

Gravity Waves in Shear and Implications for Organized Convection

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Workshop on Modelling Monsoon Intraseasonal Variability

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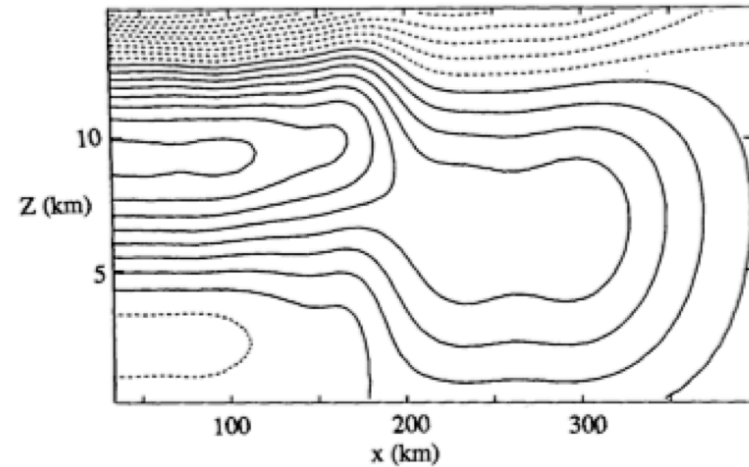
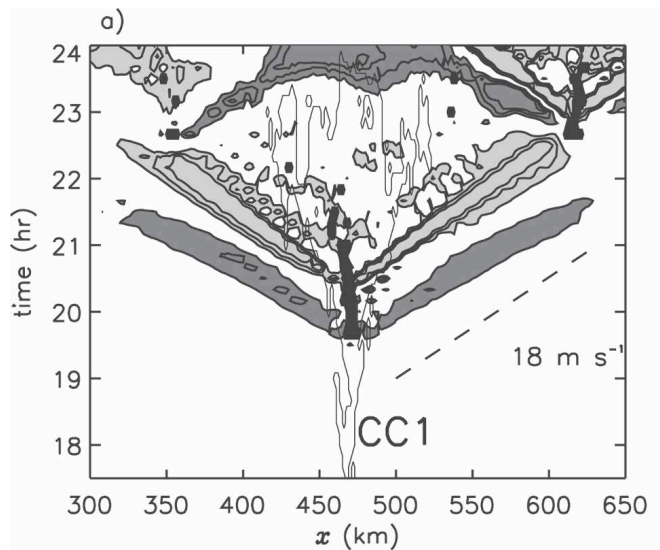
INTRODUCTION

Abstract

- Gravity waves can trigger/favor the formation of new convection
- If new convection forms repeatedly on a preferred side of preexisting convection, then a *convectively coupled wave* is formed
- **What creates a preferred side? Hypothesis: wind shear**
- Design a simple model for interactions of gravity waves and wind shear
- Results:
 - Wind shear can create a preferred side
 - Jet shears create the greatest difference in favorability between two sides
 - *Predictions of preferred propagation direction of convectively coupled waves in a given background wind shear*
- Other application:
 - Formation of new cells within an individual mesoscale convective system
 - *Resonance* renders upstream more favorable than downstream

Gravity waves and organized convection

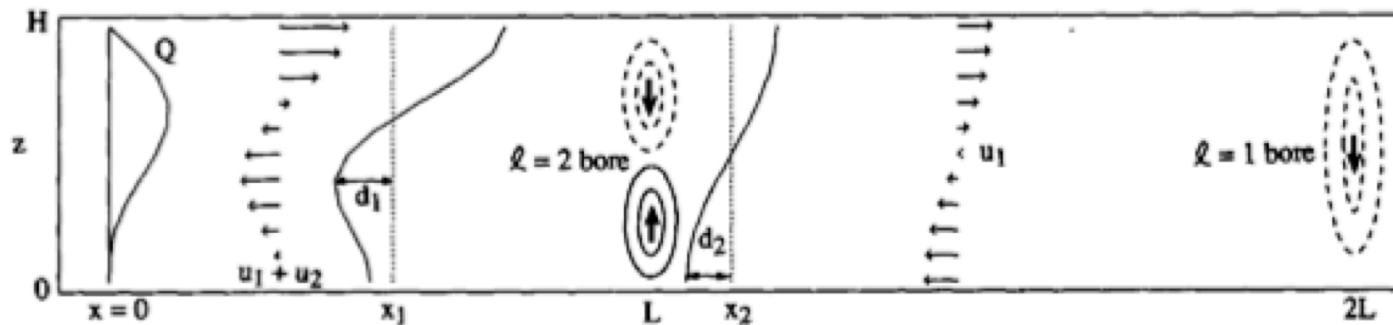
Buoyancy anomalies excited by top-heavy heating



Convection can excite gravity waves

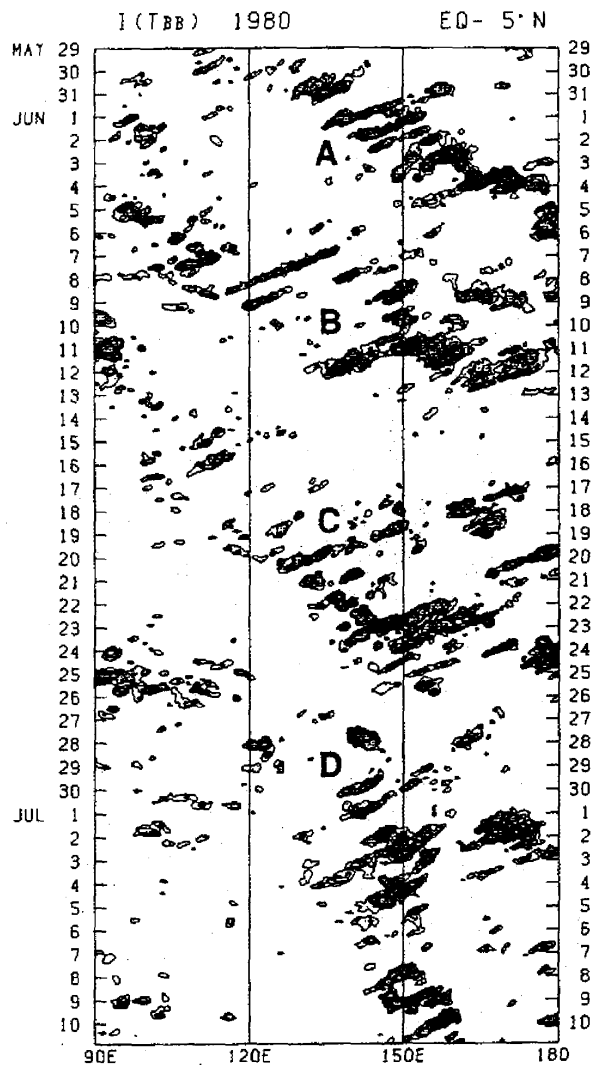
Gravity waves can excite new convection

Theory: the role of (1) deep convection and (2) stratiform heating



from Tulich and Mapes (2008) and Mapes (1993)

Convectively coupled waves: Envelopes of mesoscale convective systems



- Embedded cloud systems propagate in opposite direction of wave envelope
 - New cloud systems tend to form on a preferred side of preexisting cloud systems
-
- What causes wave trains to form preferentially (rather than scattered convection)?
 - What determines the preferred propagation direction of the convectively coupled wave?
-
- Hypothesis: interactions of gravity waves with wind shear

from Nakazawa (1988)

SIMPLIFIED MODEL

Starting point: Hydrostatic Boussinesq equations

$$\frac{\partial U}{\partial t} + U \frac{\partial U}{\partial x} + W \frac{\partial U}{\partial z} + \frac{\partial P}{\partial x} = 0$$

$$\frac{\partial P}{\partial z} = g \frac{\Theta}{\theta_{ref}}$$

$$\frac{\partial \Theta}{\partial t} + U \frac{\partial \Theta}{\partial x} + W \frac{\partial \Theta}{\partial z} + W \frac{d\theta_{bg}}{dz} = 0$$

$$\frac{\partial U}{\partial x} + \frac{\partial W}{\partial z} = 0$$

U = horizontal velocity

P = pressure

W = vertical velocity

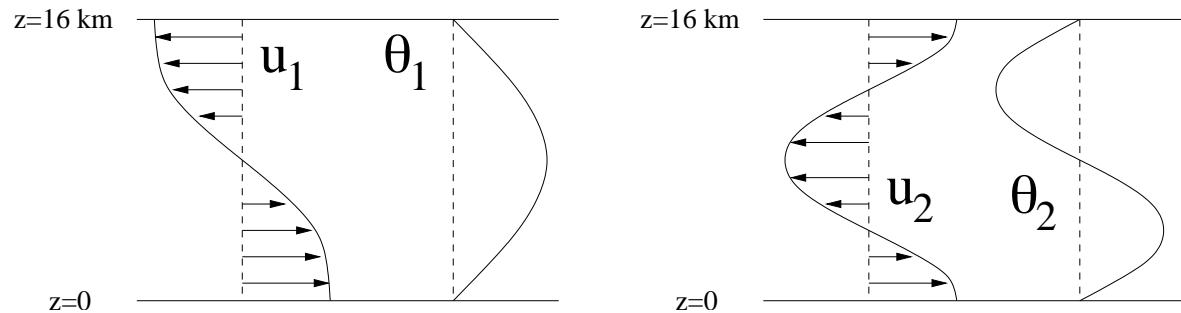
Θ = temperature

Gravity waves in the tropical atmosphere

Linear waves:

- Independent vertical modes: $U(x, z, t) = \sum_j u_j(x, t) \cos jz$, etc.
- Shallow water system for each vertical mode j :

$$\frac{\partial u_j}{\partial t} - \frac{\partial \theta_j}{\partial x} = 0$$
$$\frac{\partial \theta_j}{\partial t} - \frac{1}{j^2} \frac{\partial u_j}{\partial x} = 0$$



Nonlinear waves:

- Project **nonlinear** equations

$$\partial_t U + U \partial_x U + W \partial_z U + \partial_x P = 0$$

onto vertical modes

$$U(x, z, t) = u_1(x, t) \cos z + u_2(x, t) \cos 2z$$

- The result is ...

2-Mode Shallow Water Equations

$$\left\{ \begin{array}{l} \frac{\partial u_1}{\partial t} - \frac{\partial \theta_1}{\partial x} = -\frac{3}{\sqrt{2}} \left[u_2 \frac{\partial u_1}{\partial x} + \frac{1}{2} u_1 \frac{\partial u_2}{\partial x} \right] \\ \frac{\partial \theta_1}{\partial t} - \frac{\partial u_1}{\partial x} = -\frac{1}{\sqrt{2}} \left[2u_1 \frac{\partial \theta_2}{\partial x} + 4\theta_2 \frac{\partial u_1}{\partial x} - u_2 \frac{\partial \theta_1}{\partial x} - \frac{1}{2} \theta_1 \frac{\partial u_2}{\partial x} \right] \end{array} \right.$$

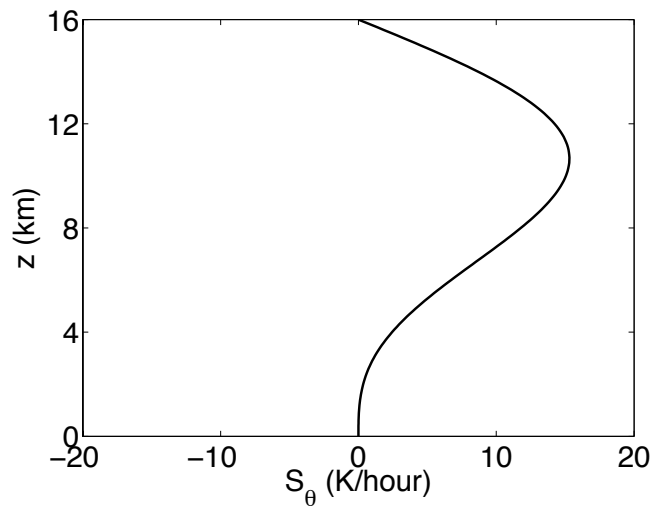
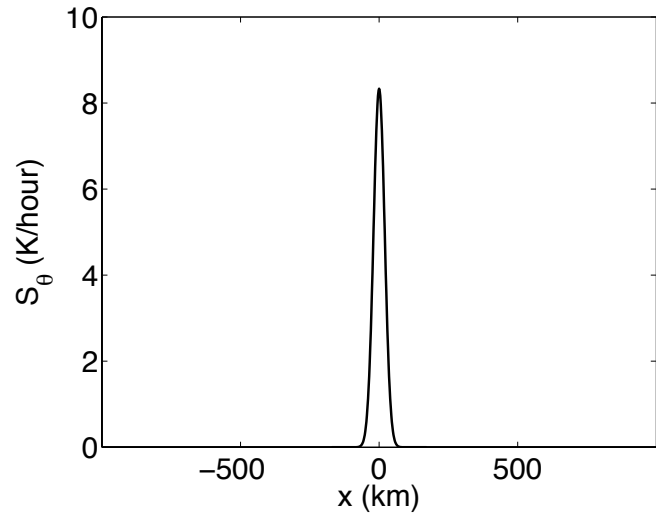
$$\left\{ \begin{array}{l} \frac{\partial u_2}{\partial t} - \frac{\partial \theta_2}{\partial x} = 0 \\ \frac{\partial \theta_2}{\partial t} - \frac{1}{4} \frac{\partial u_2}{\partial x} = -\frac{1}{2\sqrt{2}} \left[u_1 \frac{\partial \theta_1}{\partial x} - \theta_1 \frac{\partial u_1}{\partial x} \right] \end{array} \right.$$

- Nonlinear, hydrostatic internal gravity waves **with effect of background shear**

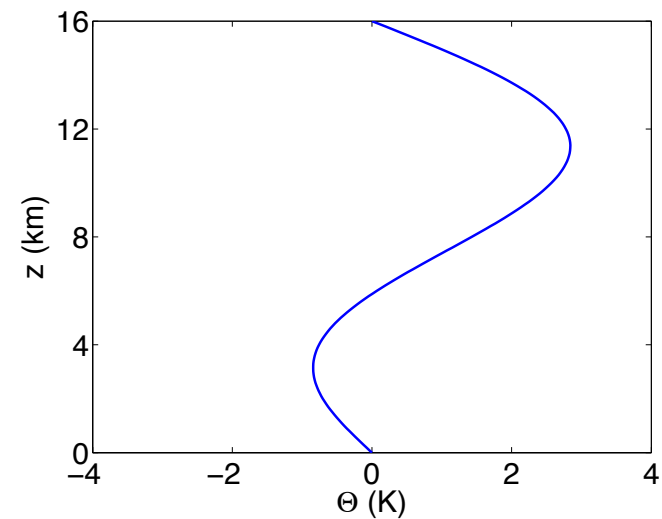
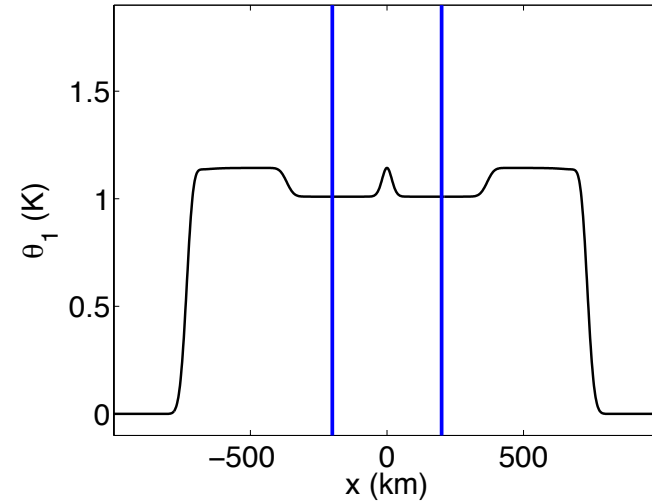
CONVECTIVELY COUPLED WAVES

Numerical experiment WITHOUT wind shear

Forcing

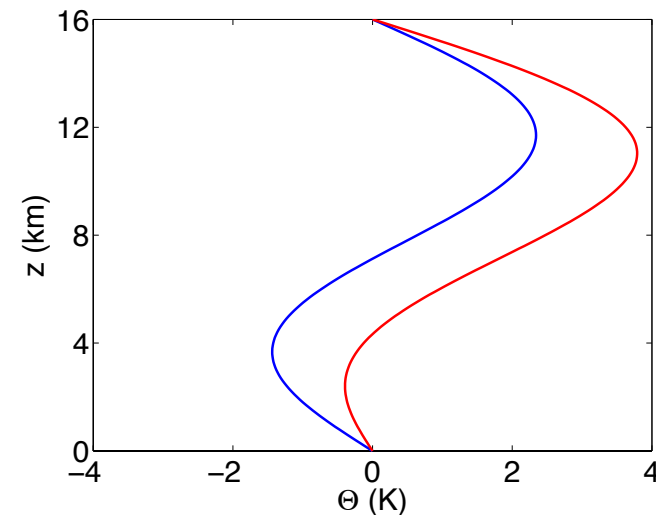
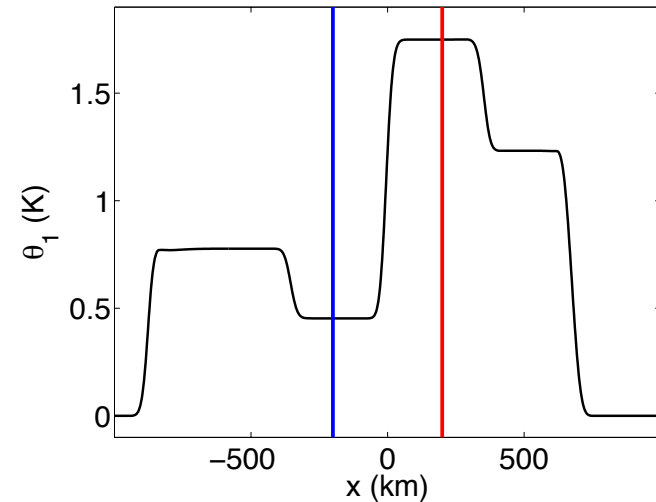
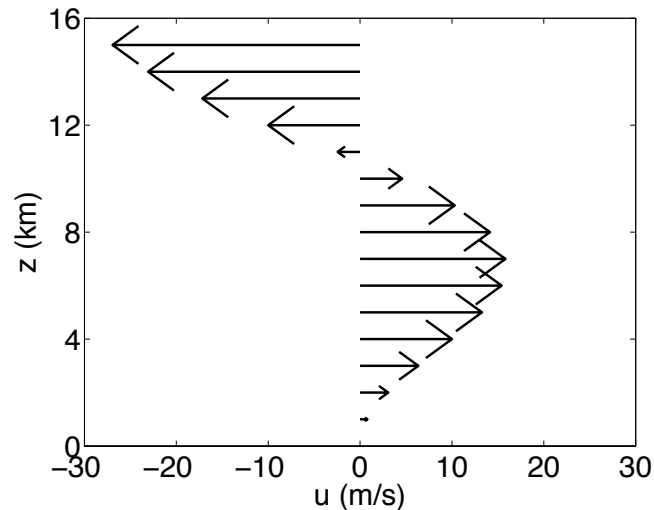


Potential temp. response



Results **symmetric** to east and west of forcing

Numerical experiment WITH wind shear



- **West** of forcing is more favorable for new convection than **east**
- Agrees with observations for this wind shear (Wu and LeMone, 1999)
- Consistent with features of CCW envelope and embedded cloud systems
 - Individual cloud systems propagate *eastward*
 - Convectively coupled wave propagates *westward*

Optimal shears for east–west asymmetry

A measure of the east–west asymmetry due to wind shear:

- the jump in θ across the source, $[\theta] = \theta^+ - \theta^-$

Which shear profiles $\bar{U}(z)$ maximize $[\theta_1]$?

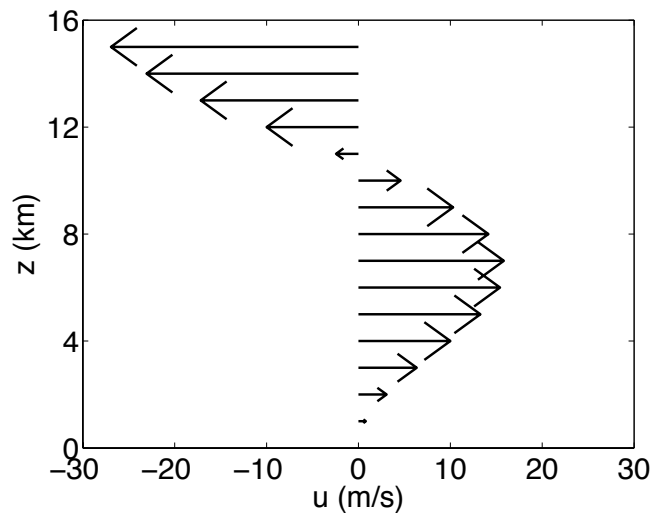
Which shear profiles $\bar{U}(z)$ lead to $[\theta_1] = 0$

Use linear theory with singular source term:

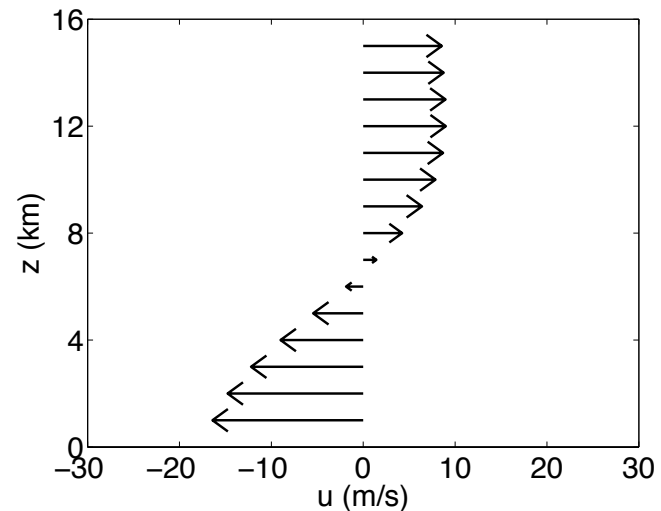
$$\partial_t \mathbf{u} + A(\bar{\mathbf{u}}) \partial_x \mathbf{u} = \mathbf{S}^* \delta(x)$$

Results:

Jet shears maximize $[\theta_1]$

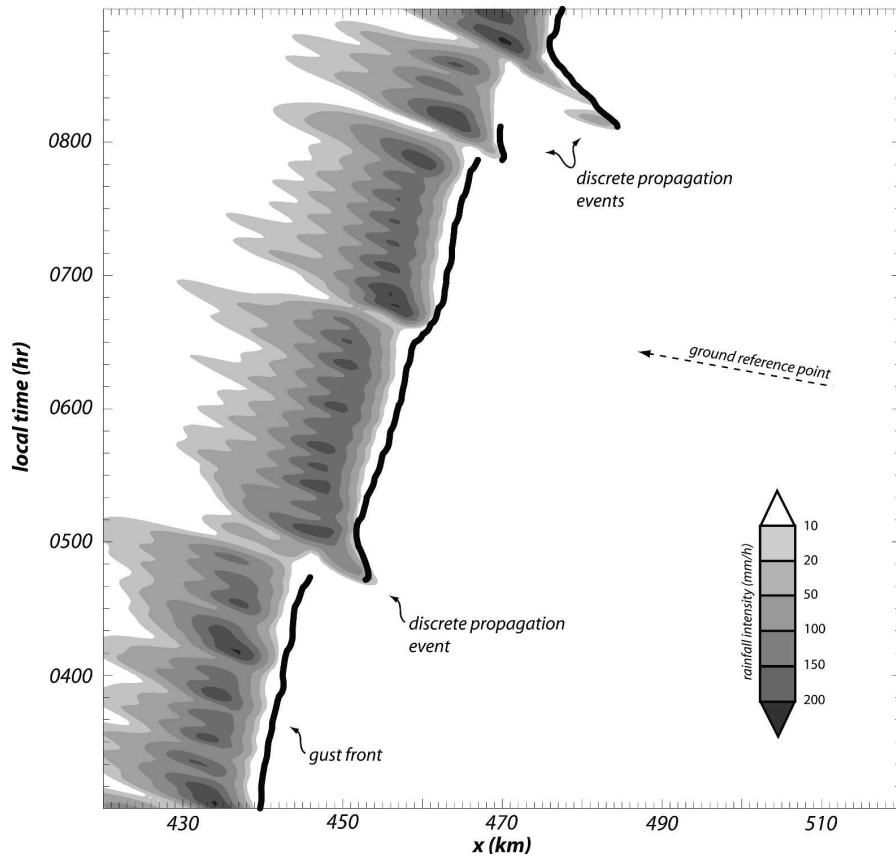


Profiles with zero shear at upper levels lead to $[\theta_1] = 0$

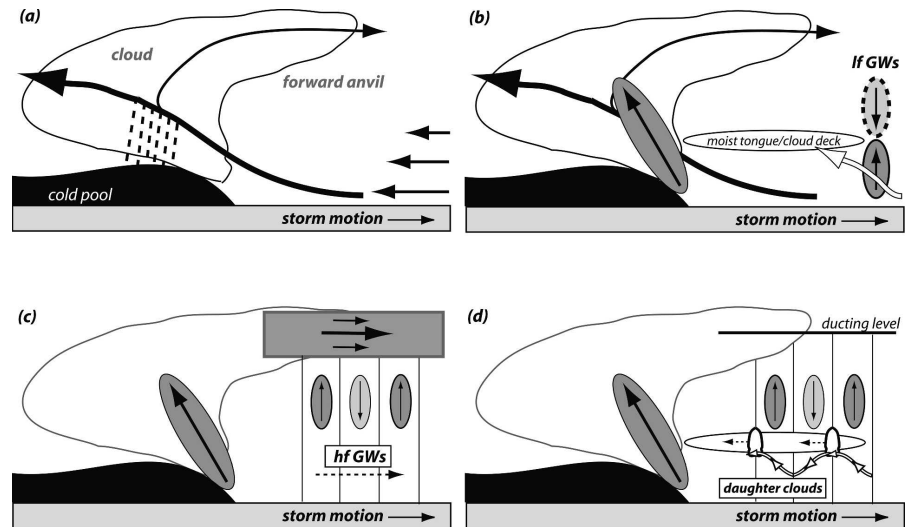


MESOSCALE CONVECTIVE SYSTEMS

Gravity waves can excite new convective cells

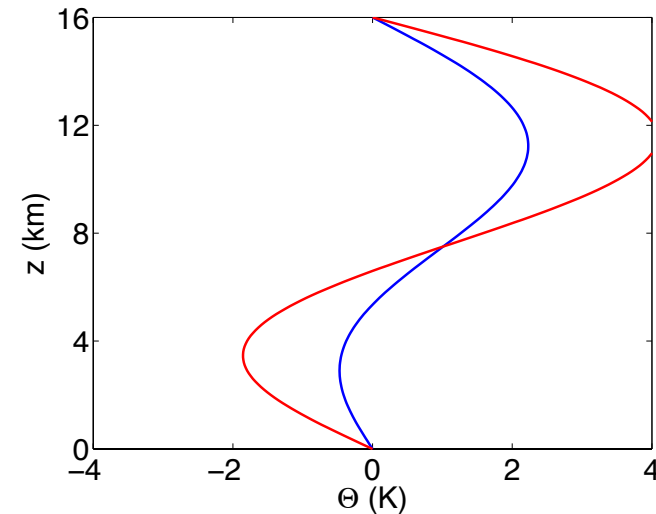
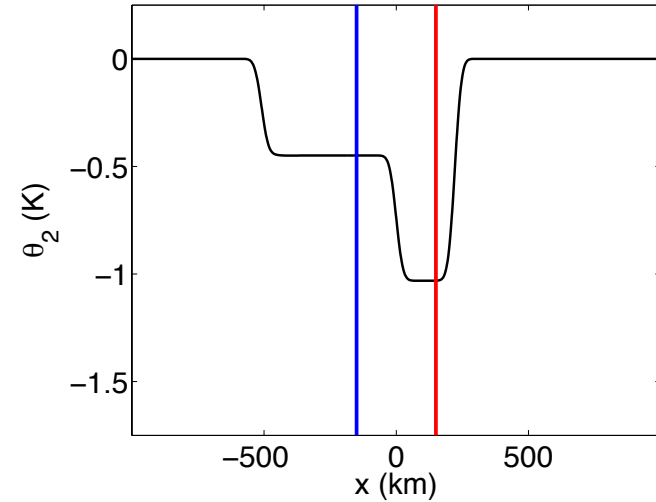
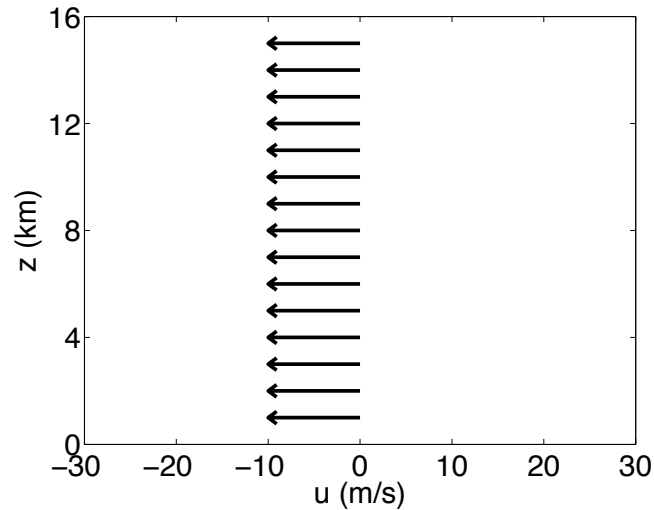


from Fovell et al. (2006)



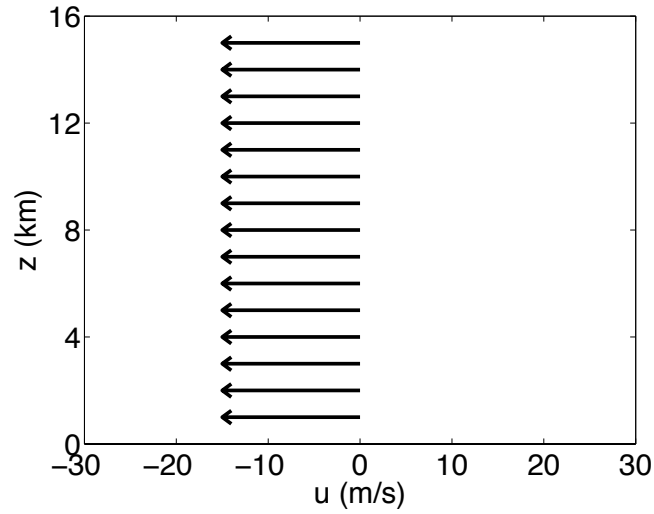
- New cells initiated *ahead of* existing squall line due to gravity waves
- New cells merge with existing squall line
- *What are the physical mechanisms involved?*

Numerical experiment with headwind

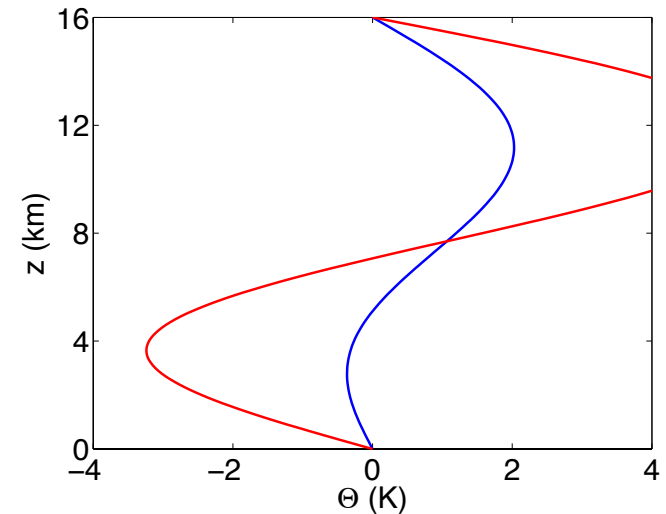
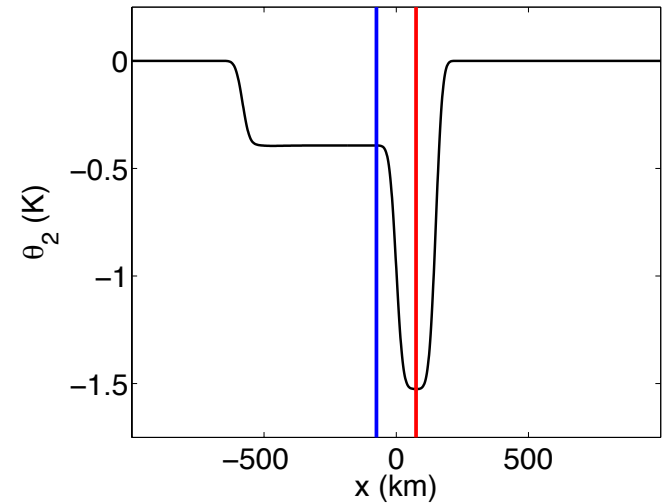


- Repeat earlier jet-shear experiment with headwind added
- Headwind is equivalent to propagating source (i.e., propagating squall line)
- Headwind confines upstream wave to vicinity of source
- **East** is more favorable for new convection than **west** at low levels

Numerical experiment with stronger headwind



Source propagation speed = 15 m/s
2nd baroclinic wave speed = 25 m/s
 \Rightarrow near resonant forcing



- Faster propagation leads to more favorable environment *at low levels*
- If squall line propagation speed \approx gravity wave speed, then *wave amplitude is large due to near-resonance*

Conclusions

- 2-mode shallow water equations:
 - simplified nonlinear model for waves interacting with wind shear
- Predictions of preferred propagation direction of convectively coupled waves in a background wind shear
 - wind shear can lead to east–west asymmetries in favorability for new convection
 - jet shears lead to largest east–west asymmetries
 - linear theory is accurate to within 10 % (usually)
- Initiation of new convective cells ahead of individual convective system
 - Propagation of source leads to *near-resonant forcing and amplification* of upstream waves

Stechmann and Majda (2009), in *J. Atmos. Sci.*

Stechmann et al. (2008), in *Theoretical and Computational Fluid Dynamics*

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