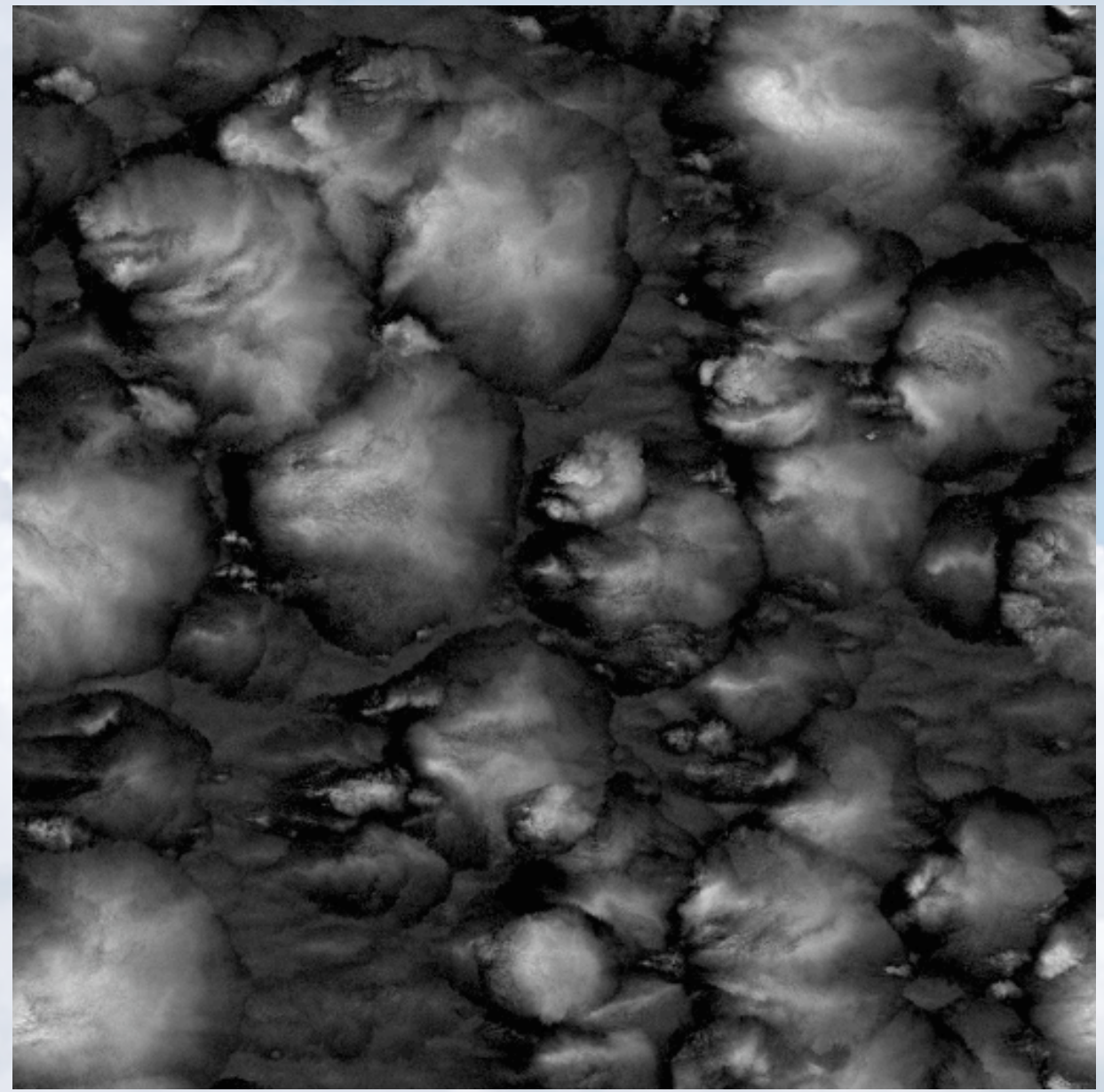
(C) Marat Khairoutdinov, 2008



liquid/ice water static energy near the surface

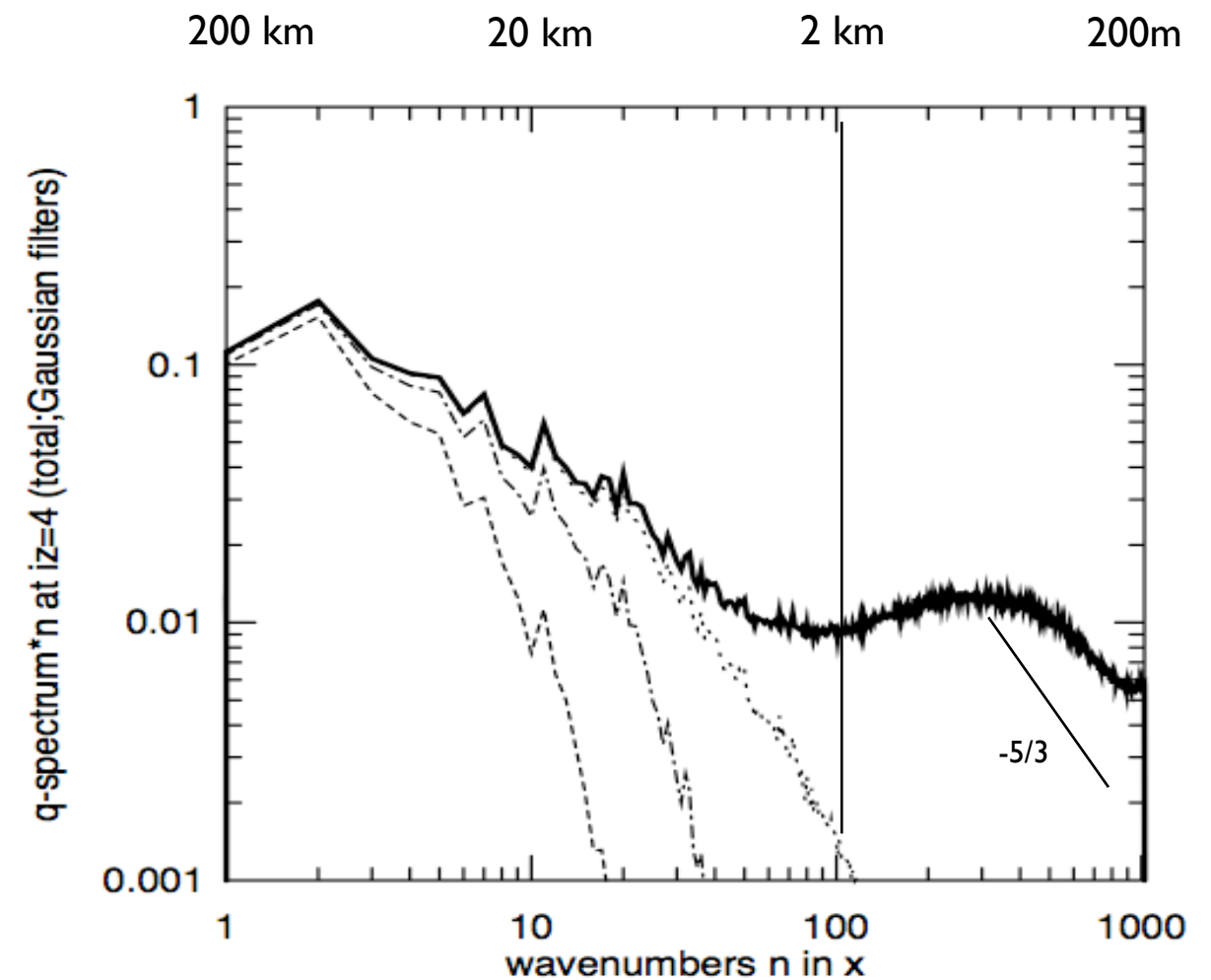
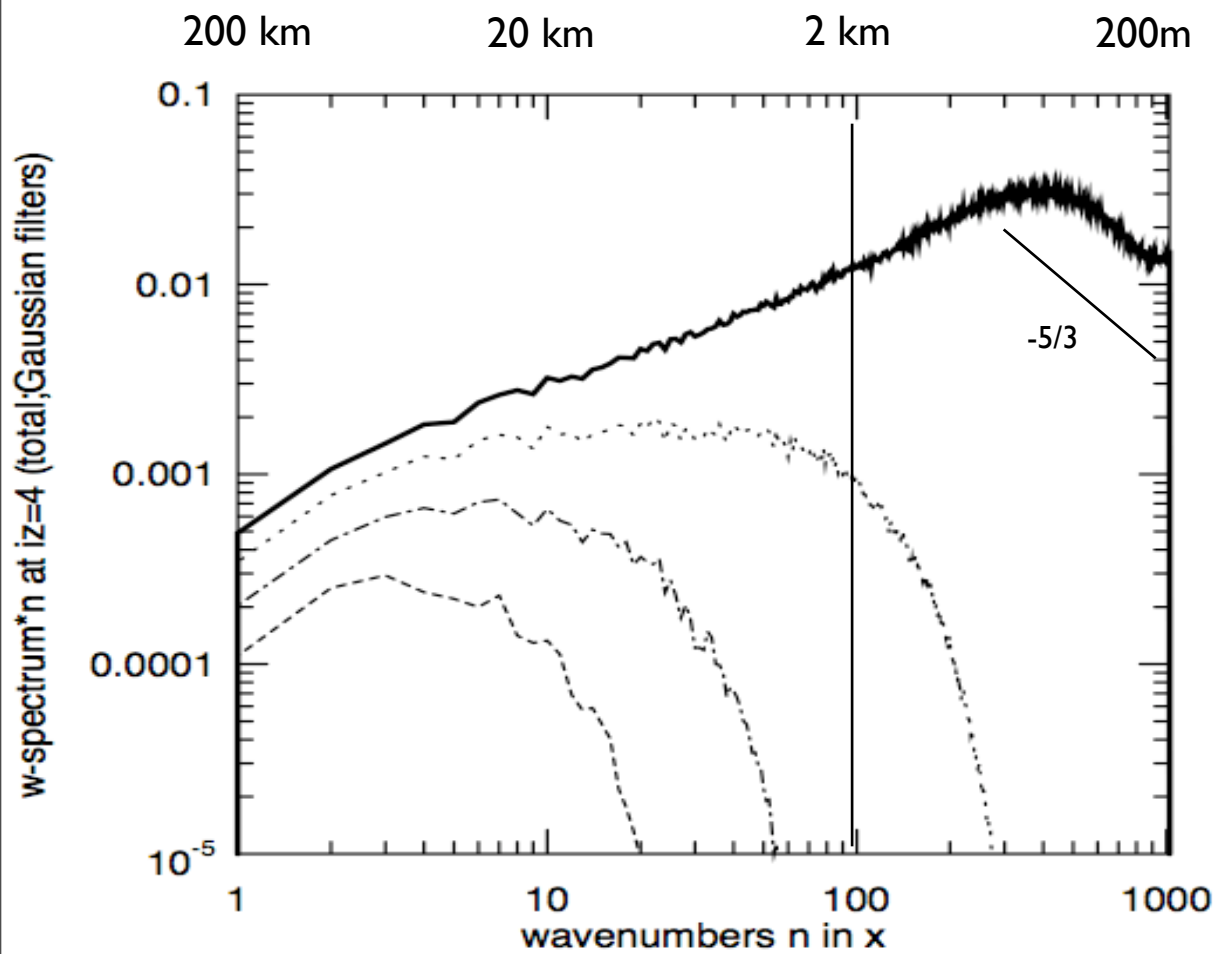


water vapor near the surface





# w-spectra      z=200m      q-spectra

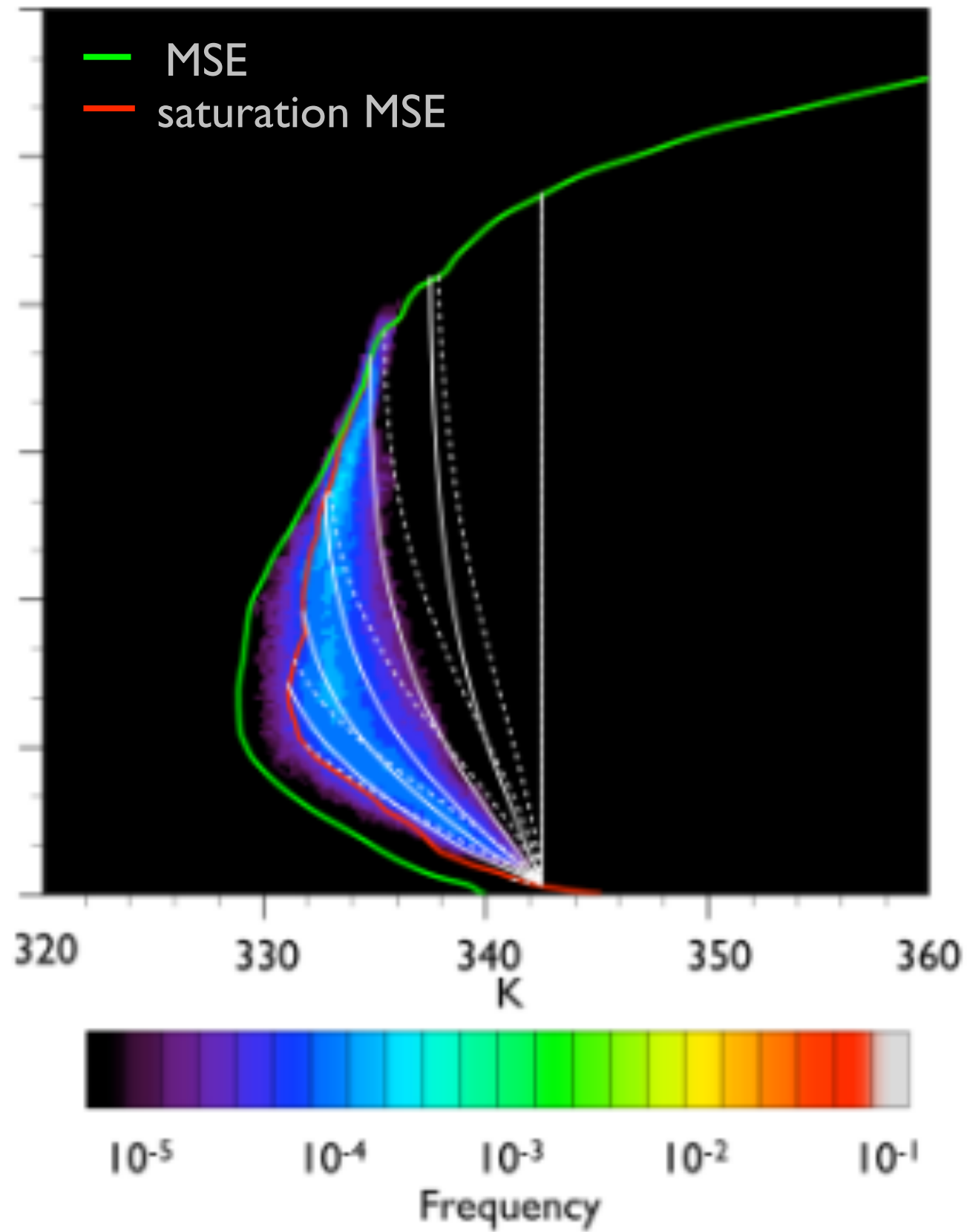


- No filter
- Gaussian filter with  $\Delta=1\text{km}$
- - - - -** Gaussian filter with  $\Delta=4\text{km}$
- · · · ·** Gaussian filter with  $\Delta=10\text{km}$

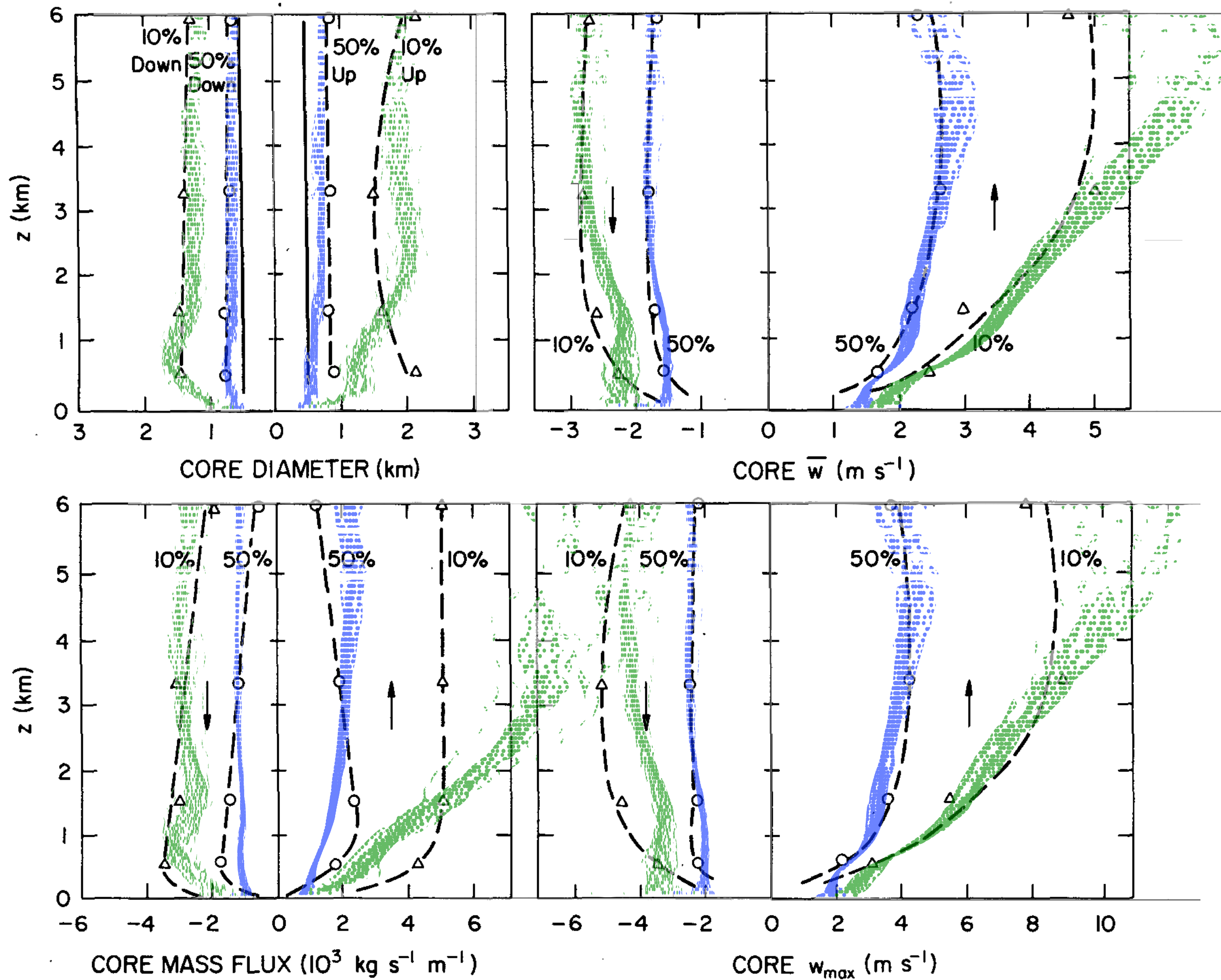
From C.-H. Moeng



# Updraft Cores



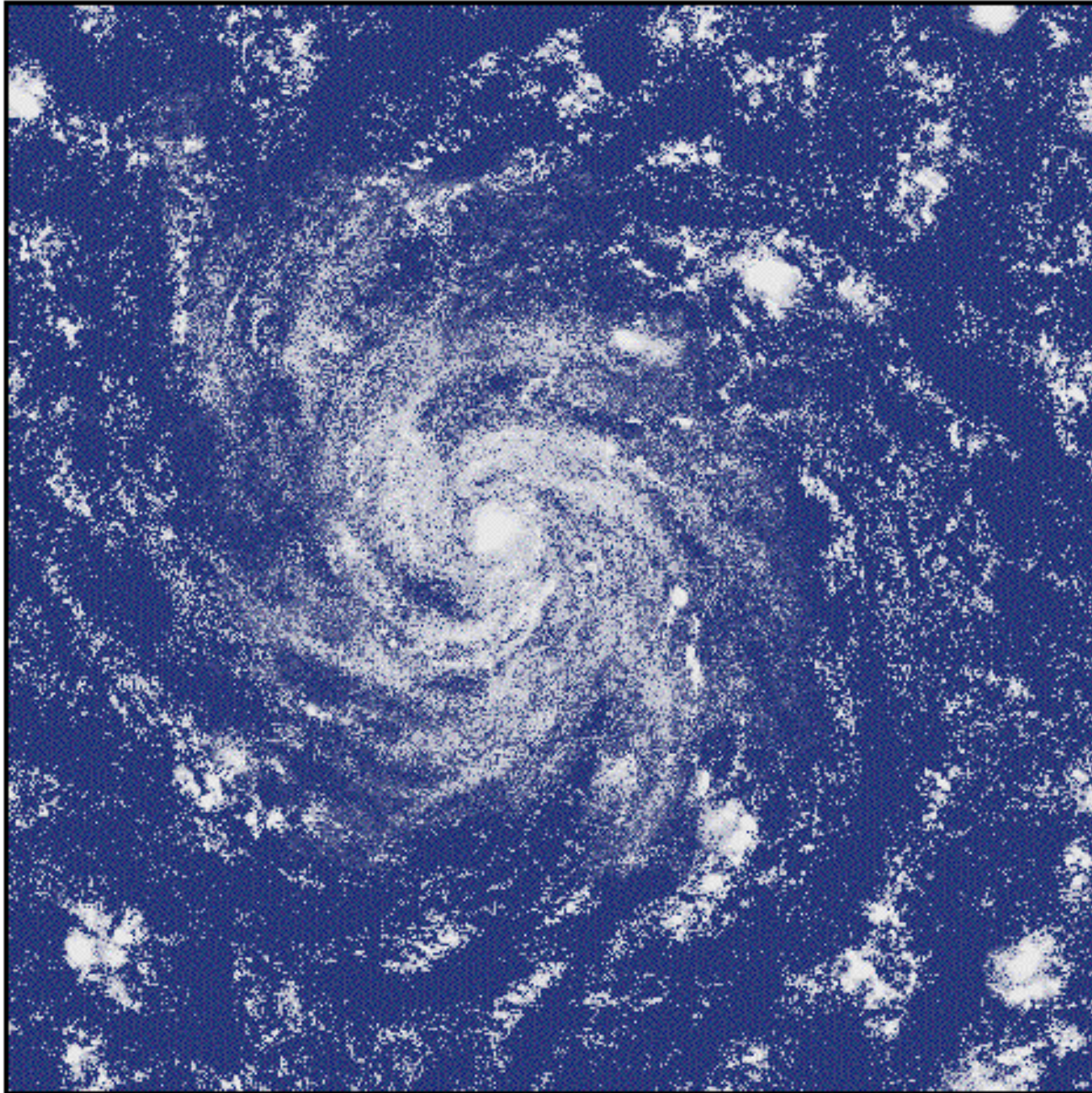






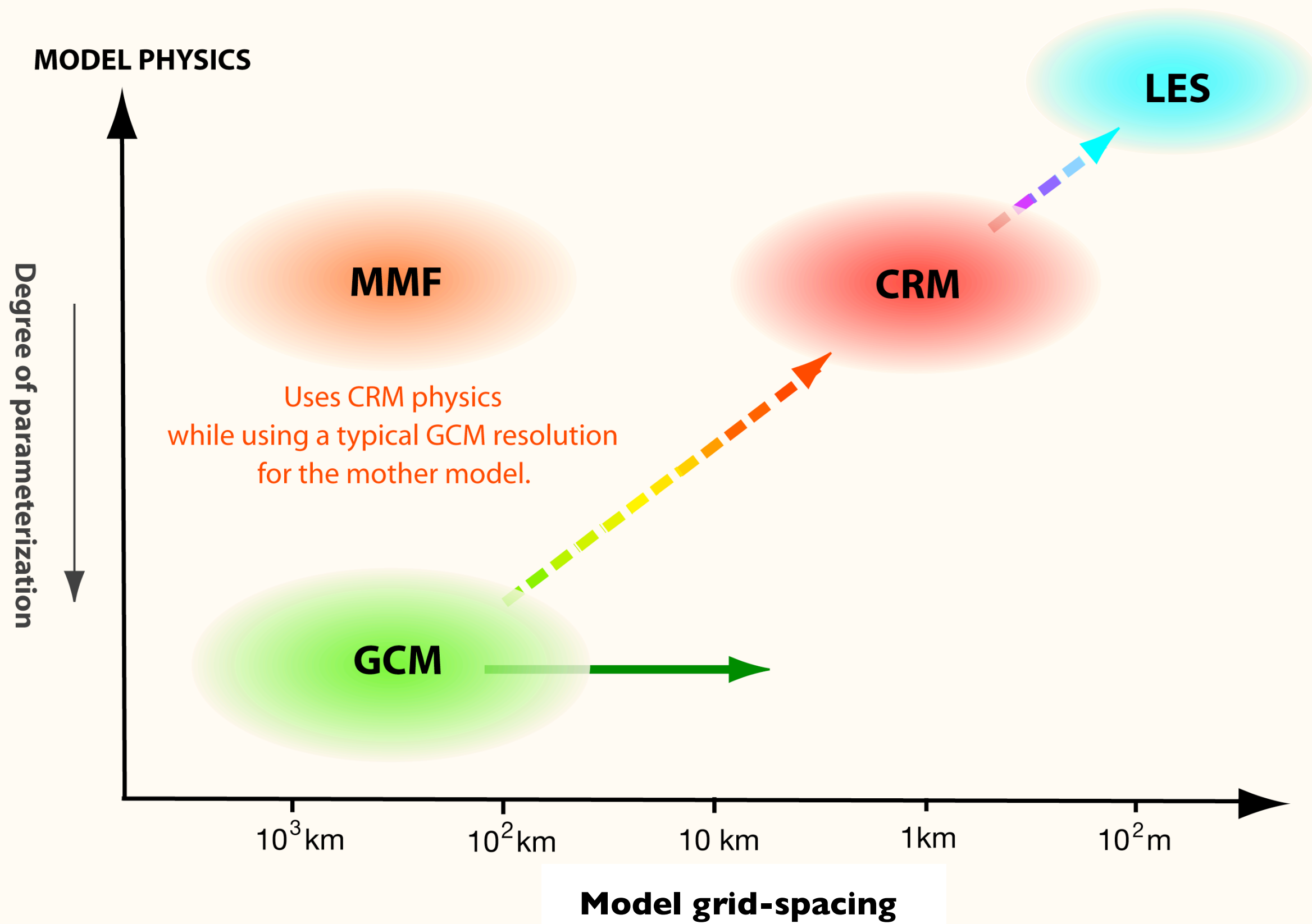
# Idealized TC intensification

1000



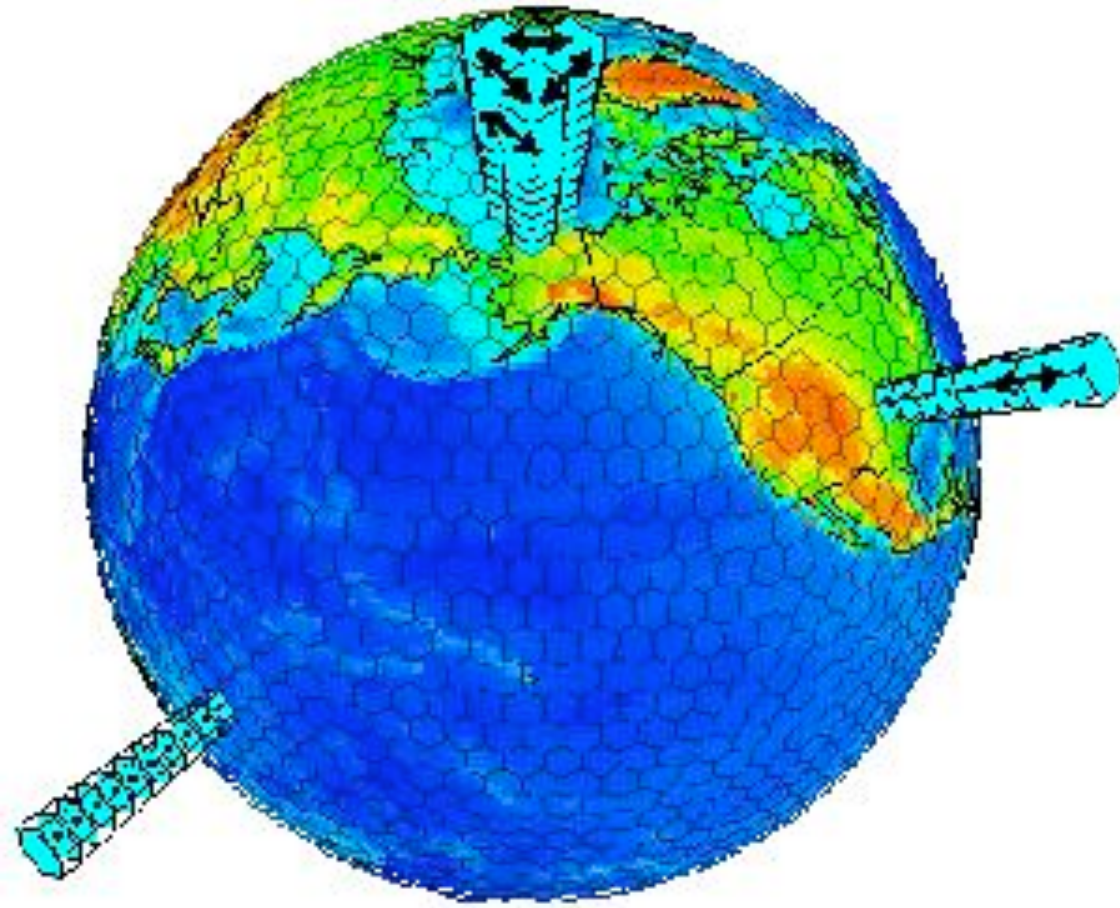
(C) Marat Khairoutdinov, 2008







# The root of MMF is in Single-Column Modeling



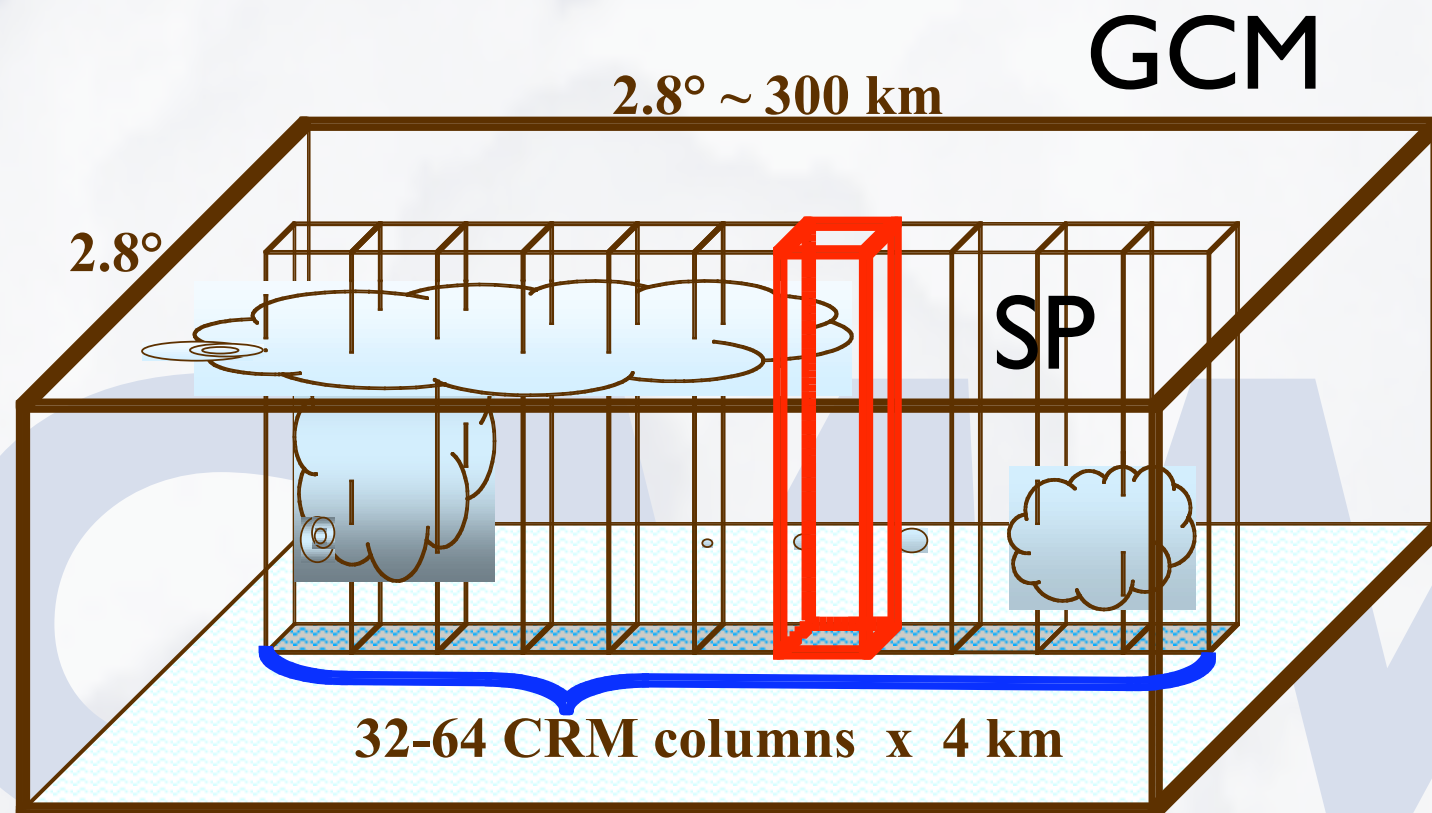
$$\frac{\partial \bar{s}}{\partial t} = \underbrace{-\overline{\nabla_s V} - \frac{\partial \bar{s} \bar{\omega}}{\partial p}}_{LSForcing} + \underbrace{Q_1}_{Param's}$$
$$\frac{\partial \bar{q}}{\partial t} = \underbrace{-\overline{\nabla_q V} - \frac{\partial \bar{q} \bar{\omega}}{\partial p}}_{LSForcing} - \underbrace{Q_2/L}_{Param's}$$

Traditionally, the large-scale forcing data would come from observations (GATE, TOGA, ARM, KWAJEX, etc.)

In MMF approach, large-scale forcing for each grid-column is explicitly computed by a GCM



## Prototype MMF Approach:



SP forcing:

$$-\overline{\nabla \phi V} - \frac{\partial \bar{\phi} \bar{\omega}}{\partial p} = \frac{\bar{\phi}^* - \bar{\phi}^n}{\Delta t}$$

SP Source/Sink:

$$Q_n = \frac{\bar{\phi}^{n+1} - \bar{\phi}^*}{\Delta t}$$

- Works only for two-level GCM time schemes.

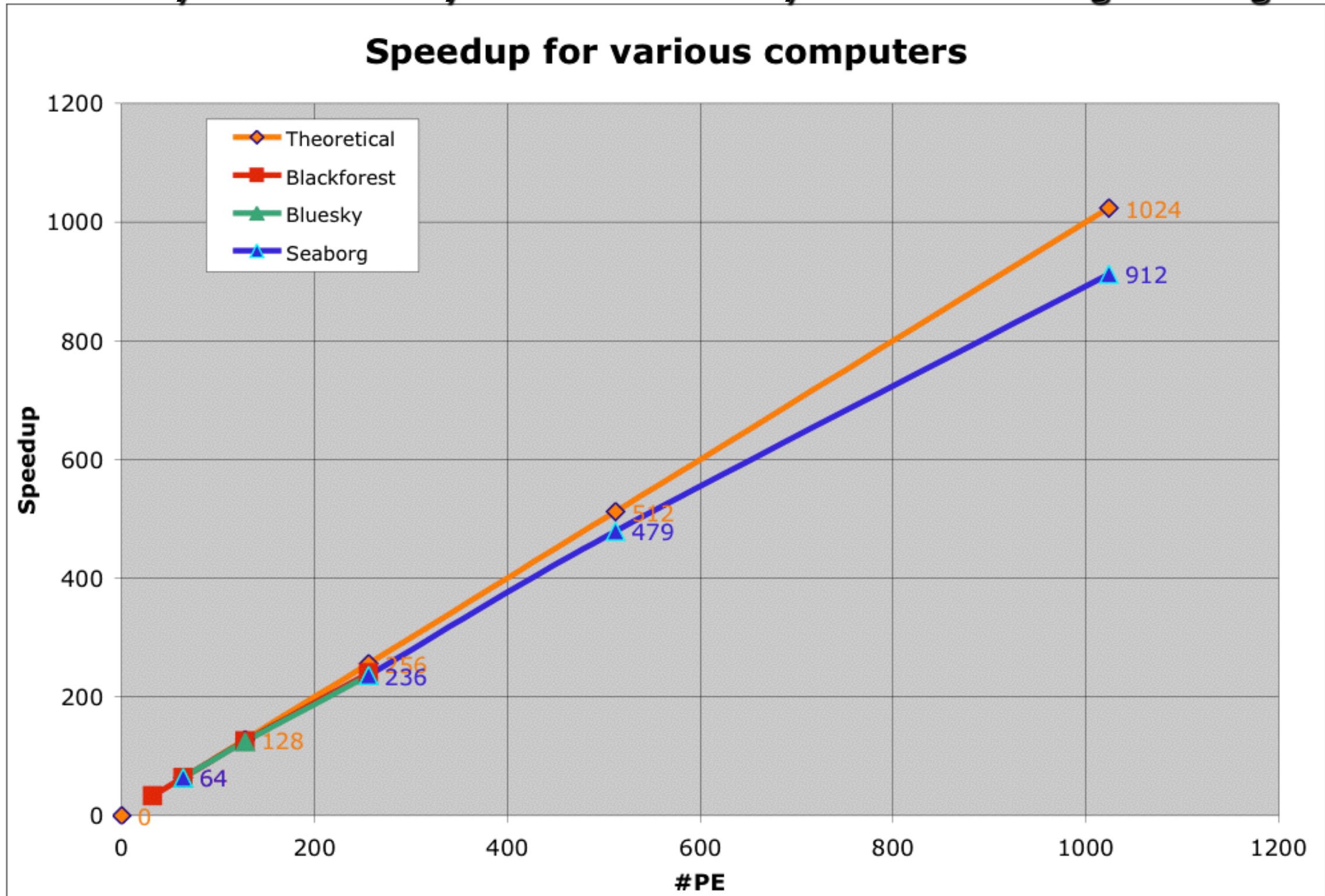
- GCM horizontal grid is not known to SP; consequently
  - ★ SPs run independently from each other (great parallelism!)
  - ★ SP should have periodical domain
  - ★ LS Forcing is applied horizontally uniform
  - ★ SP domain size is not generally determined by GCM grid size

*Reach for the sky.*

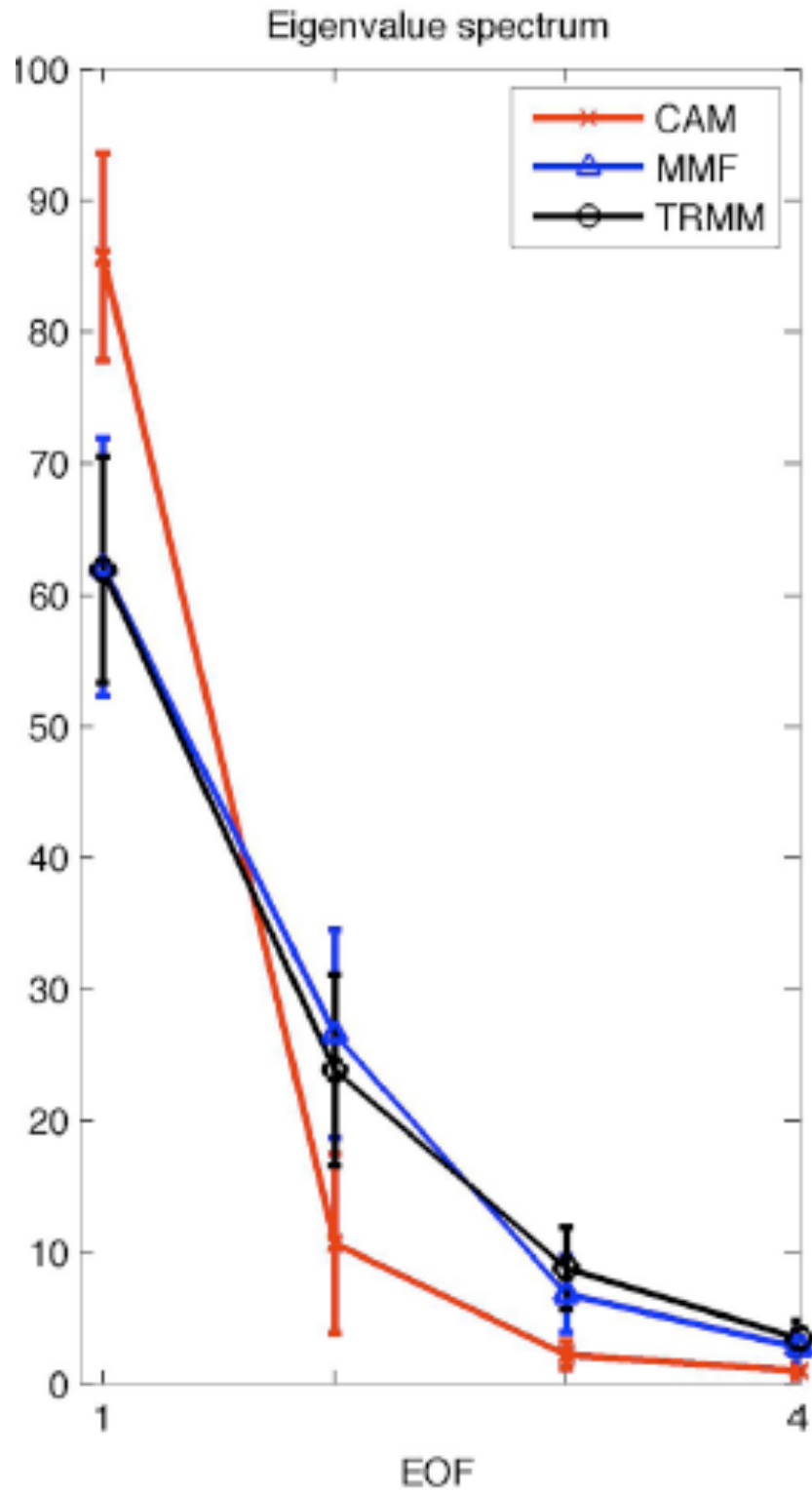


**MMF is 100-200 times slower than the CAM when run on the same number of processors. However, due to very small portion of time spent on inter-processor communications, MMF can run on about 10 times more processors just as efficiently.**

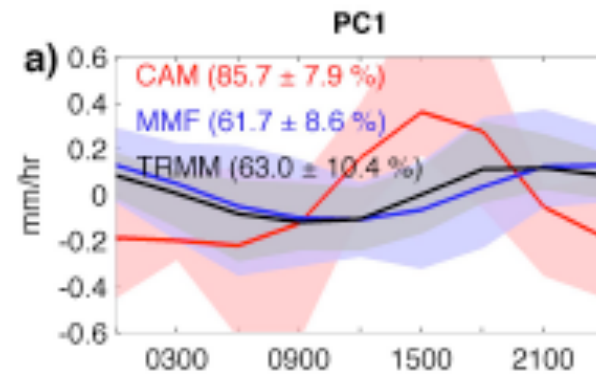
**It currently takes one day to simulate one year on 2.8x2.8 grid using 1024 PEs.**



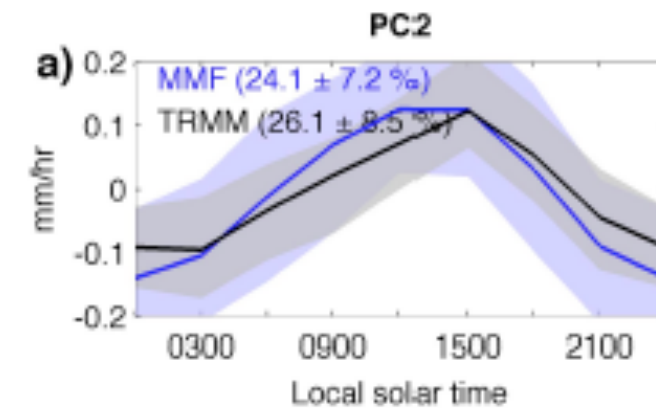
# Diurnal Cycle of Precipitation: EOFs



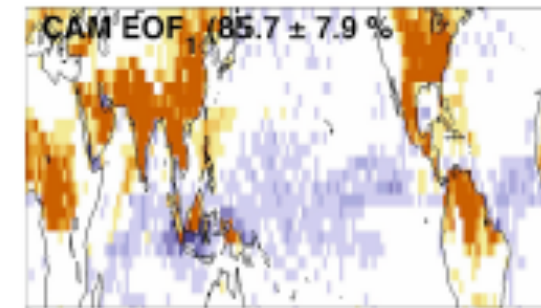
## Leading mode



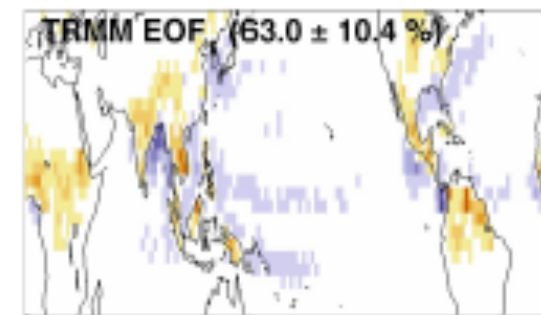
## Secondary mode



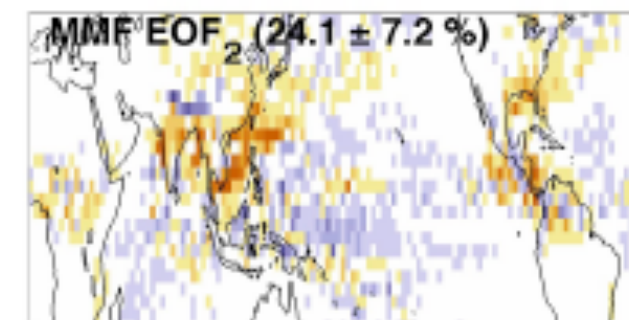
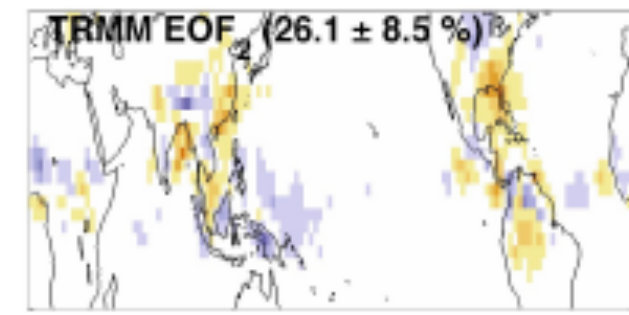
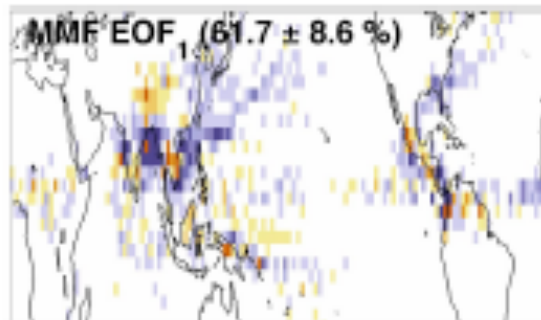
CAM



TRMM



SP-CAM



from Michael Pritchard (Scripps, UCSD)

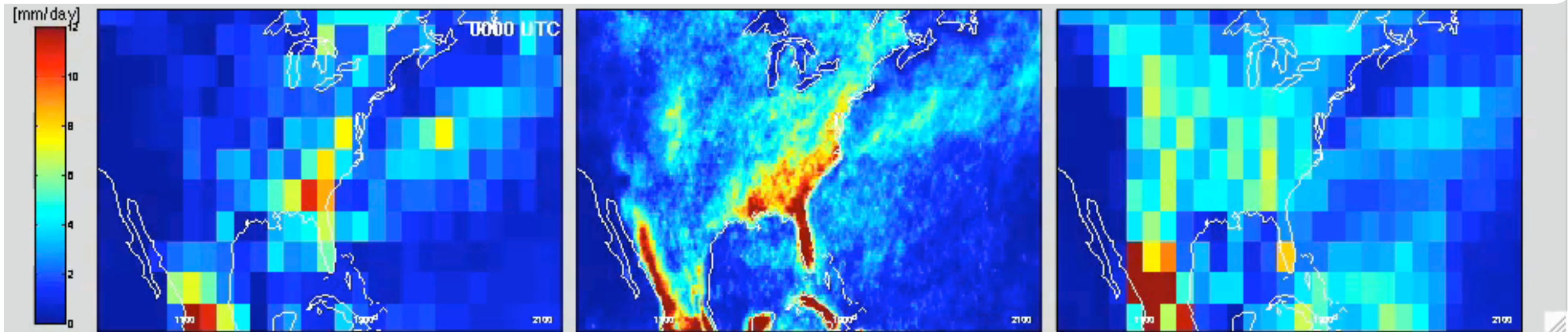


# Composite Diurnal Cycle of Precipitation: North America

SP-CAM

TRMM

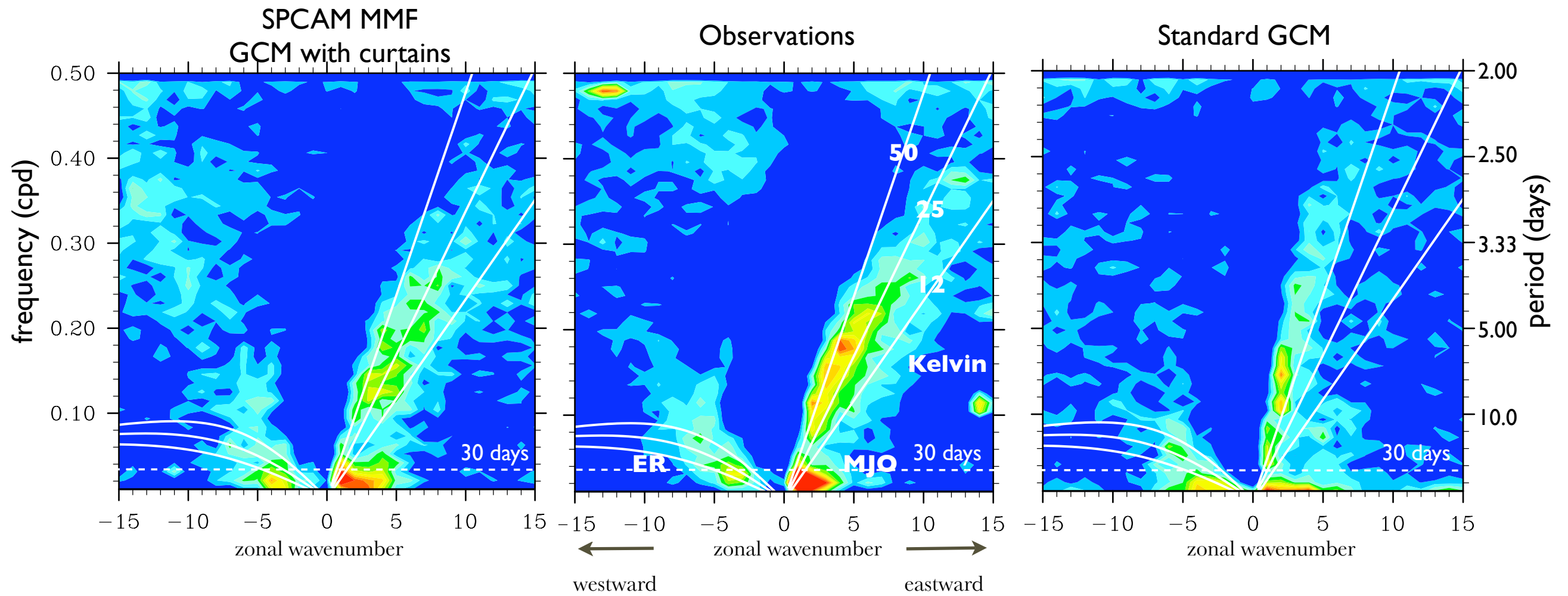
CAM



from Michael Pritchard (Scripps, UCSD)

# Simulated sub seasonal tropical variability

## OLR



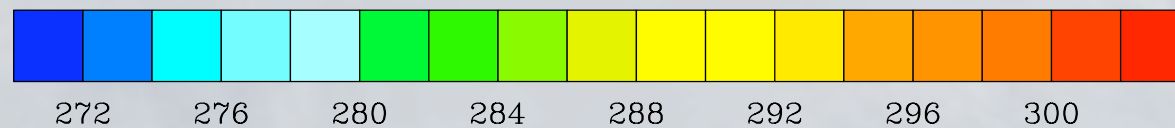
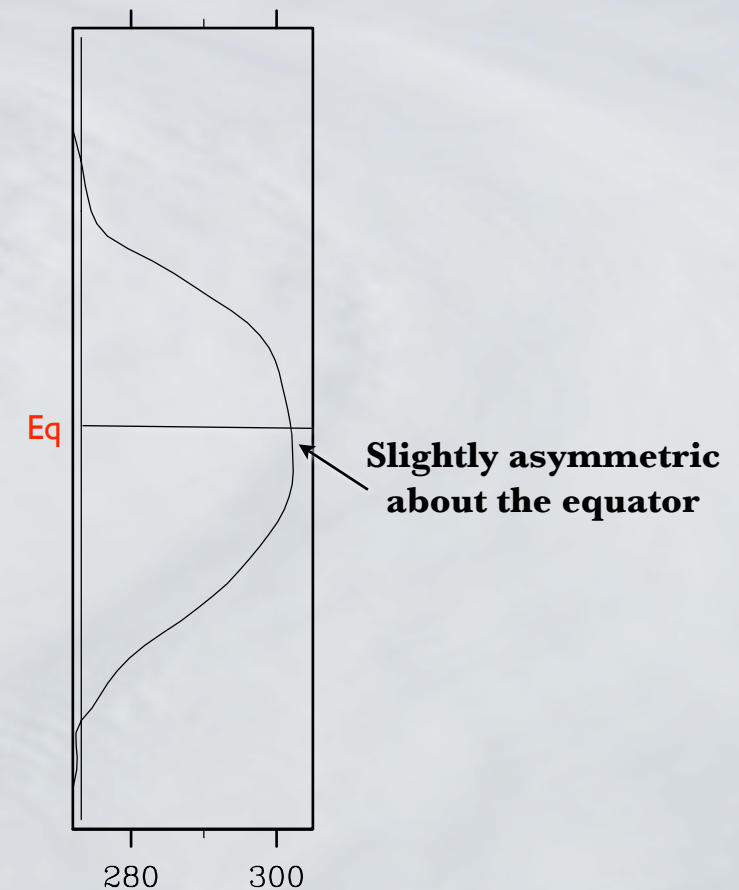
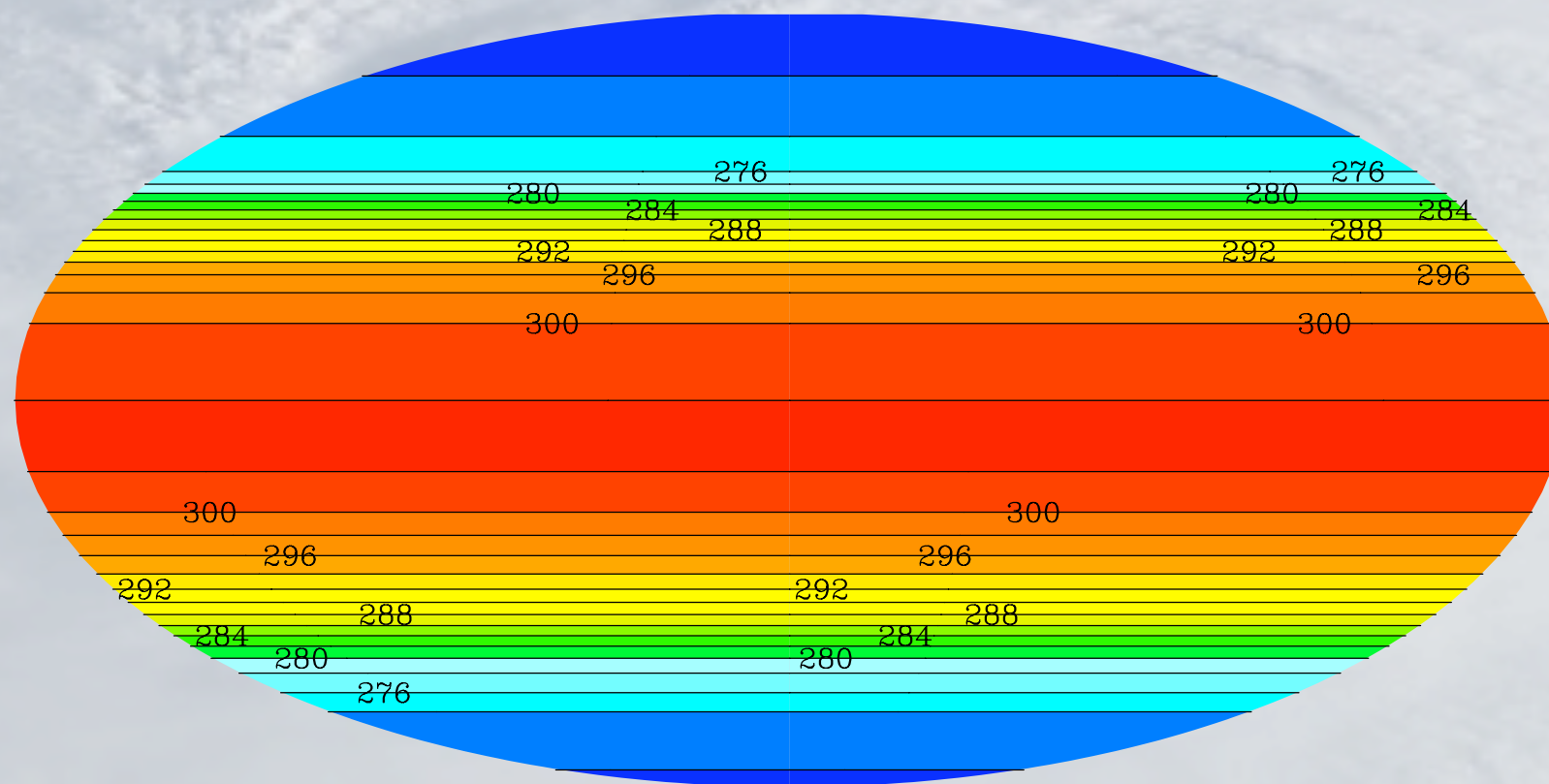


# MJO on Aquaplanet/Water World

## Why Aquaplanet? Simplicity!

- no polar ice, no land/soil/terrain complications
- heat and vapor fluxes over water - we know how
- zonally symmetric (zonal means direction of latitude circle)
- perpetual equinox - no seasons
- Mean climate resembles the Earth's
- Still has MJO and main equatorial waves

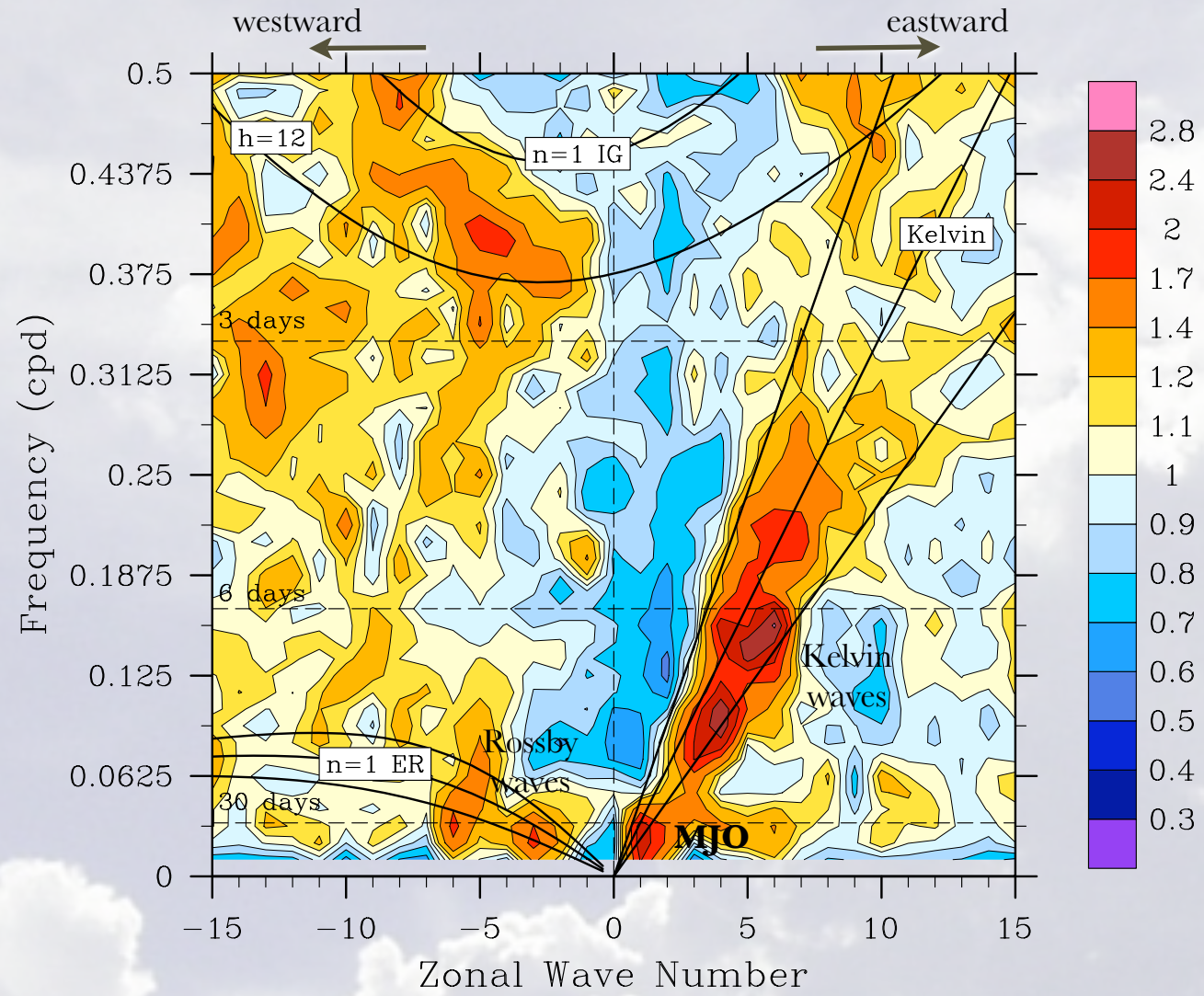
Sea Surface Temperature (prescribed)





Simulated tropical variability on Aquaplanet looks similar  
to subseasonal tropical variability on Earth

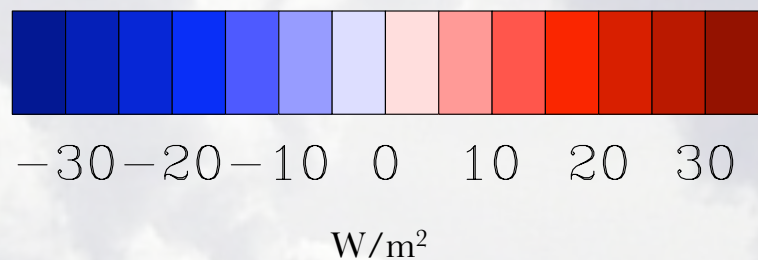
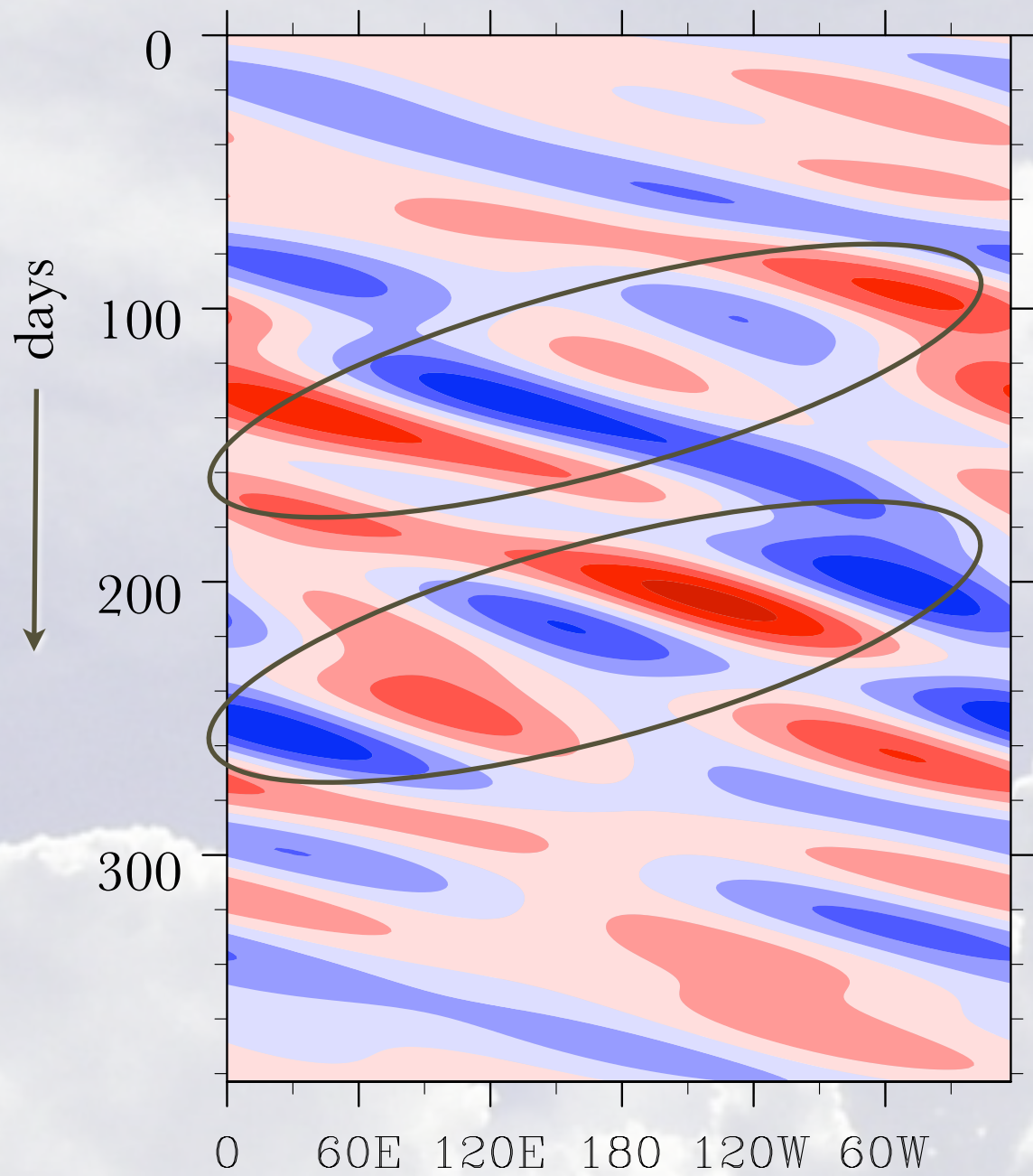
OLR



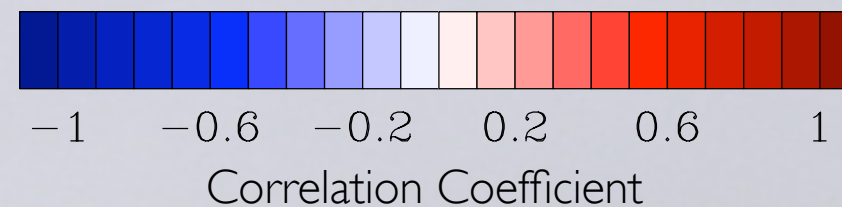
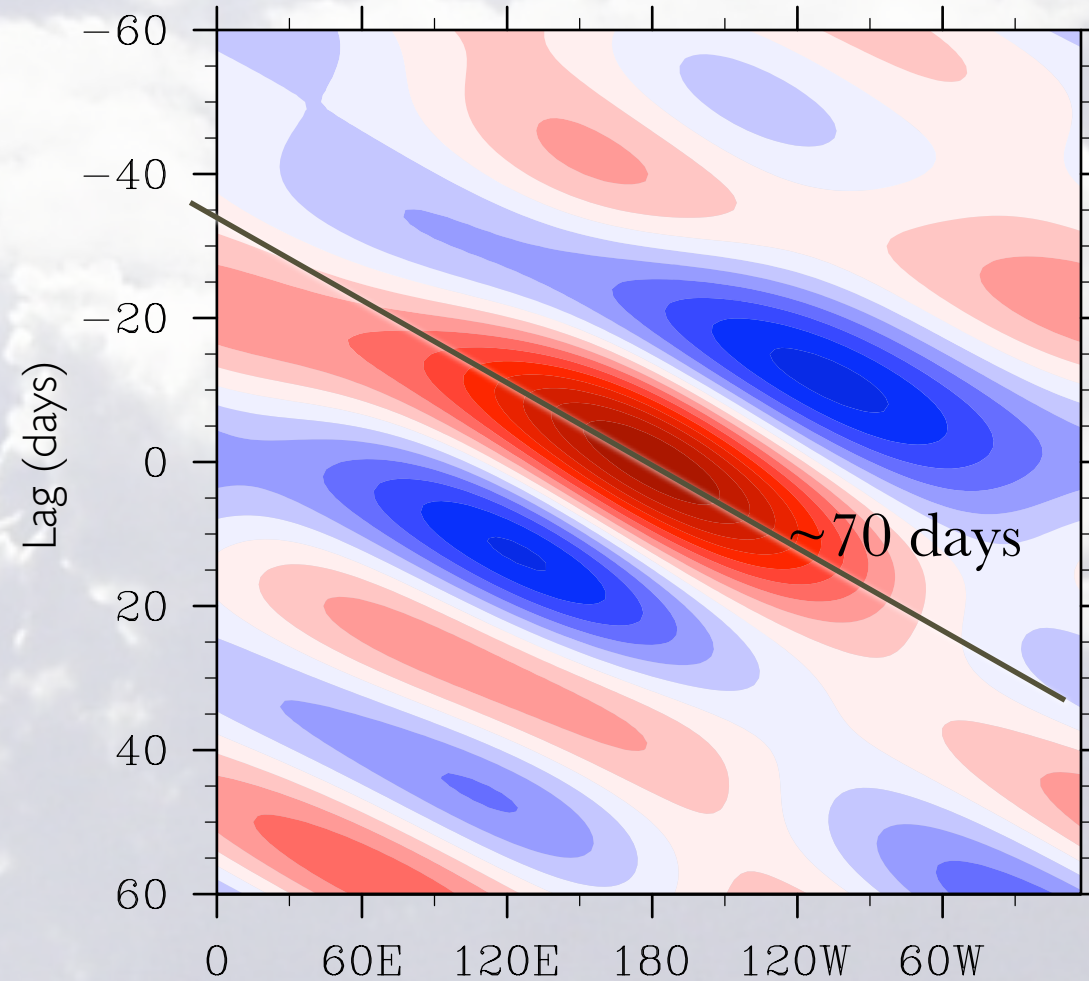


# MJO forms 'wave packets' on Aquaplanet

384-day sample



Aquaplanet MJO-filtered OLR



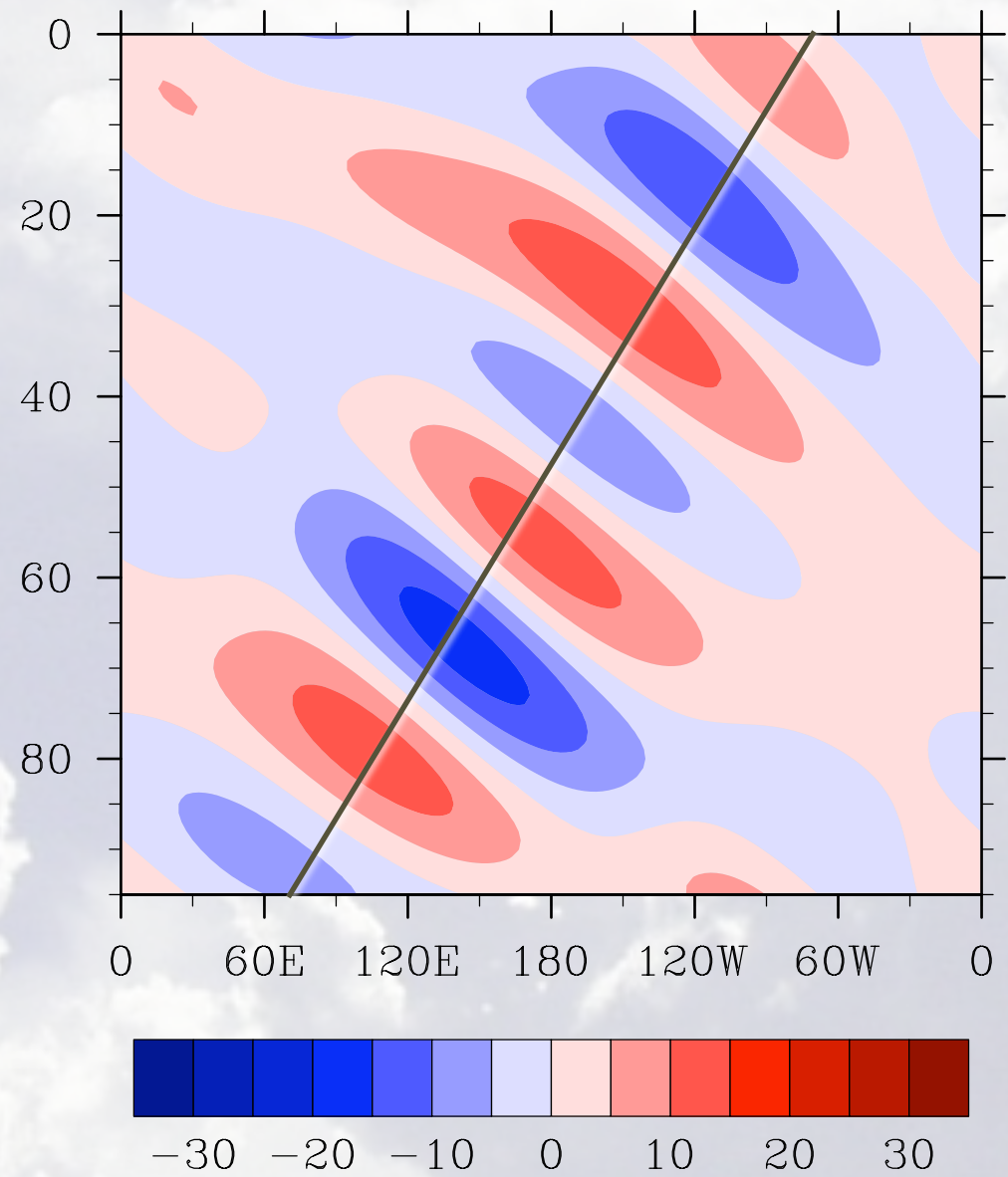
W/m<sup>2</sup>

Text



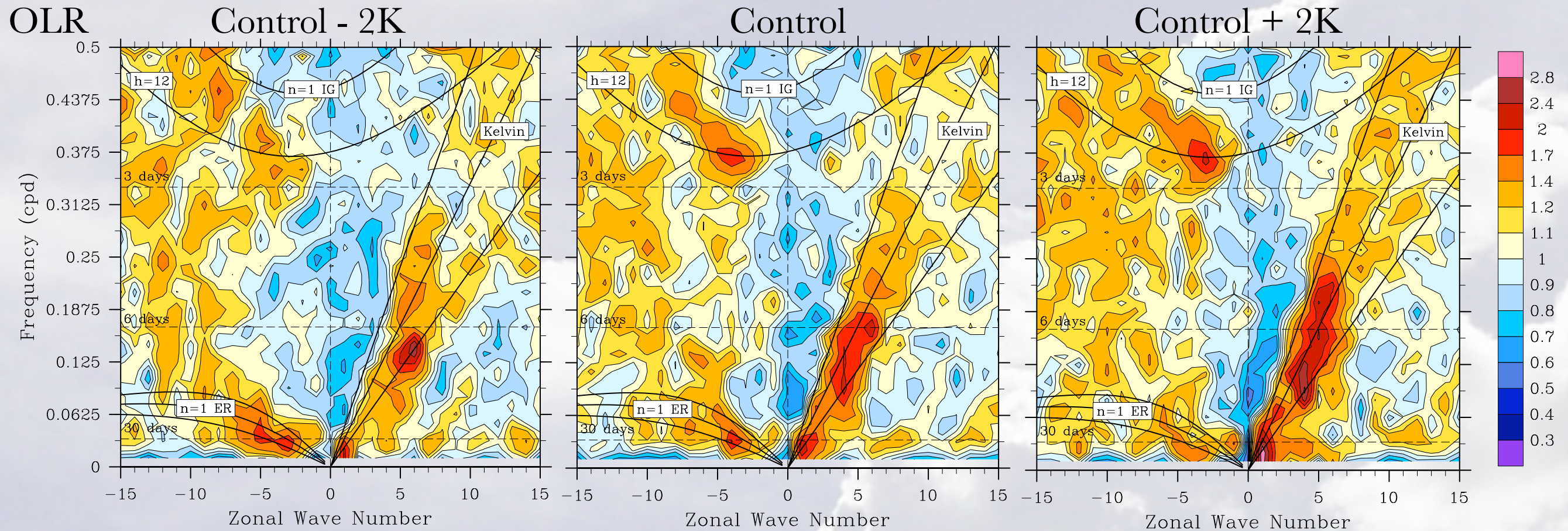
on Earth too  
NOAA OLR  
(Satellite observations, not a model)

Strong El Nino year  
1997.81

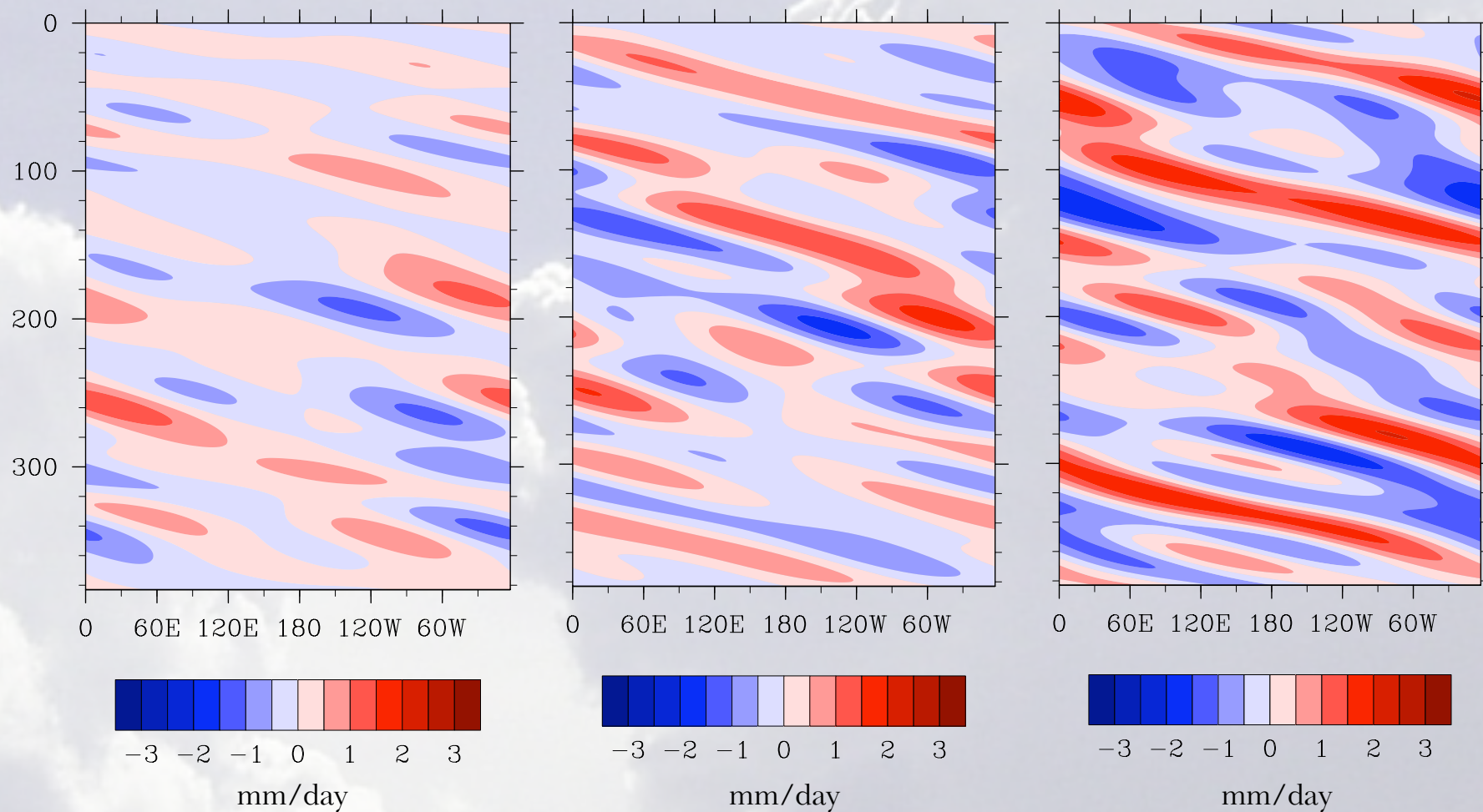




# Sensitivity to Aquaplanet Sea Surface Temperature



## MJO-mode Precipitation



**Positive SST perturbation causes stronger MJO and Kelvin modes, but weaker Rossby mode.**

**Rising CO<sub>2</sub>**

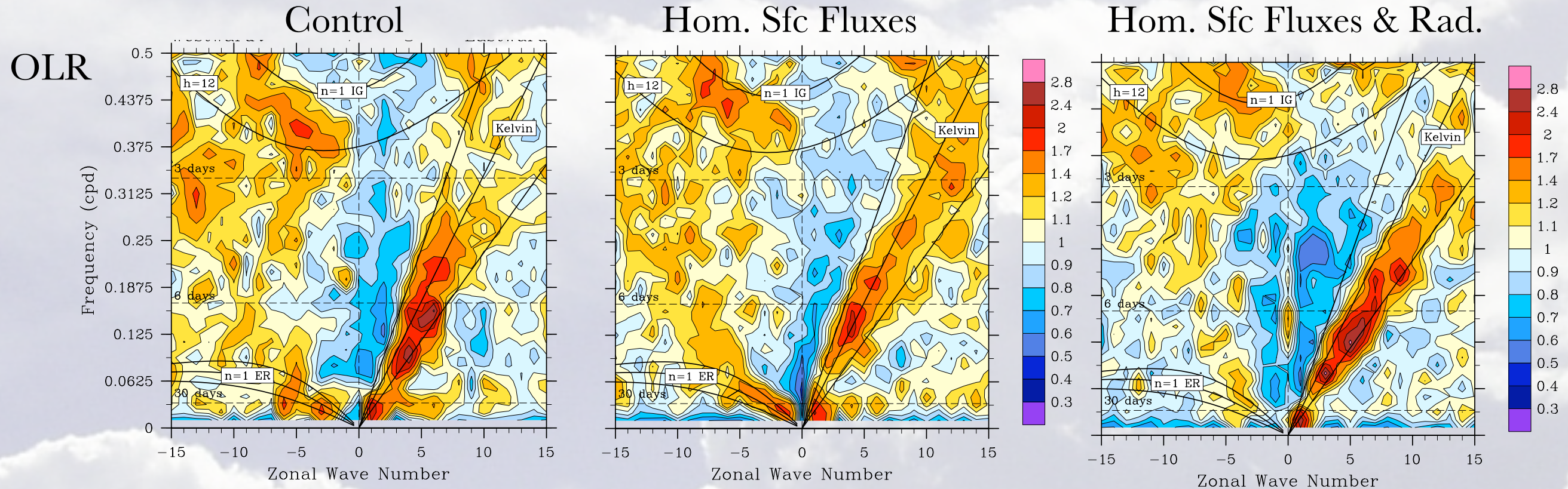
↓  
**Warmer Tropical SST**

↓  
**More extreme rain associated with MJO in the future?**

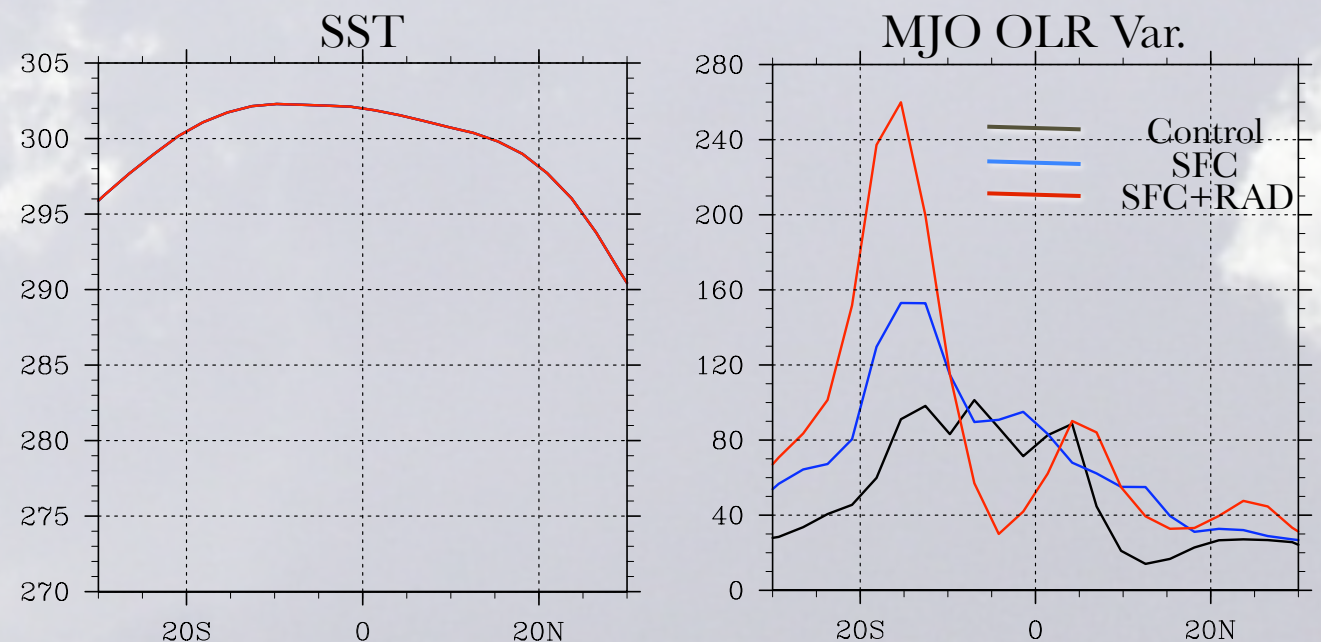


# Sensitivity to Zonally Homogenized Surface fluxes and Radiative Heating

- Zonal Homogenization: Compute surface fluxes and radiation heating rates as usual but apply them zonally averaged.



- Homogenization shifts MJO towards warmer SSTs and concentrates MJO into narrower band;
- Zonal variation of surface fluxes and radiative heating doesn't seem to be essential for maintaining simulated MJO;





# Sensitivity to Zonally Homogenized Water Vapor

- Zonal Homogenization: Nudge (relax) water vapor to zonally averaged values over diurnal time scale;

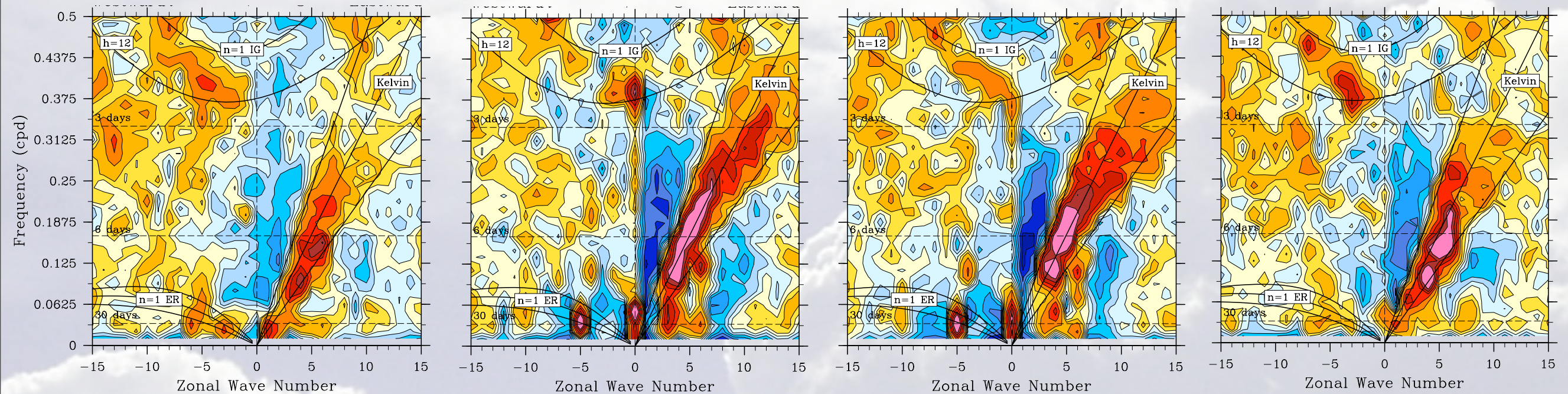
OLR

Control

Hom. all levels

Hom. above 850mb

Hom. above 400mb



- **The existence of mid-to-low troposphere (but above PBL) water-vapor anomalies is the key for the existence of simulated MJO.**
- **The anomalies are associated with the shallow and congestus cloud activity - hence, representation of those cloud types in GCMs is the key for MJO simulation**



# First ever coupled-MMF simulation

Cristiana Stan<sup>1\*</sup>, Marat Khairoutdinov<sup>2</sup>, Charlotte A. DeMott<sup>3</sup>, V. Krishnamurthy<sup>1,4</sup>, David M. Straus<sup>1,4</sup>,

David A. Randall<sup>3</sup>, James L. Kinter III<sup>1,4</sup>, and J. Shukla<sup>4,5</sup>

Lead: Center for Ocean-Land-Atmosphere studies (COLA)

		<b>SP-CCSM</b>	<b>CCSM3.0</b>
<b>Atmospheric Model</b>	Horiz. Res.	T42 (sld)	T42 (sld)
	Vert. levels	30	26
	Deep conv.	CRM	ZM
	Shallow conv.	CRM	Hack
<b>Ocean Model</b>	Horiz. Res.	gx3v5	gx3v5
	Vert. levels	25	25
<b>Sea-Ice Model</b>		CSIM4	CSIM4
<b>Land Model</b>		CLM3.0	CLM3.0



DJF Precipitation

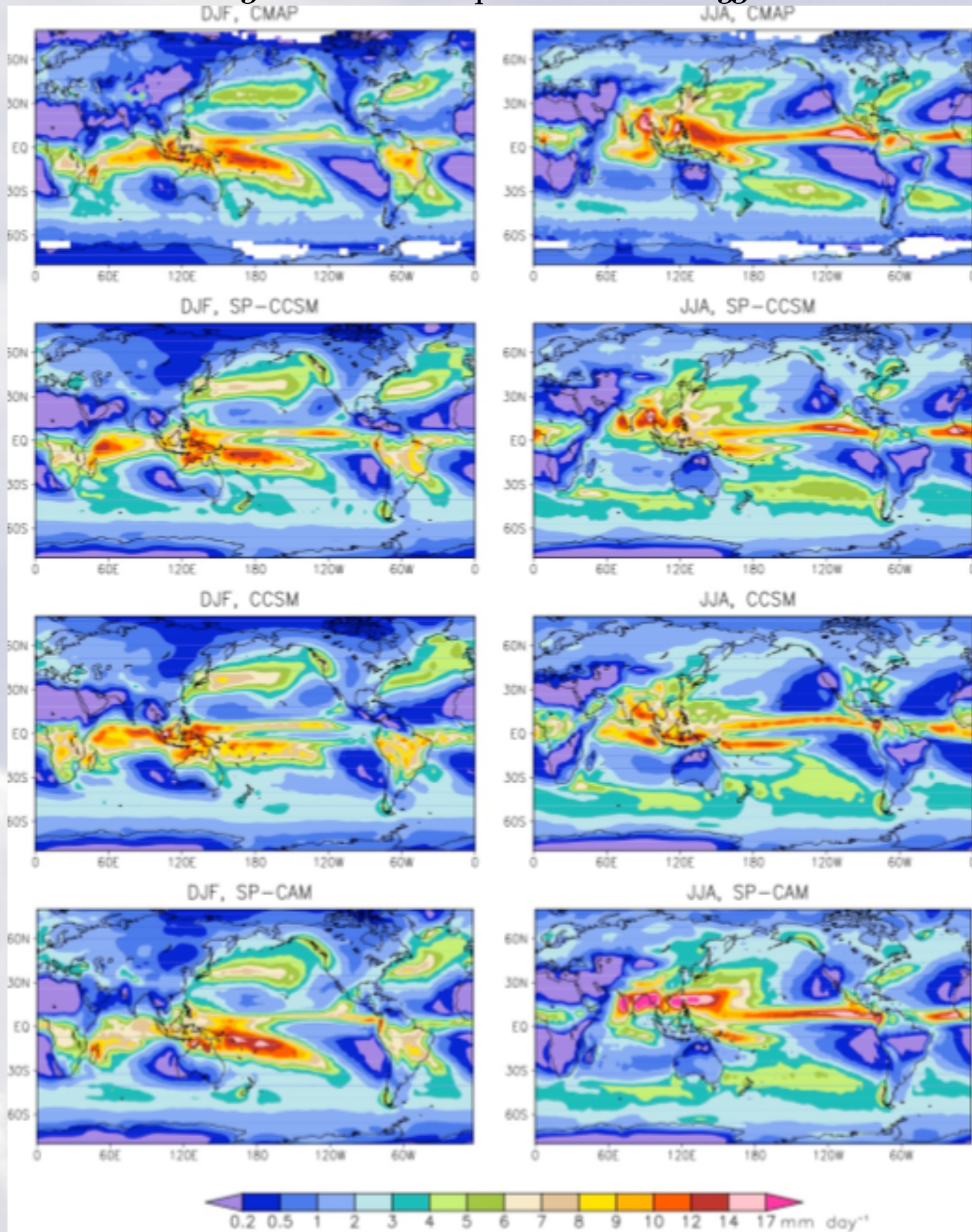
JJA

CMAP

SP-CCSM

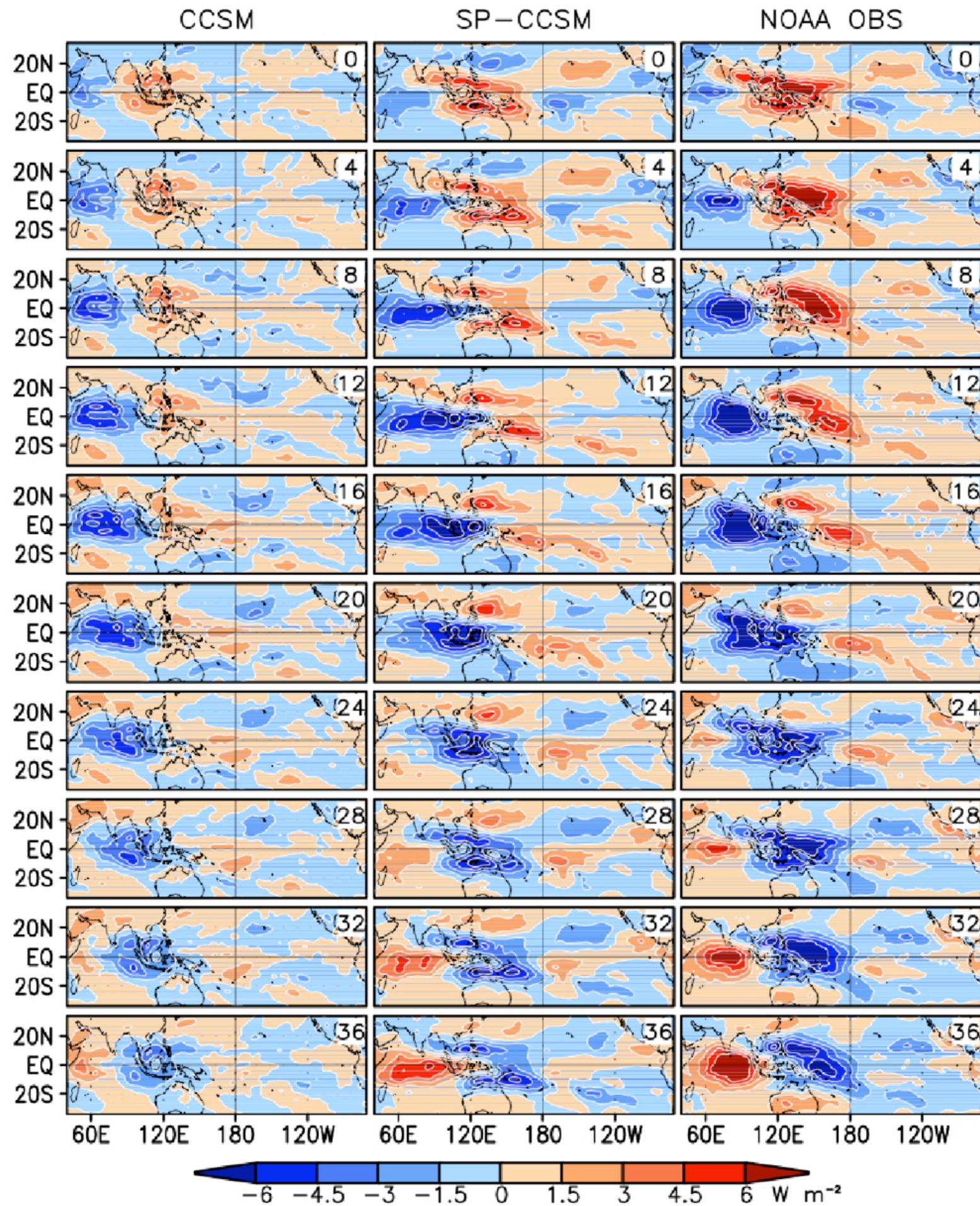
CCSM

SP-CAM





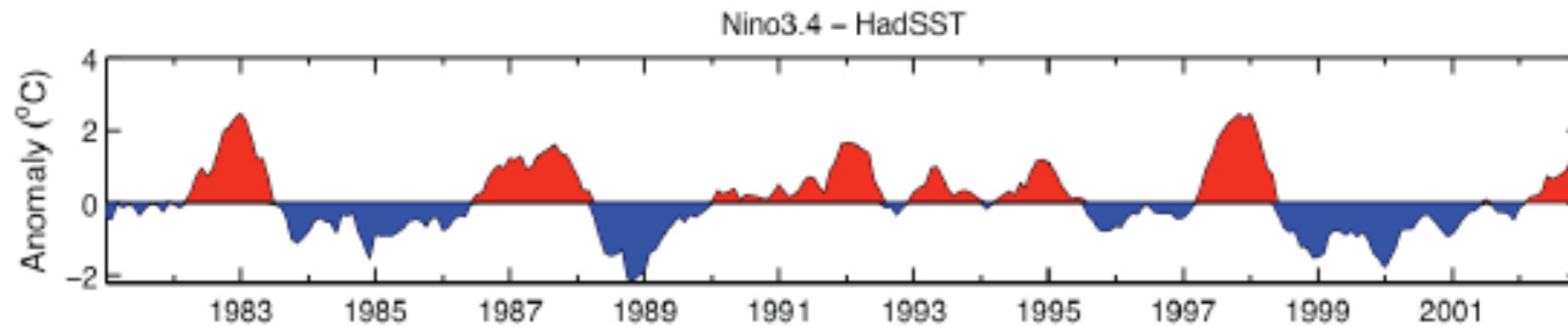
# Phase composites of OLR dominant MJO mode



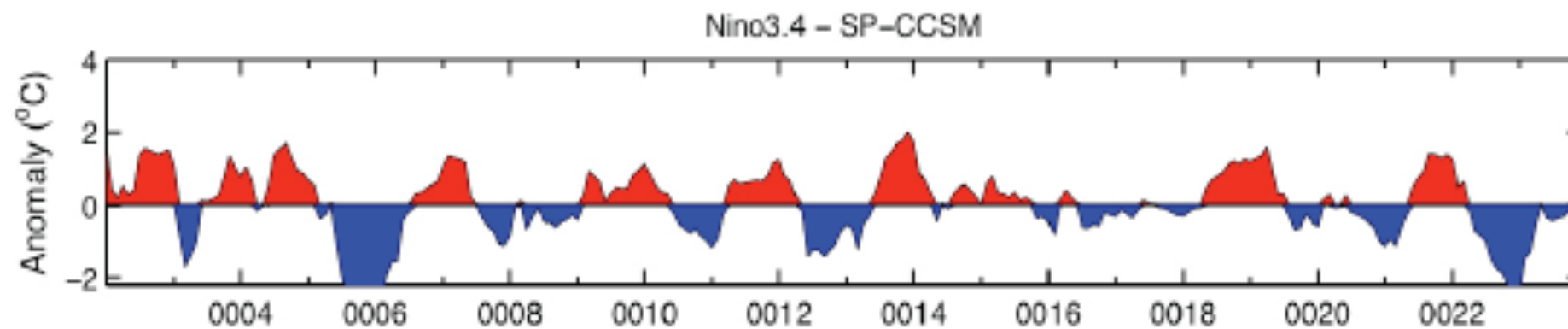


# ENSO Simulation

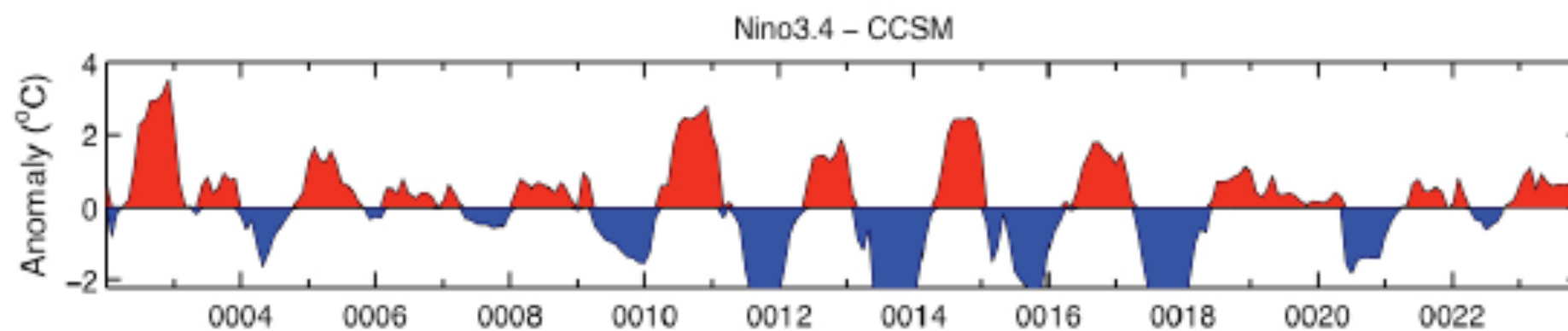
## Niño 3.4 (5S-5N,170W-120W)



**OBS**



**SP-CCSM**



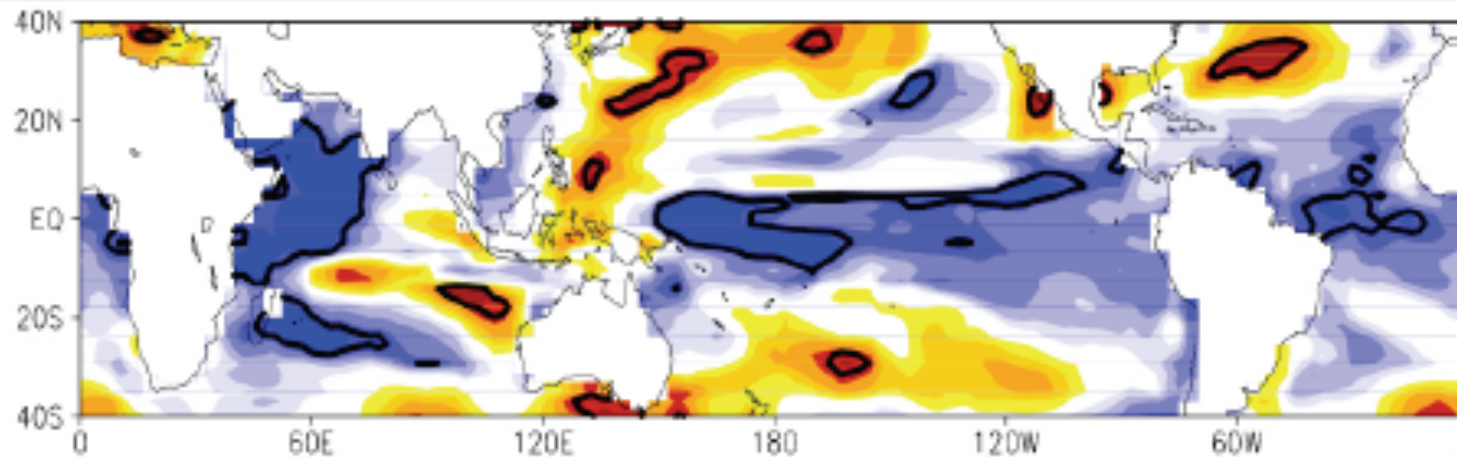
**CCSM**



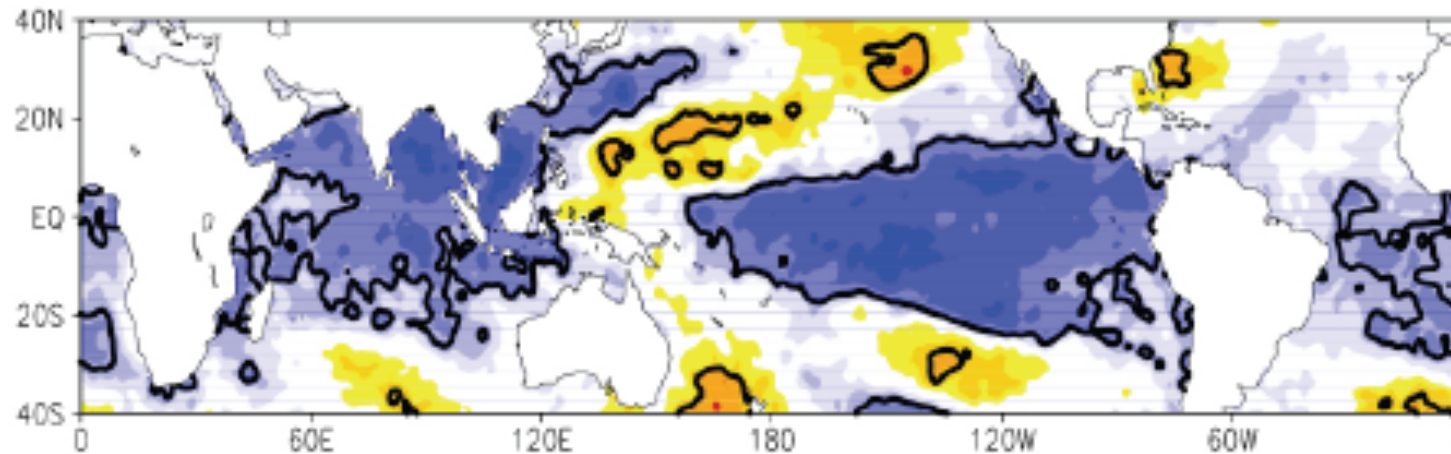
# ENSO-Monsoon Relationship

## IMR (JJA), SSTA(DJF-1)

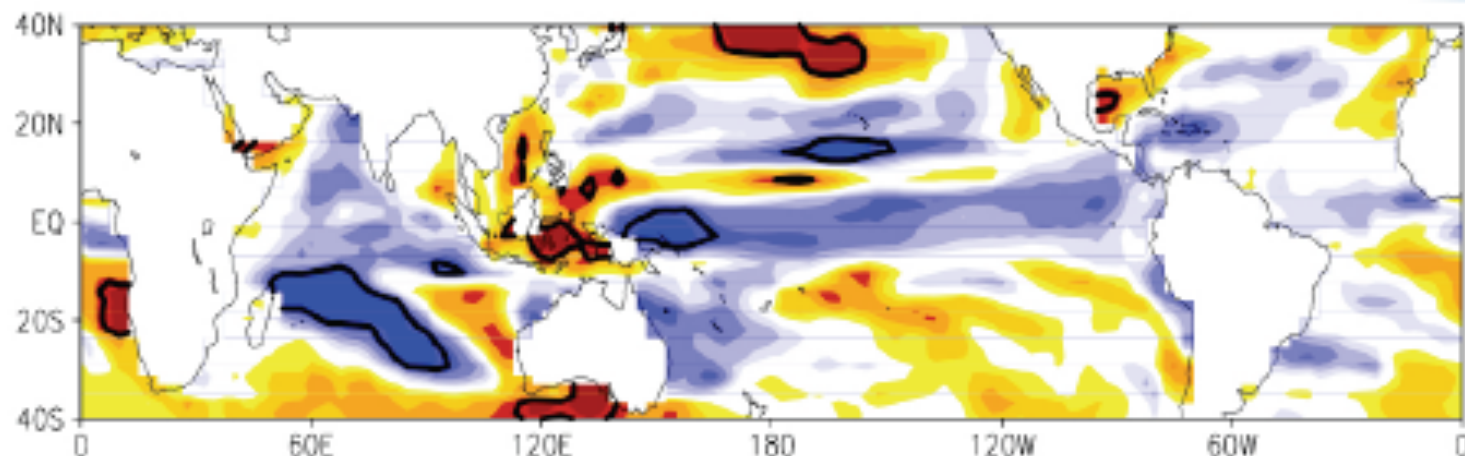
**SP-CCSM**



**OBS**



**CCSM**



# Conclusions on SP-CCSM run

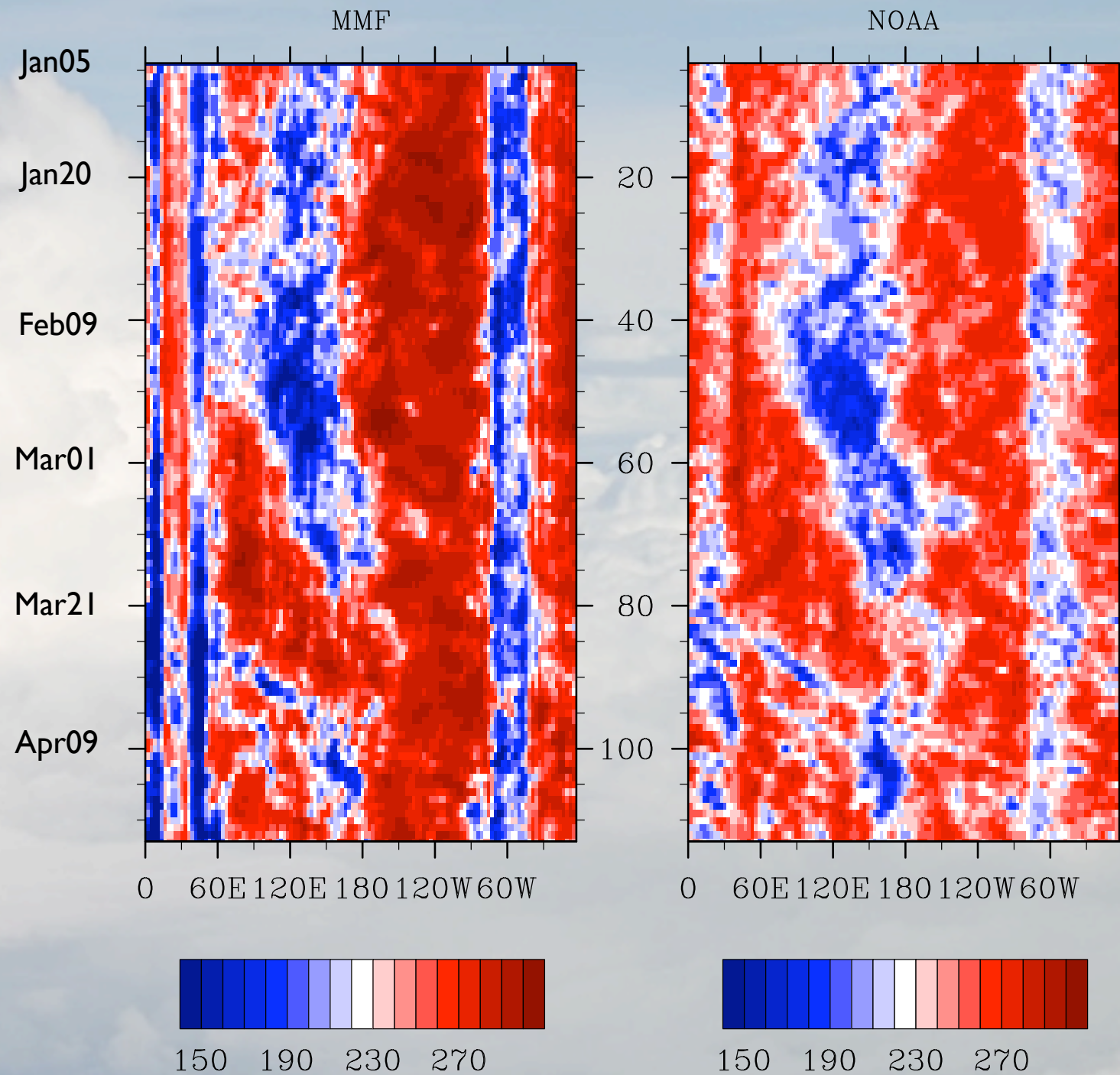
- Improved ENSO;
- Improved relationship between Indian Monsoon and SST anomalies in the Eastern Pacific the following season;
- Robust and realistic MJO simulation



# Towards hindcasts using SP-CAM: Generating initial conditions

OLR,  $\tau=6h$

$u, v, \omega, T, q$ , and  $ps$  are nudged to ERA40





# Proposed Contributions to YOTC

- Hindcasts over YOTC period starting from ECMWF analysis;
- Free-runs (ensemble?) with prescribed SSTs;
- Couple runs with SP-CCSM over YOTC period