

Tropical cyclones in high-resolution ($\Delta x \sim 10\text{km}$) climate simulations: Successes and remaining challenges

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*Thanks to: Warren Washington, Jim Hack and DoE
for computer time at ORNL*

YOTC AMY Symposium, Beijing, May, 2011

Outline

Overview

Storm tracks, intensities, and structure

– *Short runs with CAM5 and GEOS5*

Emphasis on potentially informative diagnostics ...

Sensitivities to Model Physics

Deep Convection Scheme on/off

Condensate loading

Precipitation Objects

A promising approach to understanding cyclogenesis?

Summary and Future Directions

Simulations of Global Hurricane Climatology, Interannual Variability, and Response to Global Warming Using a 50-km Resolution GCM

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NOAA/Geophysical Fluid Dynamics Laboratory, Princeton, New Jersey

(Manuscript received 23 January 2009, in final form 1 June 2009)

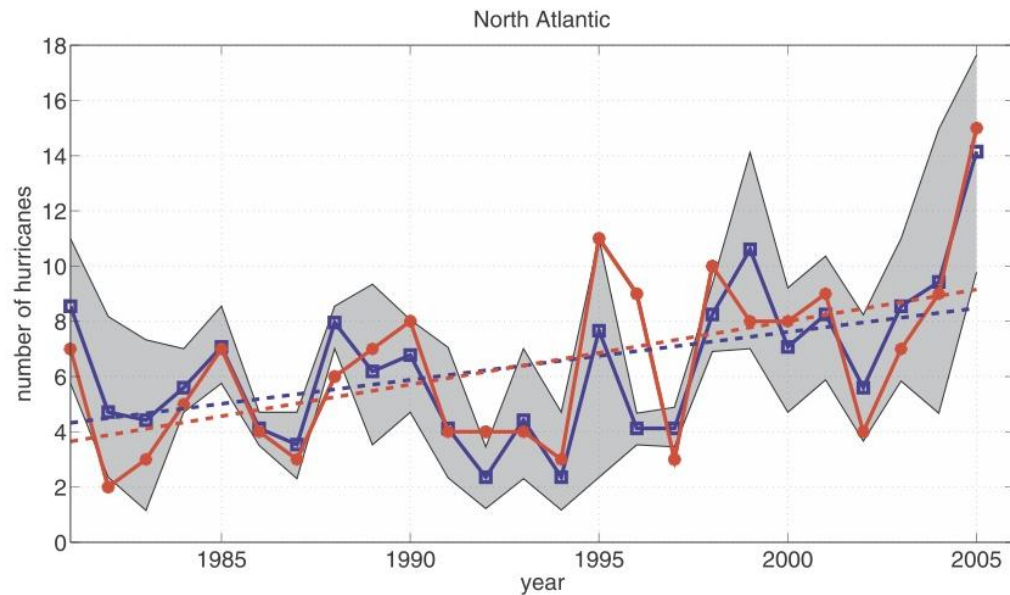


FIG. 7. Interannual variation of hurricane numbers for North Atlantic from 1981 to 2005. IBTrACS observations (Kruk et al. 2010) (red) and four-member ensemble mean (blue); shaded area shows the simulated maximum and minimum number for each year from the four-member integrations. Model time series are normalized by time-independent multiplicative factors so as to reproduce the observed total number. Dotted lines show observed and model (ensemble mean) linear trends.

Models

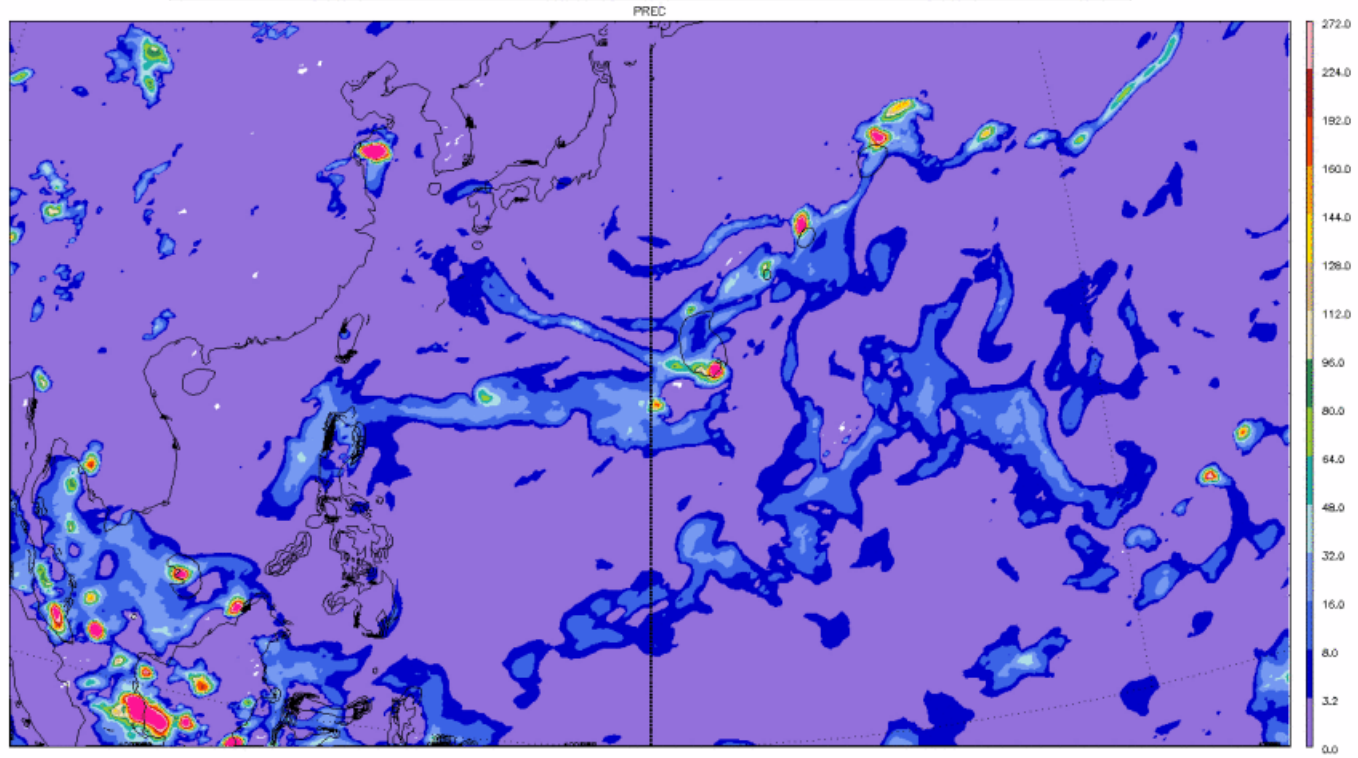
CAM-5

Zhang-Macfarlane deep convection, Park-Bretherton shallow convection and moist PBL, FV-latitude longitude dycore. Resolution shown here 0.23x0.31

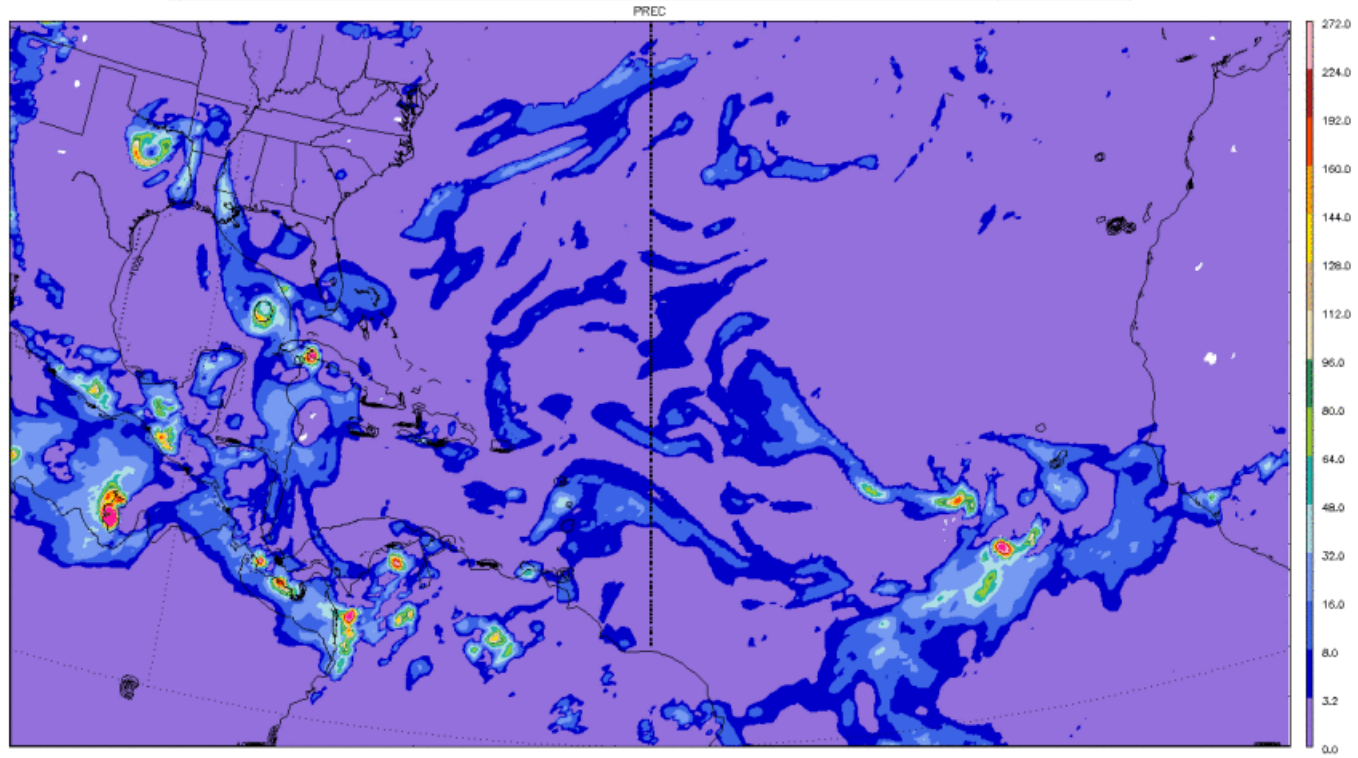
GEOS-5

Relaxed Arakawa-Schubert deep convection with stochastic Tokioka limits on plume entrainment, Lock et al. PBL, FV-cubed sphere. Resolutions shown here ~28,14,7 km

qdLL03: 2005-09-01 :00:00Z



qdLL03: 2005-08-01 :03:00Z



TC Numbers and Tracks

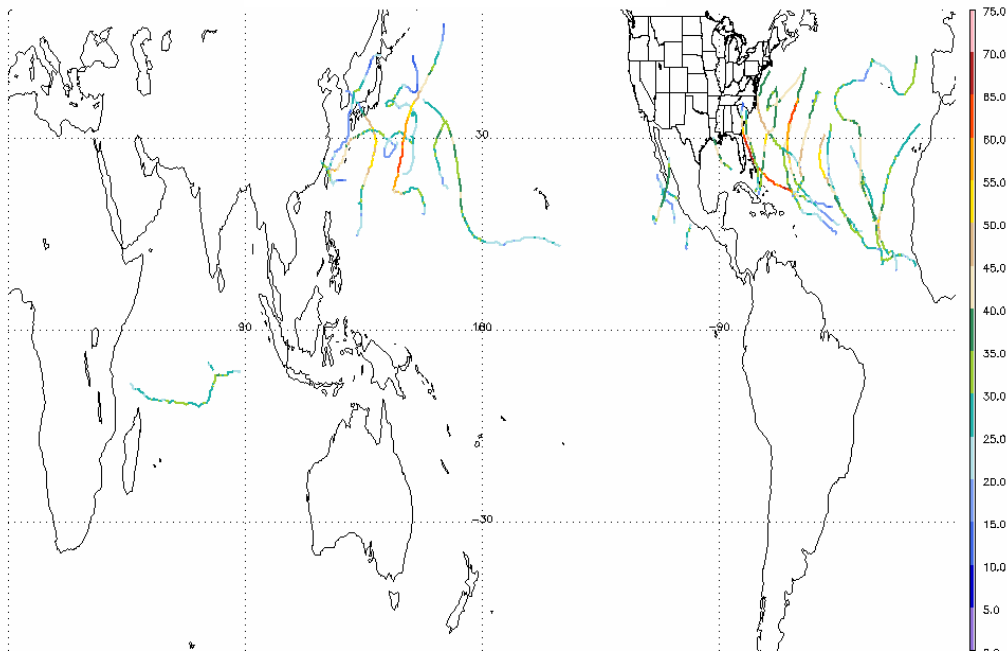
Based on short runs CAM5 and GEOS5 roughly capture correct numbers of storms and distinction between active (2005) and quiet (2006) Atlantic seasons.

Distribution of tracks seems shifted to the east.

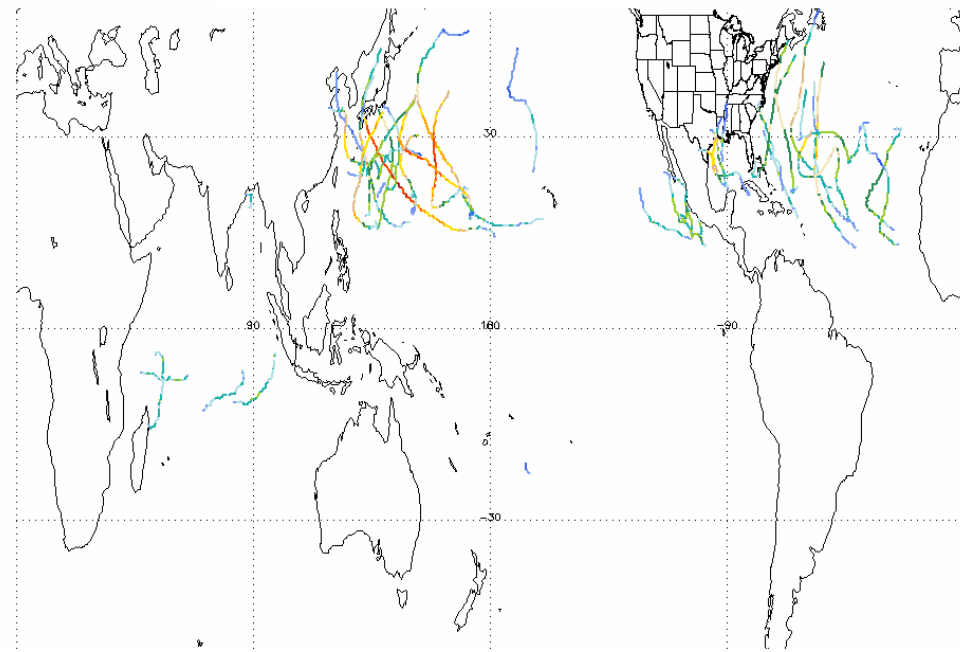
Storms with $U > 33 \text{ ms}^{-1}$: June 1 to Nov 1 2005

200

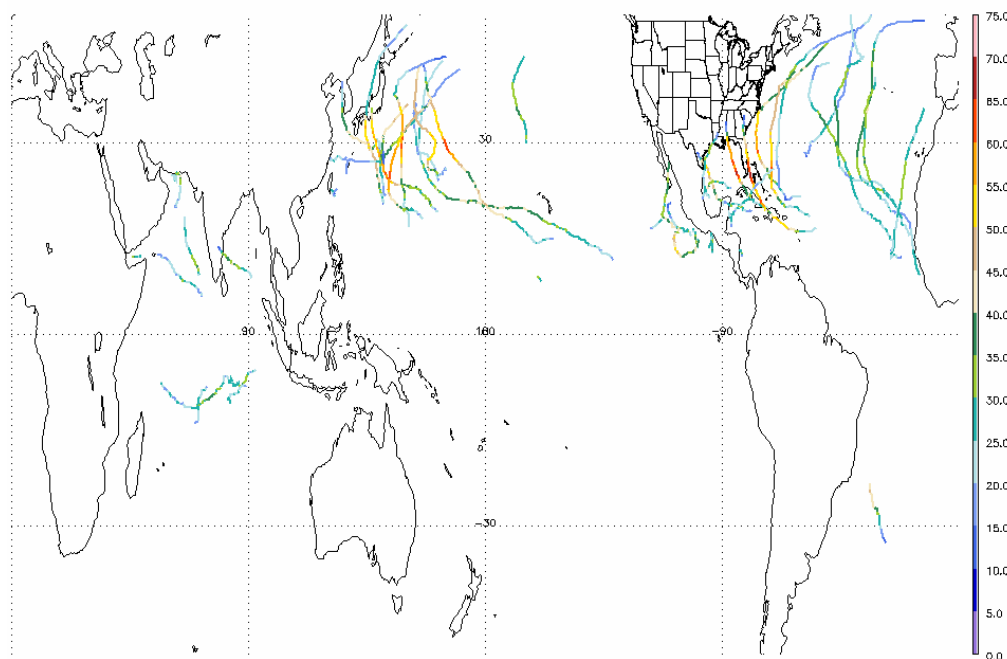
Std CAM5 #1



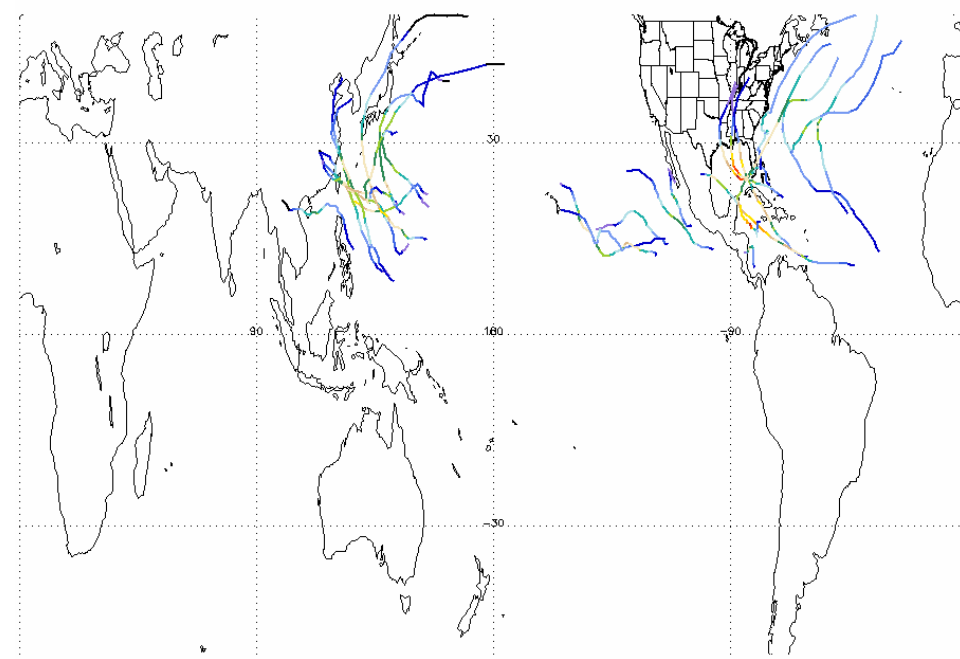
Std CAM5 #2



Std CAM5 #3

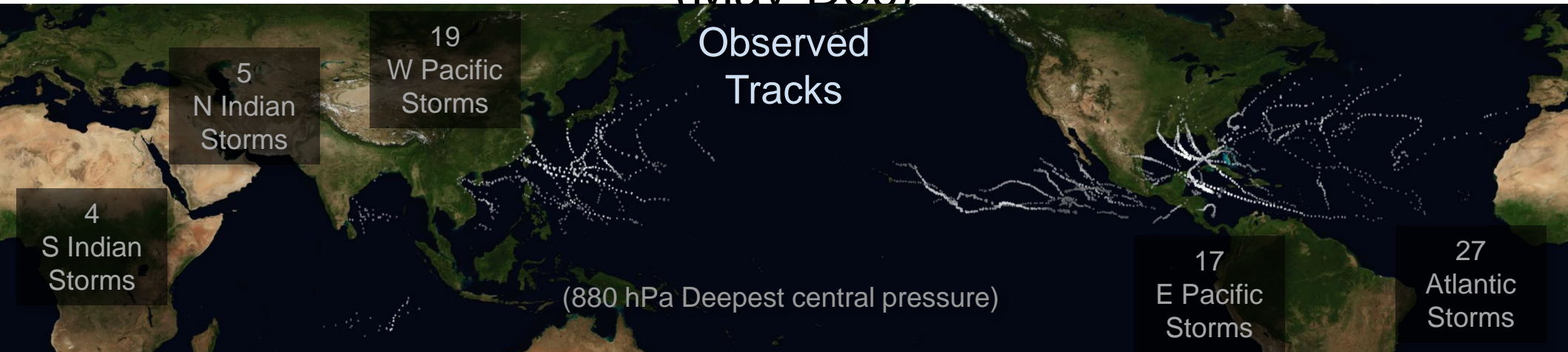


IBTraCS $d > 33 \text{ m/s}$

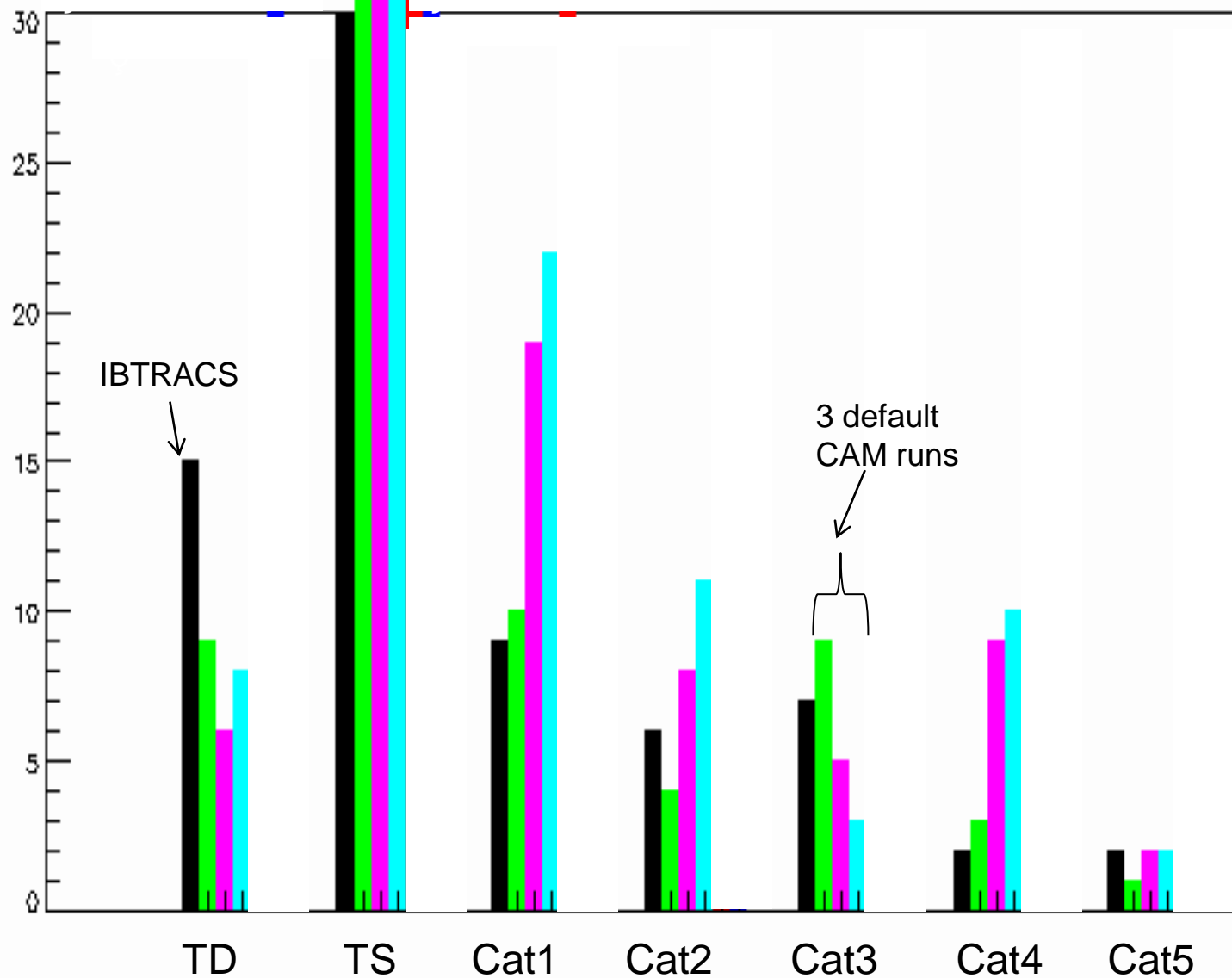


Global Tropical Cyclone Tracks 2005

(May-Dec)



Number of Storms in CAM (*June-Nov 2005*)



Tropical Cyclones 2005-2006 (May-Dec)

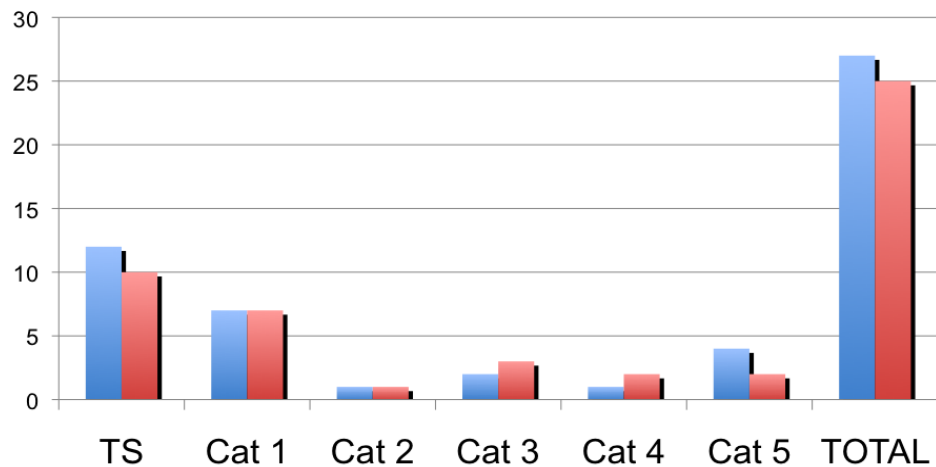
Atlantic Basin Intensities

- GEOS-5 captures the natural variability of tropical cyclone formation
 - In terms of intensity as well as number of storms
- GEOS-5 at high resolution (14-km) shows a capability to develop the most intense tropical cyclones
 - 7 Major Hurricanes (category 3, 4, or 5) in 2005 [7 observed]
 - 2 Major Hurricanes (category 3, 4, or 5) in 2006 [2 observed]

Above Normal Season

2005 Atlantic Hurricanes

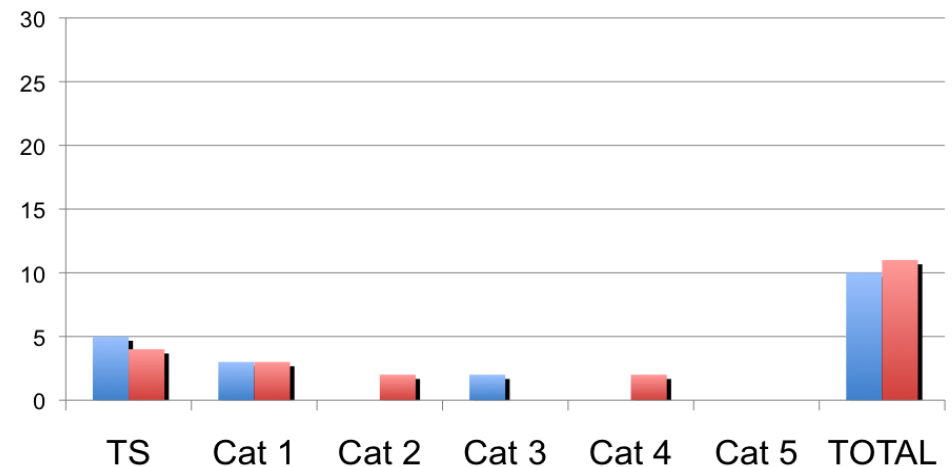
■ Observed ■ 14-km GEOS-5 Nature Run



Below Normal Season

2006 Atlantic Hurricanes

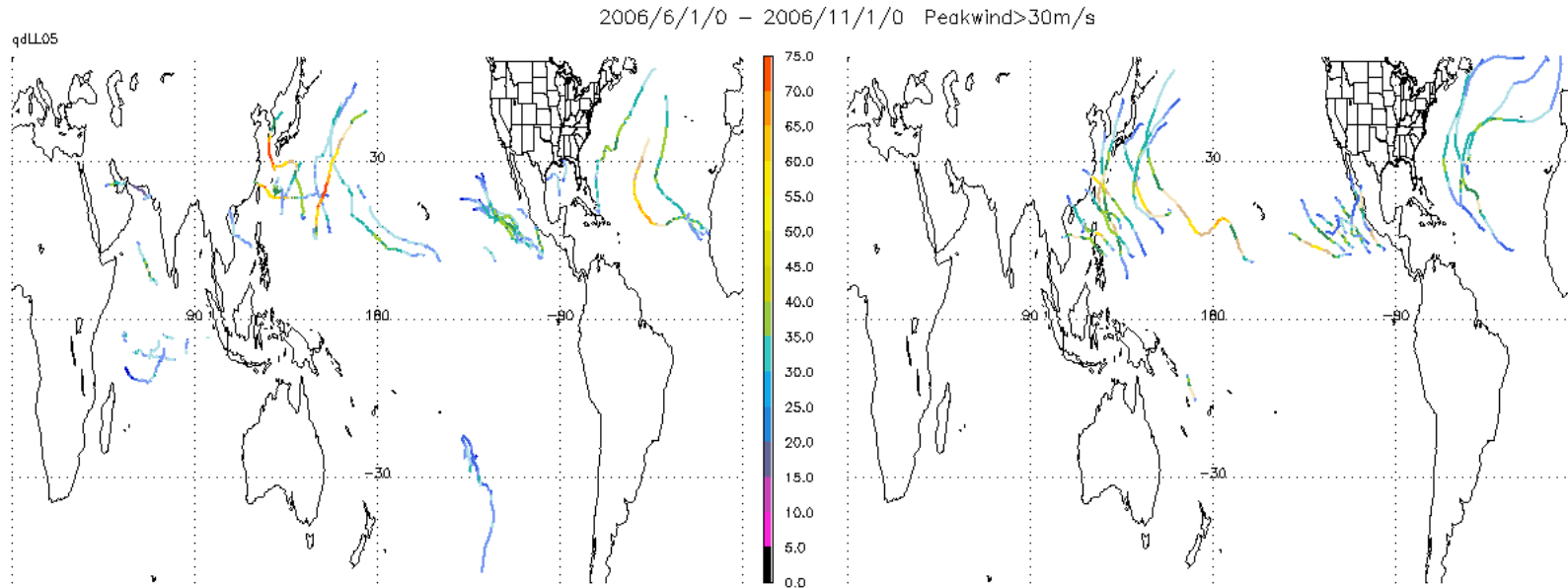
■ Observed ■ 14-km GEOS-5 Nature Run



2006 Season June-November

CAM5

Obs. (IBTRACS)



CAM also captures weak Atlantic season – somewhat stronger E. Pacific in 2006 season

TC Intensities

Intensity is more difficult to capture

Depends on “slowly” varying BCs like SST, and on noisier atmospheric quantities, e.g., shear, dry plumes ...

Models tend to produce storms that are more similar to each other than they are in nature (Zhao et al 2009)

Number of Storms in C/

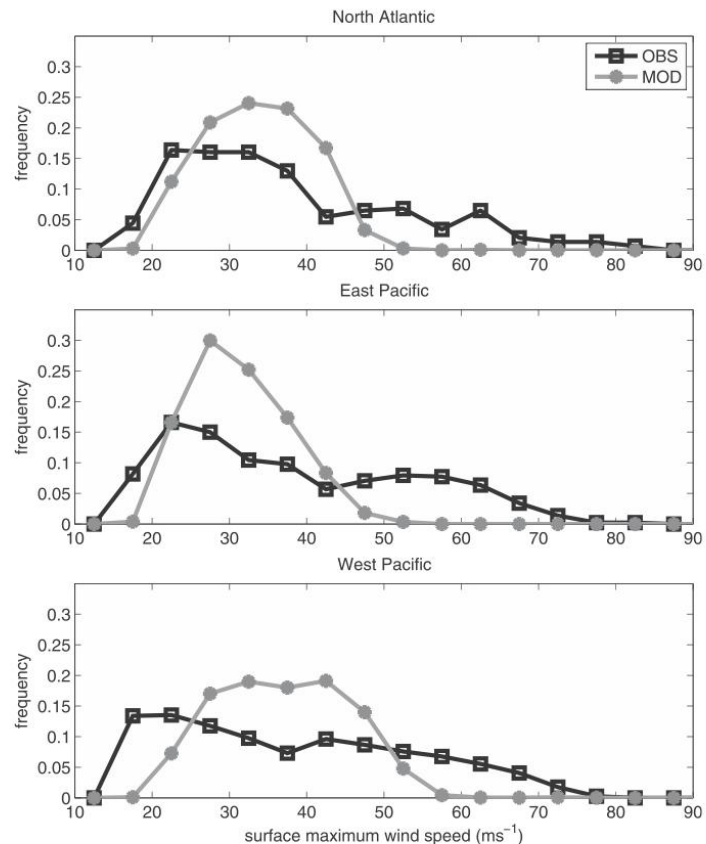
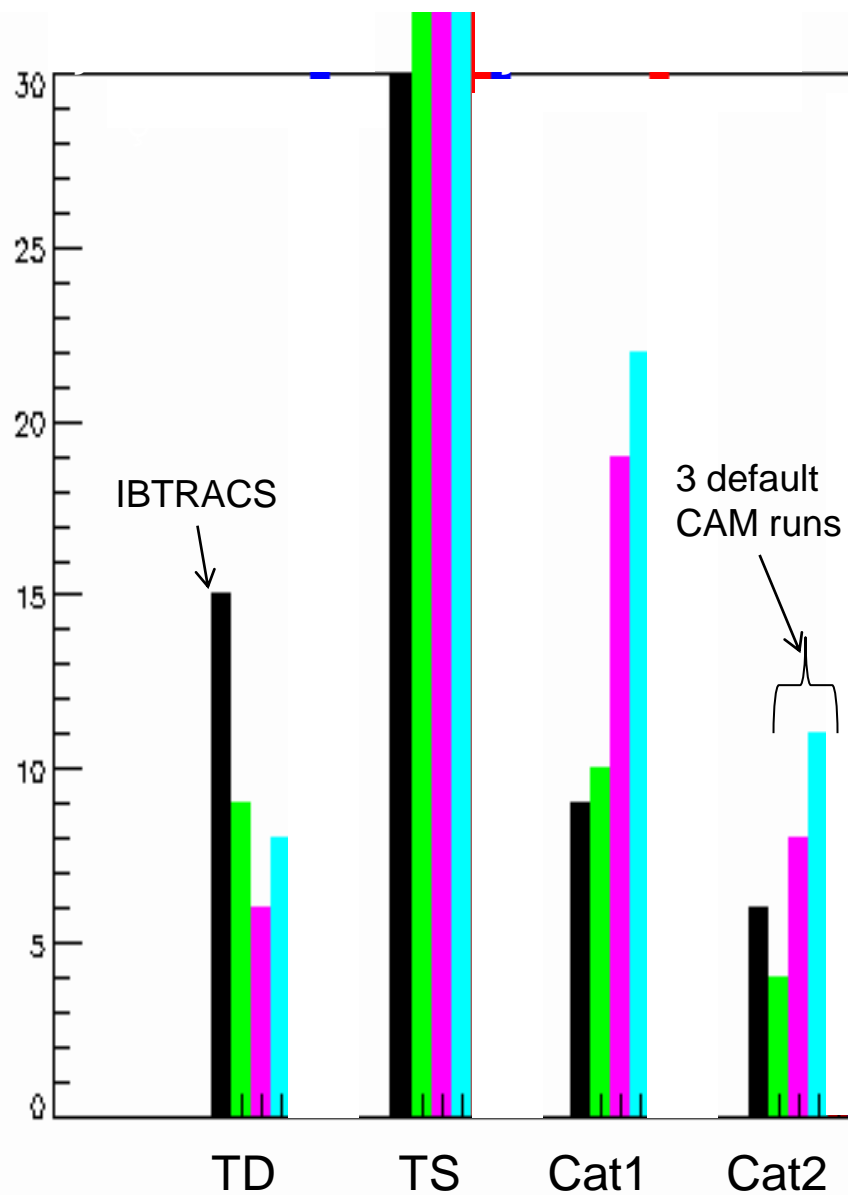
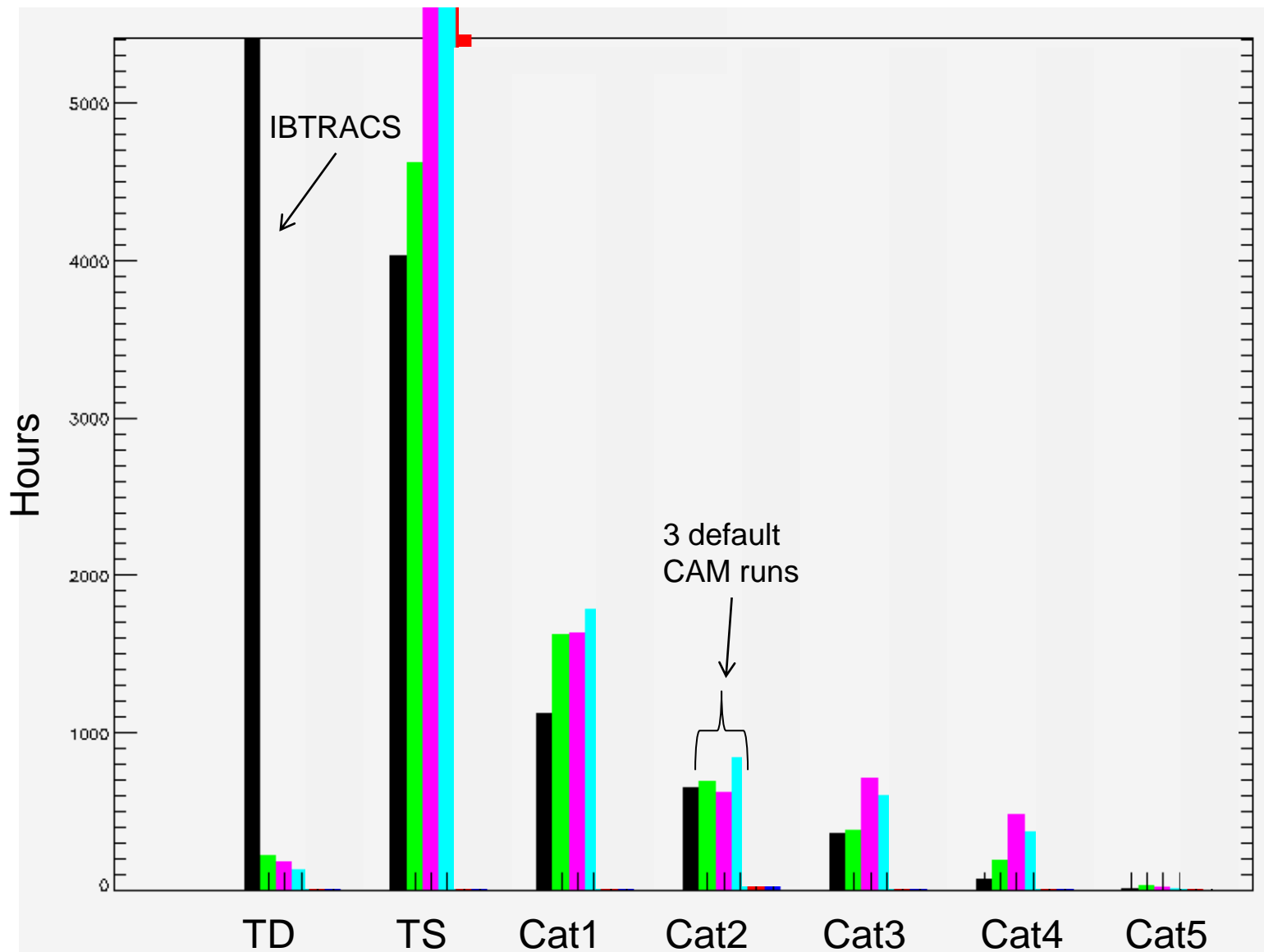
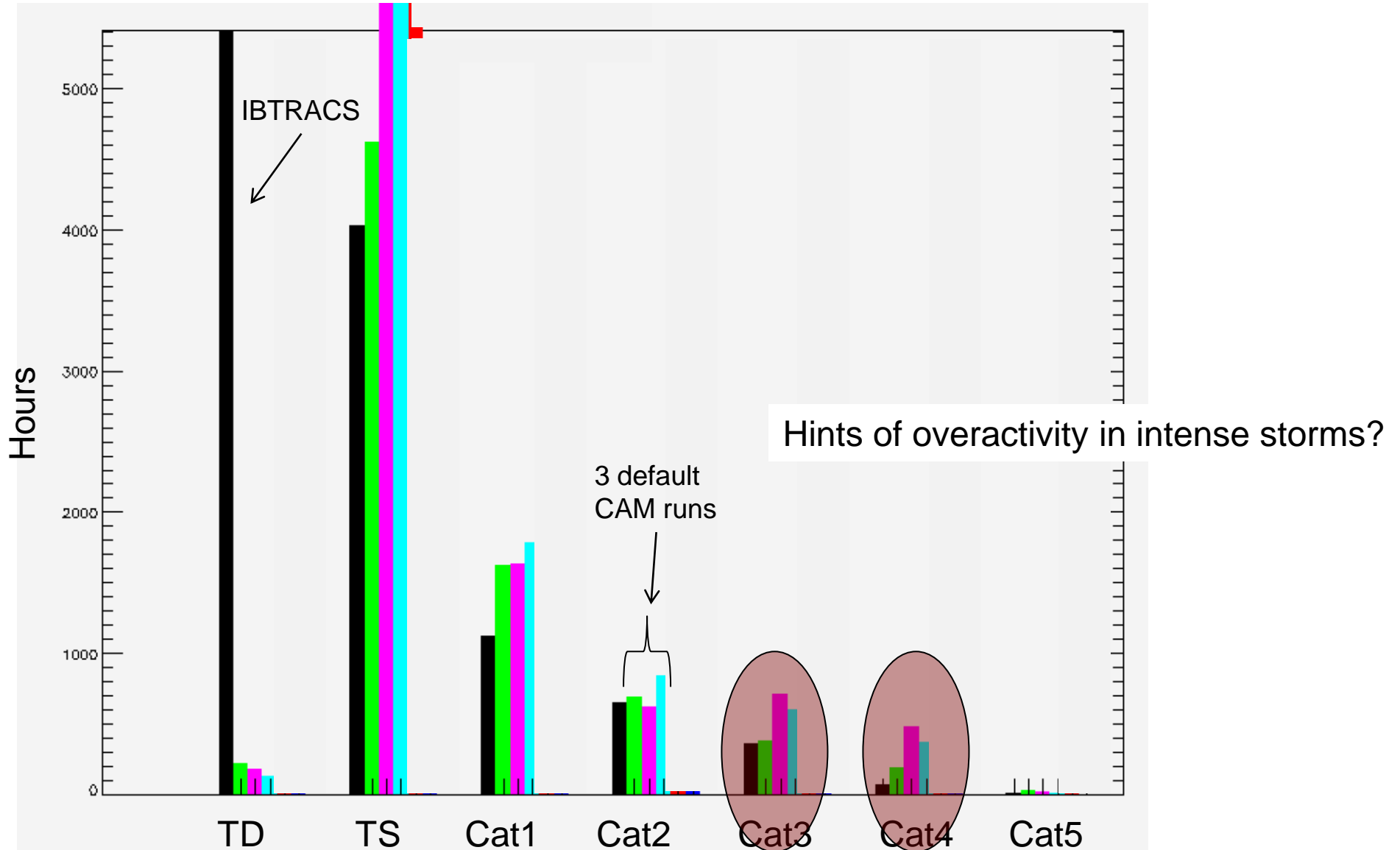


FIG. 6. A comparison of observed and model-simulated tropical storm intensity distribution as characterized by the surface maximum wind speed for the (top) North Atlantic, (middle) east Pacific, and (bottom) west Pacific. IBTrACS observations using 1-min maximum sustained wind at 10 m (black). Model simulation using 15-min (model time step) winds at the lowest model level (gray).

Time (hours) spent at Category (June-Nov 2005)



Time (hours) spent at Category (June-Nov 2005)



Maximum Intensity from Emanuel 1992 ...

THE THEORY OF HURRICANES

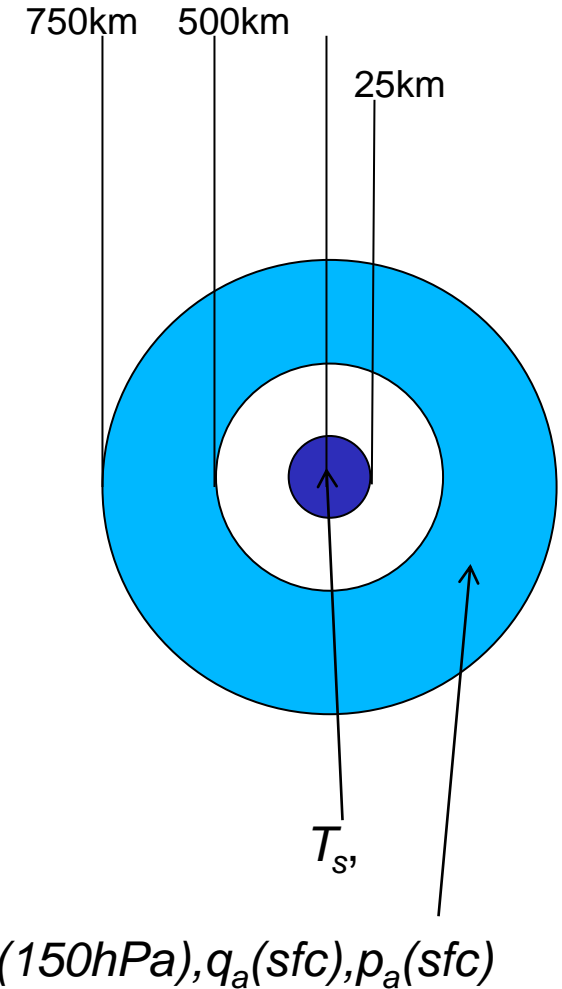
Kerry A. Emanuel

$$T_s \Delta s_{\max} = RT_s \ln \frac{p_a}{p_c} + \frac{L_v}{T_s} (q_c^* - q_a),$$

$$q_c^* = \frac{3.802 \text{ mbar}}{p_c} \exp \left[\frac{17.67 T_s}{243.5 + T_s} \right],$$

$$\varepsilon \equiv \frac{T_s - T_o}{T_s},$$

$$\varepsilon T_s \Delta s = RT_s \ln \frac{p_a}{p_c} + \frac{1}{4} f^2 r_a^2,$$



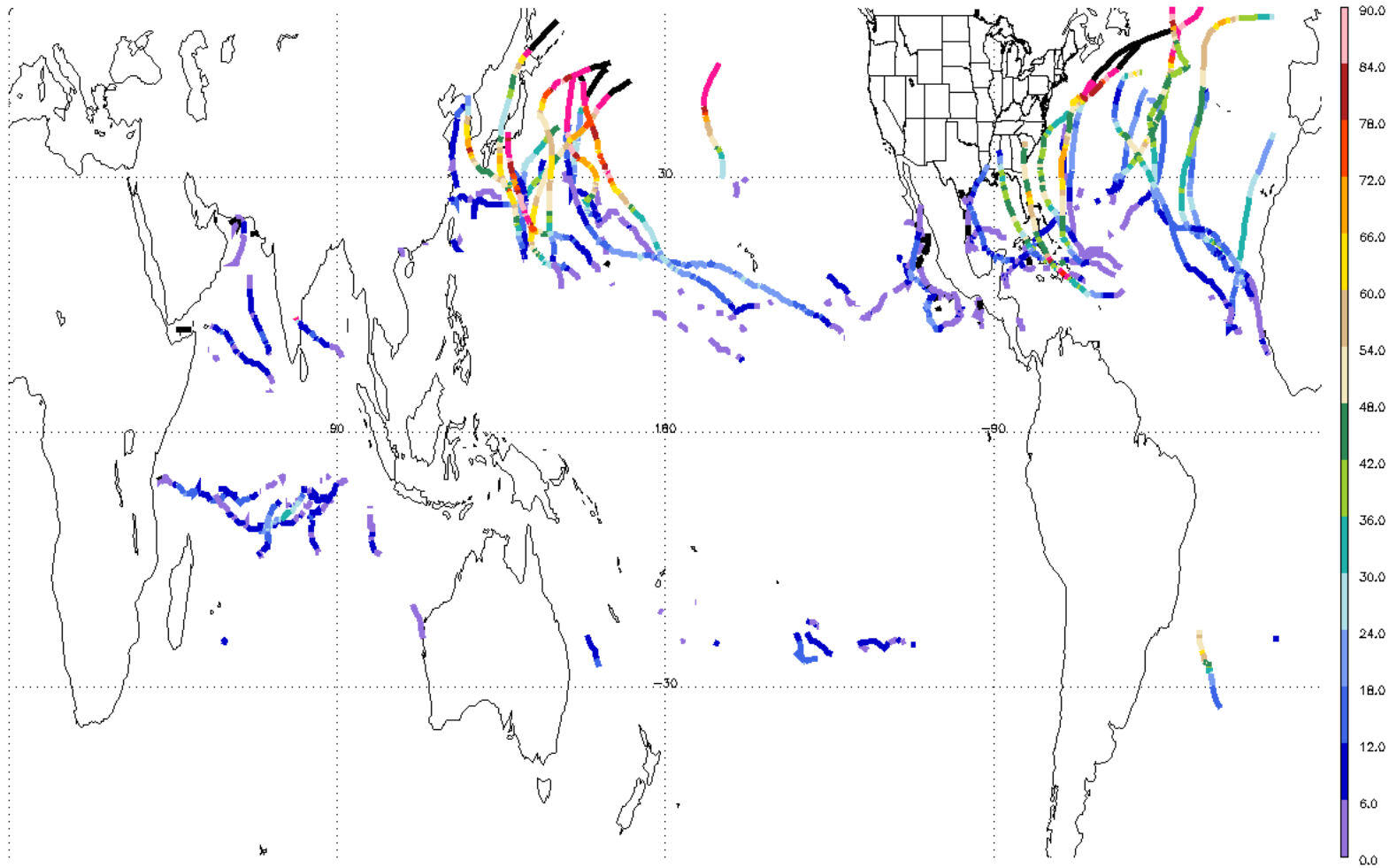
Minimum possible central pressure depends on a few parameters:

- ambient surface pressure
- ambient low-level humidity
- ambient upper-tropospheric temperature
- surface temperature in core region

Percent Maximum Potential Intensity

2005/6/1/0 - 2005/10/31/23 Peakwind>20m/s

$$\frac{p_a - p_c}{p_a - p_{\min}} \times 100$$

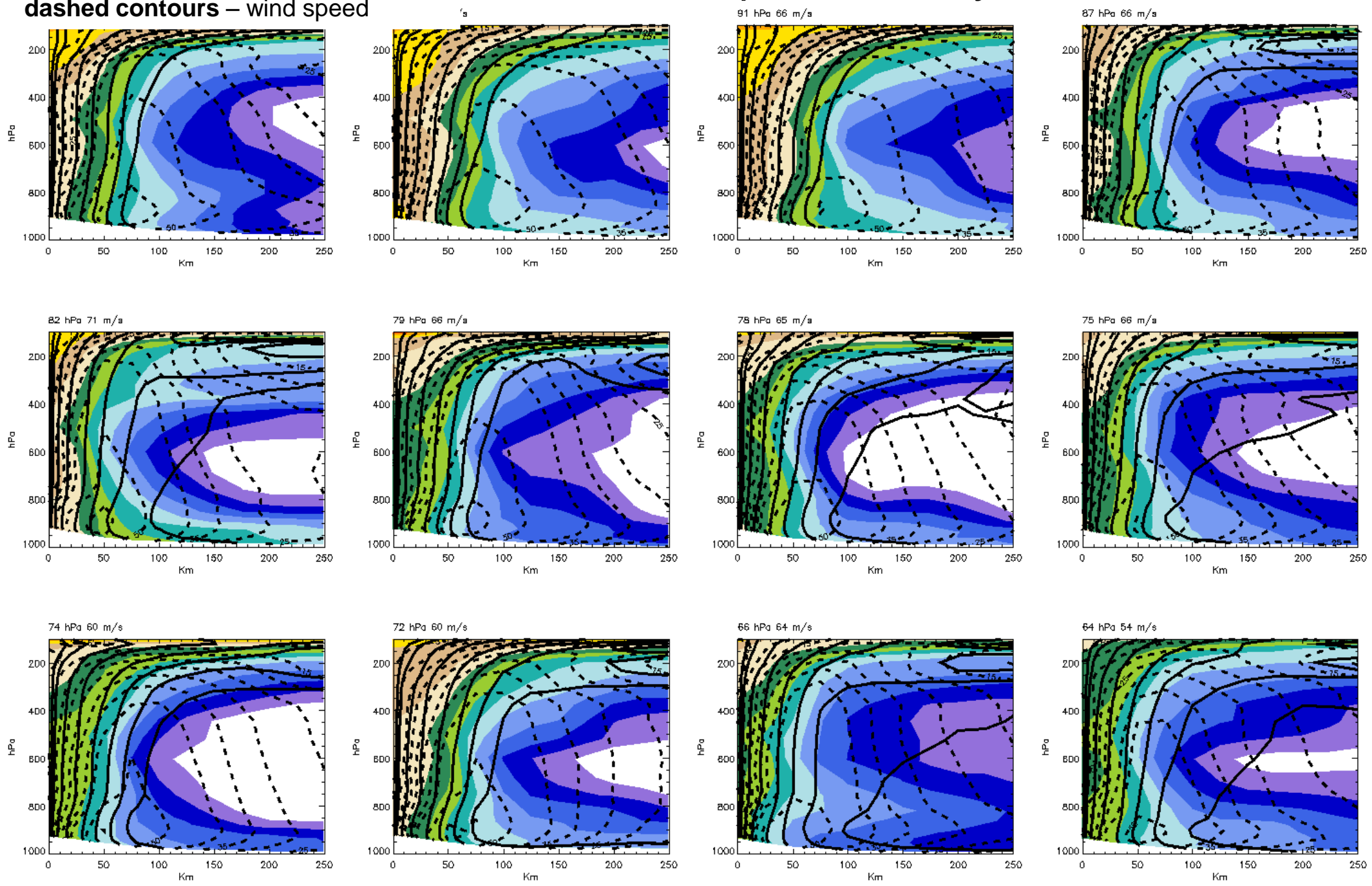


This diagnostic should be relatively straightforward to calculate from re-analysis data as well as model output

Azimuthal means of s , V and angular mom. $s = c_p \log(T) - R \log(p) + \frac{Lq}{T}$ (e.g. Emanuel, 2003)

Shading - moist entropy s ;
solid contours - angular momentum
dashed contours - wind speed

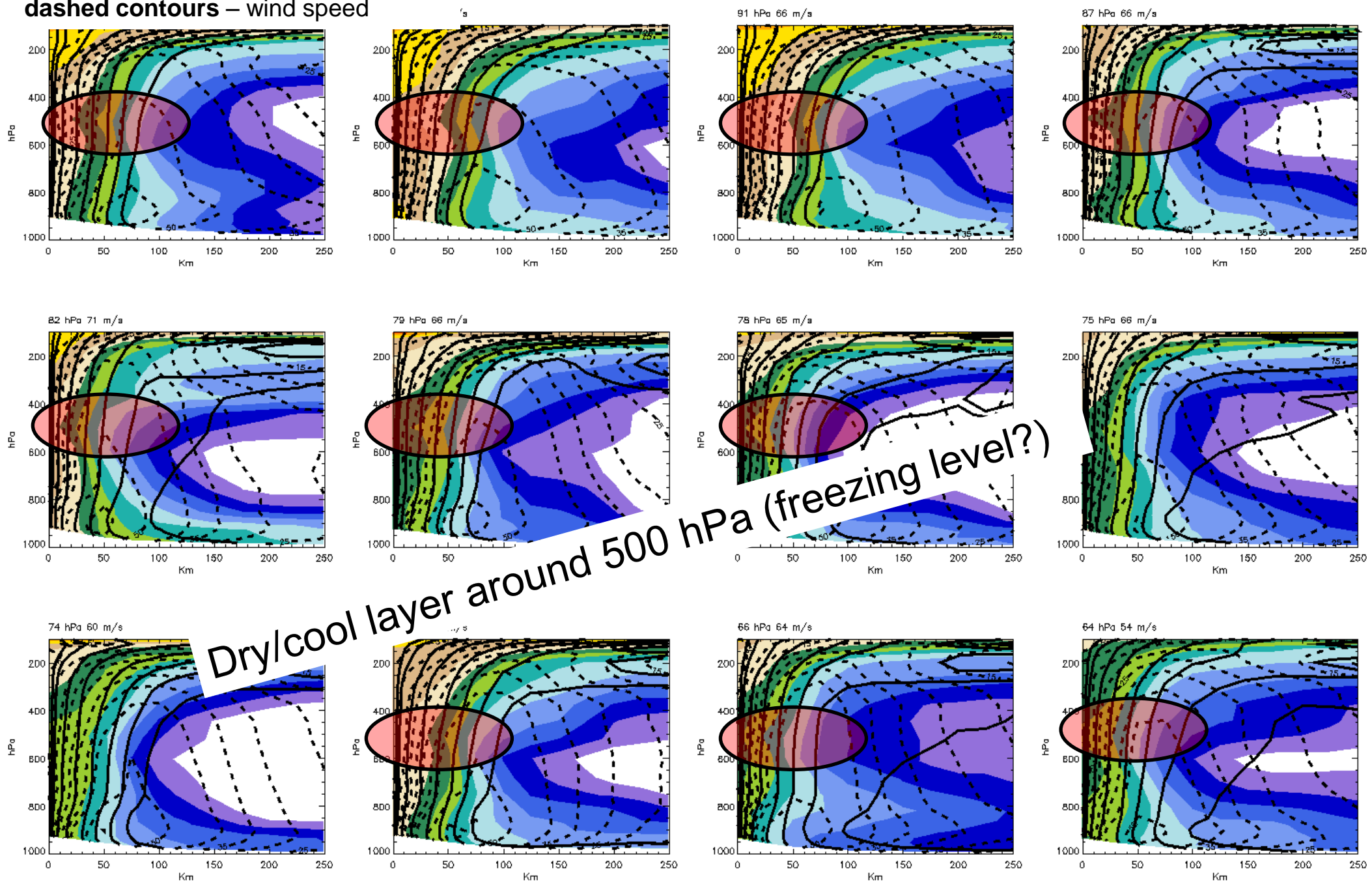
Structure at peak intensity



Azimuthal means of s , V and angular mom. $s = c_p \log(T) - R \log(p) + \frac{Lq}{T}$ (e.g. Emanuel, 2003)

Shading - moist entropy s ;
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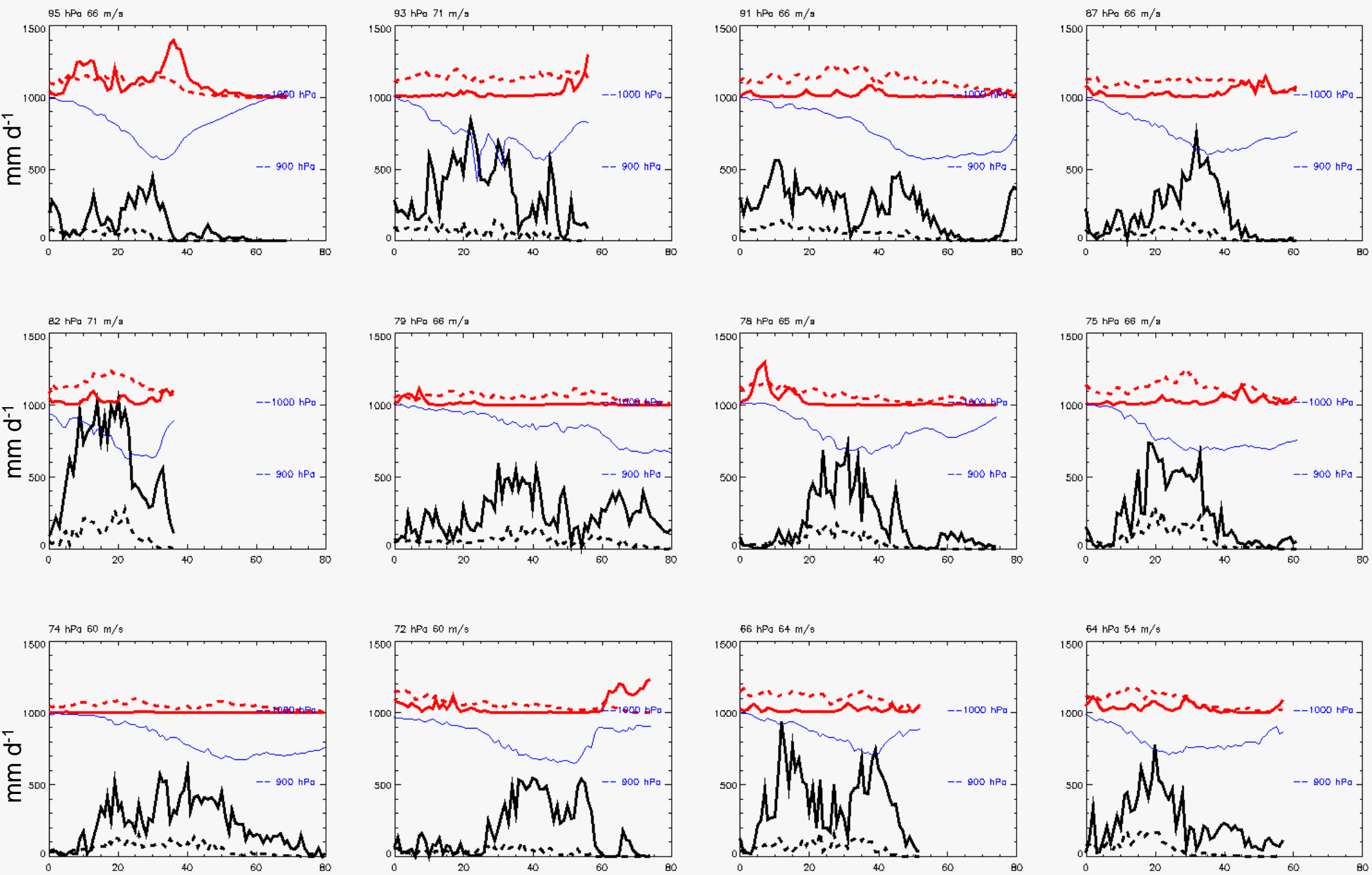
Structure at peak intensity



Precipitation time series in storm core (black), storm exterior 250-500km+1000 mm d⁻¹ (red). Convective precip (dashed), Large-scale precip¹ (solid).

Thin blue lines show surface pressure.

Note overwhelming dominance of LS in cores



Sensitivity to use of Deep Convection Scheme

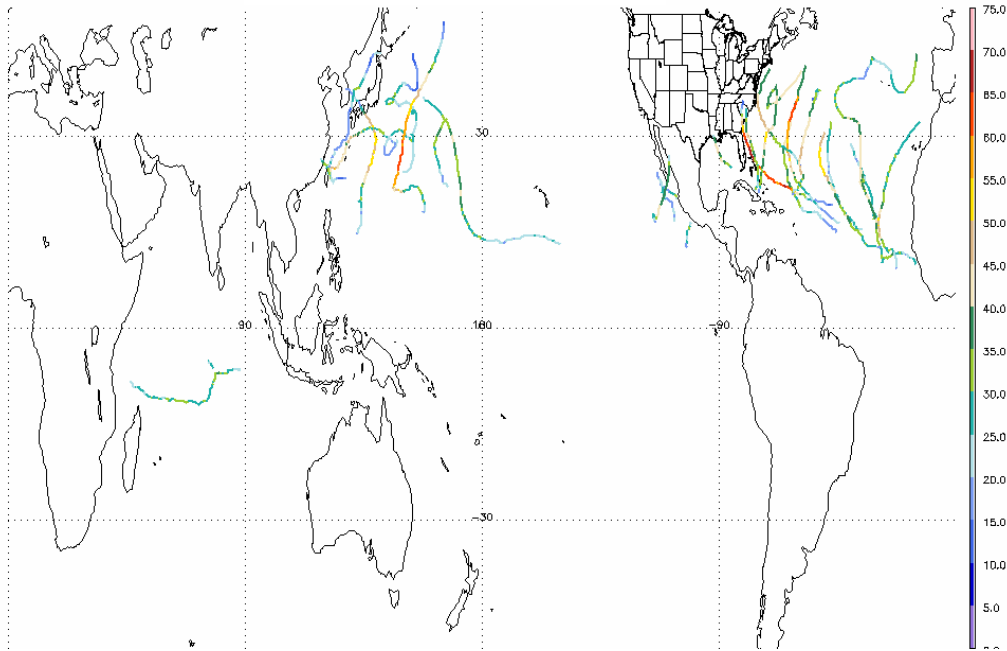
GEOS-5 attempts to hobble deep convection scheme via entrainment limits. GFDL eliminates deep scheme (with tuned shallow scheme). CAM5 precip in TC cores dominated by large-scale.

What happens if deep scheme is removed from CAM?

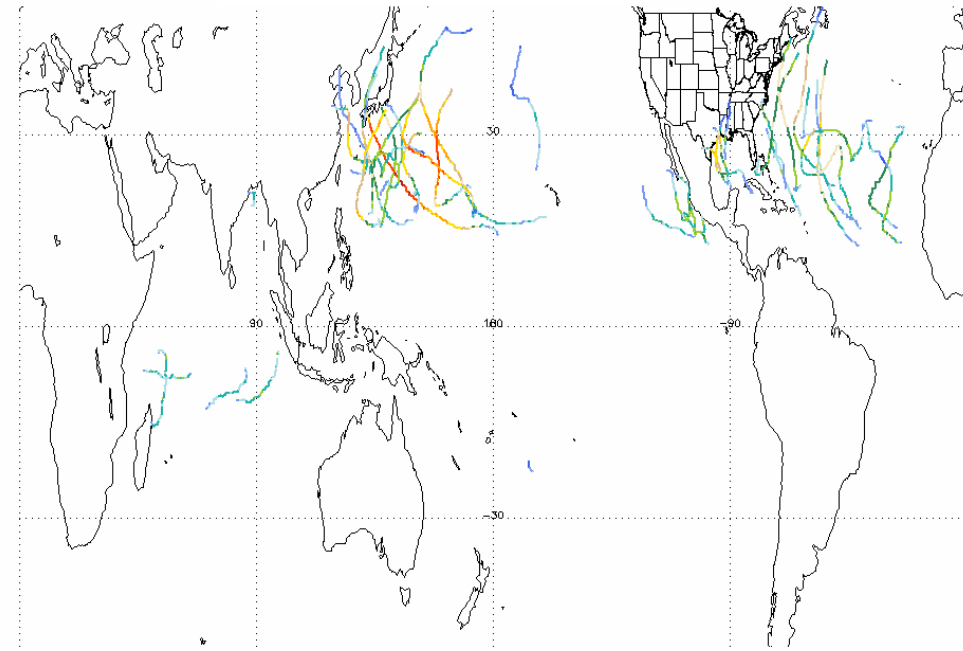
Storms with $U > 33 \text{ ms}^{-1}$: June 1 to Nov 1 2005

200

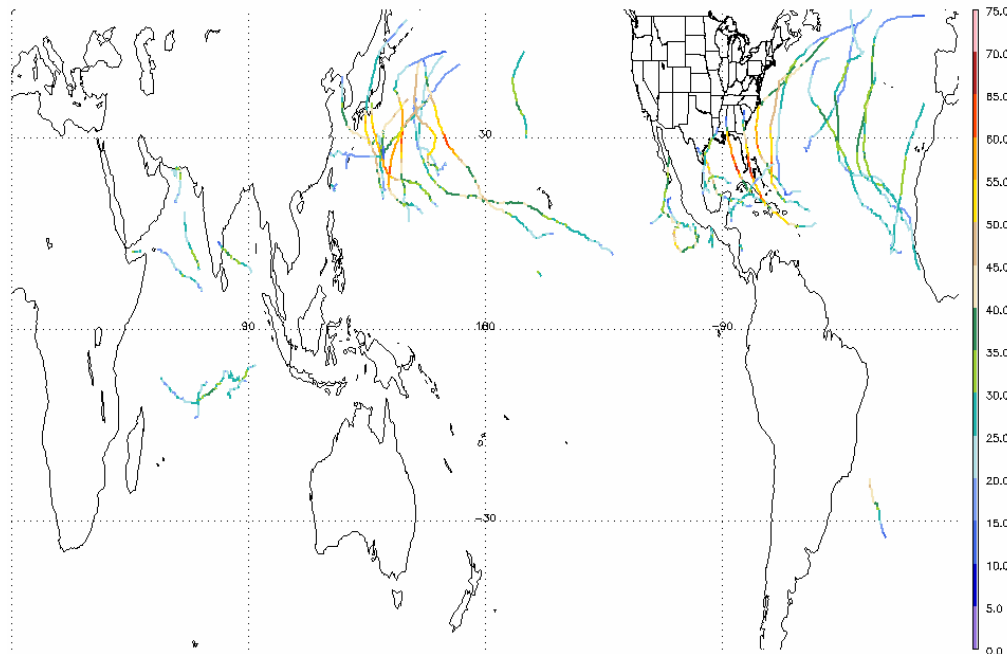
Std CAM5 #1



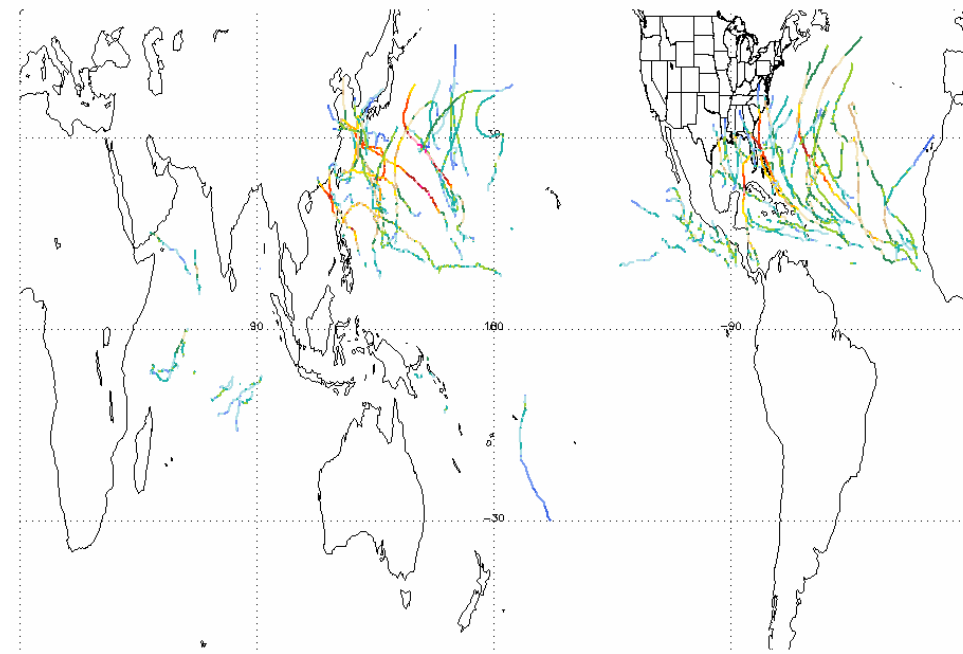
Std CAM5 #2



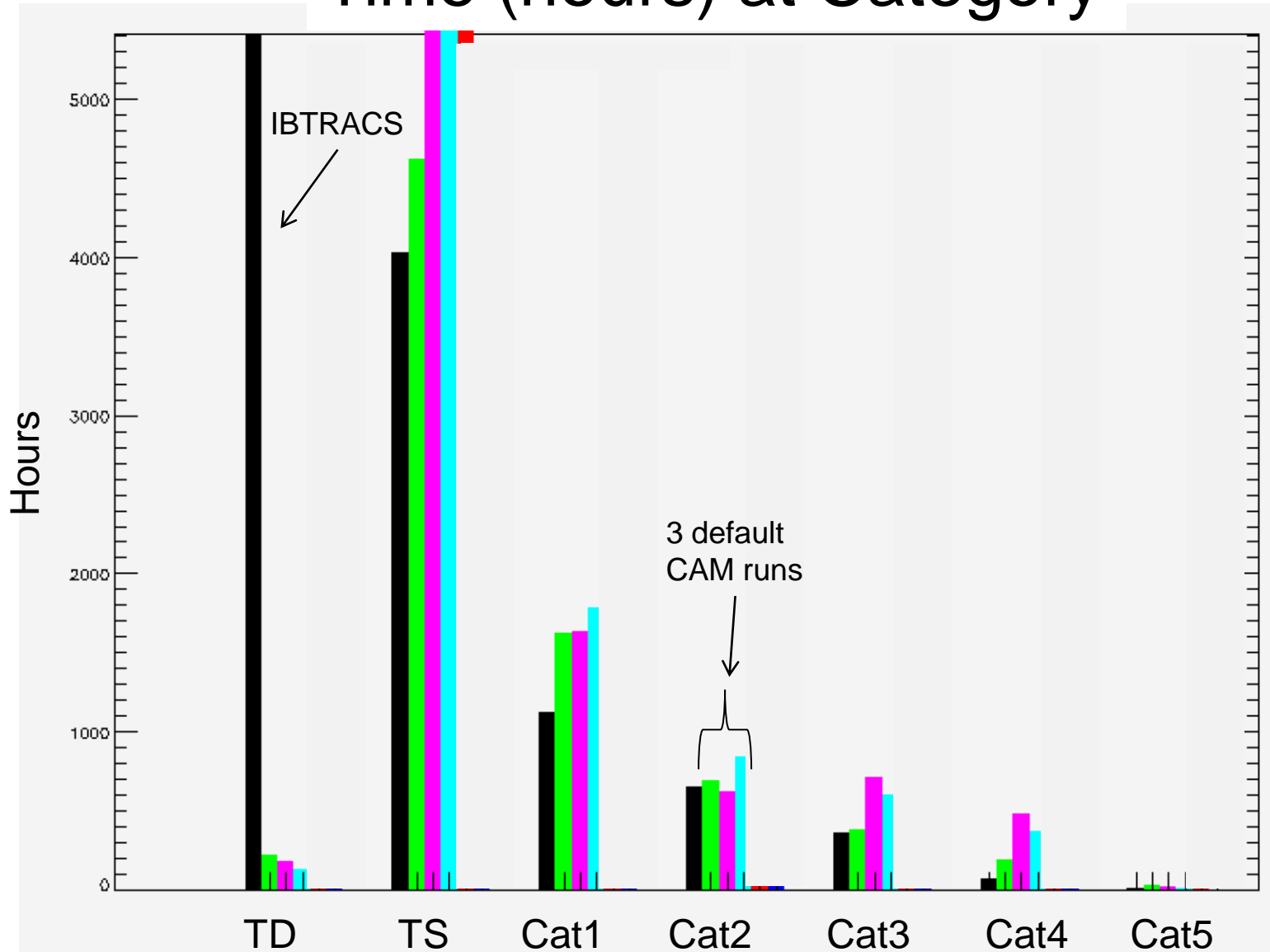
Std CAM5 #3



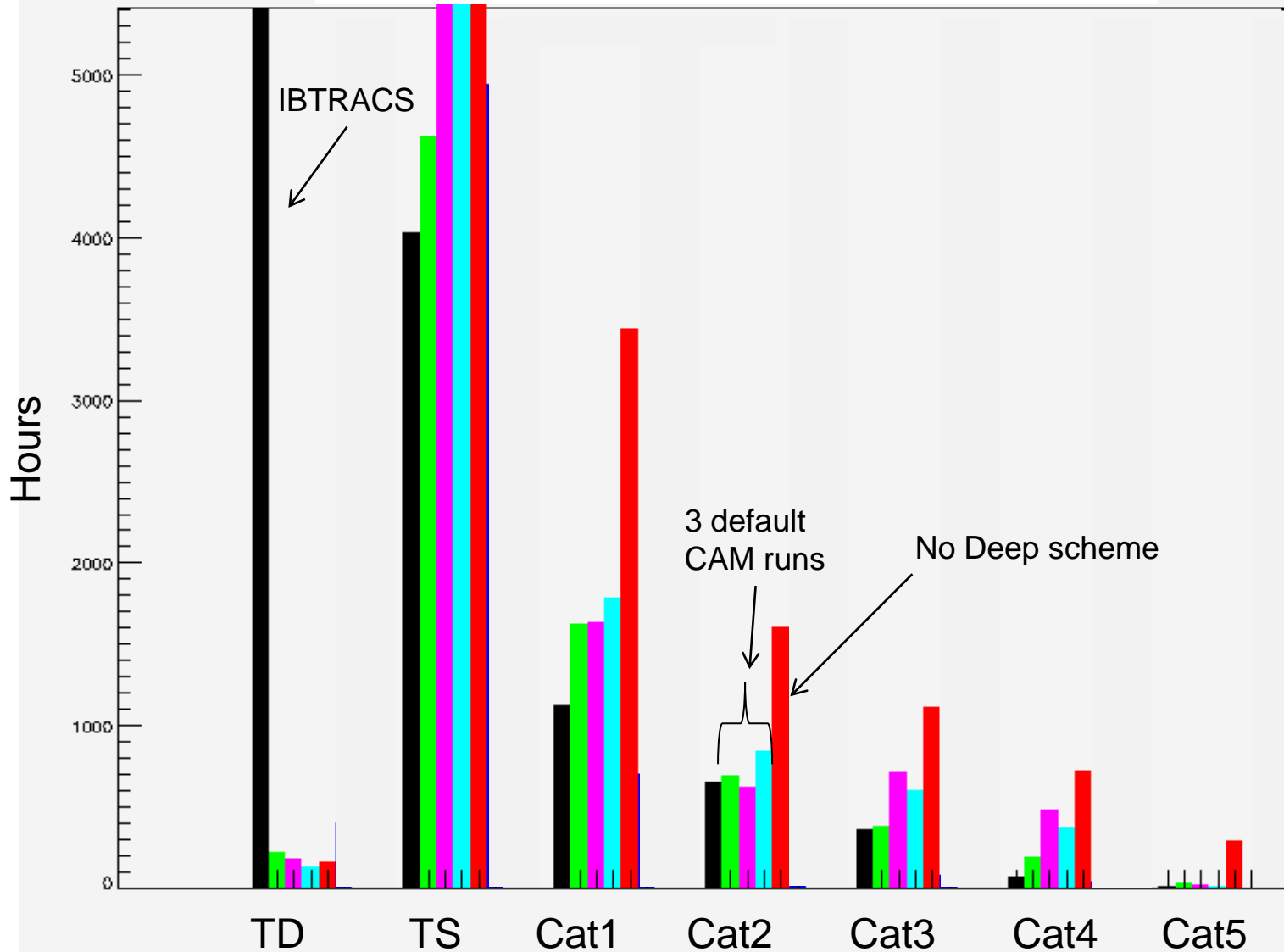
No Deep Convection Scheme



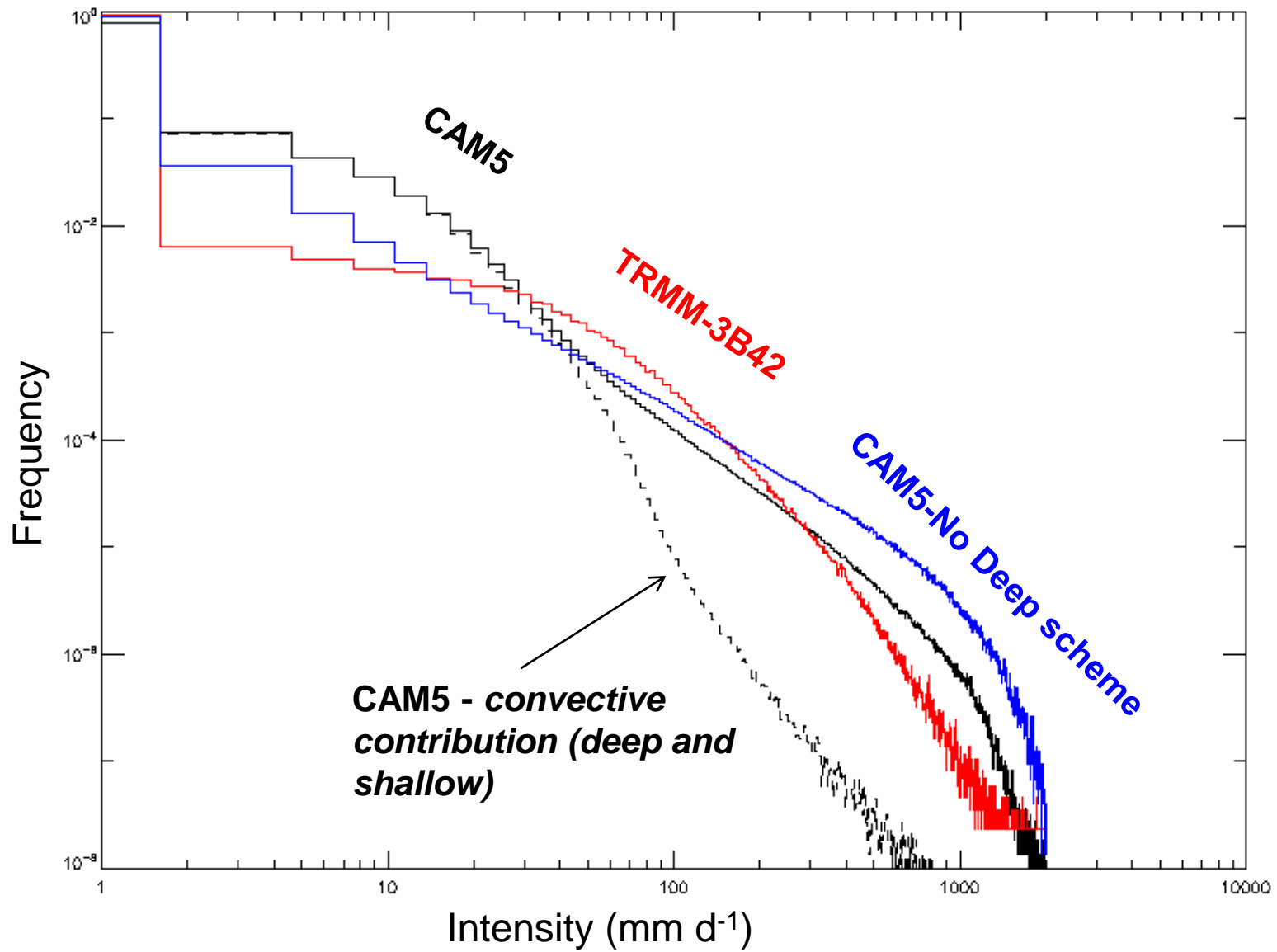
Time (hours) at Category



Time (hours) at Category



PDFs of tropical precipitation (30S-30N) rates Aug 2005

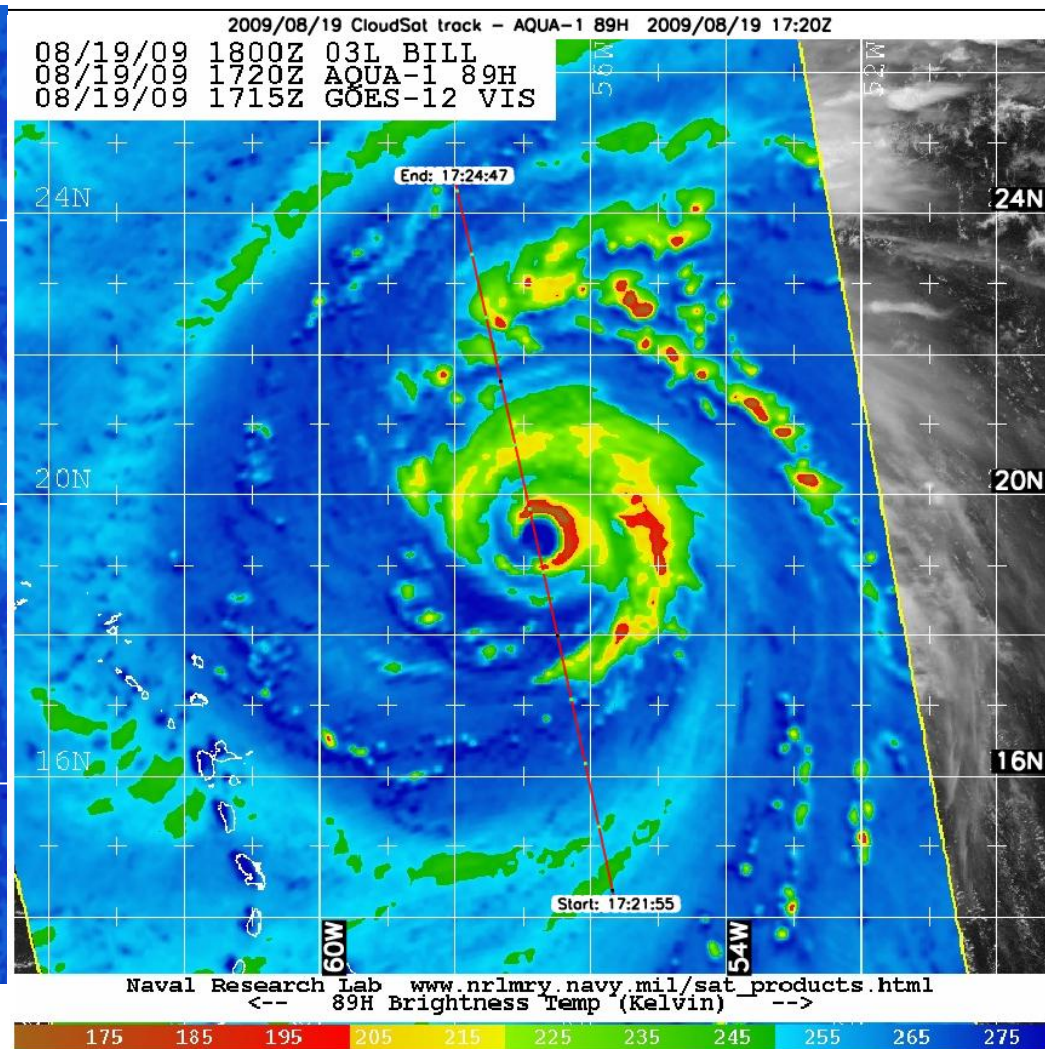
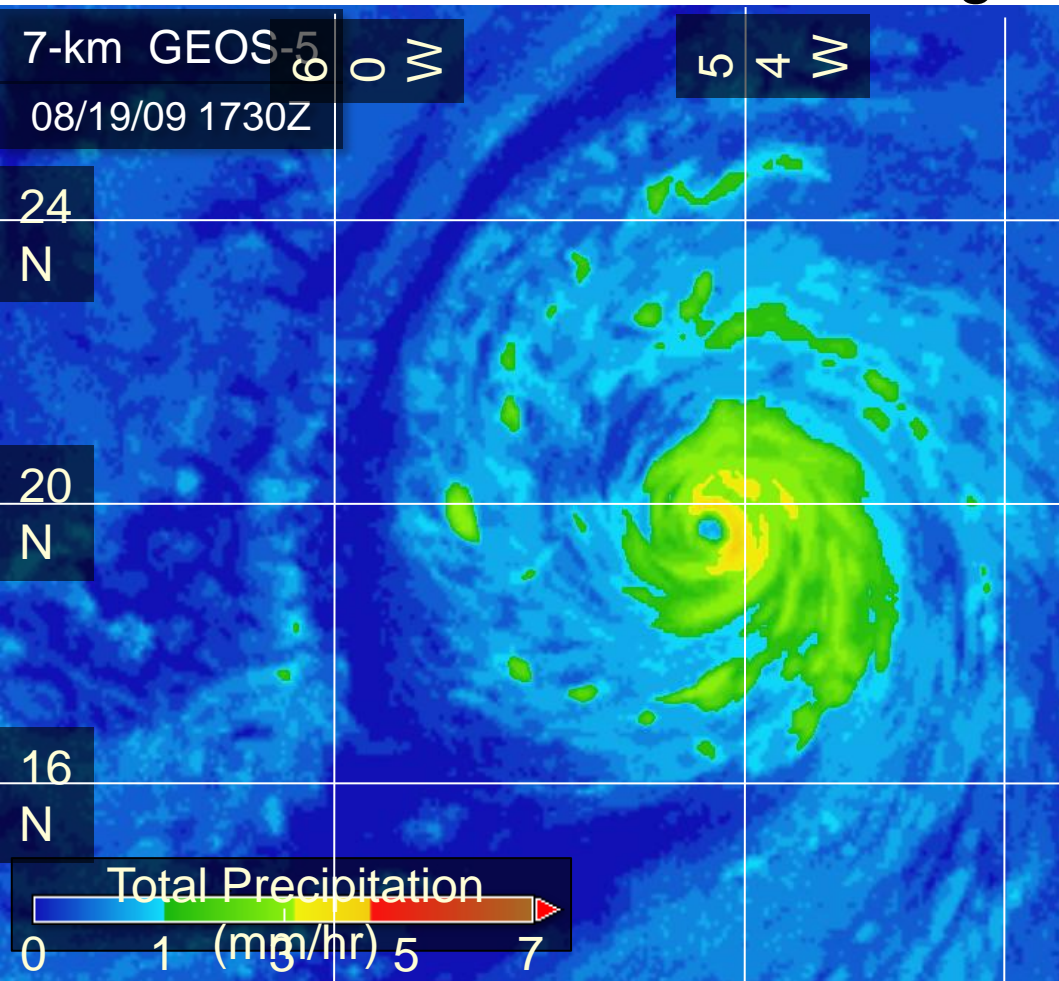


Hurricane Bill

69-hr forecast Initialized

August 2009

2009-08-16 21z



The **strong influence of RAS**, and a **15 minute time-step** for the moist physics leads to problems within the circulation of Hurricane Bill at 7-km resolution:

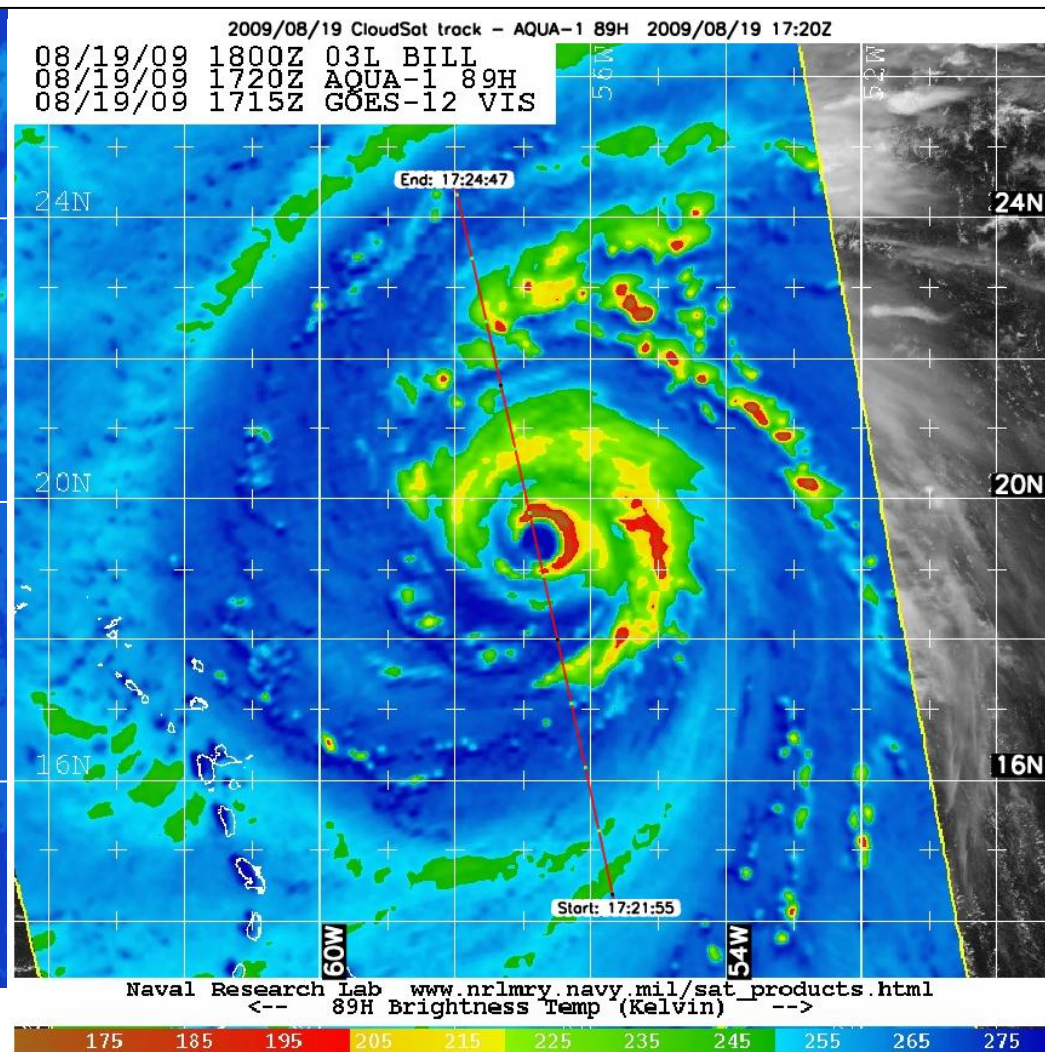
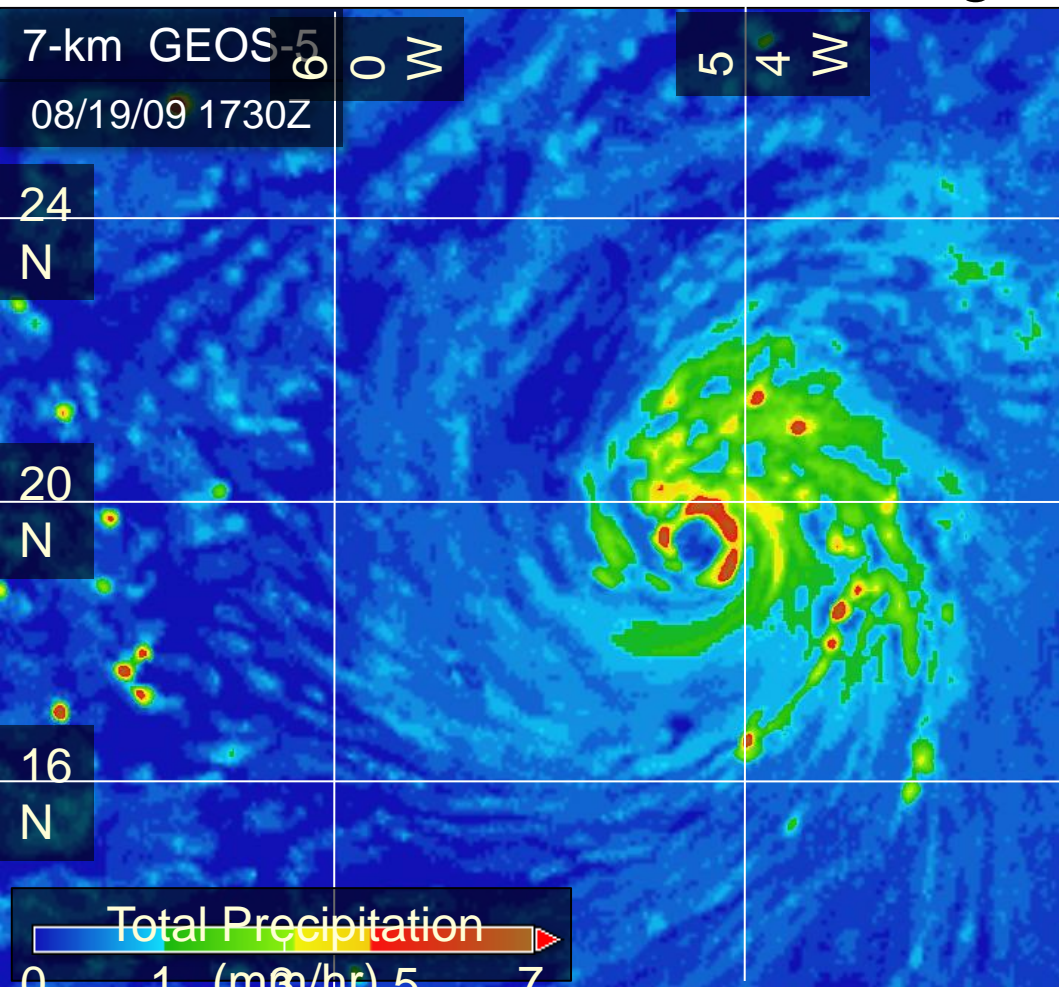
- a lack of deep convective (heavy) precipitation
- an excess of shallow precipitation
- a very small eye, filled with drizzle

Hurricane Bill

69-hr forecast Initialized

August 2009

2009-08-16 21z



Using Tokioka limiter to reduce RAS, and a 120 second time-step for the moist physics improves the convection within Hurricane Bill:

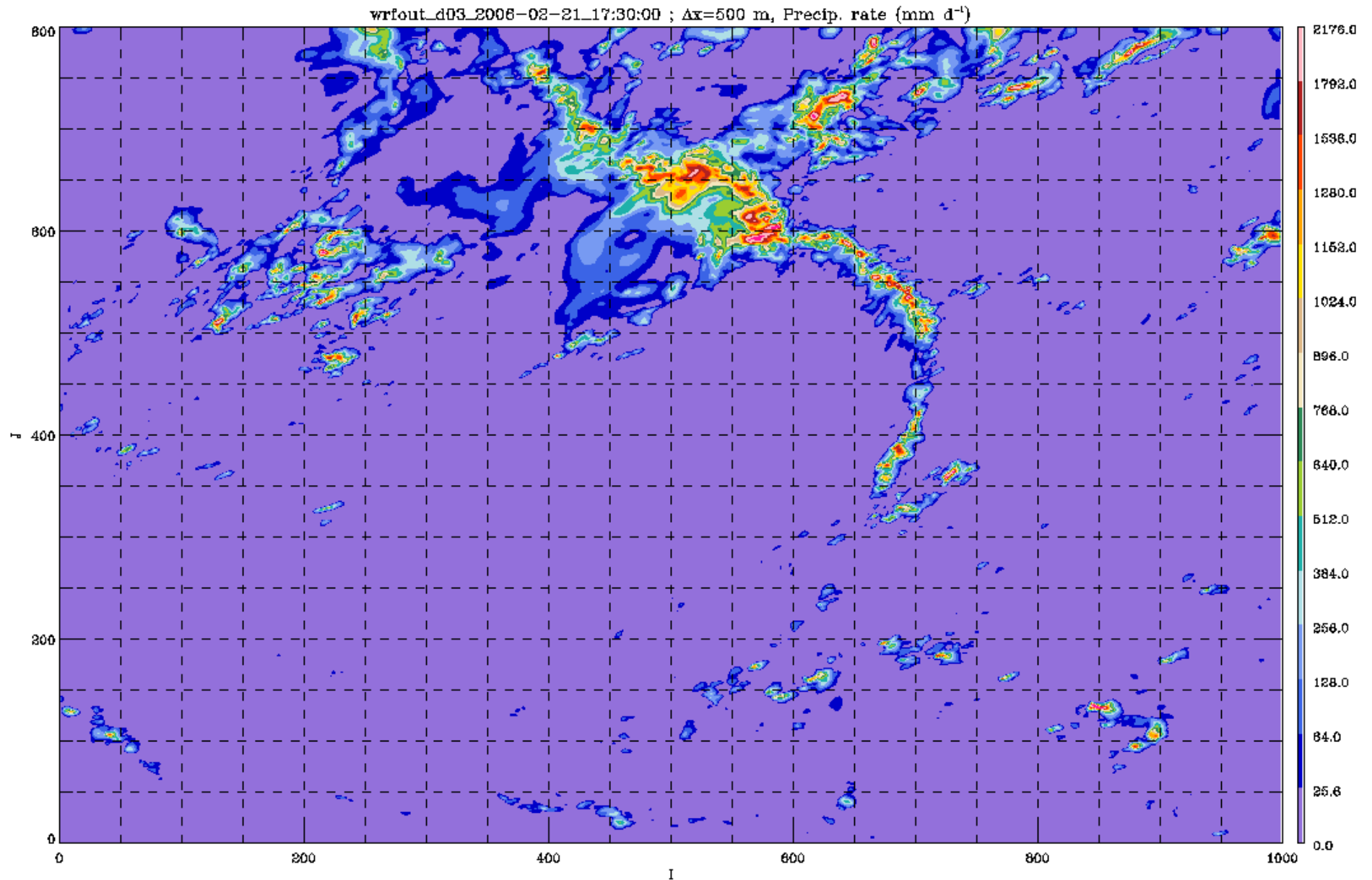
- deep convective precipitation within the eye wall
- Banded structure with embedded convection
- Realistic eye diameter, clear of precipitation

Effects of Condensate Loading

**Assessed using 0.5x0.5 km non-hydrostatic
WRF simulation**

Tropical ocean convective case (TOGA domain Feb 2006)

15-min average precipitation rate (*Hong and Lim 2006 microphysics*)



Dashed lines show 50x50 gp (25km x 25km) squares used to coarse grain WRF fields to produce “high-res AGCM” fields

Hydrostatic Balance w/ and w/out condensate loading

$$\pi_{hyd} = \int_z^{z_{top}} \frac{g}{c_p \Theta_{\{v,cond\}}} dz' + \pi_{top}$$

$$p_{hyd} = p_{00} \pi_{hyd}^{1/\kappa}$$

w/out loading:

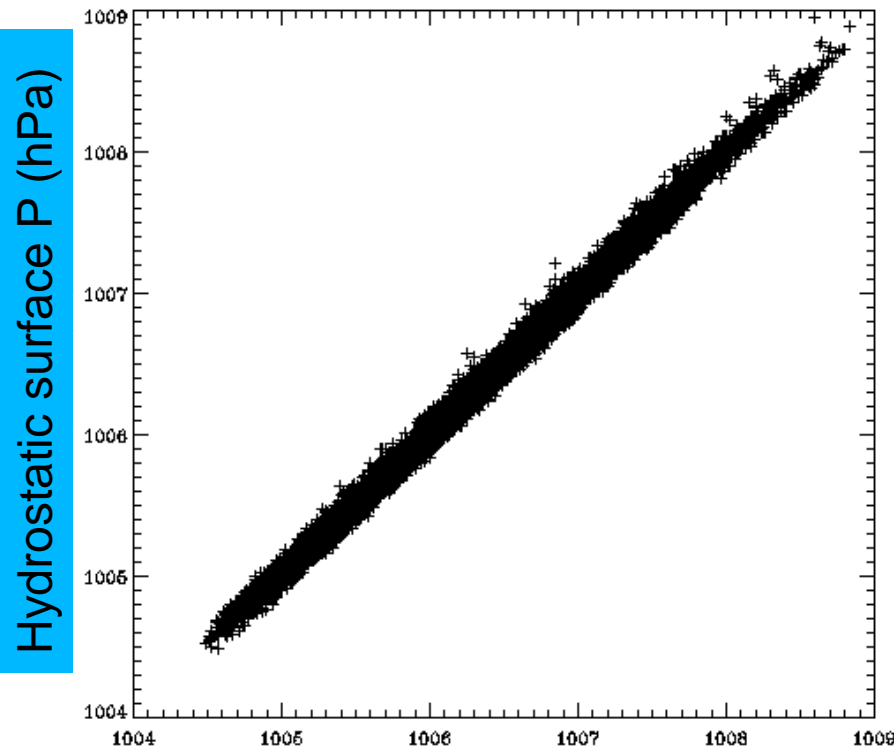
$$\Theta_v = \Theta(1. + 0.61q)$$

with loading:

$$\Theta_{cond} = \Theta(1. + 0.61q - q_{liq} - q_{ice} - q_{rain} - q_{graup} - q_{snow})$$

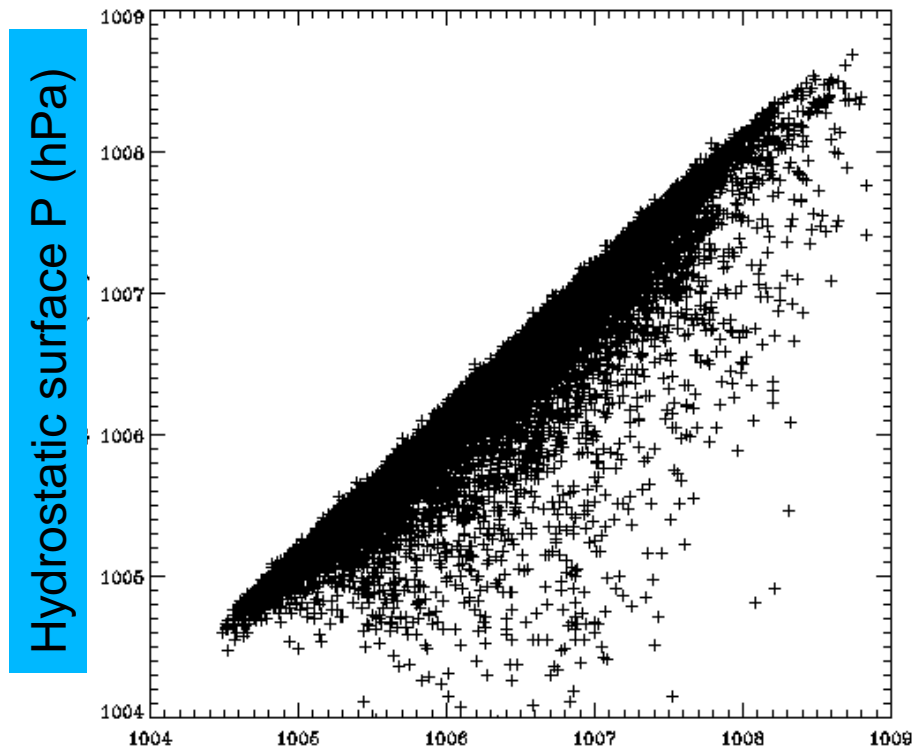
Coarse-grained to $(25 \text{ km})^2$

with loading:



Coarsened WRF surface P (hPa)

w/out loading:

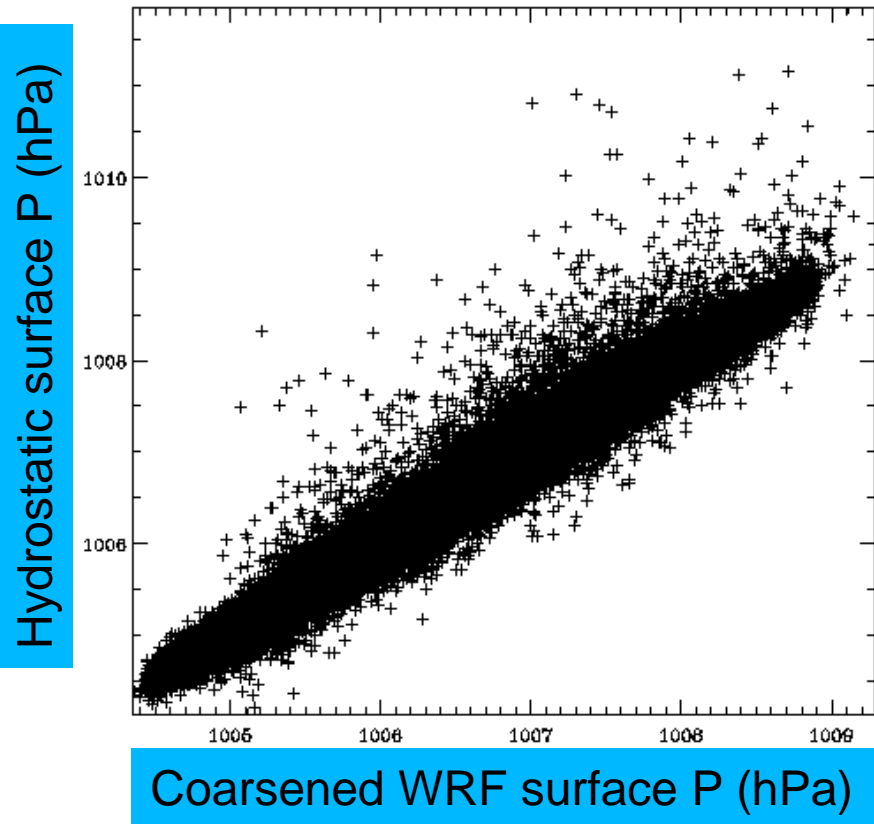


Coarsened WRF surface P (hPa)

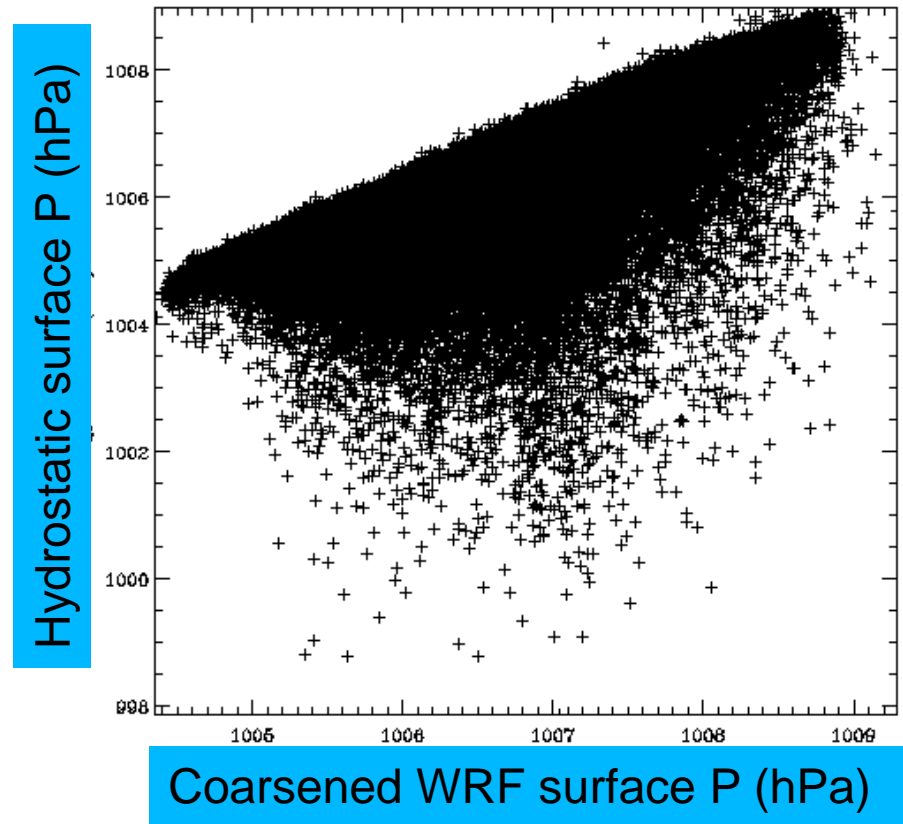
Condensate loading matters – even in $(25 \text{ km})^2$ grid boxes

Coarse-grained to $(5 \text{ km})^2$

with loading:



w/out loading:



Non-hydrostatic effects become detectable

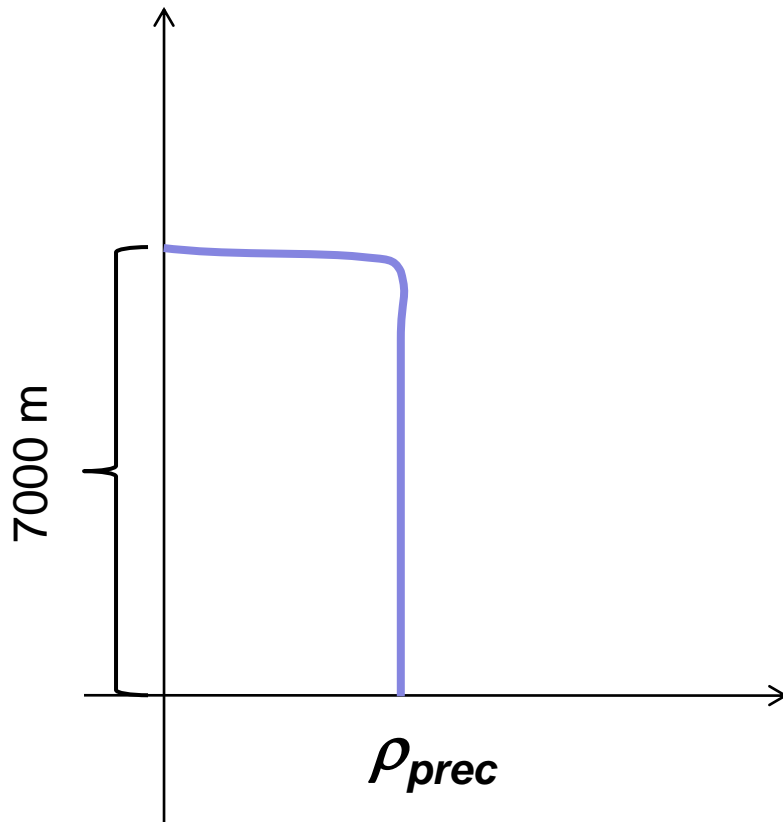
Parameterized precipitation loading

surface precip rate \mathcal{R}_{surf} used to
diagnose precipitating condensate
density ρ_{prec}

for $z < 7000m$

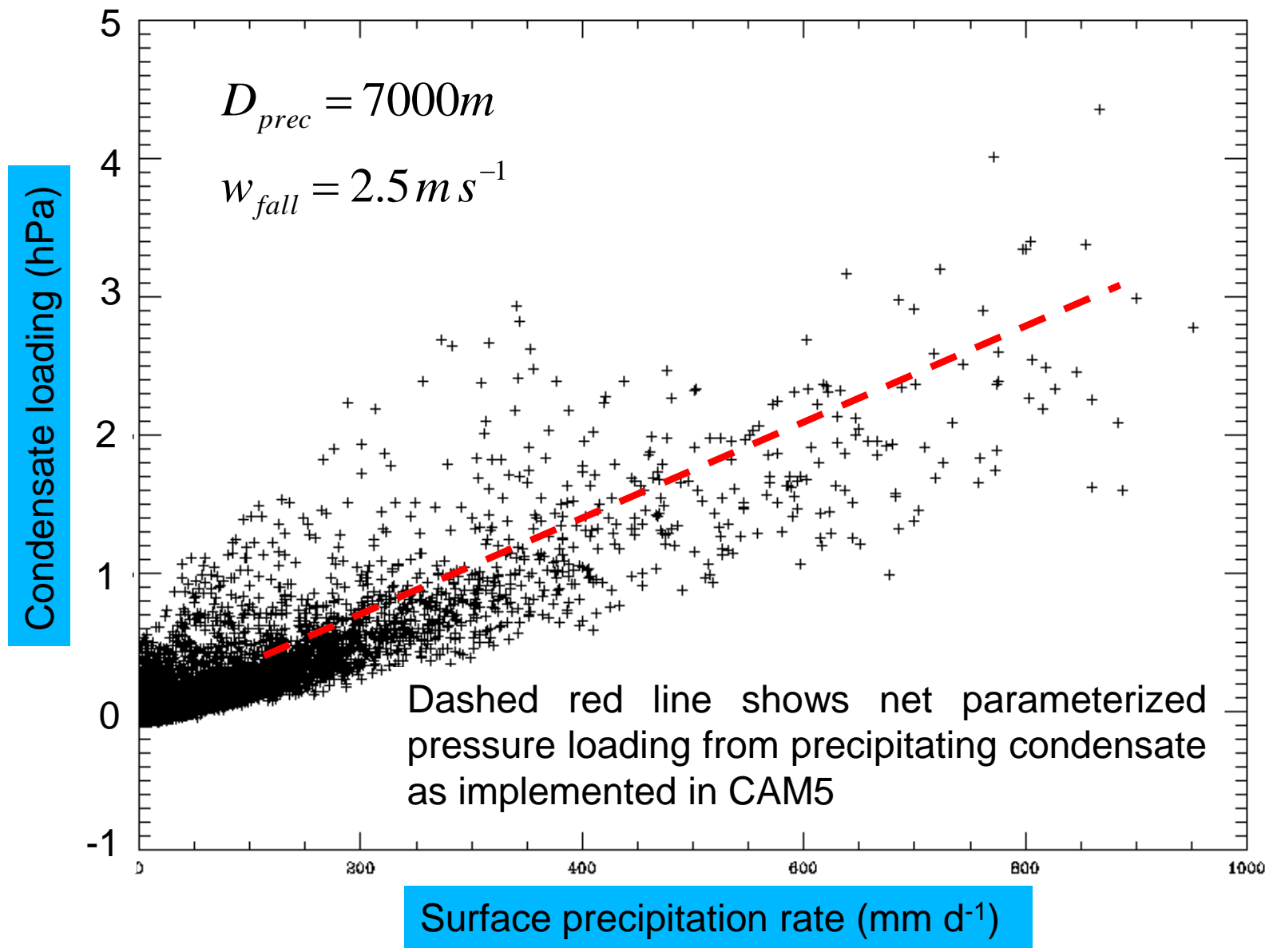
$$\rho_{prec}(x, y, z, t) = \frac{\mathcal{R}_{surf}(x, y, t)}{w_{fall}}$$

$$p_{prec}(x, y, z, t) = \int_z^{7000m} g \rho_{prec} dz'$$

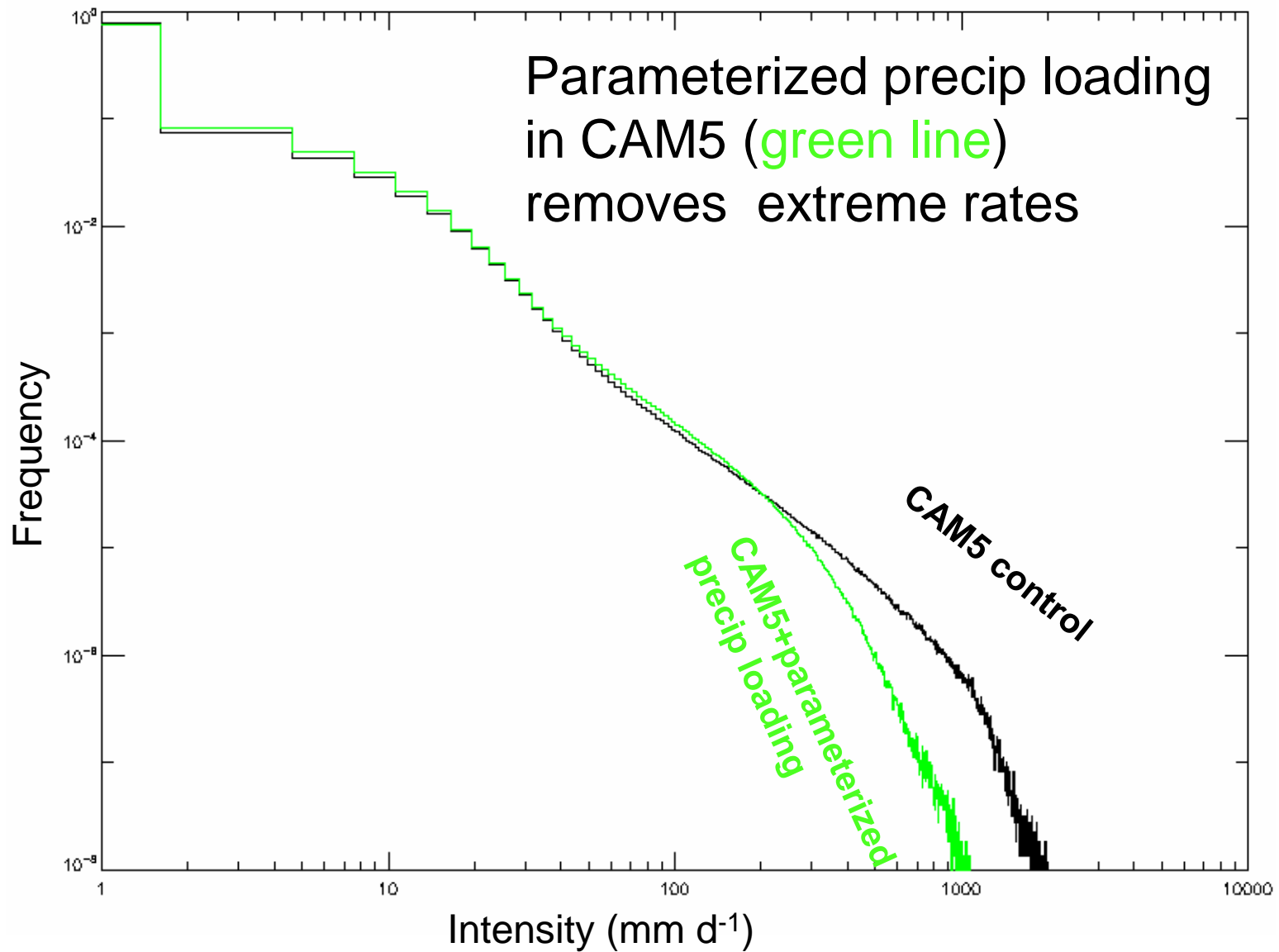


Extra condensate pressure is added to “real”
model pressure right before horizontal
gradients are calculated, then removed

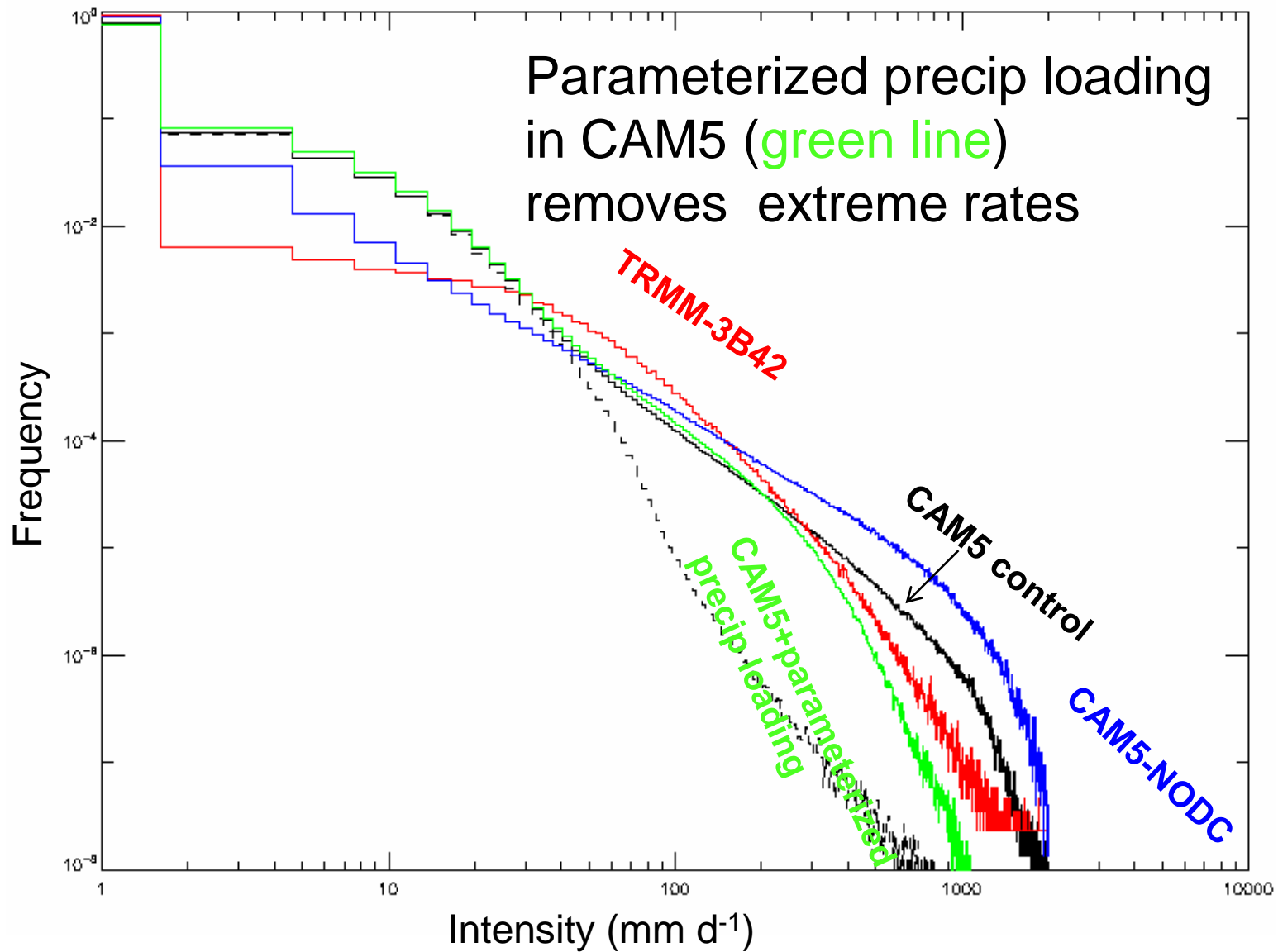
Net loading at surface (in Pa) as a function of surface precipitation rate – for $(25\text{km})^2$ resolution



PDFs of tropical precipitation (30S-30N) rates Aug 2005

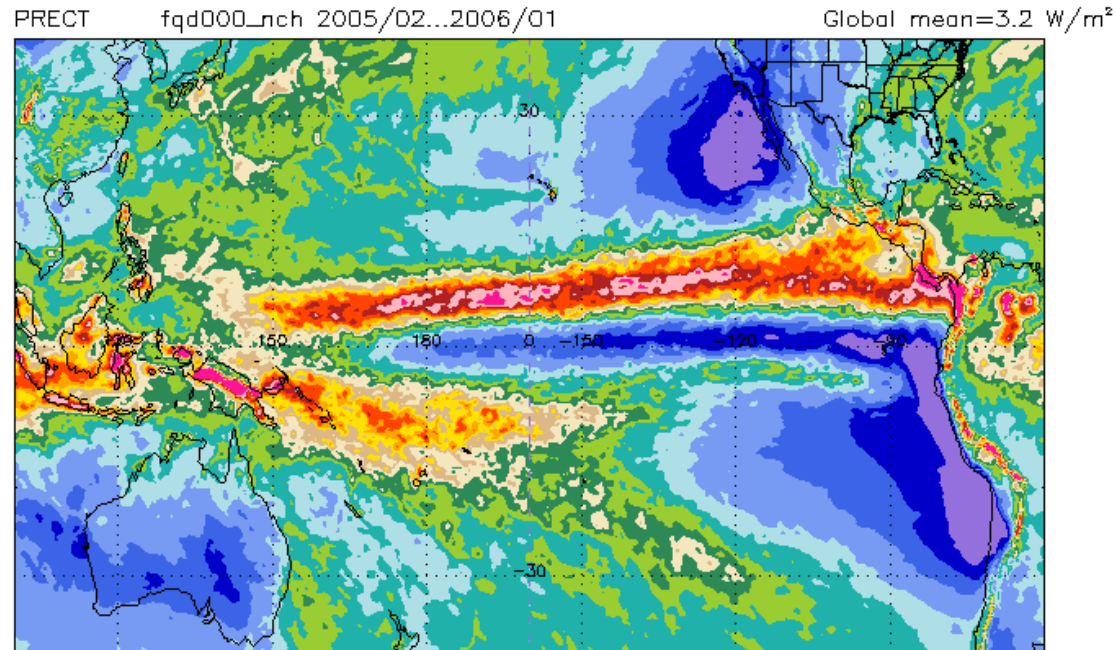


PDFs of tropical precipitation (30S-30N) rates Aug 2005

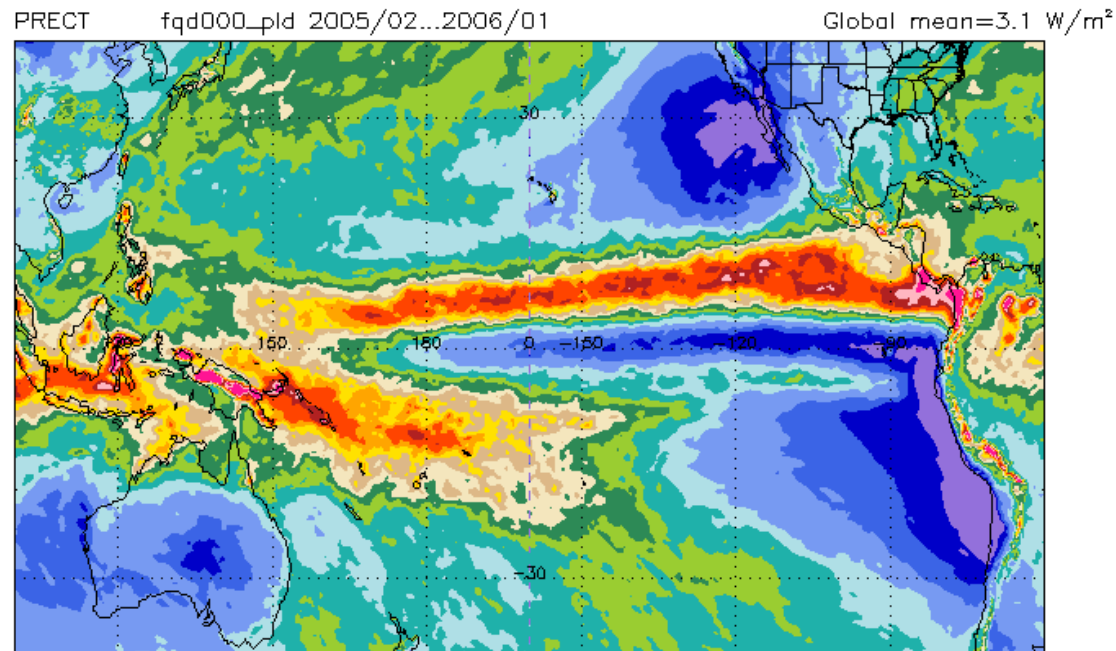


Annual mean precipitation

CAM5 control



w/ parameterized
precipitation loading

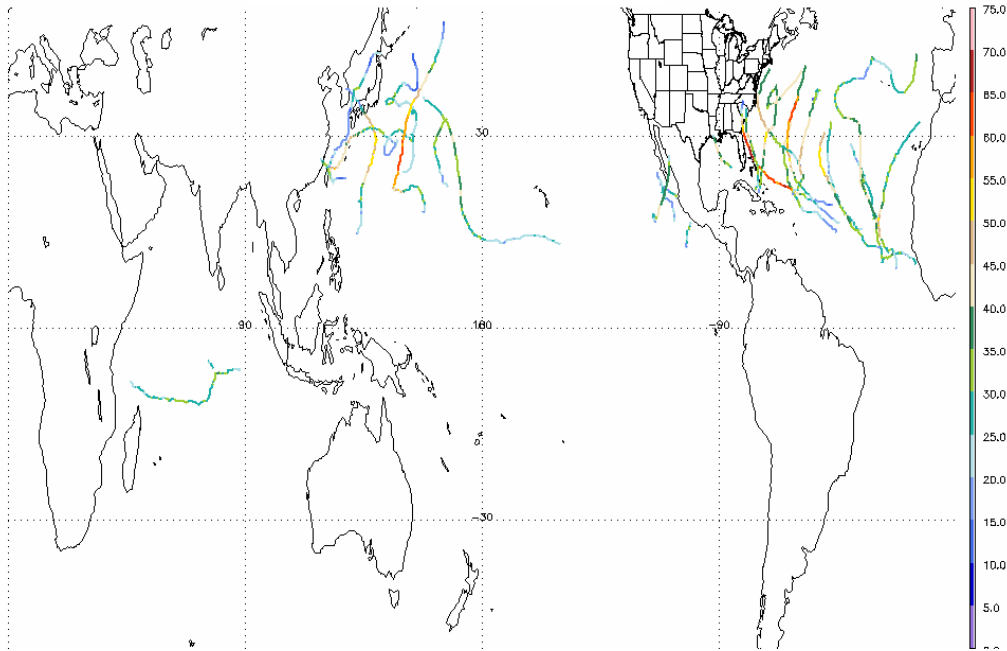


*Bad news: TC number
also decreases*

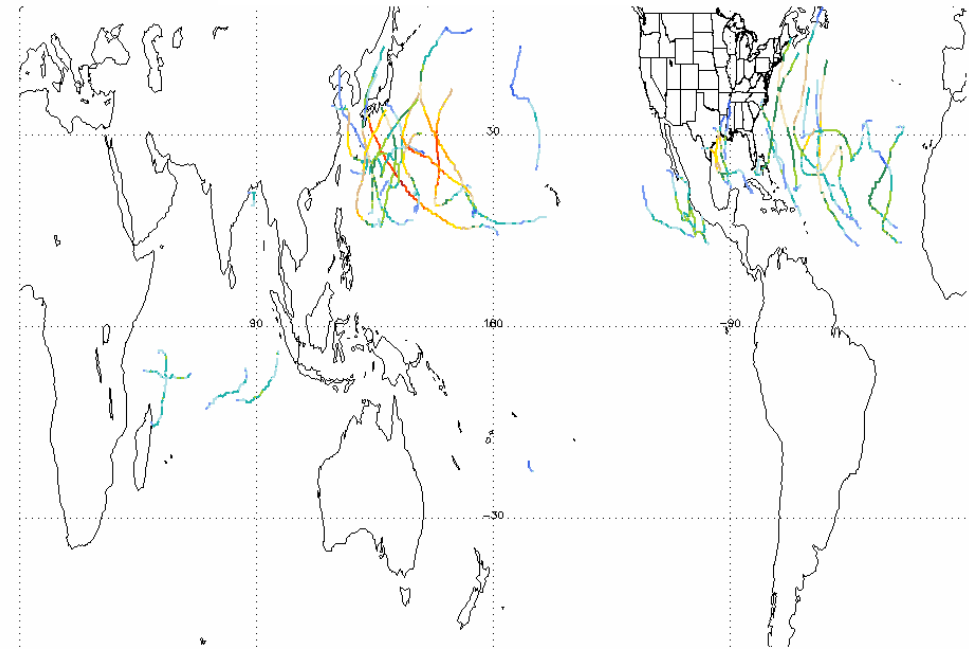
Storms with $U > 33 \text{ ms}^{-1}$: June 1 to Nov 1 2005

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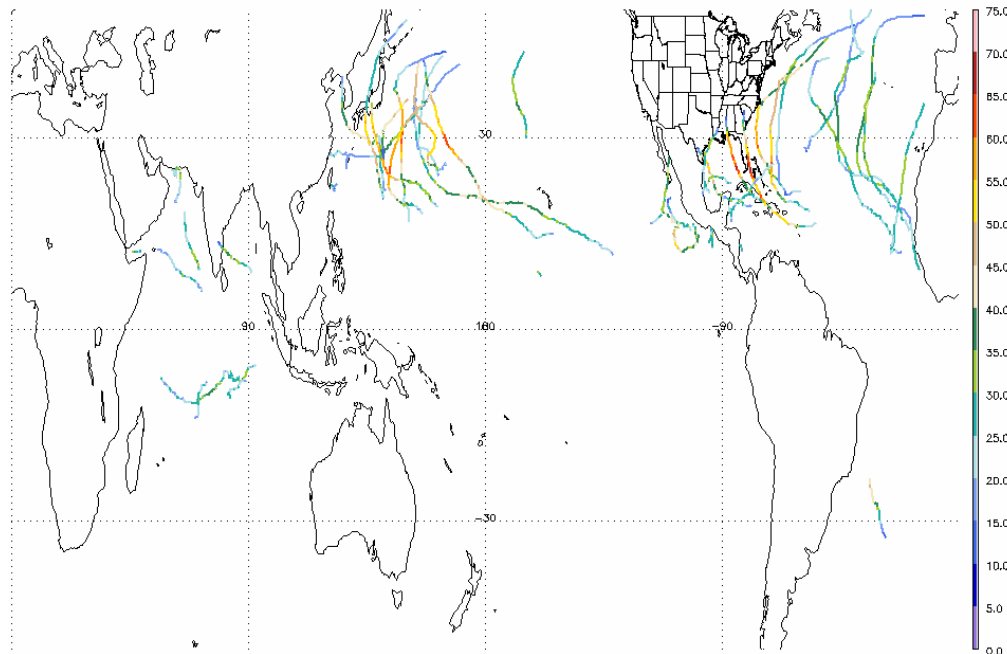
Std CAM5 #1



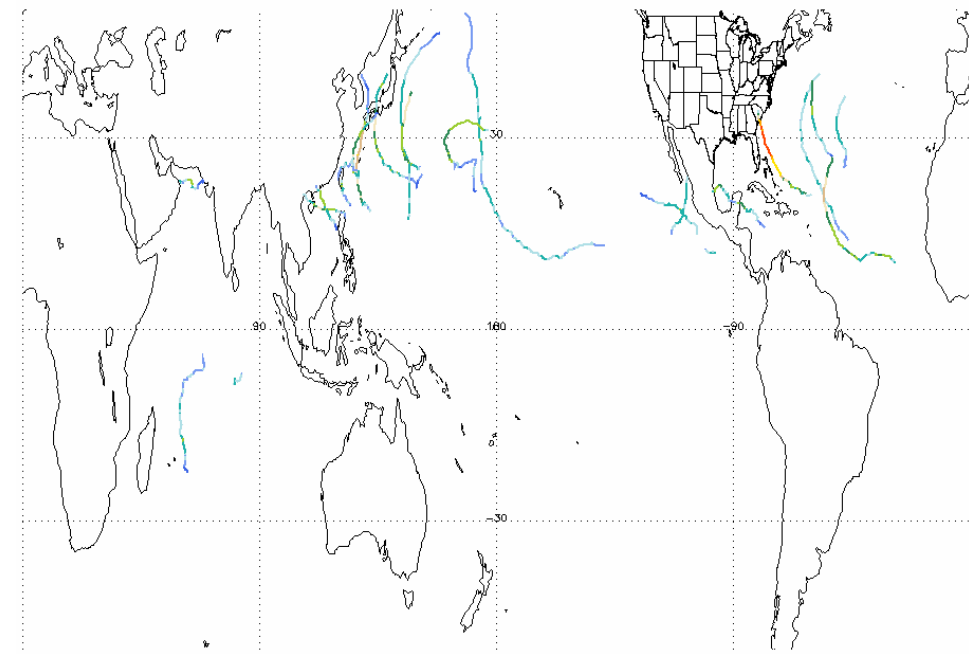
Std CAM5 #2



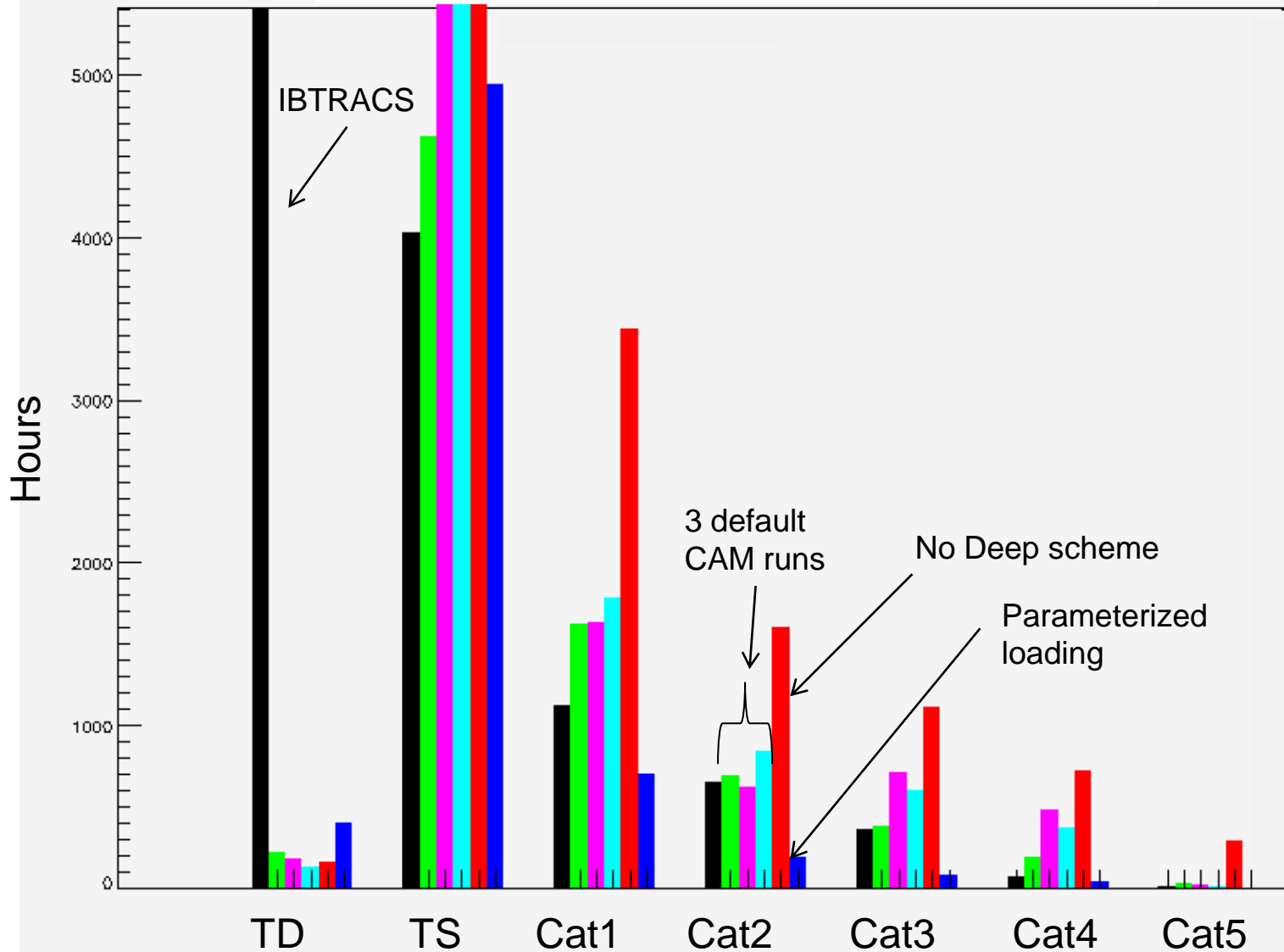
Std CAM5 #3



w/ parameterized precip loading



Time (hours) at Category



Precipitation Loading in perspective

Relatively small perturbation to pressure field in precipitating regions seems to have large effect.

Same pressure increase could be achieved by re-evaporating about 1/8-th of condensate column

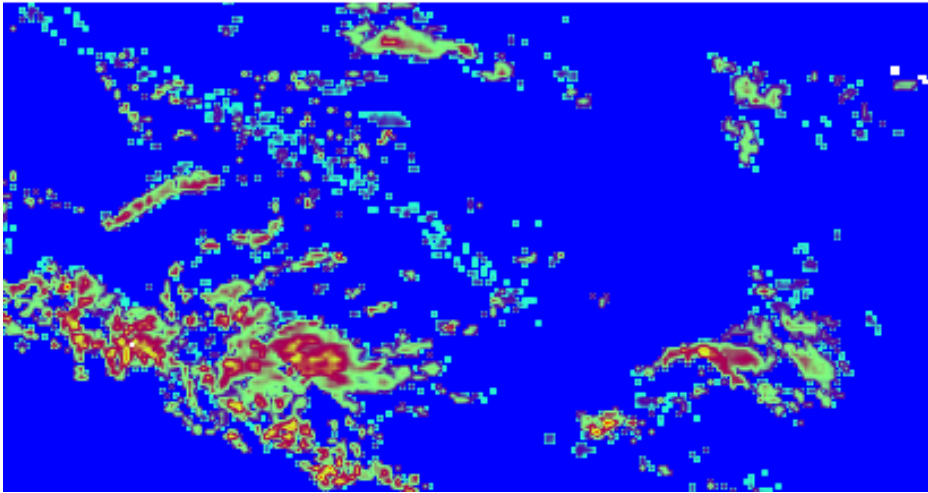
Analysis using Precipitation Objects

with

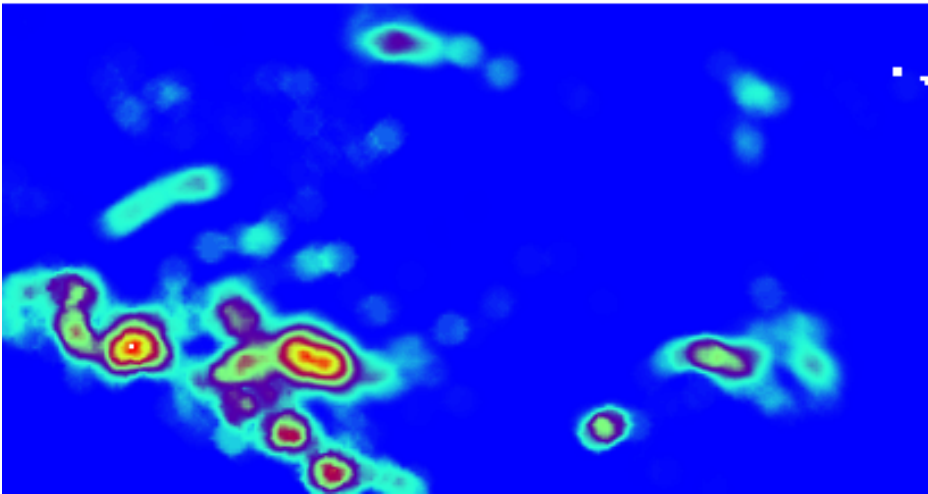
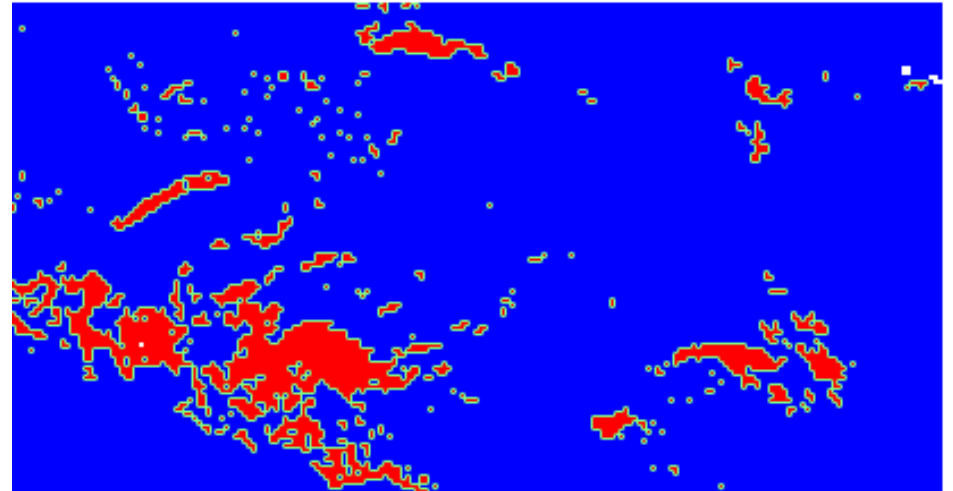
Gregor Skok,
University of Ljubljana

Joe Tribbia
NCAR AMP/CGD

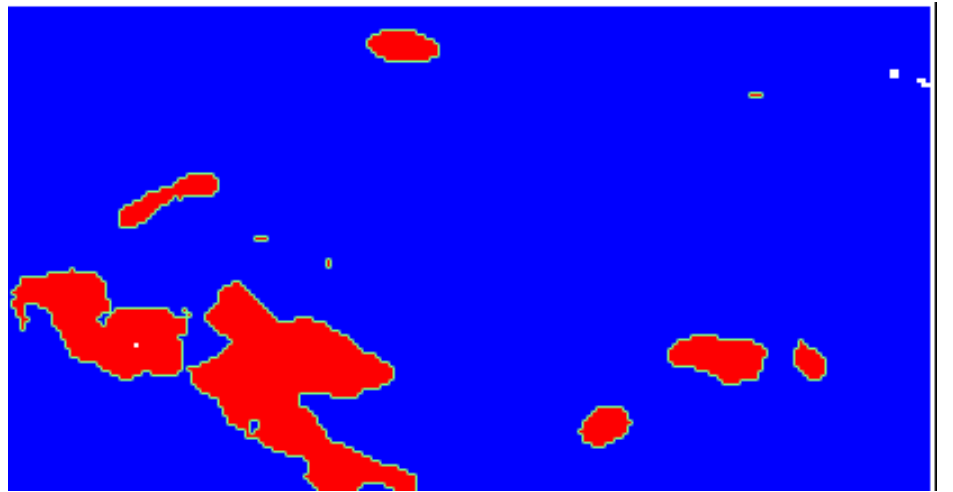
Raw precipitation (snapshot)



Thresholding only

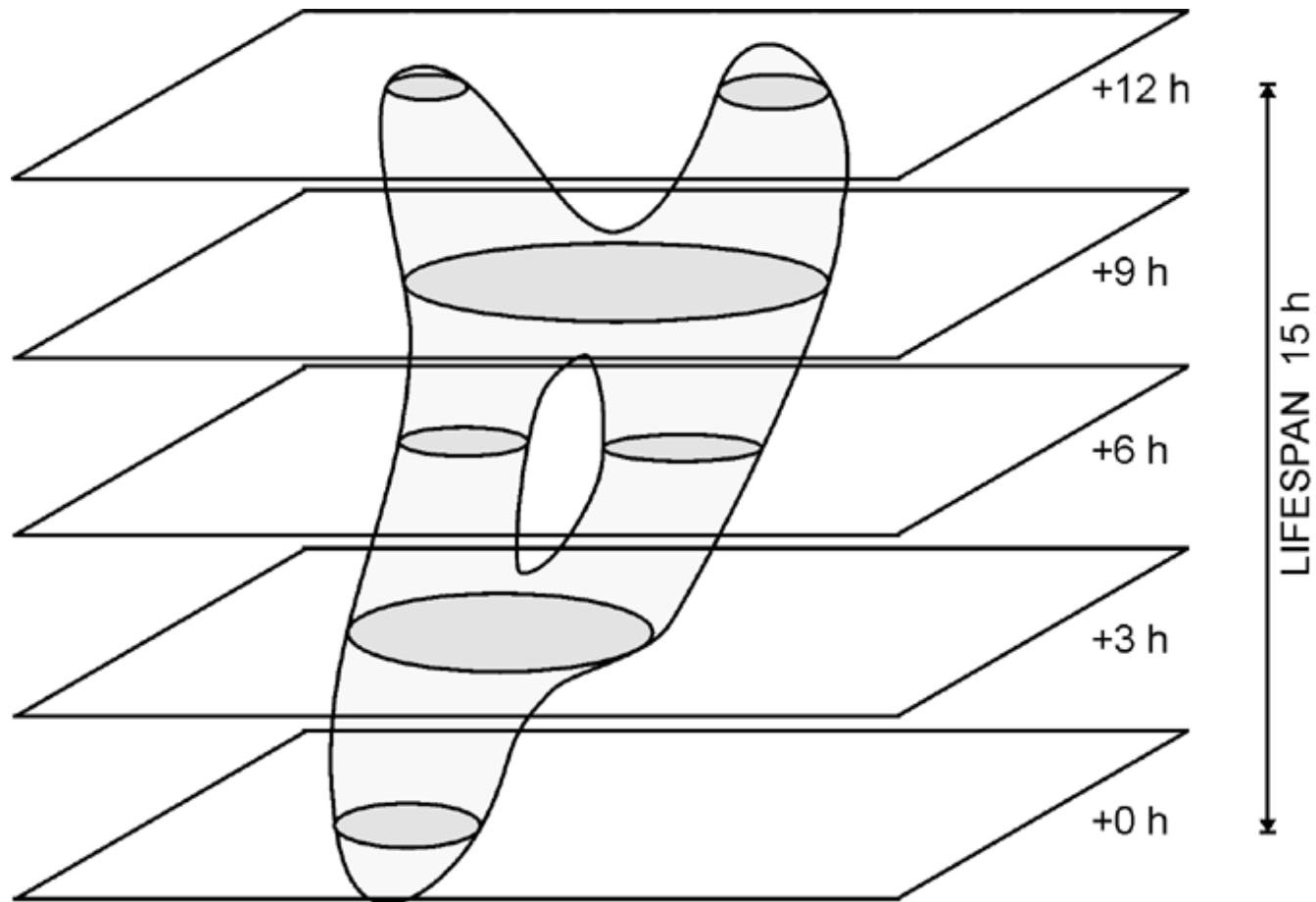


Smoothing with convolution



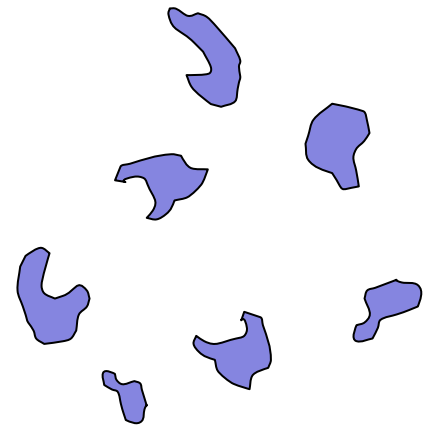
Thresholding after convolution

Objects are tracked in time using overlaps. Could be modified to include search radius.

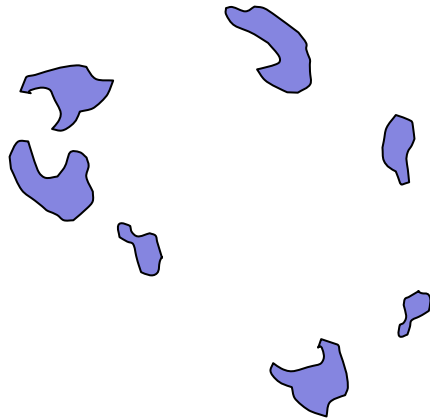


Hope to see transitions from clusters of convection to tropical cyclones (eventually). Analysis just begun.

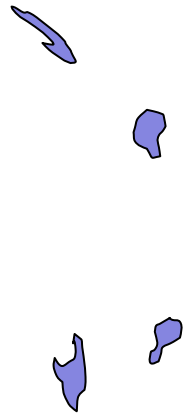
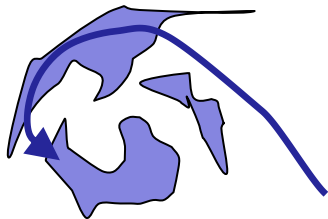
Time 0



Time 1

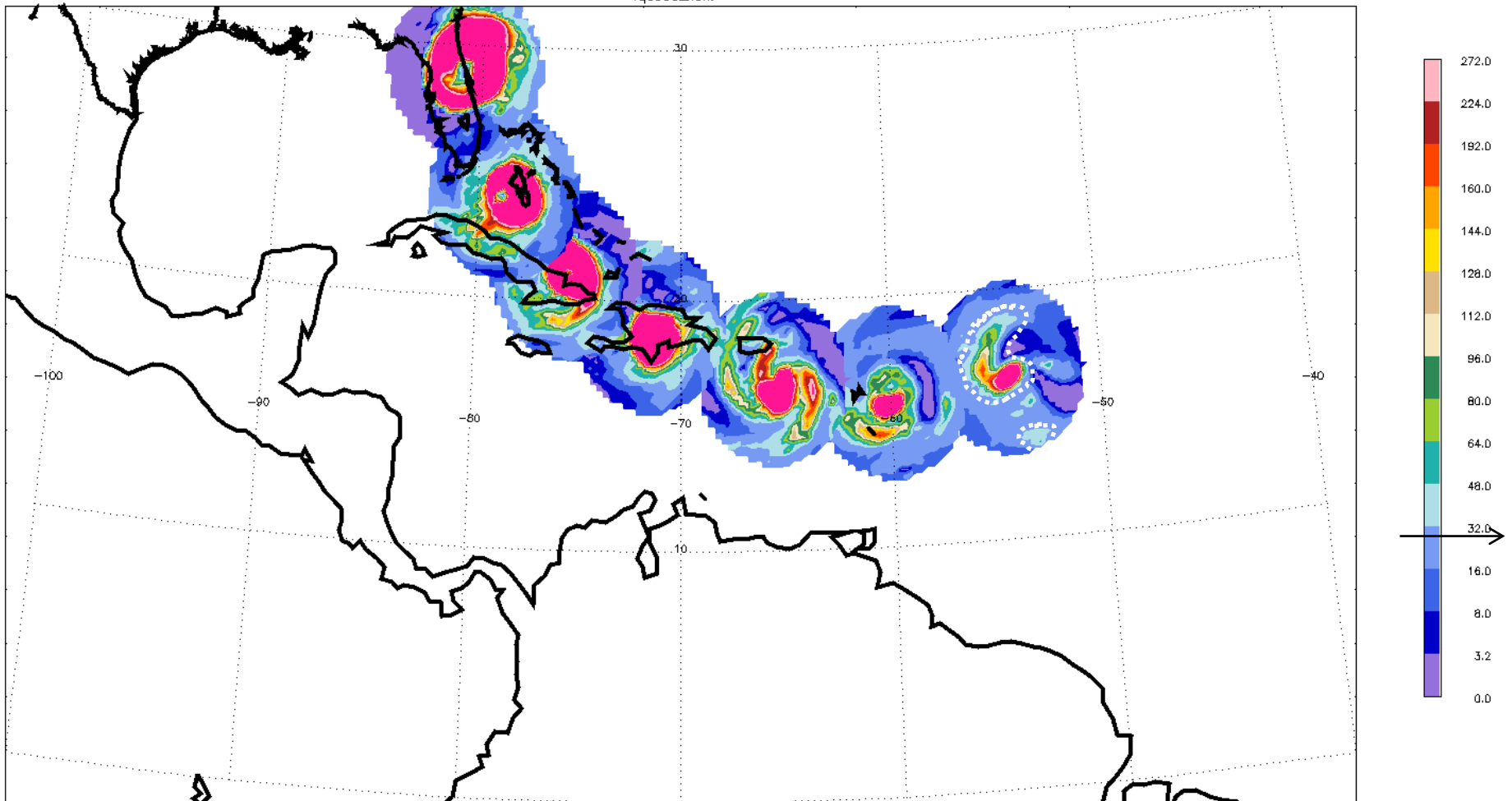


Time 2



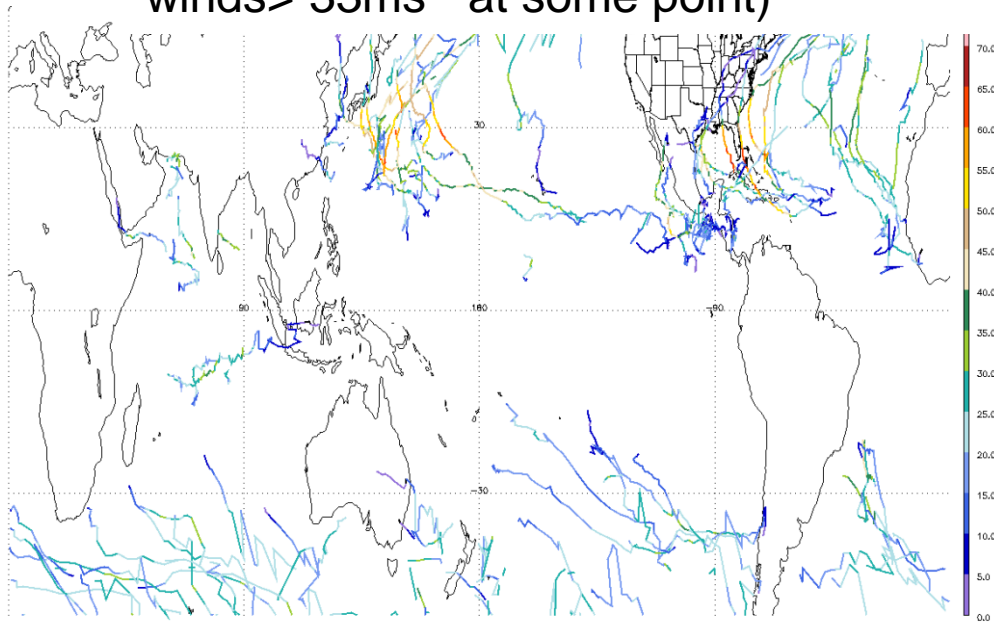
2005_08_28_54000 Local time=10:20 GMT=15:00

fqd000_nch:

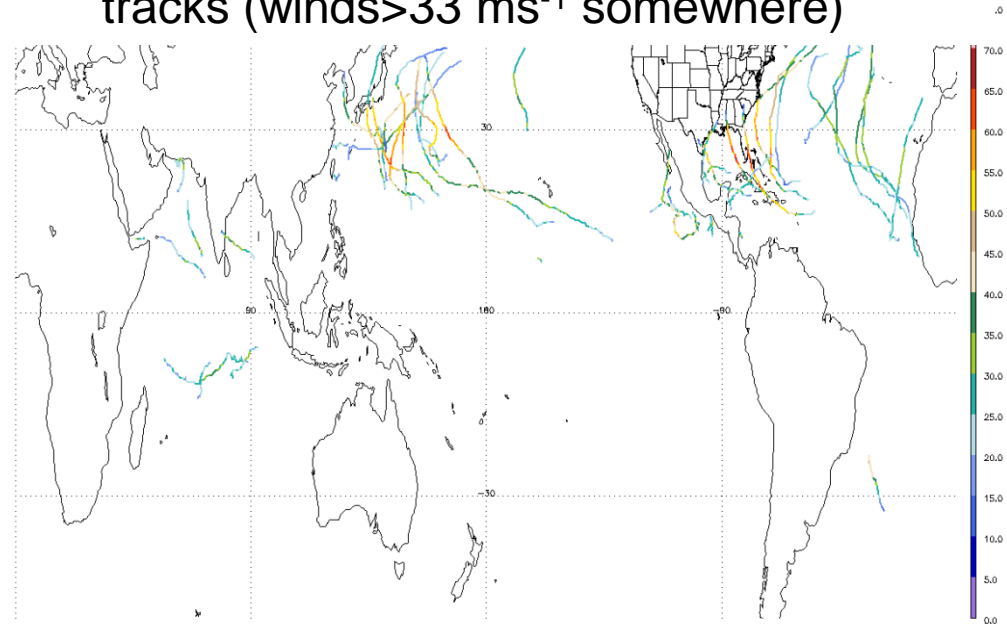


“Feature” tracks June 1 to Nov 1 2005

Precipitation object tracks (that contained winds $> 33\text{ms}^{-1}$ at some point)



Surface pressure based hurricane/TC tracks (winds $> 33\text{ms}^{-1}$ somewhere)

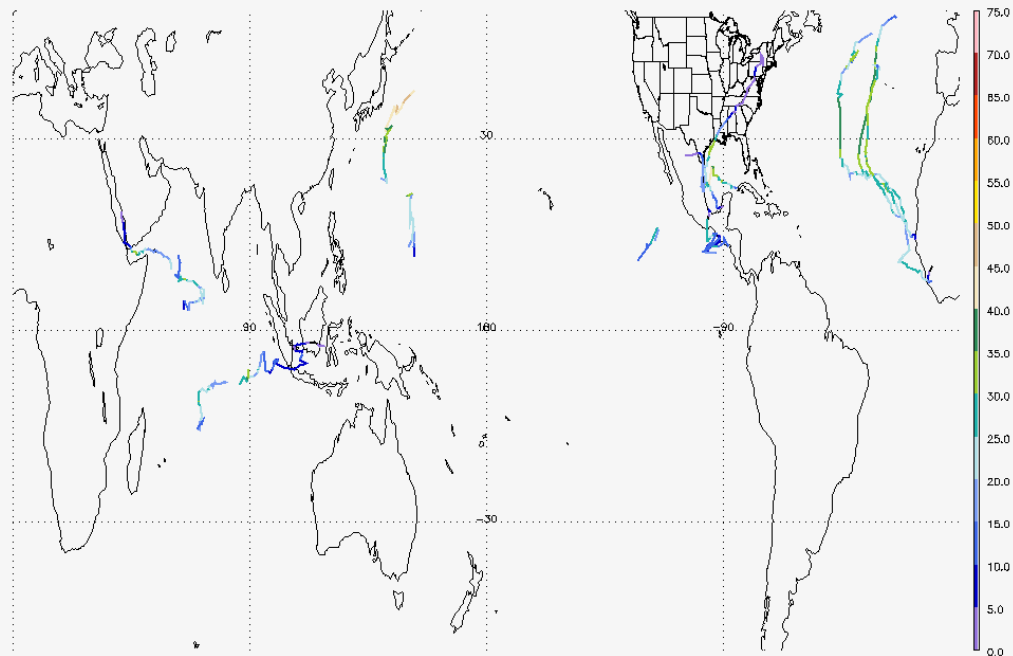
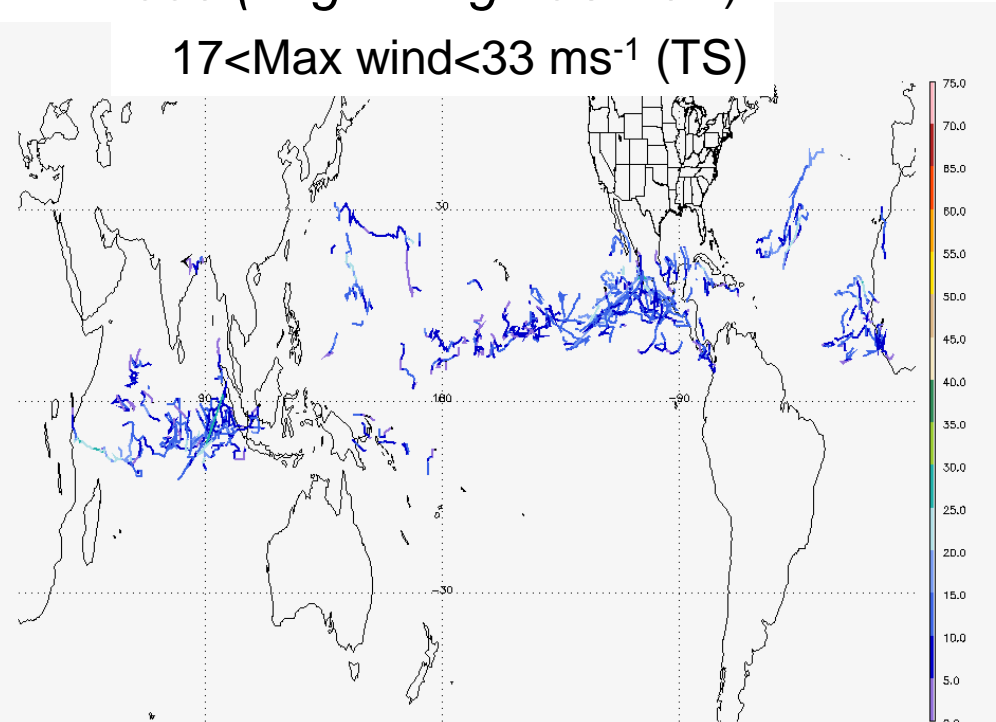
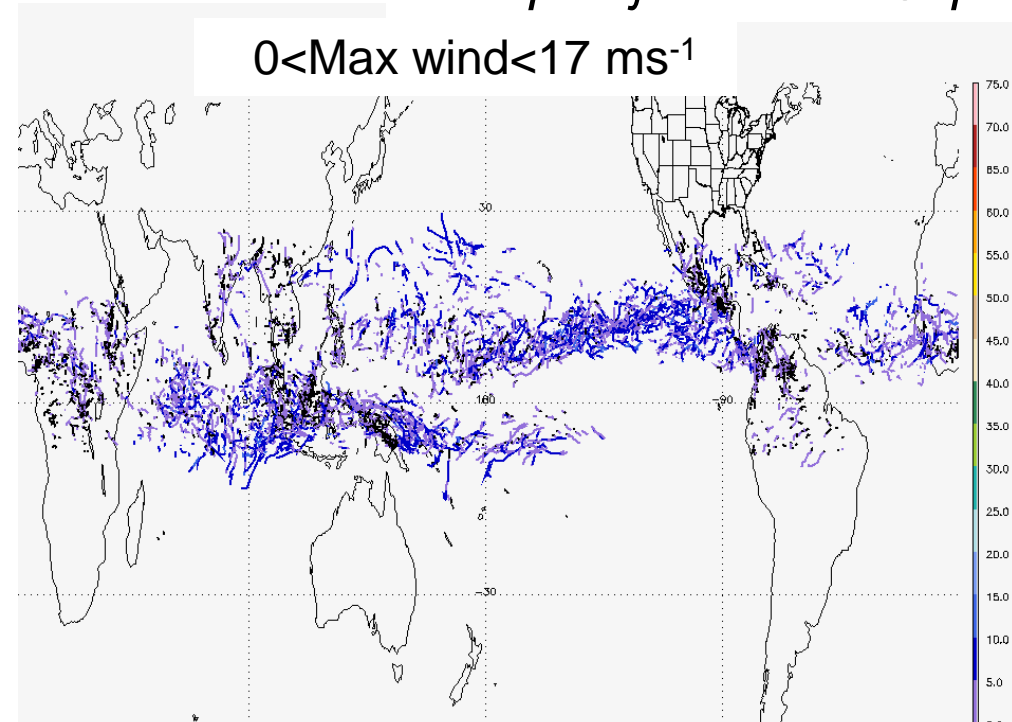


Precip object tracks September 2005 (originating 10S-25N)

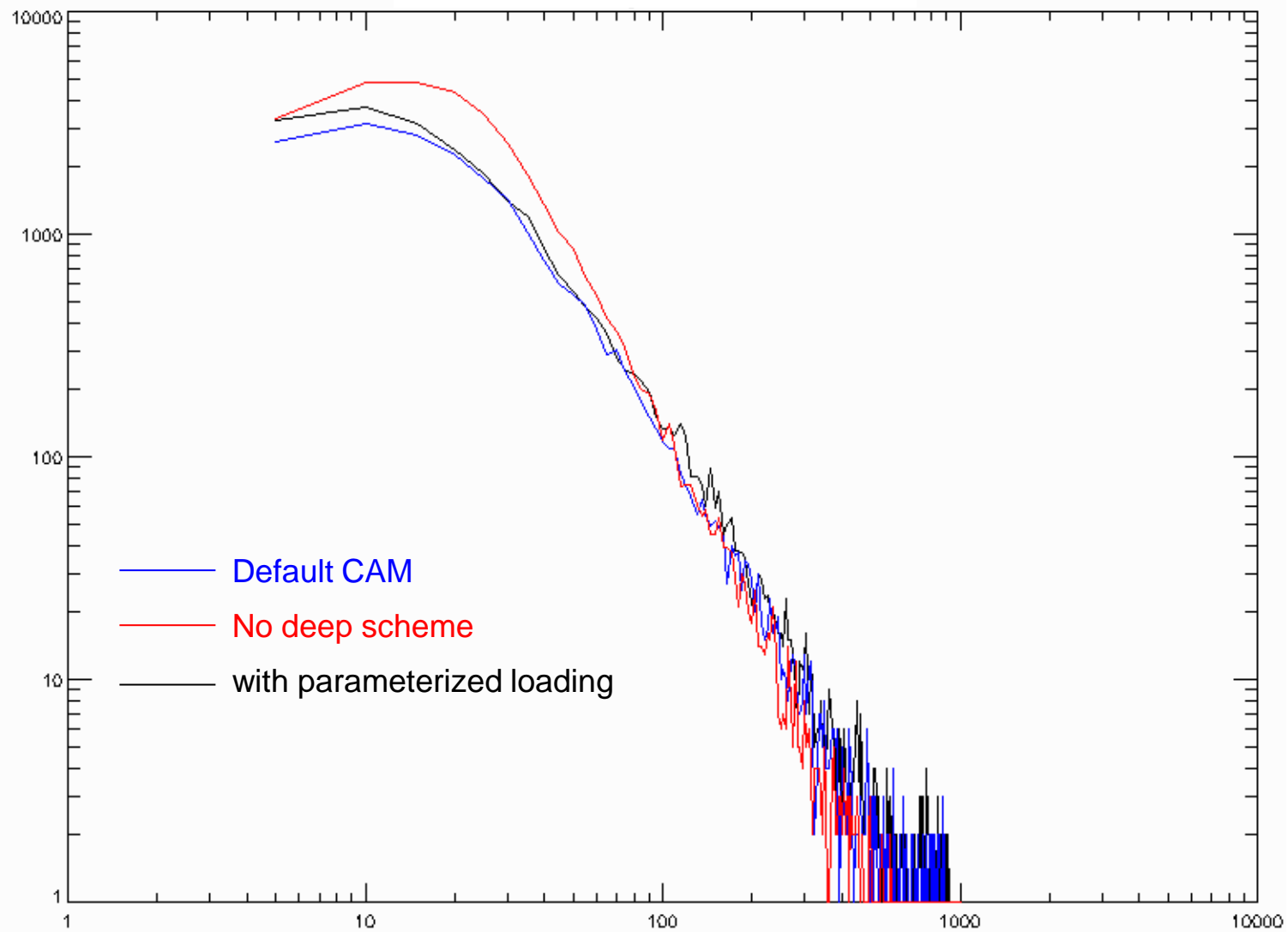
0 < Max wind < 17 ms⁻¹

17 < Max wind < 33 ms⁻¹ (TS)

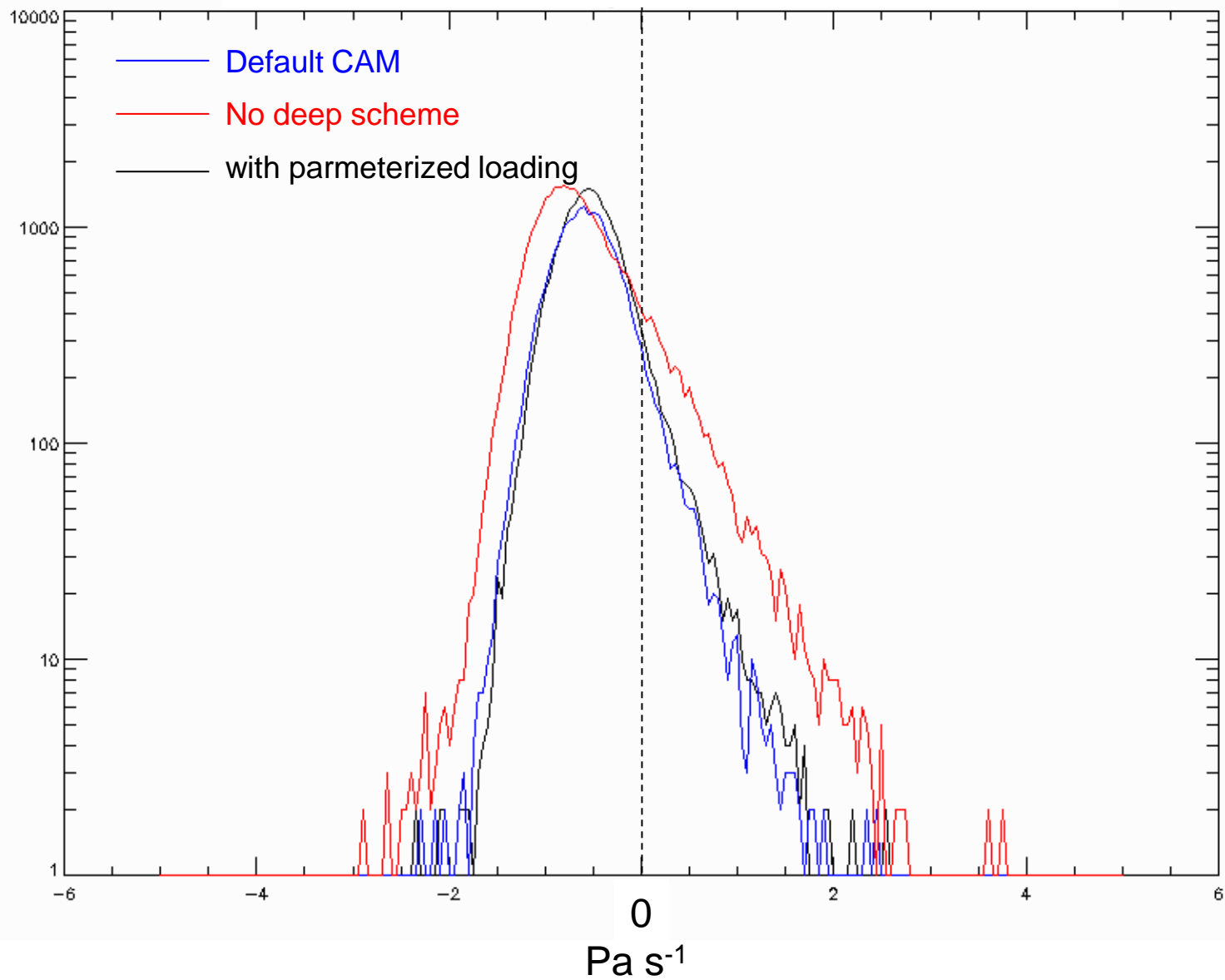
33ms⁻¹ < Max wind (TC/hurricane)



PDFs of object area



PDFs of object mean vertical motion



Summary and future directions

Models with $\Delta x \sim 10\text{km}$ **can** capture many important aspects of TC climatology ***but answer depends sensitively on physics tuning***

Time to focus on cyclogenesis processes in ~ 10 km models

Tracks and intensities may have similar biases in current models, e.g., eastward shift in tracks, not enough variability in intensity

Convection and Clouds in the Tropics

2009-Aug-20 21z

72-hour Forecasts

East Pacific
ITCZ

Atlantic
Hurricane Bill

African
Deep
Convective
Clusters

Indian
Monsoon

Pacific Warm
Pool





THANK YOU

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**To advance understanding of weather, climate, atmospheric composition and processes;
To provide facility support to the wider community; and,
To apply the results to benefit society.**

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